

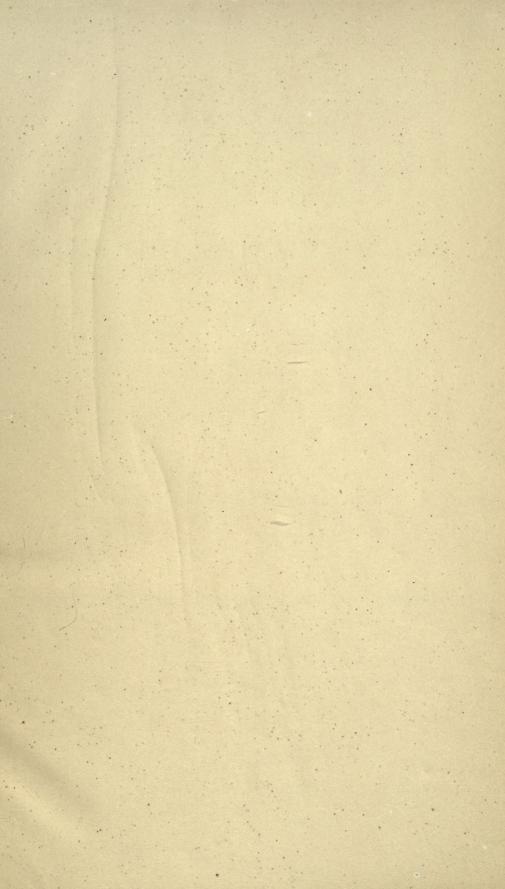




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IRON ORES

OF

MISSOURI AND MICHIGAN

BY

RAPHAEL PUMPELLY,

T. B. BROOKS,

LATE DIRECTOR GEOLOGICAL SURVEY OF MISSOURI; STATE GEOLOGIST LAKE SUPERIOR IRON DISTRICT;

AND

ADOLF SCHMIDT,

ASSISTANT IN THE GEOLOGICAL SURVEY OF MISSOURI.

WITH

ILLUSTRATIONS IN THE TEXT

AND

TOPOGRAPHICAL AND GEOLOGICAL ATLAS

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PREFACE.

A WORD of explanation touching this work will suffice. In the reports of the State Geological Surveys of Missouri and Michigan, published in 1873, the geographical and geological distribution and chemical characteristics of the iron-ores of those States are fully discussed.

As the districts covered by these Surveys supply the ore for nearly one-third of the iron-product of the United States, and as all the ore-deposits that had been discovered up to 1873 are described in the reports, it is thought desirable to combine in one volume such portions of them as address themselves specially to those immediately interested in the great iron-industry of the country.

R. P.

PART I

GEOLOGY OF PILOT KNOB AND ITS VICINITY

BY

RAPHAEL PUMPELLY

ANALYSES OF IRON-ORES, PIG-IRONS, AND FUELS
BY

REGIS CHAUVENET AND A. A. BLAIR

IRON-ORES OF MISSOURI $$_{\rm BY}$$ ADOLF SCHMIDT, Ph.D.



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- 1a. Topographical Section of Pilot Knob District.
- 2. Geologic and Magnetic Map of Pilot Knob.
- 3. Topographical Map of Iron Mountain.
- 4. Ore-mine Map.
- 5. Geological Map of Northern Missouri.

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CHAPTER I.

NOTES ON THE GEOLOGY OF PILOT KNOB AND ITS VICINITY.

BY RAPHAEL PUMPELLY.

THE region represented on the topographical map of Pilot Knob and its vicinity (Pl. I., Atlas) consists of a group of four masses of porphyry, separated from each other by valleys of the lowest Silurian Limestone known in Iron County. The porphyry forms the entire substructure of the region.* It had been eroded into hills and valleys before the deposition of the limestones.

By reference to the map it will be seen that the limestone does not often rise above the 200-foot contour-line. But in places, as in the north-west extension of Cedar Hill, we find a heavy deposit of clay, with chert and mammellated quartz ("mineral blossom") and sandstone lying over the limestone and rising to the 350-foot line. This level (350 feet above the datum of the map) is about the normal height for these clay and chert terraces through this portion of the country; but they have generally, within the area of this map, been removed by erosion, except where they remain in place as terraces and patches on the slopes of the hills.

The porphyries are older than the Silurian, and belong to the Archæan (Azoic) formation, of which they may be the youngest member in Missouri. They are the near equivalents in point of age of the great iron-bearing rocks of Lake Superior, New Jersey, and Sweden. They are stratified on an immense scale, but owing

^{*} No granitic rocks occur within the area of the map; but a few miles to the eastward there is an extensive development of granite, apparently chiefly chloritic and syenitic, in Madison County. The few observations, bearing on the question of relative age, made by both Dr. Norwood and myself, should seem to indicate that the granitic rocks are older than the porphyries. In the north-western part of Madison County several granite hills are capped with porphyry.

The red granites may be an exception to this, supposed, rule.

The fact that the granites contain numerous dykes of hornblendic rocks, while none are known to occur in the porphyries, may go toward proving the greater age of the granite.

to the rarity of interstratified beds of other rocks, the unravelling of the internal structure of the district is a difficult problem.

On Pilot Knob the strike of the formation is S. 50° E. from the true meridian, and throughout the southern half of the map the strike may be generally assumed to vary between S. 40° E. and S. 60° E., and the dip of the strata to be to the south-west.

While all the porphyries of Iron County probably contain a greater or less percentage of free silica, this is not always visible to the naked eye, nor do all the varieties exhibit the feldspars in distinct crystals.

While it would be difficult to make an absolute classification covering all the transitional forms, we can recognize (within the area of Map, Pl. I.) three very distinct varieties characterized by external features.

- a. Porphyry with few or no Crystals.—Gray, pink, flesh-color, and brown are common colors in this variety; the rock is compact, very hard, striking fire abundantly with the steel, and breaking with a conchoidal fracture. It frequently contains grains of smoky or limpid quartz, and sometimes very isolated minute crystals of either pink or white feldspar, the latter generally striated and apparently oligoclase; the rock is frequently banded in very thin layers, and may be very massive or have a columnar or tabular structure. It weathers through pink to a dirty yellow or white.
- b. Porphyry containing Crystals of Feldspar without Grains of Quartz.—The usual colors are light and dark brown and purple, more rarely black, gray, and pink; the matrix is very compact and tough, breaks with a conchoidal fracture, strikes fire with the steel, and is often banded. Its distinguishing characteristics are the absence of quartz in grains, and the abundance of crystals, one-sixth to one-fourth inch long, of white or pink feldspar, which is generally triclinic. Although quartz is not present in the form of grains, it sometimes predominates in the alternate layers of the banded varieties.
- c. Porphyry abounding in Grains of Quartz and Crystals of Feldspar.—The matrix of this variety varies in color from purplish gray to dark purple or black, and the feldspar crystals are usually white and triclinic.

Feldspar and quartz appear to be the only primary constituents in any of the porphyries of this region. But there are several

accessory minerals which occur frequently, and are evidently secondary products. The most common among these is epidote, and a soft, greenish substance resembling steatite, apparently an alteration product of the epidote. A chlorite-like mineral occurs frequently in the first stages of weathering, both impregnated in the matrix and traversing the feldspar-crystals. Magnetite and specular iron-ore occur in minute impregnations, sometimes imparting a decided polarity to the fragments of the rock. Fluorite, of a beautiful amethystine color, is found, not unfrequently, in small cavities and seams.

The Silurian rocks of the district are probably the Third Magnesian Limestone of Swallow, with, in places, remains of the Second Sandstone capping it.

At the contact of this series with the porphyries there are very generally beds of calcareous sandstone and conglomerate.

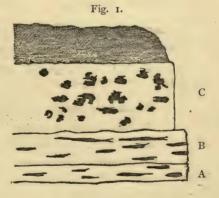
These limestones are dolomites, in which the process of dolomization has in many places run its course to completion.

The beds are generally from a few inches to six or eight feet thick. The thinner beds contain 10 per cent. to 20 per cent. of impurity, in the form of clay, and sometimes of chlorite, which gives a greenish tinge, often mistaken for a copper mineral. These thinner beds are also often pyritiferous. The more compact beds often contain layers of chert, but aside from this the insoluble impurities, in the form of clay and free silica, seem to range considerably below 10 per cent.

These heavier layers have generally a buff tinge, and contain

numerous small cavities lined with beautiful crystals of dolomite. Sometimes crystals of calcite and of copper pyrites and iron pyrites occur in these cavities.

The annexed section, and the corresponding analyses by Mr. Chauvenet, refer to the lower beds of the dolomite at Mr. Mace's quarry, Sec. 8, T. 33, R. 4, E.



A	В	С
Silicious matter 5.11	3.85	2.06
Iron, as peroxide*4.67	1.07	none
Carb. Lime47.50	52.50	54.32
Carb. Magnesia42.19	42.56	43.82
	-	
99.47	99.98	100.20

In order to compare these limestones with true dolomite, we compare the per cent. of carbonate of lime in each with that of the carbonate of magnesia, the atomic ratio being 54.35:45.65. The following figures show the amount of carbonate of magnesia required in each to make it equivalent to the carbonate of lime:—

A	В	С
39.89	 44.09	 45.62

Bed "A," therefore, contains more magnesia than the dolomite ratio calls for; the others somewhat less.

On the same property a pyritiferous limestone, from a thin bed somewhat shaly and colored green by the presence of chlorite, was examined by prospecting shafts, some years since, by persons looking for nickel-ore. This rock was analyzed by Mr. Chauvenet, with the following result:—

Silicious matter17.88
Iron, as peroxide 3.75
Carb. Lime
Carb. Magnesia34.25

The annexed section (Fig. 2) is taken at the quarry near Iron Mountain, from which the furnaces derive their flux.

A specimen from this quarry was analyzed by Dr. Litton, with the following result:—†

Residue, insoluble in dilute hydrochloric acid 6.97
Alumina, with peroxide of Iron 1.11
Carbonate of Lime50.38
Carbonate of Magnesia41.74

^{*} Alumina, with a little iron.

[†] Second Annual Report of the Geo!. Surv. of Missouri, Part II. p.[78. 1855.

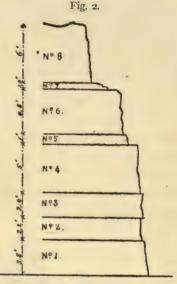
- 8. Stiff, dark, brown, foliaceous clay.
- 7. Highly-weathered dolomite.
- 6. Brown and gray dolomite, with very irregular cavities, lined with crystals of dolomite.
- 5. Very compact, hard, and fine-grained dolomite.
- 4. Similar to No. 6, with crystals of calcite porphyritically enclosed.

Greenish marly seam.

3. Similar to No. 5.

Greenish marly seam.

- 2. Similar to No. 4.
- 1. Massive, gray, and purple-gray, medium grain, with sporadic cavities.



MAGNESIAN LIMESTONE QUARRY N. W.
OF IRON MOUNTAIN,

An average sample of the rock from this quarry was taken at the Iron Mountain furnace, and analyzed by Dr. Wendel, of Troy, N. Y., for Dr. Schmidt's Report on the Iron Metallurgy of the State,* with the following result:—

Silica5.3	0
Lime27 4	8
Magnesia18.3	7
Phosphorus	03
Sulphur	

An average sample, taken in the same manner and for the same purpose, from the quarry at Pilot Knob, was analyzed by Dr. Wendel, with the following result:—

Silica	3.93
Lime	29.40
Magnesia	19.27
Phosphorus	0.02
Sulphur	0.16

^{*} To be published in the next volume.

The magnesian limestone, in some of the more massive beds, contains many cavities, from less than an inch to several feet in diameter. These have very ragged outlines, as though formed by the running together of numerous smaller cavities. They are generally lined with crystals of dolomite, but in many places, in certain horizons, and especially in the upper beds, the walls are covered with a continuous lining of crystalline quartz resting on concentric agate-like layers of quartz. This geode quartz is called "mineral blossom."

Galena and zinc-blende also occur very extensively in this rock, in some localities disseminated through it, in others occupying gashveins and narrow caves, and in others again forming extensive interstratified layers.

Residuary Deposits.—The Archæan (Azoic) rocks of southeastern Missouri are the exposed portions of the skeleton of the eastern part of the Ozark range. They appear as knobs 1,400 to 1,800 feet above the sea, and rising 300 to 700 or more feet above the valleys at their bases. They form an archipelago of islands in the Lower Silurian strata which surround them as a whole, and separate them from each other.

These rocks, consisting, as has been already remarked, chiefly of granites and felsitic porphyries, reach their most extensive surface-development in the region forming the northern part of Madison, Iron, and Reynolds, and the southern part of St. Francis and Washington Counties.

The rocks overlying them belong to the oldest known members of the Silurian, and they may be the deep-sea equivalents of the Potsdam sandstone, or even older.

As has been already stated, they are true dolomites, frequently underlaid by gritstone beds, and often covered by heavy masses (50 to 120 feet thick) of clay and chert in loose aggregations.

This region of porphyries, as well as the Ozark range, generally, through southern Missouri, has apparently been above the level of the sea from a very early period to the present time. The higher portion of the elevation does not seem to have been submerged since before the Upper Silurian period; while broad areas on the flanks of the range have apparently been dry land since the Carboniferous. The absence of the finer and coarser detrital material due to glacial action, as well as of all evidence of the direct mechanical

action of ice, prove that the region in question remained undisturbed by the various surface-modifying agencies of the Glacial period. The rocks of the Ozarks, thus exposed to the undisturbed action of atmospheric agencies, present to us in their present condition one of the most instructive records of geological history—one that is full of important facts.

Both the Archæan crystalline rocks and the Silurian strata have undergone immense changes in volume, and in other respects, under this long-continued influence. The gradual removal of the soluble constituents has left important residuary deposits of such substances as were insoluble, especially in the Silurian strata—as clay, flint, crystallized quartz, sulphuret of iron, galena, etc. The more conspicuous instances of this kind among the pre-Silurian rocks are residuary occurrences of iron-ore.

The constituents of the granitic and porphyry rocks offered a far greater resistance to the action of this process of removal than the limestone strata. Still, the amount of disintegration and of full decomposition has been very great in these older formations, although it would not be easy to say what proportion of the change has taken place since the deposition of the Lower Silurian limestones. The porphyries and granites had undergone an enormous amount of erosion before the limestones were formed; an amount at least several times as great as that they have suffered since that remote time.

In the porphyries as well as in the granitic rocks of the region we find the destructive action developed in two marked directions, respectively the resultants of influences due to local external causes, and to the structure, texture, and mineral composition of the rocks. These directions are—

(a) The forming of polygonal blocks on the surface, and the gradual disintegration and decomposition of these in place and on the talus. This is the case with most of the porphyry exposures.

A layer of 3 to 4 feet of this detritus covers the residuary clay-deposit of the Silurian limestone on the west flank of Pilot Knob; it lies on a slope of 11°.

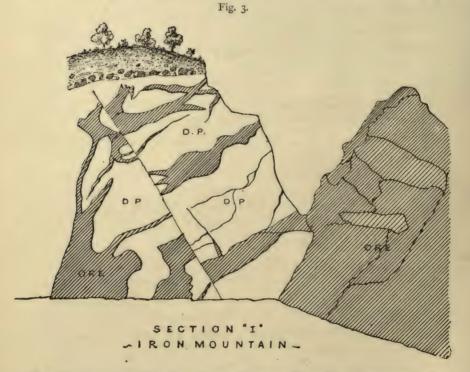
On the ridge of red granite at Ex-Governor Brown's quarry, the polygonal blocks are of great size, and their surface-disintegration in place has left picturesque, rounded masses hundreds of tons in weight.

(b) Disintegration and decomposition in mass. This, among the porphyries, is best illustrated in Iron Mountain, where the entire porphyry-hill is changed to a clay.

It is well shown, also, in Madison County, among the gray and green (chloritic) granites, where, over considerable areas, complete disintegration has taken place to a depth of certainly more than 50 feet, and possibly several times that depth.

The formation of residuary deposits of iron-ore, having their origin in the gradual removal of very resisting crystalline rocks, is one of the local results of this weathering away and decomposition of the rock, and is well illustrated in the "surface-ore" at Iron Mountain.

This hill, which rises about 250 feet above its base, is wholly covered by a mantle of ore-detritus, associated with some clay. The only knowledge we possess of its internal structure is gathered from



the mining excavations at and near the summit and on the spur called Little Iron Mountain. The appearances in these indicate deposits of the most irregular form, and which should seem, from the results of magnetic observations, to lie in zones extending north-north-east.

In all the excavations the porphyry is decomposed, generally to a clay.

At the summit an immense mass of solid ore is exposed, while the decomposed porphyry adjoining this is traversed in all directions by veins of all sizes and of the most irregular shapes, forming a reticulated network of ore and rock. The foregoing figure (Fig. 3), from a sketch by Mr. Guy, represents the exposed face of one of the great excavations. The shaded portion on the right is a part of the great ore-mass, about 50 feet high. D P is the decomposed porphyry. One of the numerous "faults" is also shown in the sketch.

Similar irregular veins and masses of ore exist in Little Iron Mountain. A remarkable dyke is exposed here, which traverses a vein of ore, and has all the appearance of resulting from the decomposition of an independent porphyry-dyke; but in its prolongation it is seen to carry in its middle plane a vein of comby quartz.

Fig. 4.

A

B

B

B

B

CITTLE IRON MOUNTAIN

A. DYKE DECOMP. PORPH.

B. DECOMP. PORPH.

C. IRON ORE

The smaller seams and veins of Iron Mountain frequently contain crystals of apatite, which seems to be a more rare occurrence in the larger ore-bodies. In the surface-ore, and in the veins near the surface, the apatite has been removed, leaving the impressions only of the crystals; and these honey-comb cavities are frequently lined with delicate quartz-prisms. This accounts for the greater freedom of the surface-ore from phosphorus, as is shown in the comparative analyses.

That the ore-seams existed before the decomposition of the en-

closing rock, is shown by the existence of similar occurrences of ore, on a smaller scale, in many of the hills of unaltered porphyry.

The surface of Iron Mountain, when discovered, was covered with a layer from four to twenty, or more, feet thick, of bowlders of pure ore, associated with ore-pebbles and ore-sand, and but little clay. The ore-detritus represents all the varieties of structure, texture, and mineral associates peculiar to the different forms of ore-masses, veins, and seams of the mountain.

As the volume of the ore-veins represents but a small percentage of the volume of the hill, the amount of the decomposed porphyry that has been wholly removed to cause such an accumulation of ore, from broken-up veins, must have been proportionately great. But the present mantle of ore-detritus represents only a portion of the concentration caused by the removal of porphyry, for the excavations at the base of the hill show heavy stratified deposits of detrital ore, having exactly the same origin, and which was washed down the slope and concentrated by the waves of the Silurian ocean.

While the present ore-mantle represents a concentrating process which has been in operation since the deposition of the Lower Silurian limestone, the bedded ore-detritus shows that the process was active before that time, and inferentially that the porphyry was even then wholly decomposed to a considerable depth.

The instance of Iron Mountain is an extreme case, where the decomposition of the porphyry in mass facilitated the separation of the ore from the rock and the mechanical removal of the latter.

There are very many points where sand and bowlders of the finest iron-ore occur on the surface, to a considerable extent. In most instances of this kind these fragments were originally isolated impregnations, segregations, or the filling of small gashes in the rock, and are the residue of a large amount of disintegrated porphyry.

In the Silurian limestone of this region, as of the Ozark range generally, the formation of residuary deposits has attained an extensive development. The long-continued wastage of strata, consisting of dolomite containing a considerable amount of insoluble substances in the form of clay, both diffused and in shaly layers, and of silica in chert-layers and nodules and quartz-geodes, has left its record in heavy masses of residuary material which cap many of

the dolomite hills, and, in the porphyry region, form terraces upon the flanks of the knobs of older rocks.

On the west flank of Pilot Knob a shaft was sunk, to a depth of about 70 feet, wholly in a residuary mass of clay containing fragments of chert, geodes of quartz, and masses of brown hematite pseudomorphous after pyrites.

On the tops of many of the flat dolomite hills there are immense quantities of mammellated crystalline quartz which originally formed the lining of the intricate geode cavities of the dolomite, and which is known as "mineral blossom," from its very general association with lead and zinc ores. There are numerous deposits of iron-ores on the flanks of the Ozarks, which owe their origin to this process of residuary concentration, and which are very uncertain as regards extent. In many places these beds of loose quartz-geodes are associated with extensive accumulations of limonite, which is, at least to a large extent, pseudomorphous after pyrite. Again, over areas of hundreds, or even thousands, of square miles, this "mineral blossom" has workable accumulations of galena and carbonate of lead for its associates in the residuary clay deposits.*

Pilot Knob.—Pilot Knob is a conical hill, nearly circular, with a north and south diameter, at the base, of about one mile. Its top is 662 feet above the datum of the map; 1,112 feet above the St. Louis directrix; 1,521 feet above tide.

On its eastern side it is connected with another group of porphyry-hills, by a neck a little more than 200 feet higher than the western base.

The rock skeleton of Pilot Knob is composed chiefly of more or less massively-bedded porphyries, porphyry conglomerates, and beds of hard, specular iron-ore.

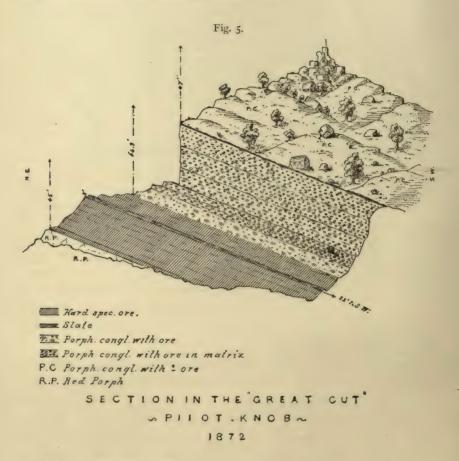
All these strata are somewhat tilted up; their strike is N. 50° W.; S. 50° E. from the true meridian, and their inclination is S. W. by S. In the eastern cut, near the summit, the inclination or dip is 21°. In the lower or westernmost cut it is 14.

By levelling between the points along the line of the greatest dip, I find the mean inclination of the ore-bed to be 13°, and shall assume this for the whole hill.

^{*} Prof. Whitney was, I believe, the first to call attention to the occurrence of residuary deposits of clay, in explaining the origin of the red clays of the upper Mississippi lead-region.

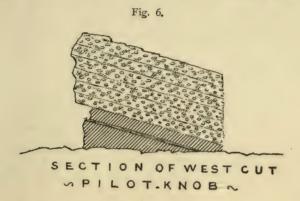
The top of the Knob consists of stratified porphyry conglomerate, with an actual thickness of 140 feet (150 feet in a vertical line).

This rock is made up of small and large, more or less angular, pebbles of porphyry cemented together by iron-ore, and containing also frequent layers and bodies of ore.



The upper portion of the series—a purple conglomerate more or less mottled with gray—has the pebbles and matrix nearly equally divided. The pebbles are rarely more than one inch in diameter, and are of brown porphyry and gray quartz, with others of a yellowish-gray substance, easily scratched, and apparently an altered porphyry. The matrix is generally a very fine-grained iron-ore, containing small grains of limpid quartz. Although the only visible

crystalline forms of the iron-ore are those of specular ore, this conglomerate matrix possesses decided polarity. A compass held close to the rock, and moved gently a few inches in any direction



parallel to the rock-surface, will have, alternately, its north and south poles strongly attracted. But the needle is not perceptibly affected at a height of four or five feet above the surface, nor does the conglomerate series, as *a whole*, affect the needle appreciably.

Toward the bottom of the conglomerate series the conglomerate structure is less marked in the eastern openings, and the lower beds have only isolated pebbles. The matrix here consists mainly of finely-divided specular ore, with a soft, greenish-white mineral, either a steatite or a clay; in this lie small grains of porphyry and of quartz. This is more properly an ore; indeed, it assumes in places a workable character, while in others it contains little ore, and becomes a nearly brown, homogeneous rock, with small grains of quartz.

Below this lies the ore-bed, with a vertical thickness of 46 feet. It is divided into two beds by a slate-seam 10 inches to 3 feet thick. This seam, which is very persistent, lies in the great cut 31 feet above the foot-wall of the lower ore-bed.

The upper ore-bed is more variable in thickness, having in the slate-seam a regular foot-wall, but having no very well defined hanging-wall, the ore often rising into the overlying rock.

The upper ore-bed is so distinctly stratified as to be a well-characterized flag-ore, and is considerably leaner than the lower bed. The slate-seam which divides the two ore-beds is variable in character, having in places the appearance of a clay-slate, in others of

talcose-slate, while it not unfrequently has all the characteristics of a porphyry.

The great ore-bed which lies below the slate-seam is a very compact, dense, and hard ore, very finely stratified in layers from I line to ½ inch thick. In places the alternate layers are somewhat silicious. This ore frequently cleaves in joint planes at right angles to the bedding-planes.

The ore-bed is underlaid by a purple-brown, indurated, clay schist, traversed by numerous threads of a steatite-like mineral.

Immediately below this lies a very compact red and brown jaspery porphyry, with minute grains of quartz, and free from feldsparcrystals. It is very hard, but contains numerous irregular-shaped patches one-eighth inch to several inches in diameter, of a soft, greenish and brownish mineral resembling steatite, which is possibly an alteration-product of the porphyry.

Below this, on the eastern slope, is a compact red jaspery porphyry, with a decided tendency to spherulitic structure. The spherules are from the size of a pin-head to that of a pea; they are of the same color as the matrix, and usually contain a small nucleus of a soft, greenish-white mineral—probably steatite. The same substance that forms the nuclei, coats also the surfaces of the spherules, and is distributed very generally through the rock in such a manner as to give it a brecciated appearance. The matrix contains very isolated minute grains of limpid quartz.

Further down the eastern declivity are outcrops of a compact, pinkish-gray porphyry, rarely containing visible grains of quartz. It contains an iron-ore, both minutely disseminated and also more concentrated in dark, irregular streaks. This rock has a tendency to tabular cleavage, and weathers through pink to yellowish white.

Underlying this we find, on the neck which connects the Knob with the hills to the eastward, a pinkish-brown porphyry with conchoidal fracture, containing grains of limpid quartz and scattered small crystals of feldspar.

A few hundred feet further east, and forming the next lower rock, is a very hard, flesh-colored porphyry, slightly mottled through the presence of irregular spots containing steatite and minute grains of iron-ore. The rock contains small grains of quartz and crystals of feldspar, which are more or less riddled with holes containing a chlorite.

The next lower rock is that which forms the western declivity of the hill next east of the Knob. It is a pinkish-gray, slightly-banded porphyry, containing innumerable small cavities filled with a chlorite, and has a well-marked columnar structure.

The inclination of the strata forming Pilot Knob being nearly the same as that of the surface of the western slope, we should expect to find the same series and thickness by boring at any point on the west slope. But there have been disturbances, which probably render this impossible, on the lower half of the western declivity.

Below the 400-foot contour-line, the rock-surface has a steeper slope than that of the hill. A shaft sunk at about the 325-foot contour, after going through about 3 feet of loose rubble of ore and porphyry, was sunk through nearly 64 feet of clay without striking rock; whereas, if there had been no disturbance, the shaft should have been wholly in the conglomerate beds which overlie the ore, and the present bottom of the shaft should be very nearly in the hanging-wall of the upper ore-bed. The absence of these rocks is due to the erosion which the south-west flank of the hill had suffered before the deposition of the Silurian rocks.

There are evidences of another disturbance. This is a zone of fracture visible in both the great cut and in the lower cut. Its course is about S. 30° W., and it is marked by a breccia of broken-up and re-cemented fragments of ore and porphyry. It does not appear to have "faulted" the ore-beds.

From the foregoing data, taken in connection with the form of Pilot Knob as determined by the topographical survey, we can delineate on the map the shape of the remaining portion of the orebed, assuming, of course, that the dip and strike remain the same, respectively, as in the upper part of the hill, and that the ore-beds are also persistent, and neither replaced by other rock material nor thrown out of position by faults.

As the western limit is probably buried beneath 75 to 90 feet of clay and chert, its delineation is only roughly given.

The south-western and northern limits are probably near the truth.

The area thus indicated is about 200,000 superficial yards as a maximum for the extent of the ore-bed.

The indication of the limit on the map may be of service in determining the best points to attack the ore-bed on the west or southern flank.

The two best points for exploration are:-

- 1. On the north-west, where the ravine intersects the limit-line, about 200 feet north of the incline and near the 375-feet contour.
- 2. On the limit east and south-east from Station 6, on the survey-line running south from the summit, there is (west of Station 6) an old opening in the conglomerate series that belongs above the ore. No line of levels was run to it; therefore, in placing its position at 20–40 feet above the ore-bed, I am making an estimate only.

The great thickness of the clay-deposit on the south-west flank would render it undesirable to attempt to approach the ore from this direction, before proving the bed near the two points above indicated.

The shaft that was started for exploration, about 500 feet southeast of the western turn-table, has its collar a little more than 60 feet above where the ore-bed should be.

Before leaving Pilot Knob, I will remark that the clay-deposit on the west flank may have a technical value. It consists of a very pure white variety and a rusty variety. The two kinds could probably be separated. Their composition, as shown by the accompanying analyses made by Mr. Chauvenet, should give, especially to the white, a very decided value as material for the fabrication of chinaware:—

ANALYSIS OF CLAYS.

I. Opalescent	White Clay.	II. Brown Clay.
Silica	63.50	57.22
Alumina	24.55	22.89
Peroxide of Iron	none	7.81
Lime	1.60	1.10
Magnesia	0.48	0.46
Water of composition	7.30	7.95
Hygroscopic Water	2.20	2.90
	99.63	100.33

Cedar Hill.—The rocks forming the south-western flank of Cedar Hill are the north-western extension of the conglomerates and orebeds of Pilot Knob.

At the time of my visit to this point, the working had not progressed far enough to indicate with certainty whether the ore after being taken out corresponds to the ore-bodies distributed irregularly in the conglomerate on Pilot Knob, or to the great ore-bed proper.

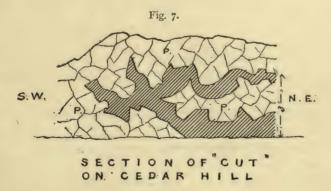
It is quite possible that the real equivalent of the Pilot Knob ore-beds has not yet been reached on Cedar Hill.

I collected a thoroughly average sample of the stock-pile at the mine, which was analyzed by Mr. Blair, with the following result:—

Insoluble silicious matter	5.62 per cent.
Peroxide of Iron	93.54 "
Sulphur	none
Phosphoric Acid	0.090 "
Equal to—	
Metallic Iron	64.47 "
Phosphorus	0.039 "
Sulphur	none

The ore opened in the cuts on Cedar Hill, at the time of my visit, is very hard, dense, and heavy, and very uniform in character. It breaks in prisms, with sharp edges, owing to a well-marked columnar structure. The ore contains minute grains of limpid quartz. It has all the appearance of a porphyry in which the matrix has been wholly replaced by iron-ore; and the irregularity of its mode of association with the porphyry heightens the resemblance.

The annexed cut, taken in one of the openings, will show what irregular forms the deposit assumes:—

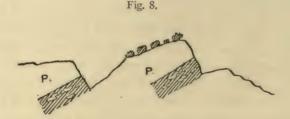


The shaded portions are ore. P is the mottled, altered porphyry associated with the ore.

In the N. W. ¼ of the S. E. ¼ of Sec. 28, east of Pilot Knob, there is another zone of iron-bearing beds. At one point, shown to me by Mr. Crane, there are beds of jaspery slate, containing more or less specular ore. The indications are such as would make it desirable to have some exploration done.

We found at another point (half-way between Station 4 on line 12 and Station 7, line 14), a bed or beds of ore, which deserve some work in the way of "proving."

It is a rich granular ore, somewhat resembling that of Shepherd Mountain. The outcrop is not sufficiently exposed to enable one to give an opinion as to the extent of the deposit. All that can be seen without digging are what seem to be two beds, one and one-half to two feet thick, of rich ore, more or less mixed with somewhat decomposed porphyry.



The deposits of manganese-ore and of manganiferous iron-ore form a very interesting feature of the porphyry region, and promise to be of considerable economic importance.

So far as I can judge, in our present limited knowledge of the porphyry district, these deposits belong considerably higher in the series than the Pilot Knob beds.

The most characteristic occurrence is that of the Cuthbertson-Buford hill, which, beginning in Sec. 19, T. 33, R. 4, E., runs W. N. W. into Sec. 13, T. 33, R. 3, E.

The northern side of this hill, near the top, seems to consist of a very dark, almost black porphyry, with numerous small crystals of white feldspar and grains of quartz.

The southern declivity, the rocks of which overlie those of the northern side, consists of a much-altered bedded rock of fine grain, which has in places the appearance of an indurated sandstone, in others that of an altered porphyry. It contains numerous broad and flat cavities partially filled with a red, ochreous substance. Fresher fragments exhibit a brown, fine-grained rock, consisting largely of minute grains of quartz cemented together by a quartz or porphyry matrix.

There are no natural exposures of the rock in place, but on the Cuthbertson tract the surface is bestrewed with large and small fragments of manganese-ore, and, in places, of specular iron-ore. A cut in the Cuthbertson tract exposes a bedded deposit consisting of exceedingly ragged tabular masses of manganese-ore, separated by a red, ochreous clay.

The following are analyses of this ore and of the specular ore, by Mr. Chauvenet:—

Manganese Ore.

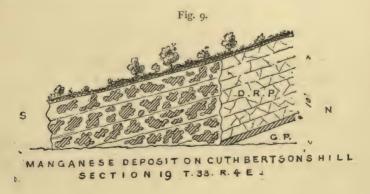
Insoluble	silicious	matter.	 			0.44
Peroxide	of Iron		 	 		3.30
Manganes	se as prot	toxide	 	 	6	58.02

The manganese exists in the ore as sesquioxide, with a probable admixture of binoxide.

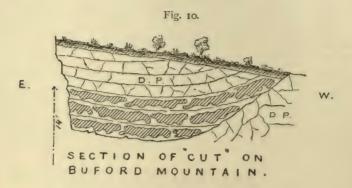
Metallic manganese52	.47
Specular Ore.	

Insoluble		 2.45	,
Peroxide of In	ron	 97.85	,
Manganese		 trace	2

It lies upon the bedded rock described last above.



Further west, at the Buford bank, a larger cut gives a better exposure.



Here we find a bedded deposit of the same nature as that on Cuthbertson's tract, with the difference that this one is a manganiferous iron-ore of a very superior quality. The cut exposes a stratified deposit of a ragged black iron-ore, overlaid by beds of pink, altered porphyry.

The whole thickness of the ore-bed was not visible, but I estimate it at nearly 14 feet in the cut, though it thins out at the outcrop on the hill-side.

A sample consisting of a large number of chippings was taken by myself as a close average of the ore. This was analyzed by Mr. Chauvenet, with the following result:—

Insoluble 8.54 pe	er cent.
Peroxide of Iron68.30	**
Manganese as protoxide15.84	66
Sulphur 0.017	"
Phosphoric Acid 0.102	"
Equal to—	
Metallic Iron	6.6
Metallic Manganese12.32	"
Sulphur 0.017	44
Phosphorus 0.044	**

This is a remarkably fine ore for the manufacture of *Spiegeleisen*. Of the extent of the deposit nothing can be known until more work is done.

Near this point there is a very irregular deposit, apparently a pocket, of ore nearly free from manganese, as appears in the annexed analysis by Mr. Chauvenet:—

Insoluble	2
Peroxide of Iron85.7	6
Manganese as protoxidetrac	е
Metallic Iron	3

Near Cuthbertson's, but apparently belonging, geologically, a little higher, is the manganese deposit on Mr. Marble's land. It forms an interstratified layer, 3-5 inches thick, in a decomposed porphyry.

A specimen of this was analyzed by Mr. Chauvenet, with the following result:—

Insoluble
Peroxide of Iron14.22
Manganese as protoxide51.06
Lime 2.75
Magnesia 0.43
Water 3.89

Mr. Marble also sank a pit in his wood-lot, near the last-named locality, and after passing the soil, found about 3 feet of the ore in ragged masses, and below these a reddish, manganiferous, soft hematite (I.), associated with another variety (II.) containing less manganese. In samples of these Mr. Chauvenet made the following determinations:—

	I.	~	11.
Insoluble	17.66		35.96
Peroxide of Iron	49.34		58.70
Manganese as protoxide	21.18		3.77
Metallic Iron	24 52		
Metallic Manganese	16.44		

This is a soft, maganiferous hematite of excellent quality.

At all of the points mentioned on this hill the strike of the beds is N. 65°—80° E., and the dip, more or less gentle, toward the south.

Overlying these manganiferous beds we find a metamorphic lime-

stone. This remarkable occurrence, which was first brought to my notice by Mr. Gage, is well exposed on Mr. Huff's land; it is a more or less thinly-stratified rock, with essentially two characteristic constituents. In one extreme we have a pink to greenish-pink, crystalline limestone, containing irregular layers, one-sixtieth of an inch to several inches thick, of a dark-brown, fine-grained material, which strikes fire readily, and shows under the glass a large percentage of quartz in minute grains, cemented by a quartz or porphyry matrix.

While the limestone bands effervesce strongly with acid, the dark bands are not acted upon except where they contain very thin layers of the limestone. The limestone is, in places, tinged green, probably from the presence of a chlorite.*

In the other extreme, the main body of the rock consists of the silicious material just described, containing very subordinate layers of the carbonate. In places, this laminated structure is highly contorted; the carbonate layers are often broken up, and the fragments distributed very irregularly in the brown silicious rock.

On the weathered surface the removal of the limestone gives rise to the same flat cavities, filled with ochrey earth, that we found in the rock accompanying the manganese-ores on Cuthbertson's tract.

On the south-west ¼ of Sec. 20, T. 33, R. 4, E., and over a considerable portion of the valley between this point and Huff's, there is a very compact, dark-brown, almost black porphyry, containing small crystals of white feldspar and grains of quartz, with some disseminated magnetite. In places the feldspar and quartz are abundantly crystallized, but more generally they are hardly visible except under the glass. The surface weathers dirty yellow, and then displays a distinctly-banded structure, which is rarely apparent on fresh surfaces. This is probably either the equivalent of the banded rock at Huff's, or it is younger.

Another interesting instance of the occurrence of manganese-ore in porphyry was examined by assistant P. N. Moore, on Section 16, T. 33, R. 2, E., in Reynolds County.

It occurs in one of the members of a series of bedded porphyry rocks, which may be here briefly described.

1st. A flesh-colored porphyry with a very compact matrix,

^{*} For an analysis of this limestone, see p. 26.

abounding in grains of smoky quartz and crystals of feldspar, is overlaid by

2d. A black porphyry with very hard matrix, abounding in grains of smoky quartz and crystals of a triclinic feldspar. In this occurs the black oxide of manganese; it is in narrow, comby strings, which are in places isolated, in others united to form a reticulated network throughout the mass; in this form the rock resembles a conglomerate, the ore representing the cement. In other instances the manganese has wholly replaced the matrix, the crystals of feldspar and grains of quartz alone remaining intact. Finally, in portions of the rock the replacement has been complete; here no traces of the porphyry, either crystals or matrix, remain, while a more or less porous, semicrystalline mass of the manganese-ore takes their place.

A specimen representing an intermediate variety of this series was analyzed by Mr. Chauvenet. It is very compact and hard, striking fire with the steel. The matrix is jet black, with metallic lustre, and has the same texture and fracture as the parent porphyry, and, like this, contains grains of quartz and crystals of feld-spar—in this instance not striated. The whole appearance of the specimen is identical with that of the adjacent porphyry, except as regards the color and lustre of the matrix.

Insoluble45.55
Peroxide of Iron 5.48
Manganese det. as protoxide37.04
Lime 2.73
Magnesia 0.81
The analysis of the insoluble portion gave—

		_	
Silica			74.98
Alumina			14.69
Iron			none
Lime and Magn	lesia		traces
Potash			9.64
Soda			0.67
			Separate Sep
			99.98

The insoluble portion is very clearly a typical felsitic porphyry with orthoclase for its feldspar. It is evident that the manganese

and the other soluble constituents take the place of more than half of the porphyry in this specimen. That the resemblance of the ore to the porphyry is not merely accidental, is shown by other instances of a similar nature.

On the land of Mr. Ackhurst (Sec. 19, T. 33, R. 4, E.) is a dark porphyry containing grains of quartz and abundant crystals of white feldspar; the matrix is frequently, in places, a compact manganese-ore.

It would seem that we have, in these occurrences, instances of replacement; but it is difficult to imagine a direct substitution of manganese oxides for the decomposition products of a porphyry, and all the more so, in this case, from the fact that the analysis shows the remaining porphyry, which is intimately associated with the ore, to have its normal constitution.

In this connection the metamorphic limestone at Huff's (near the manganiferous ores on Cuthbertson's, Marble's, and Ackhurst's land) may be of additional interest. As was mentioned before, this is nearly wholly changed into a porphyry or jasper rock, it having here a schistoid structure, in which the alternate laminæ are an impure, compact carbonate of lime, having, according to Mr. Chauvenet's analysis, the following constitution:—

Insoluble35.81
Peroxide of Iron 5.35
Lime31.62
Magnesia
Carbonic Acid25.83
99.71

Here is a member of the porphyry series which was originally, unquestionably, a limestone, but in which the original physical and chemical characteristics have almost wholly disappeared. It should not seem impossible that the manganiferous rocks which have been described may have had a similar origin, and that the manganese and iron oxides owe their present existence to a former replacement of the lime-carbonate by iron and manganese salts. The porphyry, which now surrounds these ores, may be due to a previous contemporaneous or subsequent replacement of the lime-carbonate by silica and silicates.

But there is so strong a resemblance between the mode of occurrence of these ores and that of certain of the iron-ores—that of Cedar Hill, for instance—that any hypothesis explaining the one must probably satisfy the requirements of the other.

3d. Geologically above this manganiferous bed lies a porphyry conglomerate or breccia, consisting of pebbles of a red, compact porphyry, containing grains of quartz and crystals of feldspar, cemented by porphyry of a similar character. This rock resembles very strongly the Calumet conglomerate on Lake Superior. Overlying this is

4th. A bright, red, compact jasper, apparently an altered sandstone which passes upward into

5th. A dark-brown, compact, banded porphyry, abounding in minute crystals of feldspar and equally small grains of quartz. This rock contains here and there layers of very small pebbles, parallel to the bands. Still higher this rock becomes darker colored, while some of the bands become green, from the presence of epidote. Higher in the series occurs a mottled-red and greenish-white rock. The red portions are easily scratched, and the white are still softer, having about the hardness of limestone. Grains of quartz occur indifferently through the red and white spots, while crystals of triclinic feldspar are frequent in the red spots and rare in the white. At first sight, the rock has the appearance of a variegated marble.

The greenish-white substance was analyzed by Mr. Chauvenet, with the following result:—

Silica 65.61
Alumina 20.52
Protoxide of Iron
Lime 1.97
Magnesia 2.37
Potash (trace of Soda) 7.93
100.39

This rock is clearly an altered porphyry, and the white portion represents the more advanced change. It is interesting as an instance in which the change appears to have been accompanied by a removal of silica and of iron, while the alkaline constituents were apparently not affected.

Next above this comes

6th. A porphyry with brown matrix, containing crystals of triclinic feldspar, with few or no grains of quartz, but with numerous ½ to ¾ inch spherical, crescent-shaped and irregular cavities filled with quartz, often showing free crystallization in the centre.

REPORT OF MR. REGIS CHAUVENET.

PROF. RAPHAEL PUMPELLY,

Director Missouri Geological Survey:

DEAR SIR:—I herewith submit the results of such of this year's chemical work as is fit for tabulation, together with a brief description of the methods of analysis used, and a short discussion of some of the more important coals, in regard to their prominent chemical and physical characteristics.

Respectfully yours,

REGIS CHAUVENET,

Chemist of the Geological Survey.

St. Louis, April 23, 1873.



CHAPTER II.

ANALYSES OF FUELS, IRON-ORES, AND PIG-IRONS.

BY REGIS CHAUVENET AND A. A. BLAIR.

THE analyses of coals given in these tables are what are commonly known as "proximate" analyses. The four constituents, viz., Water, Volatile Matter, Fixed Carbon, and Ash, can be subdivided into all the elementary bodies contained in the coal, but in most instances no more elaborate analysis was undertaken. We may class these four into combustible and incombustible material. Water and ash coming under the latter head, their sum will represent the percentage of weight of the coal unavailable for heating purposes.

The method followed in these examinations was one which has received the approval of the best experimenters. A weighed quantity of the coal, reduced to powder, was placed in a platinum crucible, and kept at a constant temperature of 110° C. until it ceased to lose weight. An hour was always sufficient time for this operation. The loss indicated water. The crucible being now closed, though not tightly, it was heated in a Bunsen gas-burner until the flame of the escaping gas was no longer seen. The full heat of a gas blast-lamp was then directed upon it for three minutes. After cooling, the loss of weight gave the volatile matter. In estimating ash, a fresh portion of coal was sometimes taken, but more frequently the coke left from the last operation was ignited in the same crucible, with free access of air. The loss of weight by this ignition gave fixed carbon, and by deducting the weight of the empty crucible from the last weight, the amount of ash was obtained. The color of the ash was noted, as a rough indication of its composition: the white is usually calcite; reddish-browns are due to iron; slate and chocolate indicate shale.

A few coals were analyzed by the combustion furnace (organic analysis) for total carbon and hydrogen, as well as nitrogen in a

few instances. In the table giving the results, the last column is headed "Deficiency." The sum of the total carbon, hydrogen in the volatile matter, water, and ash, will always fall short of 100 per cent. by a figure varying greatly in different coals. This is called the "deficiency," since it cannot be all ascribed to oxygen, though probably oxygen and sulphur (and nitrogen in those cases where it is not given) would fully supply the "deficiency." The low deficiency of the coals used at the St. Louis Gas-works is worthy of notice. The only Missouri coals besides these which seem to promise well are, unfortunately, found in very small seams. Westlake's coal, in Pettis County, has been used for gas at Sedalia. As coking gas-coals, this coal, Linn's (Chariton County), George's (Cass County), and possibly Munn's (Henry County), promise tolerably well. But no coal in the whole number of those examined is equal to that now in use at the St. Louis Gas-works.

In using the combustion furnace, bi-chromate of lead was substituted for oxide of copper, and with good results. The few nitrogen determinations were made by mixing the sample with soda-lime in the tube, and collection of the ammonia in hydrochloric acid, which was then treated in the usual manner, the result always being calculated from the metallic platinum obtained from the double chloride of platinum and ammonium.

Sulphur was not determined in many coals. In fact, in order to get a correct idea of the amount of sulphur, especially when it is in the form of pyrites, large samples are necessary, representing various parts of the bed, and these should be broken up and thoroughly mixed. Such elaborate sampling was not possible in the case of these coals. After a few determinations of sulphur, the system was followed of examining, for sulphur, only those coals in which no pyrites, or very trivial quantities, were visible to the naked eye. The impression being very common that sulphur in coal exists only in the form of iron pyrites (Fe S₂), several experiments were made with a view to testing the accuracy of this idea. A table will be found illustrating these experiments, and the single example of the "Baker" coal will show the fallacy of the belief. Mr. John W. Meier, of St. Louis, who has made trials of various coals, states that there are other coals similar to the "Baker" in containing sulphur and no iron. The question is one of some interest for coke-manufacturers. While a large proportion of the

pyrites in coal-slack can be removed by washing, it is not probable that sulphur not united with iron could be thus eliminated.

The method adopted for the determination of sulphur was as follows: One gramme of the coal, well mixed and finely powdered, was mixed with twelve grammes of carbonate of soda, and from four to six of nitre. (With bituminous coal, four is enough.) mixture is made in a capacious platinum crucible, and heafed in the flame of a Bunsen burner until in a state of tranquil and complete fusion. After cooling, the mass is dissolved in water, the solution acidulated with muriatic acid, and evaporated to perfect dryness, to render any silica insoluble, re-dissolved in very dilute muriatic acid, filtered, and the filtrate treated with chloride of barium, and allowed to stand for twenty-four hours. It may here be stated that the very best brands of "C. P." carbonate of soda contain sulphur in some form; hence it is absolutely necessary to make a separate test of that re-agent, and the simplest way is probably to repeat the operation, without the coal, and subtract the result from that first obtained.

This method is also the best for sulphur in coke, a determination which is frequently of importance. The only difference is, that six or seven parts of nitre should be used instead of four, as with bituminous coals.

Specific Gravity was determined by the bottle. The sample was always left in contact with the water for not less than twelve hours, before weighing, in order to thoroughly expel the air.

In glancing at the general results obtained, a few coals seem to call for special notice. The Pacific Coal Company's mines, near Knob Noster, Johnson County, are extensive, at least as compared with most of the developments in that region, and the coal, of which large and well-averaged samples were obtained, is remarkably free from sulphur, being in this respect superior to most Illinois coals. This coal is much used on the Mo. P. R. R., and although its ash never falls below eight per cent. and sometimes exceeds ten, it is probable that the trouble this would cause in burning it in locomotives, is more than counterbalanced by its freedom from sulphur, and the consequent saving in boiler-repairs. Over this coal is a curious shale, decomposing rapidly upon exposure to the air, and consequently never left in the workings. It is very bituminous, is used to run the hoisting-engine, etc. at the mine, and when thrown

in heaps, heats and finally takes fire spontaneously. It is of small value as a fuel.

Near Warrensburgh are some excellent coals, but in thin seams, and not worked upon any extensive scale. As a class, these coals show a very low ash. South of the town, and some miles from any other opening, is Grove's coal, which, from the character of its "top," its somewhat greater thickness, and its analysis, seems to belong to a different seam from the other Warrensburgh coals, but when visited was just opened, and but little could be determined about it.

Linn's Coal (Chariton County) is quite remarkable for its low percentage of ash (1.64), being the best coal examined in this respect. Equally curious are some of the Ray County coals in their high per cent. of water, one sample from Hayson's (Swanwick) mine giving 12.55 per cent.

No true "cannel" coal was examined. Many bituminous shales are known as such in their respective neighborhoods, but their ash (25 to 40) is too high, and their use must remain very limited.

Several analyses of ores, clays, and limestones are appended to these tables. Though these are given in the text, under their proper heads, it may be convenient, for some purposes, to have all the analytical results brought together. While the tables here given do not include all the chemical work done during the past year, it is believed that no result of interest or importance has been omitted.

RAY COUNTY.

	14111	COULT	11.		
NAME OF COAL.	Water.	Volatile.	Fixed Carbon.	Ash.	COLOR OF ASH.
Smith's	.10.05	38.55	45.40	6.00	White.
Howell's	. 8.05	41.85	45.80	4.30	Do.
Oberhultz	.11.02	32.48	46.30	10.20	Gray.
Hughes	. 8.15	37.60	46.35	7.90	Light brown.
Godfrey	. 7.20	30.30	37.30	25.20	Brown.
Camden Mines	. 10.33	37.73	42.04	9.90	Gray.
Swanwick Hayson's, top	. 10.00	37.85	48.30	3.85	Light brown.
Do. middle	12.55	37.05	46.65	5.75	White.
Do. bottom.		38.50	46.70	3.60	Nearly white.
Coke made from Camden coal	3.25	4.88	83.37	8.50	*
	DETT	TC COL	ATTEST		

PETTIS COUNTY.

Newport's	3.95	33.10	46.26	16.69	Red.
Westlake's	4-47	39.19	51.73	4.61	Gray.

^{*} This coke was from "Collins's" coal.

SAINT LOUIS COUNTY.

		Fixed							
NAME OF COAL. Water,	Volatile.	Carbon.	Ash.	Color of Ash.					
Parker & Russell 9.17	38.49	43.19	9.15	Light gray,					
Do 9.55	38.28	42.99	9.18	Do,					
		. ,,							
HEN	RY COU	NTY.							
Jordan, top 3.47	42.18	45.85	8.50	Purple gray.					
Do. bottom 5.14	37.91	46.82	10.13	Do.					
Owens, near Clinton 8.30	36.95	48.65	6.10	White, faint purple.					
Do. middle 7.14	38.66	48.35	5.85	Nearly white.					
Williamson 7.76	44.77	43.32	4.15	Dark purple.					
Munn's	37.30	50.75	4.45	Red.					
H. Neff 5.89	38.01	39.97	16.13	Light chocolate.					
Ogan's	33.96	43.16	14.40	Chocolate.					
Britt's, top 2.89	28.55	50.71	17.85	Reddish.					
				Light gray.					
Do. middle	43.67	43.42	9.03						
Osage Coal Co	36.95	41.87	15.53	Do.					
Do. bottom 4.86	41.74	37.24	16.16	Light chocolate.					
T A T2 A 37	ramar c	CTTATUST							
LA FAY	EIIE (COUNTY	•						
Franke's 5.55	42.95	44.08	7.42	Light brown.					
Payne's, top 8.85	37.25	44.80	9.10	Cream.					
Do. middle	37.67	39.66	15.65	Very light brown.					
Do. bottom	34.05	40.03	18.17	Light chocolate.					
Ennis & Cundiff, top	42.61	43.42	7.02	Do.					
Do. near bottom 7.03	40.72	47.11	5.14	Chocolate,					
*Lexington Coal Co., top 5.79	36.03		10.87	Light brown,					
* De middle 2 7		47.31		Do.					
* Do. middle 8.15	34.71	47.29	9.85	Yellow brown.					
	36.28	47.80	9.56 8.68						
Do. Dottom 0.25	35.03	50.04		Very light brown.					
*Tilden Davis 8.21	37.56	46.84	7.39	Nearly white.					
*Graham's	35.19	47.46	10.85	Very light brown.					
101137	7027 00	TTATEST							
JOHN:	SON CO	UNIY.							
Mrs. Wingfield's 7.31	41.88	46.36	4.45	Pale red-brown.					
Sylvester Orr's 5.87	40.06	43.45	10.62	Dark gray.					
Bruce's 5.31	43.65	43.12	7.92	Pale gray.					
Tapscott's	36.85	33.05	26.80	Light slate.					
B. Owsley	43.07	38.37	11.16	Pinkish gray.					
†Pacific mines 4.28	40.30	47.22	8.20	Nearly white.					
	40.24	47.27	8.20	Do.					
Do. 4.29 Do. 4.85			10.00	White.					
	39.85	45.30	-	Do.					
Do. 4.60	42.50	44.55	8.35	Very pale slate.					
Zimmermann's 6.77	45.10	44.01	3.62	Pale brown.					
† Do./ bottom 7.09	42.14	47.15	4.16	Pale gray.					
‡Zoll's, top 5.39	45.89	45.56		Do.					
Do. bottom 6.32	45.38	44.98	3.32 6.10	Reddish slate.					
‡Grove's 7.80	34.90	51.20		Slate.					
‡Gillum's	42.27	46.95	3.49	Pale brown.					
‡Goudy's 5.60	44.95	44.45	5.00	Reddish slate.					
§Mineral charcoal 1.59	15.63	71.28	11.50	Reddish state.					
		TATIONAL							
LINCO	LN COL	JNTY.							
Link's, top 8.17	32.58	46.50	12.75	Light gray.					
Do. next top	32.75	46.25	12.65	Do.					
		46.33	10.05	Do.					
Do. middle 8.40 Do. below middle 8.25	35.22	47.50	9.70	Do.					
	34.55	49.00	9.70	Light brown.					
Do. bottom 7.90	33.90	49.00	9.20						

^{*} From the Lexington coal-bed. ‡ Warrensburgh coals.

[†] Near Knob Noster. § From Zoll's bank, Warrensburgh.

Name of Coal. Water	. Volatile.	Fixed As	h. Color of Ash.						
Baker's 8.50	39.50	46.45 5.5							
Meadows's, top 6.30	39.20	44.30 10.2	3						
Do. bottom	36.80	42.00 14.4							
Upson's,	36.40	49.60 8.0							
2.00	, 30.40	43.73	21811 104 510 1111						
CARROLL COUNTY.									
Jas. Goodson's 2.97	36.36	47.83 12.8	4 Light brown. [specks.						
Jos. Meddlin's 2.0 "Little Compton" 4.37	29.94	47.03 20.0							
		47.21 3.8	4 Reddish brown.						
	INE COU	JNTY.							
L. Bohn	40.33	42.09 11.5							
Haynie [Miami]	31.22 48.30	35.18 31.0 42.39 5.7							
PITTSBURGH (I									
Average sample	36.61	54.17 7.9	I Faint gray.						
BIG MU	DDY CO	AL (ILL.).							
Sample No. 1 5.86		57.48 3.5	8						
Do. 2 5.88	32.81	57.66 3.6	5						
CHAR	ITON CO	DUNTY.							
Linn's 5.82	38.01	54.53 1.6	4 Salmon.						
	REW CO								
			a Dad bussess						
Niagara Creek 8.94	34.75	45.38 10.9	3 Red brown.						
LIVIN	GSTON C	OUNTY.							
Graham's 5.38	3 42.27	44.98 7.3	7 Pale brown.						
NODA	WAY CO	OUNTY.							
J. C. Smith's 3.53		40.71 13.0	4 Very light brown.						
	ASS COUR								
George's 7.80		55.75 3.2	Rich brown.						
			5 Rich Brown,						
	CON COU		T) 1 ' 1						
"Bevier"12.00	40.75	43.50 3.7	o Pinkish gray.						
	AWAY C	OUNTY.							
Fulton Coal Co 7-4		45.85 7.8							
Nesbitt's 5.00	33.95	40.73 20.3	2 Dark purple slate.						
NEBRASKA	(STATE),	NEAR NEMAH	IA.						
Omaha Coal Mining Co 4-9	38.17	49.44 7.4	6 Red brown.						
Constitution		Cara							
SULPHUI	R AND IRO	N—COALS.	Sulphur required						
NAME OF COAL.	IRON.	SULPHUR.	Sulphur required by iron to form Fe Sz. Specific Gravity.						
Smith's, Ray Co	0.84	2.41	0.96 I.249						
Godfrey's, Ray Co	2.83	4.179	3.23 1.293						
Newport's, Pettis Co	3.99	4.406	4.56 1.347						
Zoll's, Johnson Co Pacific Mines, Johnson Co	0.40	2.916 0.759	1.20 1.243 0.56 1.350						
Link, Lincoln Co	0.70	2.036	0.80 1.255						

^{*} Locality not known. Used at St. Louis Gas-works.

Name of Coa			Iro		SULP	HUR.	by ir	ur requ on to fo Fe Sz.	rm E	pecific ravity.
Hines, Lincoln Co			5.2	I	2.:	230		5.95		
Meadows, Lincoln Co			4.4	4		910		5.07		
Baker, Lincoln Co			noi	ne		532		none		1.165
Pittsburgh (St. Louis Gas-	works)		. 0.5	6		770		0,64		
Howells, Ray Co	,					702				
Oberhultz, Ray Co				*.						1.257
Hughes, Ray Co				•		509				1.277
Mrs. Wingfield's, Johnson	Co		• • • • •			170				1.328
Westlahele Pottis Co	00			•		504				1.252
Westlake's, Pettis Co	• • • • • •			•	2.0	670				1.319
	PECIFI	C GRA	AVITY	OF A	FEW C	COALS.				
NAME.		Sp.	Gr.			NAMI				Sp. Gr.
Smith's, Ray Co		I . :	249	Zimn	nerman	n's, Jo	hnson	Co		.1.225
Howell's, Ray Co		I .:	257	Mrs.	Wingf	field's,	Johnso	on Co.		1.252
Oberhultz, Ray Co		I . :	277	Sylv.	Orr's,	Johns	on Co.			1.377
Hughes, Ray Co		I.	328	Bruce	s, Jol	mson (Co			1.271
Godfrey, Ray Co		I . 2	203	Taps	cott's,	Tohnso	n Co.			1.520
Newport, Pettis Co		I.	347	Pacif	c Mine	es. Joh	nson (1.350
Westlake, Pettis Co		I.	310	Georg	ge's, C	ass Co				. 1. 261
Groves, Johnson Co		T	312	Link	Linco	In Co.				1.255
Goudy's, Johnson Co		T. 4	228	Bake	r, Line	oln Co				1.165
Zoll's, Johnson Co		T 4	242	Mine	ral cha	rcoal				T 802
2011 5, Johnson Co			43	2421110	tur Circ	icom.		••••	• • • • •	.1.003
	ULT	IMATE	ANAL	YSES	OF CO.	AL.				
	ď			.g	g .	11				
	Fixed Carbon		Fotal Carbon		in nat-	Hydro-				
NAME OF COAL.	Jar	Combined Carbon.	ari	Hydrogen Water.	Hydrogen Volatile ma ter.	Ly	'n.			Deficiency
Ziiiiiiiiiiiiiiiiiiiiiiiiiiiiiiiiiiiii	d d	bin rb	01	rog Zat	rogatil		Nitrogen.	er.		Cie.
	xe	Ca	ta	yd	lyd Vol: ter.	Total gen.	itro	Water,	Ash.	efic
	E	Ö	Ĕ	H	HA	To	Z	*	A	A
D 10 151 71 0									0	
Pacific Mines, Johnson Co.	47.22	23.11	70.33	0.47	4.66	5.13	1.45	-		11.08
Mrs. Wingfield's, "	46.38	25.98	72.34	0.81	4.81	5.62		7.31	4.45	11.09
Tapscott's shale, known	1									
ac 66 commul 22										
as cannel		24.96	58.01	0.37	4.99	5.36		3.30	26.80	6.90
as "cannel" Zoll's, Johnson Co., (War-	33.05	24.96	58.01	0.37	4.99	5.36		3.30	26.80	6.90
Zoll's, Johnson Co., (War-	33.05					5.36				6.90
Zoll's, Johnson Co., (Warrensburgh)	33.05				4·99 5·47			3·30 5·39		
Zoll's, Johnson Co., (Warrensburgh)	33.°5 45.56	28.72	74.28	0.60	5.47	6.07		5.39	4.16	10.70
Zoll's, Johnson Co., (Warrensburgh). Zimmermann's, Johnson Co.	33.05 45.56 44.01	28.72	74.28 74.76	0.60	5·47 5·58	6.07	1.56	5·39 6.77	4.16	8.77
Zoll's, Johnson Co., (Warrensburgh)	33.05 45.56 44.01 46.95	28.72 30.75 29.56	74.28 74.76 76.51	0.60 0.75 0.81	5·47 5·58 4.80	6.07 6.33 5.61	1.56	5·39 6.77 7.29	4.16 4.12 3.49	8.77 6.35
Zoll's, Johnson Co., (Warrensburgh). Zimmermann's, Johnson Co Gillem's, Johnson Co Grove's, "	33.05 45.56 44.01 46.95 51.20	28.72 30.75 29.56 17.32	74.28 74.76 76.51 68.52	0.60 0.75 0.81 0.86	5.47 5.58 4.80 4.81	6.07 6.33 5.61 5.67	1.56 1.72	5·39 6.77 7.29 7.80	4.16 4.12 3.49 6.10	8.77 6.35 11.05
Zoll's, Johnson Co., (Warrensburgh). Zinmermann's, Johnson Co Gillem's, Johnson Co Grove's, " George's, Cass Co	33.05 45.56 44.01 46.95 51.20 55.75	28.72 30.75 29.56 17.32 10.50	74.28 74.76 76.51 68.52 66.25	0.60 0.75 0.81 0.86 0.86	5.47 5.58 4.80 4.81 4.35	6.07 6.33 5.61 5.67 5.21	1.72	5·39 6.77 7.29 7.80 7.80	4.16 4.12 3.49 6.10 3.25	8.77 6.35 11.05 18.35
Zoll's, Johnson Co., (Warrensburgh) Zinmermann's, Johnson Co Gillem's, Johnson Co Grove's, " George's, Cass Co Westlake's, Pettis Co	33.05 45.56 44.01 46.95 51.20 55.75	28.72 30.75 29.56 17.32 10.50	74.28 74.76 76.51 68.52 66.25	0.60 0.75 0.81 0.86 0.86	5.47 5.58 4.80 4.81 4.35	6.07 6.33 5.61 5.67		5·39 6.77 7.29 7.80	4.16 4.12 3.49 6.10 3.25	8.77 6.35 11.05
Zoll's, Johnson Co., (Warrensburgh). Zinmermann's, Johnson Co Gillem's, Johnson Co Grove's, " George's, Cass Co Westlake's, Pettis Co Pittsburgh, used at St.	33.05 45.56 44.01 46.95 51.20 55.75 51.73	28.72 30.75 29.56 17.32 10.50 17.50	74.28 74.76 76.51 68.52 66.25 69.23	0.60 0.75 0.81 0.86 0.86 0.50	5.47 5.58 4.80 4.81 4.35 4.99	6.07 6.33 5.61 5.67 5.21 5.49	1.72	5·39 6.77 7·29 7.80 7.80 4·47	4.16 4.12 3.49 6.10 3.25 4.61	8.77 6.35 11.05 18.35 15.19
Zoll's, Johnson Co., (Warrensburgh). Zimmermann's, Johnson Co Gillem's, Johnson Co Grove's, " George's, Cass Co. Westlake's, Pettis Co Pittsburgh, used at St. Louis Gas-works	33.05 45.56 44.01 46.95 51.20 55.75 51.73	28.72 30.75 29.56 17.32 10.50 17.50	74.28 74.76 76.51 68.52 66.25 69.23	0.60 0.75 0.81 0.86 0.86 0.50	5.47 5.58 4.80 4.81 4.35 4.99	6.07 6.33 5.61 5.67 5.21	1.72	5·39 6.77 7.29 7.80 7.80	4.16 4.12 3.49 6.10 3.25	8.77 6.35 11.05 18.35 15.19
Zoll's, Johnson Co., (Warrensburgh)	33.05 45.56 44.01 46.95 51.20 55.75 51.73 54.17	28.72 30.75 29.56 17.32 10.50 17.50	74.28 74.76 76.51 68.52 66.25 69.23	0.60 0.75 0.81 0.86 0.86 0.50	5.47 5.58 4.80 4.81 4.35 4.99	6.07 6.33 5.61 5.67 5.21 5.49	1.72	5·39 6.77 7·29 7.80 7.80 4·47	4.16 4.12 3.49 6.10 3.25 4.61	8.77 6.35 11.05 18.35 15.19
Zoll's, Johnson Co., (Warrensburgh). Zimmermann's, Johnson Co Gillem's, Johnson Co Grove's, " George's, Cass Co. Westlake's, Pettis Co Pittsburgh, used at St. Louis Gas-works	33.05 45.56 44.01 46.95 51.20 55.75 51.73 54.17	28.72 30.75 29.56 17.32 10.50 17.50 23.45	74.28 74.76 76.51 68.52 66.25 69.23	0.60 0.75 0.81 0.86 0.86 0.50	5.47 5.58 4.80 4.81 4.35 4.99 5.31	6.07 6.33 5.61 5.67 5.21 5.49	1.72 1.51 1.84	5.39 6.77 7.29 7.80 7.80 4.47	4.16 4.12 3.49 6.10 3.25 4.61 7.91	8.77 6.35 11.05 18.35 15.19

HYDRAULIC LIMESTONE, LINCOLN COUNTY, OVER MEADOWS'S COAL.

Silica	21.35
Peroxide of iron	1.79
Lime 4	12.16
Magnesia	0.66
Carbonic acid 3	34.14

100.10

CLAYS FROM LINCOLN COUNTY.

		ler Coal in er's Shaft.	Morris's Shaft.	° C-IIt-
Silica	Dak	34.40	72. 35	Colbert's. 65.35
Alumina	. 4	18.62	18.11	23.25
Lime		15.27	1.00	0.52
Magnesia	• • • •	6.25	1.48	1.27
Hygroscopic water.		0.25	1.46	2.14
Water of composition		23.00	3.05	4.83
The state of the s		23.09	3.03	4.03
LINCOLN COLL	NTV IPON-()	DEC PAD	TIAL ANALYSES.	
Morris's Shaft. 1.	Morris's Shaft. 2.	Morris's Shaft. 3.	Morris's Shaft. 4. Humphi	rey's. Murphy's.
Insoluble silicious	26.98	15.42	7.74 11.0	
Peroxide of iron 4.72	63.12	79.64	84.30 86.	
(Metallic iron) 3.30	44.18	55.75	59.01 60.	
(4.200.0000 0.000)	440	33.73	39.01	39 00.72
A ROWLDER OF	RED HEMA	TITE ERON	T. 49, R. 1, E.	
		IIIE PROD	1 1. 49, 10. 1, 12.	
Section 12, Lincoln County, go	we-			
Insoluble silicious				7.55
Peroxide of iron				
(Metallic iron)				
Sulphur				
Phosphoric acid				
(Phosphorus)				
Cı	AYS FROM I	PILOT KNO	B.	
			Brown.	· White.
Hygroscopic water			2.90	2.20
Water of composition			7.95	7.30
Silica			57.22	63.50
Alumina				24.55
Peroxide of iron			7.81	none
Lime			I.IO	1.60
Magnesia			0.46	0.48
			(80,000,000,000,000)	to requirement out the
			100.33	99.63
		_		
LIMEST	ONE OF ST.	JOSEPH B	RIDGE.	
Silica				4.25
Peroxide of iron				
Lime			**********	30.55
Magnesia				19.09
Carbonic acid				44.61
				99.50
Per cent. of magnesia less than in	dolomite.			
Lim	ESTONES.—I	RON COUN'	TY.	
	A	В	С	Pyritiferous.
Insoluble silicious		3.8		17.88
Peroxide of iron		1.0		3.75
Carbonate of lime		52.50		43.52
0 1 -0 1		0 0	-	100

If the percentage of carbonate of lime in these analyses is assumed as one equivalent, and the corresponding amount of carbonate of magnesia calculated according to the formulæ for dolomite, CaO, CO₂, MgO, CO₂, we have—

	A	В	С	Pyritiferous.
MgO, CO ₂	39.89	44.09	45.62	36.55

^{*} With some peroxide of iron.

Carbonate of magnesia..... 42.19

43.82

34.25

42.56

 $[\]dagger$ Total loss by strong ignition, including a little carbonic acid. \ddag Iron and alumina.

by which it appears that A has more magnesia (in proportion to its lime) than dolomite—the others less.

Insoluble silicious. 35.81 Peroxide of iron. 5.35 Line. 31.62 Magnesia. 1.10 Carbonic acid. 25.83
Lime
Magnesia
25.83
·
99.71
Manganiferous Iron-Ores.
Buford Buford Mt. \(\triangle 239 \) \(\triangle 240 \) Marble's Mountain. Upper cut. Marble's. Marble's. Field.
Insoluble silicious matter 8.54 13.42 17.66 35.96 10.35
Peroxide of iron
7. 3
Lime 2.75 Magnesia 0.43
Sulphur 0.017 Water 3.89
Phosphoric acid 0.102
Metallic iron 47.81 60.03 34.54 41.09 9.95
Metallic manganese 12.22 16.34 2.91 39.38
Cuthbertson's Cuthbertson's
Manganese-ore. Iron-ore.
Insoluble silicious. 0.44 2.45 Peroxide of iron. 3.30 97.85
Peroxide of iron
Metallic manganese
"Shut-in" Ores.
Insoluble silicious matter \triangle 27. \triangle 29. \triangle 20. \triangle 29. \triangle 29
Insoluble silicious matter 1.75 2.10 Peroxide of iron 93.90 98.50
Sulphur
Phosphoric acid

ANALYTICAL LABORATORY OF CHAUVENET & BLAIR, 218 Pine St., Saint Louis, April 21, 1873.

PROF. R. PUMPELLY,

Director Missouri Geological Survey:

DEAR SIR:—At your suggestion I hereby submit, in a tabulated form, for convenience of reference, the analyses of iron-ores and pig-irons made by me for the survey, from samples taken by Dr. Adolf Schmidt. These analyses were made for the immediate use of Dr. Schmidt, and are incorporated in his exhaustive report. There are also a number of analyses made for private parties and published by their permission. In determining the amounts of sulphur and phosphorus in the ores and pig-irons, 5 grammes of the material was always used; consequently "a trace" of sulphur or phosphorus means a trace in 5 grammes.

It had been my intention to devote some space to the chemical

action of some of the iron-ores in the blast-furnace, and the nature of this action especially in relation to silicon in the resulting pigmetal; but the impossibility of getting perfectly reliable data in many cases, and the objections of iron-masters, in others, to make public the results of their experiments, oblige me to defer it to some future time.

Yours, very respectfully,

Andrew A. Blair.

		Iron	MOUNT	AIN O	RES.				
	20.	20.	26.	20.	27.	127.	31.	5.	128.
Insol. silicious matter.	4.71%		6.76%					1.88%	
Peroxide of iron			86.75		96.78%		95.15%	95.04	
Protoxide of iron	2.34		5 72					2.57	
Alumina	0.93							0.75	
Lime	0.45							0.15	
Magnesia	0.19							0.12	
Manganese	none							none	
Sulphur	none	none	none	none	0.016	0.008	0.003	0.005	0.012
Phosphoric acid	0.252	0.248	0.081	0.073	0119	0.112	0.125	0.071	0.067
							_	0.4	
	100.322						1	100.586	
	100.322						1	100.586	
Insol, silicious matter.					0.4		-		
Insol, silicious matter. Silica	3.99%				3.28%		4.54%	1.57%	••••
Insol, silicious matter. Silica	3.99%				3.28%		-	1.57%	••••
Insol, silicious matter, Silica	3.99%						4.54%	1.57%	
Insol, silicious matter. Silica	3.99%						4.54%	1.57%	
Insol, silicious matter, Silica	3.99% 0.47 0.06 0.03						4.54%	1.57% 0.04 0.02 0.017	
Insol, silicious matter, Silica	3.99% 0.47 0.06						4.54%	1.57% 0.04 0.02	
Insol, silicious matter. Silica, Alumina Lime. Magnesia	3.99% 0.47 0.06 0.03	• • • •	• • • •	• • • •	* * * * *		4.54%	1.57% 0.04 0.02 0.017 1.65	
Insol, silicious matter, Silica	3.99% 0.47 0.06 0.03	• • • •		• • • • •	67.75%		4.54%	1.57% 0.04 0.02 0.017 1.65	

²a, 2b.—Specular ore from enclosed conglomerate mass in the backbone-vein, cut A, on south side of Iron Mountain.

These two samples being practically the same, the great difference in the amounts of phosphoric acid seemed unreasonable. The analyses were therefore repeated, and the results were practically as before.

27.—Specular ore from backbone-vein, cut II, north side of mountain.

127.—Average sample of quarry-ore taken by Maj. Brooks.

care, and represent thousands of chippings, in 127 from the quarry-ore, and in 128 from the loose surface-ore. After being thoroughly powdered, portions were sent by Maj. Brooks to three chemists, viz.: Prof. Allen, of Yale College, Mr. Otto Wuth, of Pittsburgh, and Mr. A. A. Blair, of St. Louis. As the determination of phosphorus and sulphur in perfectly average samples of this great ore-deposit, is a subject of considerable technical interest, the results obtained are here given :-

²a deflects the compass-needle; 2b does not, yet 2b is more magnetic and contains more magnetic oxide than 2a.

^{31.—}Average sample of surface-ore from western slope.

^{5.-} Average sample of surface-ore from south-east slope. 128.—Average sample of surface-ore taken by Maj. Brooks.

All the samples, except 127 and 128, were taken by Dr. Schmidt.

Note.—The samples Nos. 127 and 128 were collected by Maj. T. B. Brooks with great

	No. 127—Qu	arry-Ore			
	WUTH.		ALLEN.		BLAIR.
D1 1		Double	determination.		DLAIR.
Phosphorus	0.016	0.043			0.049
Supras, it is it i			• • • •		0.008
701 1	No. 128—Sur				
Phosphorus	0.016	0.025	0.026		0.029
Durphur	• • • •				O.012
	PILOT KNO	B ORES.			[]
Insoluble silicious matter	47.	45.	45 (Rep.).	36.	16.
Peroxide of iron	84.33	5·57% 90.87		67.38%	62.88%
Protoxide of iron	0.15	1.67			• • • • • •
Alumina		0.53		• • • • • •	
Lime		0.13			
Manganese		none			
Sulphur.		0.078		0.006	0.079
Phosphoric acid	0.035	0.069	0.063	0.092	0.101
	100.365	100.677			
T I. 1.1					
Insoluble silicious matter. Silica	13.27	5.18			
Alumina and trace of oxide of ir		0.36			
3.5 / 11* *					-
Metallic iron		64.91 0.031	0.027	47.16 0.041	44.01
2 100pnorus		0.031	0.027	0.041	0.044
Denovide of iron	50.	46.	48.	40.	
Peroxide of iron		83.28% 58.29	77.02% 53.91	52. 18% 36. 52	
47. Soft ore from central cut				_	Ore from
lower strata eastern cut (C).					
Specular ore from western cut					48. Ore in
conglomerate central cut. 40. A by Dr. Schmidt.	Average sample	e of better of	re on refuse	-neaps.	All taken
	EPHERD MOU	NTAIN, ETC.	•		
20.		23.	68.	C.H.	L.M.
Insoluble silicious matter 5.19		06 700	** 20d	5.62%	15.33%
Peroxide of iron 94.8. Protoxide of iron 1.8		96.70%	79.39% 14.22	93.54	84.60
Alumina					0.32
Lime	0.35				0.38
Magnesia	**				none
Copper			none		
Sulphurnon		none	none	none	0.021
Phosphoric acid o.o.	25 0.039	0.032	0.038	0.090	0.065
	100.269				100.866
Insol. silicious matter. Silica	5% 5.98%				14.45%
Peroxide of iron					none
Alumina	0.46				0.51
Lime	O. I2				- (
					0.06
Trugitosia,	0.05			• • • • •	0.06
	<u>0.05</u> <u>6.68</u>				0.04
	6.68				0.04
	<u>0.05</u> <u>6.68</u> 2 <u>64.31</u>	• • • • •			0.04

20. Average sample upper part of central vein (B).

lower " 23. Soft ore from "68. Ore from eastern vein.

The above all taken by Dr. Schmidt.

C. H. Sample of Cedar Hill ore taken by Prof. Pumpelly, and representing average of

L.M. Sample of Lewis Mountain ore taken by Prof. Wm. B. Potter, and analyzed for Pilot Knob Company, by whose permission it is published.

SPECULAR AND RED HEMATITE ORES.

	Ι,	2.	3•	4.	5.	6.
Insoluble silisions metter		٠.				-
Insoluble silicious matter	11.19%	9.30%		0		0.87%
Peroxide of iron		76.45	27.40	87.92	97.23	98.96
Protoxide of iron	0.77		4.11			
Alumina	0.97					
Lime	O. I 2		36.01			
Magnesia	0.07					
Manganese	none					
Carbonic acid	0.48		29.51			
Water	0.46	13.65	2.45	0.03	0.47	
Sulphur	0.126	0.052	0.044	none	none	trace
Phosphoric acid	0.116	0.478	0.098	0.089	0.092	0.062
*						
	100.252					
Insoluble silicious matter.						
Silica	9.78%	6.78%	0.27%	11.32%	2.06%	
Alumina	1.33					
Lime	none					
Magnesia	0.12					
Titugirosiu						
	11.23					
Metallic iron	60.76	53.51	22.38	61.54	68.08	69.27
		0.208	-			
Phosphorus	0.051	0.200	0.043	0.039	0.040	0.027

^{1.} Purple paint-ore from Meramec mine. 2. Ochrey ore from do. 3. Red ore and yellow rock, with spathic ore, from do. 4. Hard, silicious specular ore from do. 5.

SPECULAR AND RED HEMATITE ORES.

	7a.	78.	8.	9.	10.	II.
Silica	0.69%	0.68%	8.39%			
Peroxide of iron	97.94	97.08	88.37	83 51%	90.03%	78.38%
Protoxide of iron	trace	0.31				
Alumina	1.17	1.50				
Lime						
		0.08	3.09			
Combined water	0.19	0.23 }	39			
Carbonic acid		0.21				
Sulphur		none	trace			
Phosphoric acid	0.068	0.079	0.207	0.079	0.083	0.205
Insoluble silicious matter				17.05	5.73	14.13
Metallic iron	68.56	68.19	61.86	58.45	63.02	54.86
Phosphorus		0.034	0.091	0.034	0.036	0.089

⁷a, 7b. Piece of ore from Iron Ridge, partly blue specular and partly red paint-ore; 7a, blue specular, 7b, soft, red paint-ore. These samples were taken for a comparison of the

Average sample best specular ore from do.

1, 2, 3, 4, and 5, taken by Dr. Schmidt.
6. Blue specular ore from Orchard bank, Salem District, Dent County, analyzed for Mr. O. A. Zane, and published by permission. Now worked by Salem Iron Company.

two kinds of ore, and is particularly interesting on that account. 8. Average sample from Iron Ridge mine.

The above taken by Dr. Schmidt.

9. Blue specular ore from Beaver Branch.

10. Partly decomposed ore from same.

11. Soft paint-ore from same. These Beaver Branch ores were analyzed for Missouri Furnace Company, and are published by permission.

BROWN HEMATITE ORES.

	ı.	2.	3.	4.	5.	6.	7.	8.	9. 1
Insol. silic. matter						8.66%	7.17%	7.42%	8.35%
Peroxide of iron	84.02%	77.42%	82.02%				78.73		79-53
Water			12.80						15.33
Sulphur							0.009		none
Phosphoric acid									
Manganese									
								• • • • •	
Silica	2.08	8 05	F 10		2 50				
Metallic iron			57.41				55.11		
Phosphorus	0.370	0.034	0.041	0.037	0.034	0.058	0.081	0.001	0.071
T1 -:1:	10.	II.	12.	13.	14.	15.	16.		
Insol. silic. matter	4.88%	3.00%	0.97%	4.34%	9.41%	39.22%			
Peroxide of iron							75.42		
Water						8.87			
Sulphur					trace	0.001	0.035		
Phosphoric acid				2.891	0.249	0.206	0.859		
Metallic iron	57.50	50.55	56.68	54.86	56.24	34.38	52.70		
Phosphorus									
	0/4	0,020	-, 123		9		3/3		

1. Brown and red hematite from Marmaduke bank. 2. Limonite from Sheldon bank, 3. Do. from White bank. 4. Do. (pipe-ore) from Elm Hollow bank. 5. Do. do. from Indian Creek bank. All on the Osage River. 6, 7, 8, 9, 10, 11. From Camden Co. Analyzed for Mr. H. S. Reed, and published by

permission.

12. From Perry Co. Analyzed for Mr. Reed, and published by permission.
13. From Perry Co., opposite Grand Tower. Analyzed for Big Muddy Iron Company,
and published by permission of Mr. Jas. E. Mills, Vice-President.

14. From Indian Ford, Bollinger Co. Analyzed for Mr. Wm. B. Spear, and published by permission.

15. From near Irondale, and used in the furnaces there and at Iron Mountain. Ana-

lyzed for Messrs. E. Harrison & Company, and published by permission.

16. Analyzed for Wm. E. Romer, Esq., of Grand Tower, and published by permission.

I, 2, 3, 4, and 5 samples were taken by Dr. Schmidt.

PIG-IRONS.

1	r.	2.	3.	4.	5.	6.	7-	8.	9.
Sulphur	0.024%	0.017%	0.005%	trace	none	trace	0.016%	0.026%	
Phosphorus	0.133	0.062	0.116	0.165%	0.098%	0.116%	0.136	0.196	
Combined carbon									*,* * *
Graphitic carbon	3.293								
Silicon	3.230	2.624	1.329	0.942	1.389	1.354	0.630	1.347	4.850
	10.	II.	12.	13.	14.	15.	16.	17.	18.
Sulphur	10.	none	0.045%	0.006%	14. 0. 103%	15. O. I 34%	16. 0.141%	0.061%	18. j
Sulphur Phosphorus,		0.108%	0.147	0.153	0.097	0.141	0.107	0.141	0.110%
Phosphorus.		0.108%	0.147	0.153	0.097	0.141	0.107	0.141	0.110%
Sulphur Phosphorus, Combined carbon Graphitic carbon		0.108%	0.147	0.153	0.097	0.141	0.107	0.141	0.110%

I. Made in the fall of 1872, at Vulcan Iron Works, from Iron Mt. ore alone; ½ Big Muddy coal, ½ Connelsville coke. 2. Made at Pilot Knob Iron Co.'s furnace, from 2 Pilot Knob and & Shepherd Mt. ores, charcoal and hot blast. 3. Made at Scotia Iron Works, from Scotia ores, charcoal, and hot blast. 4. Made at Meramec, from Meramec ores, 1 half hard and 1 soft, charcoal and cold blast. 5. Made at Iron Mountain, from Iron Mt. ores, with 8% of limonite. 6. Irondale, hot blast. 7. Irondale, cold blast; both charcoal, and same ores as 5. 8. Made at Moselle, from Iron Ridge and St. James ores, with 8% Moselle limonite, charcoal and hot blast. 9. Made at Pilot Knob Iron Co.'s Works, from all Pilot Knob ore, charcoal and hot blast. 10. Made at same, \(\frac{2}{3}\) Shepherd Mt. and \(\frac{1}{3}\) Pilot Knob, charcoal and hot blast. 11, 12, and 13. Made at Big Muddy Iron Co.'s furnace, at Grand Tower. 11 and 13 from \(\frac{1}{2}\) Pilot Knob ore, \(\frac{1}{2}\) Iron Ridge and St. James red hematites, using \(\frac{2}{3}\) Connelsville coke, \(\frac{3}{6}\) Big Muddy coal. 12. Same, except that the Illinois Patent Coke Co.'s coke was substituted (using a larger proportion) for the Connelsville coke. Coke Co.'s coke was substituted (using a larger proportion) for the Connelsville coke.

These analyses were made for Big Muddy Iron Co. and published by permission of James

E. Mills, Esq., Vice-President.

14, 15, 16, 17, and 18. Made at South St. Louis Iron Co.'s furnaces. Analyzed for them, and published by permission of Mr. H. S. Reed, President. 14. Glazed pig, made from all Pilot Knob ore. 15 and 16. No. 1 and 2 Foundry, from all Iron Mt. ore, made in summer of 1872. 17. Made from Iron Mt. with some Iron Ridge soft, red hematite. 18. Made from all Iron Mt. ore, spring of 1873. All with Connelsville coke and Big

1, 2, 3, 4, 5, 6, 7, and 8 were samples taken by Dr. Schmidt or sent to him for the use

of the survey.

CHAPTER III.

THE IRON-ORES OF MISSOURI.

BY ADOLF SCHMIDT, PH.D.

A. General Distribution.

MISSOURI is one of the richest States in iron-ores on the North American continent. These ores are, however, very unequally distributed over the State.

Very little iron-ore is found in the whole northern part of the State north of the fiftieth township-line, and in a range of counties on the western border. These districts are covered by the coalmeasures, which, although containing clay-ores and carbonates of iron, do not contain them in such quantities and in such positions as to make them workable. According to Mr. G. C. Broadhead's statements, these ores in the coal-measures of Missouri occur either as single nodules, or as thin beds, varying from one to twenty inches in thickness, imbedded in the carboniferous clays and slates. They lie, generally, deep below the surface, from 20 to 60 feet, and not close enough to the coal-beds to be mined conjointly with the latter. These ores are, besides, not very rich in themselves.

The only point where the region of workable iron-ore reaches, north of the Missouri River, is in Callaway County, where red, earthy hematite occurs as a stratum in the ferruginous sandstone of the subcarboniferous system.

South of the Missouri River there are, between this river and the fortieth township-line, valuable deposits, mostly of limonite, in Franklin, Osage, Morgan, and Benton Counties. This kind of ore also occurs nearly over the whole central and southern part of the State. In the southern part the counties of Stoddard, Bollinger, Wayne, Ozark, Douglas, Christian, and Greene, contain considerable deposits of it.

But by far the richest portion of the State in iron-ores is that between the 30th and 40th township-lines. Within this zone, iron-

ores abound in the greater part of the counties situated between the Mississippi in the east and the Upper Osage River in the west. Limonite banks are scattered over the whole of this vast region, being, however, somewhat concentrated in three districts. most eastern of these is composed of Bollinger, Wayne, and the southern part of Madison Counties; the second but smaller concentration is in the south-eastern part of Franklin County; while the third and most important one of this ore is found on the Middle Osage River, between Warsaw and Tuscumbia, in Benton, Morgan, Camden, and Miller Counties. This latter district extends also to the Upper Osage, above Warsaw, into St. Clair and Henry Counties. But while the limonites are deposited on the Second and Third Magnesian Limestones in the rest of the State, they here lie on the subcarboniferous limestone. The Upper Osage district also contains good deposits of subcarboniferous red hematites, occurring here in the same way as in Callaway County.

The specular ores are much more concentrated in certain parts of the State than either the limonites or the carboniferous hematites, and also occur in much larger masses. There are two important specular-ore districts, different by their geographical positions, different entirely by the mode of occurrence and the geological position of their ores, but quite similar, on the other hand, in the mineralogical character and the chemical composition of these The one of these districts is the Iron Mountain district in the east, extending only over a small area, in southern St. François and northern Iron Counties, but containing two enormous deposits, besides numerous smaller ones. The ore is here in veins, beds, and other less regular forms in the porphyry. The second specular-ore district lies more toward the centre of the State, yet mainly in its eastern half. Its principal deposits, as far as known at present, are concentrated in three counties, Crawford, Phelps, and Dent. The occurrence of the specular ores, however, extends somewhat into the surrounding counties of Washington, Franklin, Maries, Miller, Camden, Pulaski, and Shannon. The specular ore in this central ore-region is always more or less distinctly connected with the Lower Silurian Sandstones, especially with the so-called Second Sandstone. Many of these deposits are disturbed and broken, and altered in regard to their position and contents.

From all that has been said, we may infer that, according to our

present knowledge, there are three principal and important ironregions in Missouri, namely:—

- 1. The eastern region, composed of the south-eastern limonite district and the Iron Mountain specular-ore district. This region has its natural outlet, at present, over the Iron Mountain Railroad.
- 2. The central region, containing principally specular ores, and having its commercial outlet over the St. Louis, Salem & Little Rock and the Atlantic & Pacific Railroads.
- 3. The western or Osage region, with its limonites and red hematites. This region will have to establish an iron industry of its own, because it is too remote from the present ore-markets. Its present connection with these markets is down the Osage River to Osage City, and from there either over the Missouri Pacific Railroad or down the Missouri River. A railroad from Jefferson City, through Cole, Moniteau, Morgan, Benton, Henry, St. Clair, Bates, and Vernon Counties to Fort Scott, which would touch the Upper Osage districts, is partly in construction, partly under consideration.

These three principal regions, combined, form a broad ore-belt, running across the State from the Mississippi to the Osage, in a direction about parallel to the course of the Missouri River, from south-east to north-west, between the thirtieth and fortieth town-ship-lines. The specular ores occupy the middle portion of this belt, the limonites both ends of it. The latter are, besides, spread over the whole southern half of the State, while the subcarboniferous hematites occur only along the southern border of the North Missouri coal-field, having thus an independent distribution, and being principally represented in Callaway, St. Clair, and Henry Counties.

To make this distribution of ores more apparent and clear, I have added to this report the accompanying "Preliminary Map, showing the Distribution of Iron Ores in Missouri." Atlas, Plate IV.

This map contains about 280 deposits. Its title indicates that it does not pretend to be complete. It represents only a first effort toward a more complete map, and contains the results of information obtained during the summer of 1872. Such a map can, in fact, hardly ever be entirely complete, because new ore-banks are continually being discovered and opened.

Neither does this map pretend to show the exact relative character and size of the single ore-banks. It is only intended to show,

in a generally correct and clear manner, the general distribution of the ores. For this purpose it was necessary, however, to mark every single deposit that has come to my knowledge, and to indicate those deposits which, according to their present appearance, seem to be more extensive than others, by larger sizes. This led to the adoption of five sizes, with a respective estimate of yield of smeltable ore.

Below 20,000 tons,	for size,	I.
20,000 to 100,000,	"	2.
100,000 to 500,000,	. "	3.
500,000 to 2,000,000,	"	4.
Above 2,000,000,		5.

To distinguish the different kinds of ores, I use three colors, thus:—

Red, for red hematite. Blue, for specular ore. Brown, for limonite.

As regards the character of the deposits, those which, like veins and beds, are generally supposed to be more or less continuous, are marked by squares. All the other deposits, which are supposed to be either decidedly limited or very irregular, are marked in a circular form.

Another distinction made on the map is that between undisturbed, disturbed, and drifted deposits. This distinction was unavoidable, on account of the very frequent occurrence of disturbed deposits, especially in the central ore-region.

In the following chapters of this report the ores and deposits will be described under the following geographical arrangement:—

Eastern Ore-Region.

ORE-DISTRICT ALONG THE MISSISSIPPI RIVER. IRON MOUNTAIN DISTRICT.

SOUTH-EASTERN LIMONITE DISTRICT.

FRANKLIN COUNTY DISTRICT.

SCOTIA DISTRICT.

Central Ore-Region.

STEELVILLE DISTRICT.

ORE-DISTRICT ON THE UPPER MERAMEC AND ITS TRIBU-TARIES. SALEM DISTRICT.
IRON RIDGE DISTRICT.
St. James DISTRICT.
ROLLA DISTRICT.
MIDDLE GASCONADE DISTRICT.
LOWER GASCONADE DISTRICT.
CALLAWAY COUNTY DISTRICT.

Western Ore-Region.

LOWER OSAGE DISTRICT.
MIDDLE OSAGE DISTRICT.
UPPER OSAGE DISTRICT.

South-western Ore-Region.

WHITE RIVER DISTRICT.

OZARK COUNTY DISTRICT.

CHAPTER IV.

THE IRON-ORES OF MISSOURI.

BY ADOLF SCHMIDT, PH.D.

B. Description of Ores. - General Description.

It may be inferred, from the contents of the preceding chapter, that there are principally two species of iron-ores in Missouri, hematite and limonite. These can be easily distinguished from each other by their exterior appearance, the hematite being either grayish black, with a slight bluish or reddish tint, or red in various shades; while the limonite is always brown or yellow. But the best and least deceptive distinction is made by the streak of these ores—that is, by the mark they produce when rubbed against a rough, white porcelain-plate. The streak is invariably red with the hematite, and yellowish brown to yellow with limonite. If the ores are not too hard, the color of the streak can also be discovered, though less plainly, by scratching the ores with a knife. The hematite occurs in two very different and distinct varieties, the specular ore and the red hematite.

Specular Ore is bluish-black to steel-gray, with a more or less metallic lustre, and a more or less crystalline structure. Its mineralogical hardness is about 6, which is also about the hardness of hardened cutlery-steel. Thus an ordinary pocket-knife will not scratch the hardest specular ores; but it will scratch the softer kinds, though not without some effort, there being no great difference in their respective hardness. The streak of pure specular ore is cherry-red to dark-red, with a purple tint. Its magnetic qualities vary considerably in the same kinds of ore and in the same localities. Most specular ores are, however, slightly magnetic.

This description refers to specular ore in its pure and natural condition, being then composed chemically of nearly pure peroxide of iron, containing about 70 per cent. of metallic iron. But this ore is

sometimes found mixed with foreign substances, as quartz, sand, flint, porphyry, clay, pyrites, apatite, etc., which generally change somewhat its qualities and diminish its value. Specular ore also frequently undergoes physical and chemical changes by a gradual alteration under the influence of air, water, or mineral solutions. These alterations, which will be spoken of more fully hereafter, sometimes change the character of the ore completely. We find the specular ore to pass, under certain conditions, into compact red hematite, and into soft red hematite, by other influences into limonite, by others into carbonates. These various ores, as far as they are merely the products of such gradual alterations, will be described together with the specular ores from which they are derived. It is a very remarkable fact, that all the Missouri specular ores, with but a few exceptions, have pretty nearly the same mineralogical and chemical character, whether they occur in the porphyry or in the Silurian sandstone.

Red Hematite, when not produced by transformation of specular ore, but occurring as an original mineral in the subcarboniferous strata of Missouri, has a dark-red color, either with a yellowish or more frequently with a bluish tint. The fracture is uneven and dull in the earthy, somewhat conchoidal with a slight lustre in the compact, varieties. The structure is never crystalline, but either earthy and more or less porous, or compact and fine grained, or coarse grained to oölitic. The hardness is less than that of most specular ores. It varies from 5 to 6, but rarely reaches the latter figure. The streak is cherry-red to yellowish red. This ore is unmagnetic.

Red hematite, when exposed to atmospheric influence, seems to become more porous, and is altered gradually into brown and yellow limonite. These subcarboniferous hematites consist of a somewhat clayish peroxide of iron, and contain from 50 to 60 per cent. of metallic iron.

Limonite, also known as "brown hematite," has a dark, grayish-brown color. Nearly all the limonite occurring in Missouri is dull, and sometimes earthy in the fracture, amorphous. It occurs partly in porous masses, the irregular pores and cavities being filled with yellow ochre, partly in botryoidal and stalactitic forms. The hardness of compact limonite is about 6, and pretty uniform. The streak is yellowish brown. Limonite is unmagnetic. It is chemi-

cally composed of peroxide of iron, and water in varying quantity, and contains from 45 up to 60 per cent. of metallic iron.

It is sometimes clayish, and in several localities mixed with broken chert. It is invariably accompanied by soft, yellow ochre, distributed in small cavities throughout its mass, but also occurring in larger accumulations occasionally.

Limonite does not seem to undergo any material changes by exposure.

SPECIAL DESCRIPTION OF MISSOURI IRON-ORES.

a. SPECULAR ORES.

I. Specular Ores in Porphyry.

Iron Mountain Ore.—The iron-ore of the Iron Mountain corresponds in its mineralogical qualities to the general characterization of Missouri specular ore as given above, and may be considered as a type. It is very uniform in its character in the various parts of the vein. Also the surface-ore has the same appearance and qualities, with the only exception that it is in the great average a little softer, its hardness being generally slightly below 6, while that of the vein-ore is slightly above 6. The color of both is steel-gray, with a slight tint of blue. Their streak is dark red to Both have an uneven fracture, a nearly metallic lustre on fresh-broken faces, a subcrystalline to massy structure. structure is occasionally inclined to become lamellar. In this case the ore is brittle, and breaks in long flat splinters with very thin and sharp edges. The surface of fracture of such pieces is very bright, and shows indications of a coarse crystallization, the single indistinct crystals being flattened and drawn in length in the direction of the long axis of the splinter. Many of these splinters are strongly magnetic, some less so. The former show a distinct polarity. The magnetic axis, however, never coincides with or ever lays parallel to any one of the three main axes of the splinter, but it is always in an inclined position to all of them. Marks of distinct crystallization are very rare in the Iron Mountain ore. Wherever distinct crystals occur in holes or fissures, they are mostly small and micaceous. These small micaceous crystals are sometimes also distributed throughout the mass of the softer ore, filling

the very finest pores. This is more frequently the case in the surface- than in the vein-ore.

All Iron Mountain ore is magnetic. I have not been able to discover a single piece entirely free from magnetism. Some of it. besides the special variety above mentioned, is strongly magnetic with distinct polarity, the north pole of a compass-needle being attracted by one side and repulsed by the other side of the same piece, producing very strong declinations. The greater part of the Iron Mountain ore acts, however, but slightly on the needle, but shows, nevertheless, frequently distinct polarity. Some of it does not seem to act at all on an ordinary compass-needle. But when reduced to a fine powder, some parts of it are invariably attracted by a magnet of ordinary power, while other parts, although equally fine, are not attracted. This remark, which is correct even for the impure ores from the small veins of but one to two inches' thickness, as they occur in the so-called "bluff," proves the universal distribution of magnetism in the Iron Mountain ore, and besides indicates that this magnetism is a quality inherent in certain small particles only, while others are free from it. This is one step, unfortunately but a small one, toward the explanation of the inequality existing between the magnetic strength of one piece of ore and that of another. This inequality exists to such an extent that sometimes a piece of ore, whose largest dimension does not exceed one inch, is found to be in part strongly magnetic, in part very weak, as may be seen when the piece is broken and the single fragments tested. This inequality seems to be independent of the location, and shows itself in the same manner and degree in the large vein, in the smaller veins, and in the surface-ore. Pieces inclined toward a crystalline structure are more generally strongly magnetic than others, also those containing secretions of micaceous crystals. Small differences in the chemical composition do not seem to influence the magnetism. The latter seems especially not to be dependent on small variations in the amount of protoxide the ore contains.

We see from the following analyses that the sample No. 3, although strongly magnetic, contains only 2.34 per cent. of protoxide, while the sample No. 4, which is very little magnetic, contains 5.72 per cent. of it. The magnetism of these samples was tested by approaching the single pieces to the north pole of a compass-needle.

Afterward, however, some of them were reduced to a powder, and were tested by approaching a magnet to the powder. Tested in this way, the powder of No. 4 seemed to be attracted more lively and more copiously than that of No. 3. A repetition of these tests showed the same results.

These facts would invite to a closer investigation of these matters. The axis of polarity in single pieces of Iron Mountain ore is never either parallel nor rectangular to the cleavage or to the surface of fracture, and runs very frequently from one point near the edge to another point near the centre of the piece on the opposite side. No ore with active magnetism, constituting a natural magnet, and attracting iron-filings, was found on the Iron Mountain.

The following analyses, made by Mr. A. A. Blair, of St. Louis, will show the chemical composition of the Iron Mountain ores:—

		VEIN-	ORE.		Suri	FACE-ORE.	
	r.	2.	3.	4.	5.	6.	7.
Insol. silicious matter.			4.71	6.76		1.88	
Peroxide of iron		96.78	91.45	86.75	• • • • •	95.04	95.15
Protoxide of iron			2.34	5.72		2.57	
Alumina	** * * *		0.93			0.75	
Lime	****		0.45			0.15	
Magnesia			0.19			O. I 2	
Manganese			0.00			0.00	
Sulphur	0.008	0.016	0.00	0.00	0.012	0.005	0.003
Phosphoric acid	0.112	0.119	0.252	0.081	0.067	0.071	0.125
	* * . * *	••••	100.322	•••••	••••	100.586	
Metallic iron		67.75	65.78	65.16		68.53	66.60
Phosphorus		0.052	0.110	0.035	0.029	0.031	0.057
÷				ble Siliciou			
Silica		3.28	3.99	6.16		1.57	4.54
Alumina			0.47			0.04	** * * *
Lime	0.149		0.06		0.13	0.02	****
Magnesia			0.03			0.017	
Specific gravity			5.149	4.934		4.835	4.929

- 1. Average sample of the vein or quarry ore from all parts of the mountain, sampled by Major T. B. Brooks, of Marquette, in May, 1872.
- 2. Average sample of ore from the eastern part of the principal vein (cut D).
- 3 and 4. Average sample of ore from the western part of the principal vein (cut A).
 - 3. Decidedly magnetic pieces.

- 4. Pieces not acting on a compass-needle.
- 5. Average sample of the surface-ore from all parts of the mountain, sampled by Major T. B. Brooks, of Marquette, in May, 1872.
- 6. Average sample of surface-ore from the south slope of the mountain.
 - 7. Average sample of surface-ore from the north-west slope.

The metallic iron in samples 1 and 5 was determined by Dr. Otto Wuth, of Pittsburgh, as follows:—

From the above analyses we may conclude that the Iron Mountain ore is very rich and very uniform, in general; that the surface-ore is a little richer in metallic iron, and less phosphoric than the vein-ore; that both are nearly free from sulphur; and that the percentage of phosphorus is variable, though never running very high. A comparison made between the analyses 3 and 4 on the one hand, and 2 on the other hand, might indicate that the vein-ore grows richer toward the west. The number of analyses is, however, too small to warrant the correctness of such a conclusion.

It will be noticed that the specific gravity, as given above, agrees well with the results of the analyses.

The Iron Mountain ore is in the whole very pure and nearly free from mechanical admixtures of foreign matter. A few minerals, however, occur in it occasionally, namely, porphyry, apatite, and quartz.

Admixtures of porphyry never occur in such a manner as to injure the ore, the enclosures having so large a size that the porphyry is easily separated from the ore. This is therefore more of geological than of mineralogical interest.

Fine clay, probably of porphyric origin, is sometimes intimately mixed with the ore in the smaller veins in the "bluff." The thicker veins are entirely free from it. Apatite must have been very frequent formerly in those parts of the veins which are in immediate contact with, or at least in pretty close vicinity of, the porphyrywalls, to judge from the numerous large and small holes of crystalline shape to be found in the ore in such places. These holes occur generally in larger or smaller groups, in some parts of the

principal vein, as well as in the "bluff" veins and in the surfaceore. The crystals which originally filled these holes were hexagonal prisms of varying sizes, from the smallest up to a length of 3
inches and a diameter of I inch. They generally start at the wall
of the vein, or at the wall of some fissure existing in the vein, and
reach into the ore, sometimes perpendicular to the wall, but
more frequently at some angle to it, and often nearly parallel to
each other. The prisms are distinct and sharp near the wall,
where they start from, but the more they reach into the ore the
smaller is their diameter and the less distinct and sharp their edges.
Many run out into a sharp point and look almost like sharplypointed pyramids, or, when rounded off, as they frequently are,
like needles or lances. Some, however, show indications of the
basis of the prism, also at that end of the crystal which lies free in
the ore.

All these crystallic holes are very likely impressions of crystals of apatite formerly present in the ore, and some of them, though fortunately but few, are yet filled with nearly fresh apatite. The empty holes are generally clad out with a very thin layer of a fine ferruginous clay.

The third mineral found in the Iron Mountain ore is quartz. It occurs but rarely, and nearly always near the walls of the veins, especially in disturbed places, or near large enclosures of broken porphyries in the interior of the principal vein. This quartz is there evidently the result of a process of infiltration, which has taken place long after the formation of the ore. It fills small fissures or irregular cavities, or the crystallic cavities above described, which were left after the removal of the apatite. It is always more or less distinctly crystalline, and sometimes forms drusy aggregations of crystals, slightly covered with yellow hydrated peroxide of iron (yellow ochre).

Pilot Knob Ore.—The Pilot Knob ore differs somewhat in its qualities from most of the other specular ores in the State. Its color is steel-gray to pearl-gray, with a very marked tint of skyblue. Its lustre is so faint that it can scarcely be called submetallic. Its structure is crystalline to granular, with a very fine grain barely to be seen with the naked eye. Its fracture is either even or subconchoidal. It shows a plain stratification, and splits parallel to it into plates 3/4 to 2 inches in thickness. These plates break in par-

allelopipeds with sharp edges and with surfaces that are sometimes at right angles, but more frequently inclined to each other. The faces parallel to the stratification are generally even, or nearly so; the others are either subconchoidal or irregular, with sharp edges and corners. The general flight of two faces situated opposite each other is always more or less parallel. The hardness of the Pilot Knob ores is very variable, and the faces parallel to the stratification are always perceptibly harder than those lying in other directions. The former are mostly near $6\frac{1}{2}$, the latter near 6. There are, however, a few places in the Pilot Knob mine where the ore is considerably softer, about $5\frac{1}{2}$. The hard ore passes gradually into the soft. All Pilot Knob ores are very brittle. Their streak is uniformly dark red. Also the poorer ores show this

None of the Pilot Knob ores from the main body of the deposit, neither below nor above the slate-seam, disturb an ordinary compass-needle, with the exception of the uppermost layers of poor ore, in the eastern cut, immediately below the conglomerate which forms the summit of the mountain. Also, most of the fine ore which is mixed with this conglomerate has polaric magnetism. The greater part of it acts on the needle very strongly. This conglomerate has indications of stratification, and the magnetic axes run either at a right angle or parallel to it. The ore found in loose pieces in the conglomeratic detritus covering the western slope of the mountain is also polaric-magnetic, some of it strongly so. Magnetic pieces of thinly-stratified, poor ore are also found loose in the creek north-east of the mountain. All ores on the Pilot Knob, those in the deposit as well as those in the conglomerates, are slightly attracted by a magnet, when they are ground fine, and those which, as above mentioned, act strongly on the needle, are not attracted by the magnet with any greater power or in any larger quantity than those which do not seem to act on the needle at all.

The chemical composition of the Pilot Knob ores from different parts of the mine is very different, as may be seen from the following analyses, made by Mr. Andrew A. Blair, of St. Louis:—

	ORES :	IN THE MAIN ORE-BED BELOW THE SLATE-SEAM.				ORES ABOVE THE SLATE-SEAM.		
	т.	2,	3.	4.	5.	6.	7.	8.
Insoluble silicious matter			14.75	5.57				
Peroxide of iron	87.18	83.28	84.33	90.87	67.38	77.02	62.88	52.18
Protoxide of iron			0.15	1.67				
Alumina			0.75	0.53				
Lime			0.21	1.76				
Magnesia			0.14	0.13				
Manganese			0.00	0.00				
Sulphur	• • • • •		trace		0.006		0.079	
Phosphoric acid			0.035	0.069	0.092		0.101	
	••••	••••	100.365	100.677				
Metallic iron	61.03	58.29	59.15	64.91	47.16	53.91	44.01	36.52
Phosphorus			0.015	0.031			0.044	
	Insoluble Silicious Matter.							
Silica	••••		13.27	5.18	30.10	••••	28.16	
oxide of iron)			1.44	0.36				

- 1. Average sample of ore from the main ore-bed below the slate-seam, in the western cut (A).
- 2. Hard ore from the central portion of the main ore-bed, below the slate-seam, west of the soft ore (analysis 3), in the central cut (B).
- 3. Soft ore from the central portion of the main ore-bed, below the slate-seam, in the central cut (B).
- 4. Hard ore from the central portion of the main ore-bed, below the slate-seam, east of the soft ore (analysis 3), in the central cut (B).
- 5. Average sample of ore from the main ore-bed, below the slate-seam, in the eastern cut (C).
- 6. Ore in conglomerate, 5 feet above the slate-seam, in the central cut (B).
- 7. Average sample of ore from the strata above the slate-seam, in the eastern cut (C).
- 8. Average sample of the better ore on the refuse-heaps of the eastern cut (C). This ore is not smelted, nor shipped, but thrown away with the rock-refuse.

We may conclude from these analyses that there is a great difference between the richness of the ore in the main bed below the slate-seam, and that of the ore above the slate-seam, the former showing, in the lower and central parts of the deposit, an average of about 60 per cent. of metallic iron, while the latter contains only

53 per cent., or, as this latter ore is very variable in itself, we may say from 45 to 55 per cent. The ore from the upper part of the deposit, from the eastern cut (C), is decidedly poorer than that from the central and western cuts. It contains only 47 per cent. below the slate-seam, and 44 per cent. above it, and the uppermost and poorest layers stand as low as 36, and some would analyze considerably lower than this. The ores from the last-mentioned uppermost layers, in the eastern cut, are thrown away now. But a considerable part of them could undoubtedly be smelted with profit in the Pilot Knob furnace. They could not be shipped, being too poor to pay for both the transport and the smelting. But they would give a fair profit, when smelted in place, especially when mixed with some of the richer ores or with limonites, or, better, with both. As the metallurgical and commercial value of ores increases very rapidly with their richness, I have no doubt that a great gain would result for the Pilot Knob Company, in money and in reputation, if the 60 per cent. ores from the lower and central part of the main bed would be kept and sold separate. These could be shipped to Indiana, Ohio, and Pittsburgh, and command high prices, while the 50 per cent. ores could be used in Missouri and Illinois, mixed with other Missouri specular ores and limonites, and while the 35 to 40 per cent. ores would be smelted at Pilot Knob.

We see from the above analyses that the Pilot Knob ores contain very little sulphur and phosphorus. Their principal impurity is silica.

This silica is either chemically combined, or, more likely, intimately mixed with the ore in microscopic particles or grains. Mr. Blair's analyses show it to vary from 5 to 13 per cent. in the richer ores, and to be as high as 30 per cent. and over in the poorer ores.

The upper part of the ore above the slate-seam is often intimately mixed with porphyry. The main bed is entirely free from it. Besides these two kinds of admixtures, a few minerals occur occasionally in very small quantity in fissures, as the micaceous oxide of iron and a yellowish-white, lamellar, crystalline, translucent heavy-spar. These two minerals occur rarely, but generally together and mixed with each other, in fissures, in the eastern portion of the central cut (B), below the slate-seam. A mineral resembling the last described, probably also heavy-spar, sometimes

forms a thin film, composed of many crystalline lamellæ of a vitreous lustre, over the even faces of separation of the ore, especially
over those which have a nearly vertical position in the bed. This
film is, in some instances, very pure and transparent; in others it
is more grainy, of a reddish color, opaque, and up to one-eighth
inch thick. In the latter case it is not pure, but seems to be mixed
with loose and extremely small grains or crystals of quartz. Mr.
Chauvenet found the specific gravity of the sample No. 3, which is
soft ore from the central part of the main bed, = 4.386, and that of
sample No. 4, which is hard ore from the same place, = 5.019.
These results agree with the results of the chemical analyses, which
show much more silica and less iron in the soft ore than in the hard.

Shepherd Mountain Ore.—The ore from Shepherd Mountain, in its mineralogical qualities and chemical composition, resembles a little more a magnetite than any other ore in Missouri. It is, however, in the main a specular ore, very similar to that of the Iron Mountain, as above described. Its color and streak are slightly darker than those of the latter. Its hardness is considerably less, being about 51/2 in the average. Its lustre is less bright. It has no stronger tendency toward a crystalline structure. The crystalline faces, though equally indistinct, are rather smaller. Splendent lamellæ of micaceous crystals, disseminated through the mass, as they occur rarely in the Iron Mountain ore, are quite general in the Shepherd Mountain ore, and the ores from these two localities may generally be distinguished exteriorly from each other by this characteristic, as well as by the marked difference in their lustre. The Shepherd Mountain ore is also much tougher and less brittle, and breaks with less sharp corners and edges.

In some places in the depth of the northern vein (A), the ore gets nearly black in color, finely granular, and fully black in the streak, thus approaching a true magnetite. A very soft, black ore occurs in irregular masses of limited extent in the upper part of the central vein (B). These masses show a more distinct crystallization, and sometimes aggregates of lamellar crystals of specular hematite in the forms of the rhombohedric system. The glittering, small crystals of micaceous ore are very numerous in some parts of these soft masses, while other parts are quite dull. The ore in the small southern outcrop (C) has a distinct, coarsely crystalline structure, with a pretty bright metallic lustre.

When inspected through a magnifying glass, the Shepherd Mountain ore appears to be composed, on the one hand, of crystalline parts, with a color and a lustre very similar to that of the Iron Mountain ore; on the other hand, of a dull, dark-red, ground mass surrounding the crystalline parts. Sometimes the one prevails, sometimes the other. I found a specimen in which these two distinct kinds of ore form alternate, undulating layers about one-eightieth of an inch in thickness, giving the piece a striated appearance.

The magnetic qualities of the Shepherd Mountain ore are much more pronounced than those of either the Iron Mountain or the Pilot Here again we have to distinguish two different Knob ores. modes of action of magnetism, which seem to be independent of each other in certain respects and within certain limits; first, the magnetic influence of the ore on a compass-needle, and second, the attractive influence of a magnet on small particles of ore. The influence on the needle is much less dependent on the chemical composition and on the structure of the ore than on its position in the vein. The specimen No. I, amongst the following analyses, contains 1.8 per cent. of protoxide, and is much more magnetic than the No. 2, which contains 2.97 per cent. There is no decided and regular difference between the magnetic strength of hard and soft, of dull and bright, of fine-grained and coarse-grained ore; but all the ore in the upper part of the vein, from the outcrop to a depth of 30 or 40 feet, acts strongly on the needle, while the action of the ore from the lower part is much weaker, though very different in different places. Besides being stronger magnetic, the upper ore is also decidedly polaric, the axis of polarity being nearly parallel to the cleavage. Regarding the lower ore in Shepherd Mountain, it is a remarkable fact that, although acting on the needle invariably, it is, with rare exceptions, entirely unpolaric, while the Iron Mountain and Pilot Knob ores, although mostly weaker, are nearly always polaric. A piece of this lower Shepherd Mountain ore repels the north pole of a magnetic needle, and attracts its south pole, and, when the piece is turned over, instead of acting in a reversed manner, it allows the needle to return to its natural position, and does not produce any deflection whatever. In the northern vein (A) the strongly magnetic and polaric ore is more frequent, and reaches deeper than in the central vein (B). There also occurs some ore (analysis No. 4) which contains a large amount of protoxide, and has most of the mineralogical characteristics of a true magnetite, especially a deep-black streak.

All Shepherd Mountain ore is strongly attracted by a magnet of ordinary power, when either pulverized or ground coarsely. No decided difference can be discovered in this respect between the ore which strongly disturbs the needle and that which affects it but little. Neither can a difference be discovered in this respect between the polaric and the unpolaric ore. When a polaric piece is broken or ground, the single fragments are polaric. From a strongly polaric piece, I broke off some particles from the north pole and also some from the south pole. Each of these particles had itself two poles, when tried by approaching it to a compass-needle, and all these particles were attracted by both poles of a magnet. The position of each single particle, when attracted by the north pole, was however reversed, when compared to the position of the same particle when attracted by the south pole.

When the ore is finely pulverized, only a part of the powder is attracted, while the rest seems to be unmagnetic. This fact shows that the magnetism is inherent in certain very small particles only, while others are free from it, and explains the difference between the two modes of magnetic action. A piece of ore containing comparatively but few though strongly magnetic particles, may not disturb the needle; nevertheless, when the ore is pulverized, a magnet will exert a lively attraction on the magnetic part of the powder.

Bowlders of ore are sometimes found on the Shepherd Mountain which are strong natural magnets, possessing active magnetism, and attracting iron-filings. A piece of iron, when rubbed against such a natural magnet, becomes itself a magnet.

The chemical composition of the Shepherd Mountain ore may be seen from the following analyses made by Mr. Andrew A. Blair, of St. Louis:—

Du Louis.				
	ı.	2.	3.	4.
Insoluble silicious matter	5.15	6.76		
Peroxide of iron	94.84	88.56	96.70	79.39
Protoxide of iron	1.80	2.97		14.22
Alumina		1.55		
Lime		0.35		
Magnesia		0.04		
Manganese		0.00		
Copper		trace in 5 grms.		0.00
Sulphur	0.00	0.00	0.00	0.00
Phosphoric acid	0.025	0.039	0.032	0.038
		100,260		

Metallic iron	64.31	67.69	66.63
	Insoluble Silie	cious Matter.	
Silica 4.05	5.98		
Peroxide of iron	0.07		
Alumina	0.46		
Lime	0.12		
Magnesia	0.05		
Specific gravity	4.714		

- I. Average sample of ore from the upper part of the central vein (B); magnetic; streak dark red.
- 2. Average sample of ore from the lower part of the central vein (B), about 80 feet below the outcrop; slightly magnetic; streak dark red.
- 3. Soft, friable ore from the lower part of the central vein (B); slightly magnetic; streak dark red.
- 4. Hard, black ore from the northern vein (A); strongly magnetic; streak black.

These analyses show that the Shepherd Mountain ore is very uniform in its chemical composition, very rich in iron, and almost entirely free from sulphur and phosphorus. It is very nearly as rich as the Iron Mountain ore, and much purer than either this or the Pilot Knob ore.

It is, besides, nearly free from mechanical admixtures, small specks or thin seams of a soft, white clay, probably decomposed porphyry, being the only foreign matter generally found in it. The northern vein (A) contains, in a few places near its outcrop, some crystalline iron pyrites. This mineral, however, occurs very rarely, and does not injure the general quality of the ore in the northern vein, as the above analysis, No. 4, shows. The Shepherd Mountain ore is perhaps the best iron-ore in Missouri.

Specular Ores from the smaller Deposits in Iron County.

—These ores are very variable, approaching partly the Pilot Knob ore, partly the Iron Mountain ore, in their general character. Micaceous ore is of very frequent occurrence in most of these smaller deposits. On Buford Hill, 2½ miles west of Iron Mountain, micaceous oxide is found almost exclusively. It occurs in considerable irregular accumulations, mixed with quartz. The crystalline lamellæ have a black color, a bright lustre, and variable sizes, up to one-fourth inch diameter. The ore is strongly magnetic, with distinct polarity.

The ore from Cedar Hill, north-west of Pilot Knob, resembles the

Pilot Knob ore externally. It has a grayish color, with but little lustre, and is very dense, hard, and brittle. The streak is red. This ore is distinguished by the absence of all magnetism. It does not affect the needle, and is not attracted by a magnet, after being crushed or pulverized. It is sometimes mixed with specks and seams of brown porphyry. Most of it is pure and very rich in iron.

Mr. Blair found in an average sample, taken by Prof. Pumpelly from all parts of the mine—

Insoluble matter 5.62	2
Peroxide of Iron93.54	+
Sulphuro.oc)
Phosphoric Acid)0
Metallic Iron	7
Phosphorus 0.03	

The ores found imbedded in stratified porphyry in some localities, one-half mile east of Pilot Knob, are more like the Iron Mountain ore, having a darker color, a brighter lustre, and a less degree of brittleness than the Pilot Knob ores. They are mostly softer than 6. All the specimens I found there have a pretty strong magnetic polarity.

The ore from Lewis Mountain, near Arcadia, is very variable in its mineralogical qualities. Most of it looks like the Cedar Hill ore, and is unmagnetic. Other parts are softer and tougher. Wherever crystallization appears, the ore is micaceous and magnetic. The following analysis, made by Mr. A. A. Blair, of St. Louis, shows this ore to be about equal in quality to the better ores of Pilot Knob. This analysis was kindly furnished to me by Hon. Thomas Allen:—

INSOLUBLE SILICIOUS MATTER.

Silica	14.45
Alumina	0.51
Lime	0.06
Magnesia	0.04

What has been said of the Lewis Mountain ore may also be applied to that of Hogan Mountain. Masses of soft, coarsely crystalline, semi-micaceous, slightly magnetic ore, are, however, more frequent in the latter than in the former.

The Buford Mountain ore is dull, dark-colored, soft, in part earthy. It is partly massy, with indications of stratification, partly in irregular botryoidal forms, and frequently mixed with specks and seams of decomposed porphyry, of white and red clay, of hydrated peroxide of iron, and of black peroxide of manganese. The streak is therefore very variable, being red where the iron-ore prevails, and black where the manganese-ore prevails. The ore has a strong polar magnetism. An average sample, taken by Professor R. Pumpelly from all parts of the lower or main cut, was analyzed by Mr. Regis Chauvenet, of St. Louis, and gave the following result:—

Insoluble matter 8.54 pe	er cent.
Peroxide of Iron	"
Peroxide of Manganese19.46	"
Sulphur 0.011	"
Phosphoric Acid 0.102	66

As this ore is rich in both iron and manganese, and as the oxides of both these metals are, to all appearance, intimately mixed, this ore is likely to prove a very valuable material for the manufacture of Spiegeleisen, now so extensively used in the Bessemer process.

Ores with little iron and much manganese were found by Prof. Pumpelly on Mr. Cuthbertson's land in that vicinity, and analyzed by Mr. Chauvenet, showing—

Insoluble matter	0.44 p	er cent.
Peroxide of Iron	3.30	"
Peroxide of Manganese	33.56	"

The above descriptions will suffice to characterize the specular

ores occurring in the porphyries of the eastern iron-region of Missouri.

2. Specular Ores in Sandstone.—I have mentioned above, that the specular ores occurring in the Silurian sandstones in the central ore-region, have nearly the same mineralogical and chemical properties as the specular ores just described. They differ, however, considerably in one respect. I have alluded to the alterations which specular ores often undergo, passing either into soft, red hematite, or into limonite, or into carbonates. These alterations have not been spoken of in the special description of the specular ores in the porphyry, because they never take place there. We may find these specular ores, which have originated in the porphyry, in their original position as veins, beds, etc., in the solid porphyry; we may find them in veins, in the clayish "bluff" of Iron Mountain, or broken up into large and small fragments, imbedded in loose detritus in the same locality, or in outcrops, or as surface-ore being exposed to the atmosphere during an incalculable length of time. In no case can we find these ores to have undergone any material changes in their chemical or mineralogical character. On the other hand, those specular ores which have originated in the sandstone are invariably altered when broken up, or when exposed to external chemical influences. These alterations will be more fully spoken of in the following special description of the ores in which they occur, especially in that of the Scotia Iron Ridge and St. James ores. I will however say here, that they take place in three distinctly different directions, depending in each case on the character of the external influence which causes them.

First. When certain specular ores are directly exposed to the influences of the atmosphere, or when they are in such a position that both air and water may have access to them alternately, the ores become gradually changed into brown and yellow limonites.

Second. When these same specular ores are broken up in larger or smaller blocks, and covered with sandy detritus, so that water or mineral solutions have access to them, while the atmospheric air is more or less completely excluded, the ores become gradually changed into a soft, red hematite, which is in many instances greasy to the touch.

Third. When these same specular ores are broken up and in a state of gradual transformation into soft, red ores, or else when they

are already transformed into soft, red ores, certain mineral solutions containing carbonic acid seem to dissolve the iron under certain circumstances, and to depose it again as carbonate in fissures and cavities, either in the ore itself, thus changing the latter gradually into spathic ore, or in the adjacent rocks.

The last-mentioned transformation occurs but rarely, and on a small scale; the first is more frequent, though not generally very extensive; but the second is quite common, and has produced thorough-going changes in several important deposits, while no broken-up or disturbed deposit is entirely free from it.

The fact that these transformations do not occur in the specular ores which have originated in the porphyry, leads us to consider more closely the dissimilarity, however small it may be, between these ores and those originally imbedded in the sandstone. In doing so, we find that the latter are in the average somewhat softer, their hardness varying from 51/2 to 6. They are also a little less silicious, less compact, and more porous. The small, irregular cavities and cracks which are occasionally found in the former are quite universal and more equally distributed in the latter. When sufficiently large, these cavities are also clad with numerous small crystals of peroxide of iron, but they are less frequently filled with quartz or other minerals. Regarding their form, these cavities look somewhat different in the two different kinds of specular ores. While those in the porphyry-ores seem to be produced in part by the removal of minerals formerly enclosed in them, in part by a local crystallization of the mass of the ore, and by the expansion or contraction effected by crystallization, the cavities in the sandstone-ores have more the appearance as if they were produced by a very slow and uniform contraction of the ore, which may have taken place during its consolidation, perhaps by a gradual abstraction of the fluid from which the ore was precipitated.

I have the impression that the above slight differences between the two kinds of specular ores of Missouri are sufficient to account for the great difference in their capability of being altered by exterior influences. The associated rocks or soils may, however, contribute toward effecting this difference. The porphyry-ores, when broken up or exposed, are generally associated with less penetrable, clayish materials, the sandstone-ores with sandy and cherty detritus, which presents an easier passage to water or solutions.

If the existence of numerous small cavities in the mass of the sandstone-ores is not one of the principal causes of the alterations of these ores, the fact that these alterations nearly always start in such cavities certainly proves that their existence greatly facilitates and promotes transformation of any kind.

Franklin County Specular Ores.—There are a few occurrences of specular ore in Franklin County, in the neighborhood of Stanton and in the north-west corner of Washington County. The greater part of these ores is not very pure, as far as can be observed from the present condition of the ore-banks. The hard ores are either silicious in themselves or intimately mixed with sand, so much so in one locality as to constitute rather an impregnated sandstone than a real iron-ore. They have a dark, bluish-gray color, a light-red streak, and are slightly magnetic. The soft, red ores, as produced by the alteration of the specular, have a light-red color and streak, are somewhat greasy, and mixed with very fine, clayish substances. They are, however, sufficiently rich for being smelted.

Scotia Ores.—The specular ore of the two Scotia banks, on the Meramec River, in Crawford County, occurs in various forms and in nearly all stages of transformation. The hard, unaltered ore is in bowlders imbedded in the soft, red ore.

The hard ore is steel-gray, with a submetallic, and on fresh-broken surfaces frequently metallic, lustre. It has a finely-crystal-line structure, an even to subconchoidal fracture. Its hardness is 5½ to 6. It is slightly polaric-magnetic. It is pretty uniform in its appearance and 'structure, but contains those numerous and pretty equally distributed little cavities of which I have spoken above. The Scotia ore is, however, distinguished from all the other known specular ores in the State by the frequent occurrence of larger drusy cavities, which contain botryoidal and reticulated forms of ore, and are covered all over with small, highly-splendent crystals of peroxide of iron, which often have an irised tarnish, and play in all colors, presenting a beautiful appearance.

Smaller and larger, well-formed and transparent quartz-crystals, up to one quarter-inch diameter, often of a fine yellow color and of a bright, vitreous lustre, are likewise met with in these irregular cavities, which sometimes reach a length of several inches. Occasionally such cavities are filled by amorphous or subcrystalline,

wax-yellow jasper, enclosing thin seams of white quartz and fine specks of crystalline ore.

The soft, red hematite which forms the greater mass of the ore in the Scotia No. 1, as far as it is at present disclosed, is not a very uniform material. It breaks with very irregular surface, almost like a conglomerate, and is full of irregular streaks, running in the deposit more or less vertical, of similar though somewhat differently colored and composed materials. Some of these are red, crystalline, and glittering, and often silky or greasy, others yellowish brown and earthy. The main body of this ore seems to be an irregular but intimate mixture of these same two materials, which separately form the streaks just described. The hardness of the mass is only 2 to 3. Its streak is red to brownish red. It is not magnetic.

Soft, yellow iron-ochre occurs also in big seams through the oredeposit, and some large pockets of it reach into it from the surface. This ochre is generally very porous, mixed with broken chert or with red loam, and permeated irregularly by very thin seams of subcrystalline quartz.

The bowlders of hard specular ore are generally surrounded by a layer of red ore, which has not fully the same degree of softness as the rest of the soft ore. Yet the transition from the hard to the soft ore is here more abrupt than in the Iron Ridge and Meramec deposits, and pieces showing this transition very plainly are comparatively rare. Nevertheless, all appearances indicate that the red ore has been gradually formed by an alteration of the specular ore, while the ochre seems to be a later product.

The deposit called Scotia No. 2 is distinguished by the occurrence of long and fine stalactites, all of which are, in the greater part of their mass, transformed into red ore. Some of those, however, which are over an inch thick have preserved a specular kernel, though softened to about 4 in the mineralogical scale of hardness.

The following two analyses, made by Dr. August Wendel, of the Bessemer Steel Works, Troy, N. Y., will show the composition of the Scotia ores, and also the chemical difference between the hard and the soft ore:—

	I.	2.
	Hard Specular.	Soft Ore.
Metallic Iron	69.37	63.15
Silica	0.59	1.52

Alumina	0.11		0.76
Magnesia	traces		traces
Phosphorus	0.016		0.105
Sulphur	0.058	, : .	0.095
Water	0.20		7.95

- I. Average sample of the hardest speculur ore from the eastern cut (a) of the Scotia No. I. Looks very pure. Amorphous to subcrystalline. Hardness, 5½ to 6½. Slightly magnetic.
- 2. Average sample of soft, red hematite from Scotia No. 1., partly red and greasy, partly earthy and slightly greasy, partly yellow ochre, with numerous small specks of specular ore and fine seams of quartz.

By calculating the peroxide of iron from the above percentages of metallic iron, we find for No. 1, 99.1 per cent.; No. 2, 90.21 per cent.

These analyses show that the hard Scotia specular ore is nearly pure peroxide of iron and nearly free from impurities, while the mixed soft, red, and ochrey ores contain some quartz, some probably hydrated silicate of alumina, a considerable amount of water, nearly twice as much sulphur and six times as much phosphorus as the hard ore. Nevertheless, these soft ores are quite rich in metallic iron.

The chemical changes which have taken place in the specular ore, by its transformation into soft, red ore, consist principally in the introduction of four substances, namely: of hydrated silicate of alumina, which substance very likely produces the greasiness of some of the soft ore; of water in considerable quantity; of phosphoric acid, and of sulphur. The three last-named substances are probably in combination with the iron. Quartz has besides been infiltrated into seams in the ochre. From the loose and porous structure of these soft ores, it must be supposed that some oxide of iron has been removed, and was partly replaced by hydrous silicates and phosphates. I shall come back to this subject in speaking of the Iron Ridge and St. James ores, of which similar comparative analyses have been made.

The high amount of water in analysis 2 is undoubtedly due to the presence of a considerable quantity of yellow ochre in the sample analyzed.

Specular Ores in the Steelville District.—The Steelville ores resemble the Scotia ores very closely, and do not need a special description. Some of the banks of this district seem to be almost entirely composed of hard, specular ore, as the Cherry Valley banks, which, although not yet opened, present that appearance. The ores of some other banks are more or less broken up and altered into soft, red ore, which is here also conglomeratic in its fracture, but less streaky, and more even in its color and general character than the Scotia ore. Distinct crystallization is rarely seen. Stalactitic forms of specular ore, exteriorly converted into red ore, occur at the Cherry Valley banks. Ouartz is not often found in these ores. Transformation of surface-ore into brown and yellow limonite is here very frequent, and can be well observed at the Cherry Valley and Ferguson banks. Smaller pieces are often entirely changed. Large bowlders retain a kernel of specular ore with mostly straight and sometimes almost sharp limits, which limits become irregular in such places only, where the small cavities and pores in the specular ore happen to be more numerous. The limonite is much more porous than the specular ore, and a removal of iron has evidently taken place during the transformation. The crystalline particles seem to resist this transformation better than the amorphous or subcrystalline ground-mass; for the limonite contains numerous specks of small, crystalline, specular ore. But the uppermost layer of bowlders is generally free from them, which shows that also these crystals finally become altered. The porousness of the limonite increases toward the surface, and the color gets more yellow, owing to very fine seams of ochre. Close to the specular kernel, the limonite is mostly dark brown, sometimes reddish brown. The cavities in the specular ore near its limit are clad with a brown or reddish-brown, earthy film. With the formation of this film the change evidently begins. film seems to get gradually thicker and more brown. The small cavities seem to widen and often to run into each other, thus forming larger cavities and spongy masses. The limonite partly remains brown, partly is altered into yellow ochre, permeating the brown ore irregularly in extremely fine seams.

The specular ore close to the limit of the limonite is apparently as hard as ever, and no gradual softening of the ore seems to precede this kind of alteration.

I will add here two analyses of pure specular ores from the Steelville district:—

	Steelville.	Cherry Valley.
Silica	1.84	1.73
Peroxide of Iron	97.49	
Alumina	0.07	• • • •
Lime	0.34	
Magnesia	0.12	
Phosphoric Acid	0.14	
Sulphur	0.00	0.16
Metallic Iron	68.24	67.69
Phosphorus	0.061	9.039

The sample from the Steelville No. I bank was analyzed by Dr. Otto Wuth, of Pittsburgh, for the "Iron Mining Company of Missouri," and the result published in the prospectus of this company. The second analysis was made by Dr. A. Wendel, of the Bessemer Steel Works, Troy, N. Y., from an average sample taken by myself at the Cherry Valley bank. Most of the latter sample was a pure, crystalline, specular ore. All of it was magnetic, a few pieces strongly so.

These ores are, according to these analyses, very rich in iron, and sufficiently pure for any purpose. In comparing these analyses with those of the Iron Mountain, Scotia, and other ores from deposits which are opened and mined, it must be borne in mind that the former were made with samples picked up on or near the surface, where they may occasionally have taken up some phosphoric acid from the ashes of the grass and brushes, which are purposely burnt off every year in many districts of central Missouri, or some sulphur from the reducing action of decaying plants on solutions of sulphates.

Specular Ores on the Upper Meramec River and its Tributaries.—None of the specular ore-banks in this district are as yet opened to any extent, and my last remark will therefore also apply to them. To judge from the appearance of the surface-ore, some of these banks, as the Winkler, Lamb, Benton Creek, Fitzwater, and Hutchins Creek banks, are likely to contain specular ore almost exclusively, although on all of them superficial transformations into limonite are developed more or less. Others, as the Grover, Ar-

nold, and Smith banks, seem to contain much red ore besides the specular. Both kinds of ore have here about the same mineralogical character as those in the Steelville district.

Pronounced magnetic properties are however, here, more frequently met with. Some of the specular ores from Benton Creek, Fitzwater, Hutchins Creek, and Smith banks are strongly polaric, especially those which are crystalline, or which contain numerous small crystals. Amorphous ores, as they sometimes occur, rarely possess magnetism in an observable degree. Also stalactitic specimens are generally unmagnetic. Some black sandstone, strongly impregnated with iron, but giving a light-red streak, from the Benton Creek bank, has distinct polarity.

Fine stalactitic forms, "pipe-ores," occur on the Smith bank No. 1. Some of these show on their upper side, which was exposed to the atmosphere, a beginning of a change into limonite, while a thin layer on the under side, which was imbedded in loam and sand, is changed into soft, red ore.

Many of these "pipe-ores" consist of clusters of thin, hollow stalactites, regular pipes, one-eighth to one-quarter of an inch in diameter, with comparatively wide holes and thin walls. They are covered on their inside walls with crystals of peroxide of iron, and on the outside with a thin film of dark-vellow limonite. The structure of these thin stalactites, whether hollow or massive, is generally crystalline and granular, rarely radiated. But they are sometimes surrounded concentrically by larger stalactites, in whose hollow interior they lie like a casting in the mould. These larger surrounding stalactites have always a radiated structure. They never close tight to the kernels which they surround. There is always a cylindrical space between the inner wall of the large and the outside wall of the small stalactite. This space is either empty, in which case both walls are covered with small crystals of oxide of iron, or it is filled with soft, red ore, perhaps produced by the alteration of such crystals.

This cylindrical space is sometimes very narrow and nearly filled up with splendent crystals. Thus the fracture of the stalactite shows sometimes a small, round, crystalline surface in the centre, surrounded first by a thin, annular layer of more loose and much more splendent crystals, and outside of this by a thick layer of less bright ore with a radiated structure. These formations sometimes

repeat themselves. A specimen I found on the Cherry Valley bank, where they are, however, less frequent, has five layers of radiated ore, alternating with thin layers of either splendent crystals or soft, red ore. The whole stalactite, thus composed, is divided in two halves by a thin crack, running across all the layers and through the thin central kernel, and being filled with the same splendent crystals which form some of the annular layers. The whole stalactite is exteriorly converted into soft, red ore, and lies loose in a conformable cavity in a piece of specular ore, the wall of which cavity is also covered with a layer of red ore.

Stalactites split lengthways, by a thin crack, partly filled with fine crystals of oxide, have also been found at the Scotia No. 2 bank, of which I have spoken before.

I add two analyses of ores from the upper Meramec district:-

Analysis I was made by Dr. Otto Wuth, of Pittsburgh, for the "Iron Mining Company of Missouri," and published in their prospectus. The specimen was taken from one of the Smith banks, and was evidently a very clean piece of specular ore.

Analysis 2 was made by Dr. A. Wendel, of Troy, N. Y., from an average sample of the various kinds of rich and poor surface-ore, taken by myself at the Benton Creek bank. This sample consisted only of one half-inch good specular ore. The other half was mostly a dull, black, uncrystalline, very hard, brittle, and silicious ore, mixed with some hard, grainy, and sandy ore of a black color and of a weak, resinous lustre. The object of this analysis was principally to see whether these silicious materials, which sometimes occur, especially at the outskirts of specular ore-banks, are worth smelting, as I sup-

posed they were from their color and weight. The result of the analysis shows that these materials are yet quite rich in iron, though less pure in regard to phosphorus and sulphur. Clean specular ore, from this and all other localities in this district, would undoubtedly analyze as favorably as the above specimen I.

Specular Ores in the Salem District.—Most of the ores of this district seem to be unaltered specular, corresponding in their properties to the general characteristics of this ore. The small, irregular cavities are very distinct and numerous in them, having sometimes the form of short cracks, wider in the middle and thinning out toward both ends.

Alterations into soft, red ore can be observed on the Arnold, Jamison, Pomeroy, and Taylor banks, and very fine alterations into limonite on the Simmons Mountain, and on the Arnold, Taylor, and Pomeroy banks, in the latter two on a pretty large scale.

Fine specimens of specular ore with a mossy and reticulated structure are sometimes found on Simmons Mountain.

The following observations I made in this district will throw some light on the paragenesis of the various materials connected with the specular-ore deposits in sandstone:—

I found on the Arnold bank botryoidal and mammillary forms of specular ore, clinging directly to a slightly ferruginous but otherwise unaltered sandstone, composed of coarse, loose grains with hardly any cement. The grains are slightly red, apparently from a thin film of red ore deposited on their surfaces. The specular ore is superficially converted partly into red ore, partly into brown limonite.

A specimen from the Taylor bank shows wax-yellow jasper of the same description as that observed at the Scotia bank, filling irregular and reticulated cavities in the ore. Other cavities in the same specimen are filled with transparent quartz. A specimen from the Jamison bank represents a conglomerate of irregular grains of ore, each of which is either partly or wholly surrounded by a variable layer of fine-grained, yellow sandstone adhering to it. They are cemented together by a coarse crystalline, transparent or white quartz.

Infiltrations of transparent quartz in the massive specular ore are frequent on the Jamison bank and on the Simmons Mountain. This infiltration seems to be accompanied or followed by a recrys-

tallization of the ore, or by an alteration of enclosed particles of specular ore into loose aggregates of splendent ore-crystals.

The limonite on the Simmons Mountain is, as most metamorphic limonite, full of pores and of large holes. These holes frequently contain infiltrations of crystalline quartz. Splendent ore-crystals, of a flat, rhombohedric form, are found lying on the quartz. Such ore-crystals are also seen there, lying on films of limonite which cover specular ore. Some of those rhombohedric crystals are themselves again exteriorly, or thoroughly, changed into brown limonite.

The porous limonite on the Pomeroy bank contains yellow ochre in its seams and cavities.

When these observations are held, together with others mentioned previously, and with the following—that the Scotia ores contain seams of wax-yellow jasper in specular ore, seams and specks of crystalline, transparent quartz and of splendent ore-crystals in the yellow jasper, seams of crystalline quartz in yellow ochre, and rhombohedric ore-crystals lying on drusy, crystalline quartz—we come to the conclusion that the order in which these various minerals have come into existence, beginning with the oldest, is as follows:—

- I. Sandstone, white or yellow.
- 2. Sandstone, colored by, or impregnated with, oxides of iron.
- 3. Massy specular ore.
- 4. Yellow jasper, perhaps simultaneous with the latter.
- 5. Soft, red, and greasy ore.
- 6. Brown limonite.
- 7. Yellow ochre.
- 8. Transparent, crystalline quartz.
- 9. Rhombohedric, splendent ore-crystals.
- 10. Red and brown incrustations of these crystals.

The specular ores of the Salem district are decidedly more magnetic than any I have mentioned, with the only exception of those from Shepherd Mountain. As the ores in the Upper Meramec district are more magnetic than those in the Steelville district, it is evident that there is a gradual increase of magnetism in the ores, from north to south, from the northern boundary of Crawford County toward the central part of Dent County. Here, however, as in the Iron Mountain and Shepherd Mountain, the magnetism

seems to have its seat principally near the surface of the ground. Specimens taken from the south-eastern ore-shaft on Simmons Mountain, about twenty feet below the surface, are nearly unmagnetic, while the ore near the surface in the same locality is distinctly polaric, some of it so strong that pieces two or three inches thick act on the needle from a distance of one to two feet. Pieces from the surface, which are, in the greatest part of their mass, converted into limonite, and contain only irregular and rough seams of specular ore, cropping the limonite in various directions, show nevertheless distinct polarity. Wholly converted pieces are unmagnetic. Magnetic but unpolaric pieces are comparatively rare, and generally attract the south pole and repulse the north pole of a compass-needle. Also stalactitic specimens possess polarity, the polar axis being parallel to the position of the stalactites. The specular ores from the Pomeroy, Taylor, Orchard, Jamison banks, and the "pipe-ore" from the Wiggins bank, have mostly a very strong polarity. Quartz-infiltrations seem to diminish the magnetism.

In the Salem district, as everywhere else in Missouri, unaltered specular ores contain very little sulphur and phosphorus, besides being very rich in metallic iron. The following analyses are proofs of this very important fact:—

	I.	2.	3.	4.	5.
	Arnold	Simmons	Jamison	Wiggins	Huzzah
	Bank.	Mountain.	Bank.	Bank.	Bank.
Silica	4.12	1.41	0.94	0.87	2.64
Peroxide of Iron	95.24	98.14	98.62	98.96	97.26
Alumina	O.II	0.06	0.06		
Lime	0.33	0.24	0.23		
Magnesia	0.15	0.11	0.08		
Sulphur	0.00	0.00	0.00	trace .	trace
Phosphoric Acid	0.052	0.038	0.07		
Metallic Iron	66.66	68.69	69.03	69.27	68.08
Phosphorus	0.023	0.016	0.031	0.027	0.03

All these analyses were evidently made with clean specular specimens. Protoxide of iron was not determined, although undoubtedly present in determinable quantity, to judge from the mineralogical properties of these ores, and from the fact that it is invariably

found in such ores when sought. Analyses 1, 2, 3, and 5 were made by Dr. Otto Wuth, of Pittsburgh; analysis 4 by Messrs. Chauvenet and Blair, of St. Louis. The three first analyses were made for the "Iron Mining Company of Missouri," and published in its prospectus; the last two for Mr. O. A. Zane, of St. Louis, who kindly put them at my disposal.

Iron Ridge Ores.—Unaltered specular ores from Iron Ridge are similar to the Scotia and Steelville ores in their general mineralogical character. They are rather more porous, and contain frequently enclosures of crystalline, transparent quartz. They are less magnetic than the Steelville ores, and as Iron Ridge is situated north of Steelville, they present a further proof that the magnetism of the ores in the central ore-region decreases toward the north and increases toward the south. Pieces of over three inches' diameter do not deflect a compass-needle. Fine ore-powder is, however, somewhat attracted by a magnet, especially the crystalline and glittering particles. The red ore is unmagnetic. By far the greater part of the ore from the Iron Ridge No. 1 is softened, and altered into a red and frequently clayish hematite. The transition from the specular to the red ore can be observed in the bowlders which are imbedded in the soft mass of ore, which principally constitutes the deposit. When these bowlders, which are sometimes two or several feet in diameter, are broken, the interior is seen to consist of a somewhat porous but pretty hard specular ore, of bluish-gray color, and composed of a subcrystalline matrix with little lustre, and of numerous very small and splendent crystals. Toward the outside of the bowlder the matrix gets softer, more porous, and disappears by degrees, leaving finally a spongy agglomeration of glittering black crystals. This black, crystalline zone in the section of a bowlder is from one-eighth to one-half inch thick. It is surrounded by and passes into a similar crystalline and glittering zone of red color, one-half to one and a half inches thick, which itself gets gradually less bright, then more and more greasy, then mixed with specks and flakes of white clay, and finally turns into a soft, clayish, dull and pale, red hematite, of which a considerable part of the deposit is composed.

The gradual disappearance of the matrix, and the appearance of the clay, are in many bowlders very plain and unmistakable. These two changes are, however, not simultaneous. The white clay is never seen in the black crystalline, and rarely in the red crystalline, zone.

The following analyses, made by Mr. Andrew A. Blair, of St. Louis, will show the chemical changes accompanying this interesting transformation:—

	I.	2.	3.
Silica	0.69	0.68	8.39
Peroxide of Iron	97.94	97.08	88.37
Protoxide of Iron	trace	0.31	
Alumina	1.17	1.50	
Hygroscopic Water	0.02	0.08	
Combined Water	0.19	0.23	3.09
Carbonic Acid	0.00	0.21	
Sulphur	0.00	0,00	traces
Phosphoric Acid	0.068	0.079	0.207
74. F . 11* T			- 0-
Metallic Iron			61.86
Phosphorus		· · · · · ·	0.091

The samples I and 2 are taken both from the same bowlder, which had a diameter of over 3 feet.

Sample I is specular ore, pure, porous, very little softened, from the centre of the bowlder.

Sample 2 is from the outside of the bowlder, more than a foot from the central part, where sample I was taken, and consists of soft, crystalline, glittering, and somewhat greasy red ore.

Sample 3 is an average sample, taken from all parts of the mine, and containing some specular but mostly soft ores of all varieties.

We see from the first analysis that the unaltered specular ore from Iron Ridge is free from sulphur, and has but little phosphorus, and is very rich in iron.

The second analysis shows an increase in protoxide of iron, alumina, water, and phosphorus, and a small amount of carbonic acid. The latter is probably combined with the protoxide of iron, being just about as much as is necessary to make carbonate of iron. This would indicate that carbonic acid might have something to do with this transformation, perhaps by dissolving and removing the matrix. As the alumina has increased, while the silica has not increased, we must suppose that either some alumina was added and

combined with some of the silica already present, or that some silica was removed and replaced by water, so as to form a hydrated silicate of alumina, which, I think, produces the greasy appearance and touch. Some of the alumina may be or may have been in combination with phosphoric acid.

The analysis 3 proves that the soft and fully transformed ore is not nearly as pure as that in the bowlders—a conclusion which we had already drawn from the comparative analyses made with the Scotia ores.

Specular Ores in the St. James District.—The ores in the St. James district are, when fresh and unaltered, very similar to the unaltered Scotia and Iron Ridge ores. They are, perhaps, a little less porous and a little more magnetic. They occur in very variable conditions and alterations.

The Meramec bank is especially interesting in this respect. There we find very pure and clean ore, generally somewhat softened; we find very hard and silicious ore, containing in its cavities transparent, crystalline quartz and yellow jasper; we find soft, red hematite in all stages of transformation; we find greasy paintores in various colors, from light red to dark purple; we find brown and yellow ochres, and porous, soft limonites, with seams of a very fine and uniform reddish-brown clay; we find the spathic iron-ore in specks and seams in red ore and in a peculiar, very dense, yellow, ferruginous limestone; we finally find ferruginous chert-conglomerates and sandstones, impregnated with iron-ore, or intimately mixed with brick-red and yellow ochres. As the Meramec bank was found to contain so many varieties of ore, it seemed to present a fine opportunity for comparative analyses. As, however, time and means would not allow to have a large series of analyses made, five of the most characteristic specimens were selected for this purpose. They were analyzed by Mr. Andrew A. Blair, of St. Louis, with the results given in the following table, under 1, 2, 3, 4, 5.

The two analyses 6 and 7 are taken from a pamphlet, entitled "Contribution to a Knowledge of the Iron-Ores of Missouri," published in 1872 by Prof. Charles P. Williams, Director of the School of Mines of Missouri. They refer to other ores from this district.

MERAMEC BANK.										
	I.	2.	3.	4.	5.	6.	7.			
Insoluble silicious matter			11.19	9.30						
Peroxide of iron	97.23	87.92	85.95	76.45	27.40	84.463	49.245			
Protoxide of iron			0.77		4.11	0.783	1.203			
Alumina			0.97			7.278	none			
Lime			0.12		36.01	trace	0.374			
Magnesia			0.07			0.114	0.000			
Carbonic acid			0.48		29.51	trace	0.530			
Water, combined	0.47	0.03	0.46	13.65	2.45	trace	0.087			
Sulphur	0.00	0.00	0.126	0.052	0.044	0.050	0.022			
Phosphoric acid	0.092	0.089	0.116	0.478	0.098	0.153	0.109			
Manganous oxide		• • • •	0.00			0.360	0.213			
			100.050							
		• • • •	100.252		• • • •		• • • • •			
Metallic iron	68 06	61.54	60.76	F2 F7	22.29	No Maa				
Phosphorus	0.040	0.039	0.051	53.51	0.043	59.733	35.397			
2.100001.01.01.01.01.01.01.01	0.040	0.039	0.051	0.200	0.043	0.000	0.047			
	Inso	luble Sil	icious Ma	atter.						
Silica	2.06	11.32	9.78	6.78	0.27	6.686	46.330			
Alumina			1.33							
Lime			0.00							
Magnesia,		* * * *	0.12							
			11.23							

- I. Is an average sample of the best ore found in the Meramec mine. It is a somewhat softened and slightly altered specular ore.
- 2. Is a hard and silicious specular ore from bowlders in the central and upper part of the Meramec bank.
- 3. Is a soft, greasy paint-ore of purple color, from pockets in the Meramec bank.
- 4. Soft and ochrey, porous limonite, brown and yellow, with some thin seams of very fine, brown clay from the outside of some of the bowlders in the Meramec bank.
- 5. Pretty dense but soft, red ore, and hard, yellow limestone, mixed, both containing crystalline, spathic ore in numerous specks and seams, from the lower part of the Meramec bank.
- 6. Is an ore from T. 38, R. 6, Sec. 33, probably the "Santee and Clark's bank." Prof. Williams describes the specimen thus:—
- "Mammillary and concretionary, with concentric layers, the central one being the blue-specular variety, the second of the brownish-red hematite, and the outer one a thin coating of brown hematite, probably limonite."
- 7. Is from T. 38, R. 6, Sec. 29, probably the "James bank." Prof. Williams describes the specimen as "finely granular, compact, brownish-red ore."

The analysis I shows that the clean specular ore, although in this

case slightly decomposed or altered, is very rich in iron, free from sulphur, and does not contain much phosphorus.

Analysis 2 shows that the harder and more silicious ore, when it has the natural color and brightness of a true specular ore, does not contain any more sulphur and phosphorus than the ore which is not silicious.

A complete analysis has been made of sample 3, principally for the purpose of finding out whether the supposition, made above, that the greasiness is produced by a small admixture of very fine and perhaps hydrated silicates of alumina, holds good, or what else may cause it. The result of the analysis does not suggest any other cause, but seems to support the above supposition. As in the 2d of the Iron Ridge analyses, we also meet here with a small amount of carbonic acid, and with a corresponding amount of protoxide of iron, so as to suggest the probability of the presence of carbonate of iron, and to lead us to the belief that carbonic acid is one of the agencies which effect, or at least prepare, the transformation of specular into red ore. Another interesting feature in this analysis is, that the percentage of phosphorus is increased but little, when compared to analyses I and 2, and that the increase is about in the same proportion as that from No. 1 to No. 2 of the Iron Ridge analyses. It seems therefore probable that the paintore is nothing else but the crystalline and glittering red ore (similar to the Iron Ridge sample 2), in a crushed and compressed condition. The comparatively large amount of sulphur in analysis 3 is also remarkable, and explains the purple color, which is undoubtedly produced by sulphides of alumina, lime, magnesia, and perhaps of alkalies in minute quantities. Analysis 4 gives the practically important result that the yellow ochre contains a very large amount of phosphorus, much larger than any of the other hard or soft ores.

Analysis 5 proves that the peculiar hard, yellow rock in which the spathic iron-ore frequently occurs, is a very dense carbonate of lime, probably mixed with some silicate and phosphate of lime, and with some peroxide of iron. This rock, as well as its enclosures, would deserve a repeated chemical examination, separate from that of the red ore, with which it was mixed in sample 5.

Analyses 6 and 7 are interesting, because, considering the mineralogical description of the specimens, they fully agree with and

confirm the results of all the analyses of ores given in this report, and lead to the same conclusions and views regarding those ores.

Specular Ores in the Rolla District.—The Rolla ores are nearly all more or less altered, or at least softened, though many not so much as to lose entirely their specular appearance. Most of them consist of specular and of soft red particles intimately mixed. They do not attract the needle perceptibly, but are attracted by a magnet partially, when powdered fine. Their porosity is very unequal in different parts of a piece or bowlder. Some parts are dense, while others contain irregular cavities up to an inch in length and onequarter of an inch in width, clad with crystals, or, more frequently. with a film of red ore. Such ores are those from the banks on Big Beaver Creek. They pass into red ores, softening at first, and then changing their color into dark red, and finally getting earthy and lighter red. The ores in the immediate vicinity of Rolla are mostly thus transformed, and besides frequently mixed with spathic ironore in specks and seams. This spathic ore sometimes occurs in larger masses, and is then accompanied by white or light-gray clay, enclosing well-formed crystals of iron pyrites, either single or in bunches.

The following analyses were taken from Prof. Charles P. Williams's "Contribution to a Knowledge of the Iron-Ores of Missouri":—

bourt ,					
	I.	2.	3.	4.	5.
Ferric Oxide (Perox) ide of Iron)	89.684	97.572	45.968	77.905	83.275
Ferrous Oxide (Protoxide of Iron)	0.684	0.400	18.988	2.251	1.206
Manganous Oxide	0.252	0.265		0.000	0.715
Alumina	0.199	0.802			traces
Lime	2.097	0.568	0.289		traces
Magnesia	trace	0.166	trace		traces
Silicic Acid (Silica)	2.951	1.144	1.159		3.099
Carbonic Acid		trace			
Phosphoric Acid	0.249	0.035	0.281	0.033	0.315
Sulphur	trace	0.009	trace	0.094	0.000
Combined Water		trace			
		100.981			
Metallic Iron	63.306	68.611	46.944	56,283	59.220
Phosphorus	0.109	0.015	0.122		0.137

The empty spaces in the above table, as well as in all the pre-

ceding ones, indicate in each case that the respective substance was not determined. The various specimens are described by Prof. Williams as follows:—

- 1. From T. 36, R. 7, Sec. 26 (perhaps the Hyer bank). "Blue specular mixed with brownish-red hematite, and containing some limonite and spathic iron."
- 2. From T. 37, R. 8, Sec. 33 (perhaps the Beaver Creek bank). "Blue specular ore, finely granular and compact; powder gives particles attracted by the magnet."
- 3. From T. 37, R. 8, Sec. 20 (perhaps the Buckland bank). "Mixed spathic iron and limonite, with some blue specular ore; powder slightly magnetic."
- 4. From T. 37, R. 8, Sec. 21 (perhaps the Kelly bank, No. 2). No description of this sample is given. It was probably a silicious specular ore.
- 5. From T. 37, R. 8, Sec. 15 (perhaps Taylor's Rolla bank). "Brownish-red hematite, somewhat cellular, slightly magnetic."

In comparing these descriptions with the above analyses, we see that the samples 1, 3, 5, which enclosed some red hematite, limonite, and spathic ore, contain much more phosphorus than the pure specular ores 2 and 4.

Analyses I and 2 show that the specular ores from the Rolla district are as rich in iron as any in central Missouri, and analysis 2 shows that, in an unaltered state, they are nearly free from injurious ingredients. As the same observations have been made regarding the ores of all the other districts, we may infer that these observations are generally true, and generally applicable to all specular ores that have originated in the Silurian sandstones of Missouri.

Gasconade and Miller County District.—The specular ores on the Gasconade River, and in Miller and Camden Counties, are similar to those above described. Some ores near Linn Creek are almost entirely unmagnetic. The specular ores found south of Tuscumbia seem to be very pure, to judge from the following analysis made by Messrs. Chauvenet and Blair, of St. Louis, of a sample from the west bank. This analysis was kindly furnished to me by Mr. M. S. Cartter, of St. Louis:—

Insoluble	11.077
Peroxide of Iron	
Sulphur	
Phosphorus	
Metallic Iron	61.96

b. RED HEMATITES.

A general description of the red hematite found in the carboniferous formation of Missouri has been given in the introduction to the third chapter of the present report. According to that description there are three varieties of this ore, namely: one, dull-red, soft, earthy, and frequently coarsely porous to spongy, and uneven in the fracture; another, dark, bluish gray, sometimes with a slight submetallic lustre, hard, dense, and compact, with very fine grain, and with a subconchoidal fracture; a third, coarse grained to oölitic, the grains being of the compact ore, and surrounded and cemented by the earthy ore. The earthy variety is the most common. The other two varieties have only been observed in Callaway County. The earthy ore sometimes encloses spathic iron-ore. When exposed to atmospheric influences it is altered into limonite.

Callaway County Hematites.—The red hematites in the sub-carboniferous strata of Callaway County occur in the most variable forms. The "Old Digging" and "Murphy's Hill," five miles south-east of New Bloomfield, contain some handsome, coarsely oölitic ore, besides the earthy hematite.

The ores from the vicinity of New Bloomfield are mostly compact and heavy, some bluish gray, others dark red. They frequently show a thin stratification. They are also found in concretionary forms, and seem occasionally to pass into the earthy hematite. They sometimes enclose spirifera and other fossils. Some of the ore on the Henderson bank has an earthy or a finely oölitic structure.

The ore on the Knight bank, near Fulton, is in part dense and concretionary, in part earthy. The latter has a light-red color and streak, and is especially distinguished by the admixture of a large amount of spathic iron-ore in specks and seams. Concretions of this ore are hard, and sometimes reach the mineralogical hardness 6. Spongy hematites are rarely found in Callaway County.

The following analyses were made by Mr. F. Emmerton, of the Joliet Iron and Steel Works, and were put at my disposal through the kindness of Mr. A. B. Meeker, of Chicago:—

į N	EW BLOOMFI	ELD ORES.	KNIGHT BANK.
	Compact Variety.	Earthy Variety.	3.
Metallic Iron	63.87	61.17	53.00
Silica	5.80	8.63	
Phosphoric Acid	0.10	0.165	1.87
Sulphur	0.017	0.018	

The specimen 3 was of the red, earthy kind, mixed with spathic iron, as above mentioned. This ore is probably altered, and has taken up a considerable percentage of phosphorus, besides the carbonates. The analyses I and 2 show that the dense as well as the earthy varieties of these ores, when unaltered and free from carbonates, do not contain a large amount of injurious ingredients, and are very rich in iron.

Hematites in St. Clair and Henry Counties.—The red hematites in the north-eastern corner of St. Clair County, and those in the southern and eastern portions of Henry County, seem to belong exclusively to the soft, earthy, and spongy variety, and are very frequently and extensively altered into soft, spongy, brown or yellow limonites, as far as can be judged from the present appearance of the banks, none of which are as yet opened. The following analysis was made by Mr. A. A. Blair, of St. Louis, from a mixed sample of soft, spongy, partly red but mostly brown hematite, from the Marmaduke bank:—

84.02
3.08
0.861
0.171
10.98
×0.0-
58.81
0.376

This ore has a high percentage of metallic iron, but is not very pure in regard to sulphur and phosphorus. As the specimen had

to be taken from the surface, where the ore had been in contact with the soil and its vegetation for a long period, and as it was almost entirely altered into limonite, it may be expected that the ore to be found in the hematite banks of this district, when opened, will contain a much smaller quantity of these substances, and will probably also prove richer yet in iron.

c. LIMONITES.

I have given a general description of the Missouri limonites in the introduction to the present chapter (III.), to which description I must here refer. From that it would seem that two different varieties could be distinguished, the porous and the stalactitic. This is, however, not really the case. Indeed, if we inspect those limonites, which apparently consist of one coherent though porous mass, more closely and more carefully through a magnifying glass. we find them almost invariably composed of single, but denselypacked, round, stalactitic columns, and we find also that the pores, which are seen by the naked eye, are generally interstices left between such stalactites, or between less regular mammillary or botryoidal forms. These pores are distinguished from those found in the specular ores of central Missouri, by sharper and smoother outlines, curved toward the inside of the cavity, thus forming very sharp angles, which point, not in two opposite directions only, but always in several directions. We may therefore say, in general, that all the Missouri limonites, with the exception of some of those produced by transformation of specular ores, are of stalactitic structure and origin.

The appearance of the pores and cavities does not, however, present a perfectly reliable means by which the original limonites might be distinguished, in all instances, from those produced by the alteration of specular ores, for two reasons: first, because the specular ores occur themselves not unfrequently in stalactitic forms, similar to those of the original limonites; second, because both kinds of limonites, and especially those produced by alteration, often lose their original structure entirely, through the influence of atmospheric agencies, both kinds becoming thereby either irregularly streaky or more or less spongy.

Ores in the South-eastern Limonite District.—Several limonite deposits, situated along the Mississippi River, have been favorably

reported on by Dr. B. F. Shumard, in his geological reports on St. Genevieve, Perry, and Cape Girardeau Counties, which reports will be found in the second volume of the present (third) Annual Report of the Geological Survey of Missouri.

I have inspected but one locality in this district, namely, the Collins bank in Perry County. The limonite which occurs there is very dense, hard, and apparently silicious, and is frequently mixed with fine chert-breccia. It is also found as an impregnation of sandstone. The ores in the vicinity of Irondale, Washington County, are in part hard and botryoidal or reniform, in part soft and ochrey. They sometimes contain splendent quartz in seams.

The richest part of the south-eastern limonite district is in the southern portion of Iron, Madison, and Bollinger Counties, and in Wayne County. Most of the limonites, which are found abundantly in these counties, are hard, compact, and massive, yet showing distinctly their stalactitic origin in the structure. The Cornwall limonites are softer and more ochrey than the others, but pretty free from foreign mechanical admixtures.

The ores found east of Marquand and Marble Hill are hard, and in many places very pure, while in others they seem to be more silicious, and enclose fragments of white chert, and then resemble somewhat the above-described limonites of Perry County.

The following analyses will show the chemical composition of the south-eastern limonites:—

	I.	2.	3.
Peroxide of Iron	72.58	81.40	80.98
Silica	5.84	3.01	1.98
Sulphur	0.17	0.07	0.00
Water	14.96	11.78	
Metallic Iron	50.81	56.98	56.68
Phosphorus	0.34	0.15	0.123

The two first analyses were made by Dr. August Wendel, of the Bessemer Steel Works, Troy, N. Y., the third by Messrs. Chauvenet & Blair, of St. Louis. I owe the latter to the kindness of Mr. H. S. Reed, of St. Louis.

Sample I was an average sample taken at the Ford bank, near Cornwall, and consisted of one-third hard limonite, and two-thirds soft limonite, mixed with some ochre.

Sample 2 was a hard, dense ore of stalactitic structure, from the Francis bank, 6 miles south of Marble Hill.

Sample 3 was a hard limonite, from the vicinity of Marble Hill.

We see that these ores, although less rich and less pure than the specular and red ores, are, however, good and valuable. We also see that the hard limonites are purer than the soft, ochrey ones.

Franklin County Limonites.—The Moselle limonites, and the brown ores south of Stanton, have mostly a dark color, and are partly hard and dense, partly soft and not unlike a fine sponge. A few banks in the vicinity of the Moselle Iron Works are distinguished by the reniform structure of their ores. These are commonly called "kidney-ores." The single kidneys are sometimes several inches in diameter, and have walls, half an inch thick, of very hard, dark-brown limonite, passing, toward the outside, into an ochrey clay. The hollow space inside these kidneys frequently contains rounded pieces of a very fine, hardened clay of white or yellowish color.

Admixtures of chert are not often seen in this district. But in some localities heavy-spar is mixed with the ore. Also pseudomorphs after pyrites occur, some of which contain undecomposed masses of this mineral in the centre.

The following analysis was made by Dr. August Wendel, of Troy, from a specimen of hard, dark-brown limonite, containing fine, irregular pores, evenly distributed, from the Bowlen bank, south-east of Moselle:—

Peroxide of Iron	81.38
Silica	2.88
Sulphur	0.13
Water	11.70
Metallic Iron	56 07
Phosphorus	

This analysis has a great resemblance with that of the Francis bank ore, as given above.

A good specimen from the Blanton limonite bank, 10 miles south of Stanton, was analyzed by Mr. A. A. Blair, and contained—

Peroxide of	Iron.					 •			er 0	84.16
or Metallic	Iron.	 	٠,		 					58.91

Limonites in the Central Ore-region.—A look on the ore-bank map which accompanies this report will show that the central ore-region contains a number of limonite deposits, besides its numerous deposits of specular ore. These limonites resemble those of Franklin County very closely. Some such ores in the Steelville district are very clean and uniform, while some on the tributaries of the Upper Meramec are occasionally mixed with white chert. Heavy-spar has not been found with the limonites of this region. Also pyrites occurs rarely.

Limonites on the Osage River.-The western ore-region of Missouri, on the Osage River, contains almost exclusively limonites. Those on the Lower and Middle Osage River, near Tuscumbia, Linn Creek, Boulinger Creek, and Warsaw, are mostly fine, pure ores of medium hardness and of a very favorable degree of porosity. Their structure is very distinctly stalactitic, and true stalactites are frequently met with on some banks in considerable quantities; as, for instance, on the Indian Creek, on the Elm Hollow, and on some banks near Tuscumbia. These stalactites are generally round in section, long and thin. Their diameter varies from $\frac{1}{16}$ to $\frac{1}{2}$ inch. They are always massive, and are formed of a dense and hard, grayish-brown ore. They do not show a crystalline structure in their fracture; but the surface is almost invariably covered with small pseudomorphous crystals after pyrites or after marcasites. In some instances these crystals are larger, up to 1/4 inch in diameter, and constitute the main portion of the stalactitic individuals, which then appear externally to be altogether composed of aggregated crystals. The single stalactites are themselves aggregated either in bunches or in wavy sheets. The inside of bunches sometimes consists of sulphide of iron; in most cases, however, they are entirely changed into limonite.

That ore, which forms large coherent bodies, and which is much more common than the true stalactites, is also of stalactitic origin, to judge from the shape and distribution of its pores and cavities, and was undoubtedly deposited in caves, from solutions which have been infiltrated from above.

The Osage ores are mostly free from foreign matter. In a few

localities only, chert is found mixed with the ore. The walls of the small cavities are generally covered with a layer of brown and yellow ochre. Larger masses of ochre have not been discovered.

While the ores on the Lower and Middle Osage, which I have just described, occur on the Silurian limestones, the limonites on the Upper Osage, above Warsaw, are found on the subcarboniferous limestones, and are also somewhat different mineralogically. They have a very dark, sometimes nearly black, color. The stalactitic structure is less common and less distinct than many of the other limonites occurring on limestones in Missouri. On the other hand, botryoidal and mammillary forms, consisting of numerous thin and parallel layers of dark-brown ore, are very frequent. The single layers peel off from each other easily, and are sometimes naturally separated from each other by fine, empty fissures with rough surfaces. The massy ore, which is more common, is often spongy in the fracture, and is mixed irregularly with small botryoidal masses and seams of a soft but splendent limonite, of nearly black color and of vitreo-metallic lustre.

The chemical composition of the Osage limonites will be seen from the following analyses, made by Mr. Andrew A. Blair, of St. Louis. The first analysis was kindly furnished to me by Mr. M. S. Cartter, of St. Louis:—

	Lower Osage.	M	Iiddle Osage.	Upper Osage.	
	I.	2.	3.	4.	5.
Peroxide of Iron	67.07	82.02	84.10		77.42
Manganese					0.00
Silica		5.13	3.59		8.05
Phosphoric Acid		0.091	0.077	0.084	0.076
Sulphur		0.015	0.00	0.084	0.147
Water		12.80	11.60		12.49
Insoluble matter	14.27		• • • • •		
Metallic Iron	46.95	57.41	58.87		54.19
Phosphorus		0.041	0.034	0.037	0.034

- 1. Soft, earthy limonite, from the Laclede bank, near Tuscumbia.
- 2. Hard limonite, with a stalactitic structure, the pores filled with yellow ochre; from the White bank, near Boulinger Creek.

- 3. Loose pipe-ore, broken stalactites, from the Indian Creek bank, near Warsaw.
- 4. Stalactitic aggregate of pseudomorphous crystals of limonite after marcasite, from the Elm Hollow bank, near Warsaw.
- 5. Average sample of ore from the Sheldon bank, on Bear Creek; in part a moderately hard, brown limonite, containing some yellow ochre in its pores; in part a very dark, nearly black, mammillary limonite; in part a spongy limonite, with dark-brown, submetallic seams.

All these ores are rich in iron, and contain but little sulphur, while the percentage of phosphorus is so low that they can, in this respect, almost be compared with the specular ores. These limonites on the Osage River are indeed remarkable for their chemical purity. It is also apparent, from the above analyses, that there is no difference in this respect between the stalactites and the massy ore.

Analysis 4 was made for the purpose of investigating whether pseudomorphs after sulphides do not contain a considerable amount of sulphur. The result shows that such pseudomorphs may be as pure as the rest of the ore.

As a part of sample 5 looked nearly black, it was supposed that it might contain some oxide of manganese. The analysis, however, showed that this is not the case.

CHAPTER V.

IRON-ORES OF MISSOURI.

BY ADOLPH SCHMIDT, PH.D.

C. Modes of Occurrence and Descriptions of Deposits.

WE have seen in the preceding chapters, II. and III., that two principal mineralogical species of iron-ores are represented in Missouri—the hematite and the limonite—and that the hematite occurs in two very different and distinct varieties, the specular ore and the red hematite.

We have, moreover, seen in section A that the specular ore occurs either in porphyry or in sandstone; that the red hematite forms strata in the carboniferous system; and that the limonite is generally deposited on limestone. This gives us four classes of original ore-deposits. I have mentioned, in the same connection, that the deposits of specular ore in sandstone are very frequently found to be disturbed and broken, and altered in regard to their position. Some of them, according to their present appearance, seem to be broken up entirely, and their parts and fragments seem to be drifted some distance, and to be deposited a second time, either irregularly or in alternate layers, with the sandy and cherty detritus produced by the destruction of sandstones and limestones. Such is the present appearance of many specular-ore banks, as well as that of some deposits of red hematite and of limonite.

Very few deposits of this character, however, have been opened, and none of them are as yet sufficiently opened and worked to allow a clear insight into their structure and formation. It is not impossible, in some instances perhaps probable, that the working of such deposits will lead to the discovery of original deposits in the interior of the same hills on which the former appear, and that these will prove to be only outliers of such original deposits in their vicinity, and not to be beds of far-drifted ore. Their present appearances, however, do not generally indicate this, and we have to judge them

by what we can see, to avoid losing ourselves in bold suppositions and speculations. Therefore, as long as the internal structure of these deposits is not better known, we must place them in a separate class, as "drifted deposits."

Thus, for the purpose of a systematic description of the modes of occurrence of iron-ores in Missouri, we are led to assume the existence of the following eight kinds of deposits:—

- a. Deposits of specular ore in porphyry.
- b. Deposits of specular ore in sandstone.
- c. Disturbed deposits of specular ore.
- d. Drifted deposits of specular ore.
- e. Strata of red hematite.
- f. Disturbed or drifted deposits of red hematite.
- g. Deposits of limonite on limestone.
- h. Disturbed or drifted deposits of limonite.

a. DEPOSITS OF SPECULAR ORE IN PORPHYRY.

The deposits of specular ore in the porphyries of eastern Missouri, especially in St. François and Iron Counties, occur in the most varied sizes and shapes. There are very large deposits side by side with those scarcely workable. There are regular veins, as in Shepherd Mountain and Iron Mountain; there are regular beds, as in Pilot Knob and in some localities east of it; there are irregular deposits, some of which somewhat approach veins by their shape, as on Lewis Mountain; while others have proved to be isolated pockets, as on Hogan Mountain.

The principal and most common kind of porphyry of which the greater portion of the hills in the vicinity of the ore deposits, and in general most of the porphyry-hills in that whole region, are composed, is chocolate-brown to brownish-black in color, and contains numerous small feldspar-crystals, pretty evenly distributed in the matrix. These feldspar-crystals are generally not over one-eighth inch long and about one-half as wide. Some of them are colorless and transparent, others red and opaque, the latter being mostly thicker, without being longer than the former. The red crystals have the appearance of orthoclase, the transparent ones that of oligoclase.

These two kinds of feldspar-crystals occur sometimes separate, sometimes together. I will call this porphyry the brown or normal

porphyry. It is very hard and brittle. It breaks in thin and flat splinters with very sharp edges and with a subconchoidal fracture. It breaks very suddenly under the hammer, and the splinters fly off with great rapidity and vehemence.

This porphyry sometimes contains quartz either in light-gray grains or in transparent crystals, colorless or slightly yellow. Also specks of a green mineral, perhaps chlorite, and iron and copper pyrites occur in it. Large masses of a brown feldspathic rock of a dirty-brown color, containing no crystals, and being tougher and softer than the surrounding rock, are frequently met with in this normal porphyry. So are also smaller streaks and masses of red porphyry. The latter are of very variable and irregular shapes, and seem to be produced by accidental infiltrations which changed the color of the rock.

The brown or normal porphyry occurs in some localities flaked and banded, with black, brown, and red, parallel, undulating, thin stripes; the crystals cutting through the stripes in the matrix. But this porphyry is always massive, never stratified. By gradual decomposition the transparent feldspar turns opaque and white, the red feldspar light brown and yellow. When the whole mass of the porphyry is gradually decomposed by the dissolving action of the atmosphere, or of acid waters on its alkalies, it turns at first pale brown, then light red, then light yellow, and finally white. It gets at the same time more and more soft and friable, and is finally changed into a white or gray or light-yellow, loose, clayish mass, in which sometimes the original feldspar-crystals can be indistinctly recognized, the transparent ones as very white spots, the red ones as yellow spots of a darker shade than the surrounding clay.

Another kind of porphyry, which is less frequent than the former, but occurs in immense bodies, so that large portions of certain hills are composed of it, is the "red porphyry." This porphyry has a light flesh-red color, and contains generally no feldspar-crystals. In very rare instances, however, small, red crystals of a darker shade than the matrix are found in it, either single or irregularly distributed. This porphyry, when entirely intact, is nearly as hard and brittle as the normal porphyry, and breaks then with a subconchoidal fracture. It is, however, rarely found thus—probably owing to its being much more liable to be decomposed than the normal porphyry. As mostly found, it is much softer, rather tough in

breaking, and presents, when broken, an even to irregular fracture.

The red porphyry is very frequently inclined to assume a stratified appearance, and in several localities forms regular strata, apparently extending, in uniform thicknesses, over considerable areas. The thickness of the single layers varies from one-quarter of an inch to several inches. These stratified porphyries seem to contain sometimes transparent feldspar-crystals, though very rarely. But they are principally distinguished by the more frequent, though irregular, occurrence of quartz in grains and specks.

The red porphyry passes by gradual decomposition and under removal of the alkalies, at first into a peculiar, light-yellow, soft rock, somewhat resembling an uncrystalline limestone, and finally into a more or less white, loose, clayish mass. Such a clayish mass is the so-called "bluff" on the Iron Mountain. The red porphyry, at its lines of contact with the brown porphyry, either mixes with the latter irregularly, or passes into it gradually by taking up feld-spar-crystals and by assuming a darker color. Both porphyries must therefore be considered to be of a similar and simultaneous geological origin.

The red porphyry seems, however, to be in a certain relation, though not a very plain one, to the ore-deposits, and to be in a closer connection with them than the normal porphyry. The latter contains very seldom veins or seams or specks of ore, while all such are very common in the red porphyry. Large masses of red porphyry occur in close proximity to the ore-deposits on Iron Mountain, Pilot Knob, Shepherd Mountain, and Cedar Hill. Certain enclosures in the Iron Mountain veins, the distinct stratification of the ore-bed on Pilot Knob, and other facts to be mentioned hereafter, indicate that both these hills were originally composed of red porphyry.

It might be supposed from this that red porphyry owes its existence to a mere change of color produced by the same influences which caused the deposition of the ore. This is, however, not so; for the red porphyry, besides being somewhat different lithologically, occurs very frequently without ore, and veins and seams of ore do occur sometimes in the brown porphyry without producing in the latter a decided and uniform change of color. We also see from this that the existence of ore-deposits, although evidently favoring the vicinity of the red porphyry, is by no means dependent upon its immediate proximity.

The question, how iron-ore deposits like those which are found in such variable shapes in these porphyries may have been formed, is one which is treated in a general way in all manuals of geology, and more specially in numerous other geological publications. I will, however, say a few words on this subject, with special reference to these East-Missouri deposits.

It seems to me, in the first place, that these deposits of specular ore, being all of a very similar mineralogical character, being all associated with the same kind of rock, and all situated within a small area of territory, must have been produced by one and the same kind of geological action, although the chemical action may have been different in different localities. Regarding this geological action, we may consider whether the specular ores can have been brought into their present places and shapes by injection in a melted condition, or by distillation in a gaseous state, or by segregation from the adjacent rocks, or by infiltration of chalybeate waters and springs and precipitation from the same.

- I. The descriptions of the various deposits which I shall give hereafter will show the impossibility of supposing that the ores were injected in a melted condition. There is no sign of the action of very high heat on the associated rocks. All the enclosures found in the veins and beds of ore are of such a character that they would combine and smelt in a very short time, when in contact with such large masses of melted ore. The ore is found in very thin fissures, of such an extent that a force capable of filling them with melted ore would certainly have opened the fissures, and would have produced thicker veins. Smelted and chilled ore has a very different appearance from that existing in these deposits.
- 2. A distillation could rather be thought possible. But distillations of iron-salts occur only in volcanoes, and are then mixed with other distilled matters, and never form large deposits. Here, on the contrary, the deposits are very extensive, the ore very pure, and neither lavas nor any other volcanic rocks are found in the oreregion, nor does the configuration of the ground indicate the former presence of any kind of volcanic action.
- 3. These specular-ore deposits cannot be derived from segregation, by waters penetrating the adjacent porphyries, dissolving iron

out of their mass, and depositing it in the fissures; for, such an action would have altered these porphyries uniformly along the veins and beds, which is not found to be the case. Such segregation could, besides, hardly have produced such immense deposits as on Iron Mountain and Pilot Knob, and would certainly have caused a less uniform structure and frequently stalactitic forms.

4. The specular-ore deposits of eastern Missouri, therefore, must all have been formed by infiltration and precipitation from chalybeate or iron-bearing waters, similar to those which occur still in all parts of the world in the form of chalybeate springs, and are now forming ore-deposits in numerous localities.

All rocks of which the earth is composed are more or less penetrable by water; all are more or less broken or permeated by wide or narrow cracks and fissures; nearly all of them contain iron, some in very small, others in larger quantities. The iron can be dissolved and extracted from these rocks by certain solutions at certain temperatures. From the bottom of the sea, of the lakes, of the rivers, from the surface of the ground all over, waters continually filtrate into the rocks below. These waters all contain more or less carbonic acid and other substances which they take up from the atmospheric air and from the soil. As they descend they dissolve various other matters under various circumstances, from the rocks through which they flow. They follow the easiest and widest channels. They grow warm, and sometimes hot, partly by the natural warmth of the rocks, partly through the heat produced by chemical reactions. The higher temperature and the higher pressure increase their capacity for dissolving mineral matters, with which they become charged as much as the existing circumstances allow. They may contain chlorides, sulphates, silicates, carbonates; they may contain silica, alumina, alkalies, lime; they may contain zinc, lead, iron, etc.

When these solutions have reached a sufficiently high temperature, and happen to find sufficiently easy channels upward, they will rise through such channels, driven by the pressure of the colder and therefore heavier solutions which follow them, and frequently assisted by the development of gases through chemical reactions. In this course upward they will again follow the preëxisting natural channels, wide fissures, small cracks, irregular holes and pockets. When they reach porous or loose strata, they will pene-

trate and impregnate them. When they come in contact with strata of materials which they are apt to decompose chemically at the existing temperature, they will alter, transform, metamorphose them. When these materials are such as to precipitate oxides of metals from the solution, these oxides will be precipitated and ore-deposits will be formed. The same effect may result when solutions of different chemical composition meet. The same effect must result when such metallic solutions approach the surface, where their pressure and temperature, and therefore their dissolving capacity, is diminished or altered.

As the circulation of waters and watery solutions just described, although locally variable, has existed during the whole geological history of our globe, as it exists still, it is evident that the largest fissures and cavities, when kept filled with however dilute yet continuously renewed metallic solutions for hundreds and thousands of years, under otherwise favorable conditions, will finally become filled with deposits of ores.

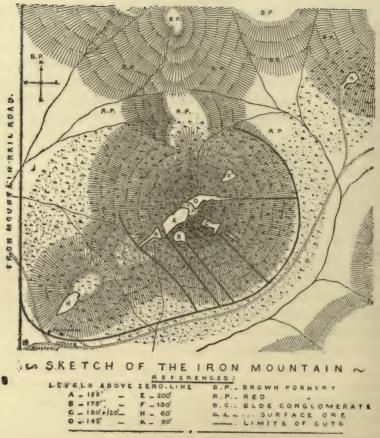
It also appears evident from the above, that the same mineral solutions can, under different local conditions, produce very different kinds of deposits—veins in one place, pockets in another, beds in a third. I have no doubt that the various deposits of specular ore in porphyry, which I will now proceed to describe, were formed in this wise.

Iron Mountain.—T. 35, R. 4, E., Sec. 31, north-east quarter, St. François County. The Iron Mountain is undoubtedly the largest ore-deposit in Missouri.

For the size and exterior description of the Iron Mountain hill, I may refer to page 75 of Prof. G. C. Swallow's Second Annual Report, where Dr. A. Litton mentions and describes it in a very lucid manner. This report was published in 1855, at which time the main part of the hill had not been opened, and no accurate knowledge could then be had of its internal geological structure. The openings which have been made meanwhile enable me to give a more detailed account of it.

The accompanying topographical sketch, Fig. 11, of the Iron Mountain and its surroundings, shows the surface-geology, the distribution of the porphyries, the extent of the surface-ore, and the position of the cuts or openings made by mining operations. The surrounding hills are composed mainly of the normal brown





Scale. 1800 ft & the inch

porphyry. A zone of red porphyry, frequently mixed with ore in larger and smaller seams and specks, runs along the north side of the Iron Mountain, across the ridge which connects the latter with the hill north of it. This zone of red porphyry thus separates the ore-deposit from the brown porphyry on the north side. The red porphyry in places becomes paler and takes a reddish-gray color. In the eastern portion of the zone the porphyry is half decomposed, forming a soft, yellow rock, which in one place, near its contact with the blue porphyry, is in distinct layers or strata from one to several inches thick, and has sometimes been taken for a limestone.

The red porphyry in the western portion of the zone seems to pass into the normal porphyry, in some places gradually, while in others it mixes irregularly with the latter, constituting a rock of a conglomeratic appearance and of a dull-brown or bluish color, with irregular flesh-red enclosures. This is the case in the place marked B C on the sketch.

The whole surface of the Iron Mountain itself is covered with surface-ore, which also extends over the south-western knob, called the Little Iron Mountain, and reaches into the valley south and west. and across the valley north-west of the mountain, and over a part of the slope of the opposite hill. This surface-ore, which, according to the preceding chapter of this report, is of the same character and composition as the vein-ore, occurs in more or less rounded bowlders and pieces, of very variable sizes, from a diameter of several feet down to a pretty fine sand, all irregularly mixed with each other, as well as with a fine, clayish or sandy detritus of a yellow or red color, and with single bowlders of half-decomposed porphyry. Pieces of a coarse-grained sandstone are sometimes found with it on the south-west slope. The surface-ore generally lies directly under a few inches of soil, and varies from one to five feet in thickness, which is, however, considerably exceeded in some places, especially on the south side, below cuts B and C, where it attains a thickness of 40 feet and over.

The main body of the hill, as far as opened, consists of a loose clayish mass, undoubtedly decomposed porphyry, known amongst the miners under the name of "bluff." This "bluff" and its origin have been described above. It is in the north-western half of the mountain, to the west of the principal ore-vein, of a very uniform and purely clayish character, while on the south-eastern half it has a more conglomeratic character. In this latter half, especially in the northern part of it, the "bluff" contains masses, which, from their partly preserved darker tint and from the numerous decomposed feldspar-crystals they contain, must be considered as being altered normal porphyry, while the north-western bluff seems to be exclusively altered red porphyry. A large mass of entirely fresh and unaltered, thickly stratified, red porphyry has been struck by the lower cut C; made on the east side through the "bluff." (See Fig. 6.) This mass, which is uncovered to a height of about 30 feet, contains numerous fine ore-seams, and evidently is not in its original position, as the stratification dips about 40° toward the

east. Large disturbed masses of unaltered brown porphyry have been struck in cut D, on the north-east side. (See Fig. 17.)

The whole immense clayish mass of decomposed porphyry or "bluff," forming certainly the upper part if not the whole of the Iron Mountain, is cut in two pretty nearly equal halves by an enormous vein of specular ore, from 40 to 60 feet thick, striking N. 53° E., which direction may be observed on the sketch by drawing a line through the cuts A E D F. Whether the main portion of this vein is in a vertical or somewhat inclined position cannot now be ascertained. It seems to be rather irregular in thickness and shape, to be split up in two branches for a part of its length, and to enclose large bodies of broken ore mixed with porphyry. This large and principal vein is called the "backbone" of the Iron Mountain.

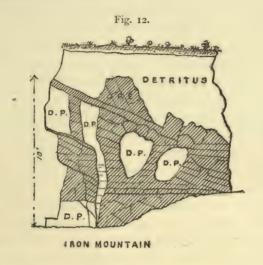
The "bluff" contains, however, besides the backbone-vein, numerous other veins of various and very irregular thicknesses, from less than one half-inch up to 6, and in places, 10 feet. These smaller veins cross the bluff in various directions, not subject to any definite rule. The limits between each of these veins and the "bluff" are very sharp, and there is nowhere a gradual transition from the ore into the "bluff."

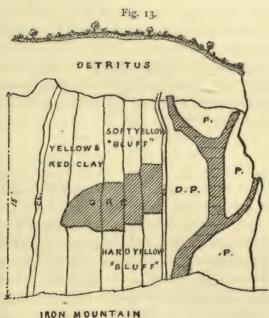
A line drawn through the cuts A E D F along the backbone, when prolonged in both directions, will pretty nearly touch the cut H on the Little Iron Mountain in the south-west, and the cut K on the hill across the valley in the north-east. As both these cuts have struck large bodies of ore, it seems probable that the principal vein extends over the whole distance from H across the hill to K, which is not much less than a mile.

The Iron Mountain ores have been described in the preceding chapter of my report.

To make the mode of occurrence of the specular ore in the Iron Mountain more plain, I will illustrate the above general description by a few sections taken in the different mining-cuts, adding explanations as far as required or desirable. The position of these cuts, as well as their elevation above the zero-line of the topographical survey, are given on the sketch, Fig. 11.

The cut H, near the blast-furnaces on the Little Iron Mountain, is not worked now; but much ore has been taken out of it and more seems to be left, especially in depth. This place is, however, at least near the surface, greatly disturbed, and the formation is to a great extent composed of displaced materials. The following two sketches present some interesting features:—





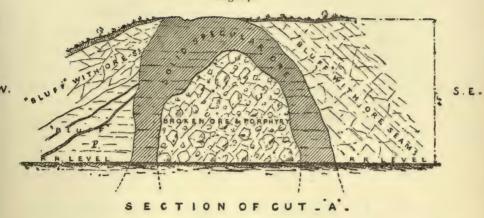
We see in Fig. 12 four irregular masses of decomposed porphyry, (D P) surrounded by formerly massive, but now broken, specular

ore. Such enclosed masses of rock are, however, rarely entirely surrounded on all sides by ore, although it looks so in the present section. In digging or blasting, perhaps only a few feet further, an entirely different section would present itself, and the apparently floating masses would be found to be in connection with, or supported by, other masses of the same nature. From the parallelism of the contour-lines of the three enclosed pieces, situated on the left side in the sketch, it must be supposed that they formed originally one mass, but were broken up and the interstices filled with ore. The breaking may have begun by the formation of thin cracks, produced by the contraction of the porphyric mass during its drying or cooling, or both. These thin cracks may have been widened, afterward, gradually by the crystallization of the ore. The jointed structure of the ore is very instructive. The arrangement of the joints shows that the ore has been formed round the preëxisting porphyry, and that the latter has offered a strong resistance to the contraction of the former, and consequently that the porphyry was still hard and fresh when the ore contracted, and that its decomposition took place afterward. Wherever there was an equal resistance on both sides, the ore separated in nearly parallel plates. the lower part, where no resistance existed, the ore contracted and separated into blocks of irregular shape but nearly equal size.

Fig. 13 represents an undoubtedly disturbed formation. portion of it which is to the right of the thin clay-seam (Cl) may have been formed as it is. We see here an upright, ramified orevein, having on the right side solid porphyry, on the left side, between the vein and the clay-seam, decomposed porphyry. The mass to the left of this clay-seam, between it and another thicker clay-seam, to be seen on the utmost left of the sketch, was evidently not formed in its present position, but must have slid into this position long after its formation. At the time when the thick mass of ore enclosed in this part was formed, the stratification in the porphyry was undoubtedly horizontal, while it is now vertical. When afterward brought in its present position, and when, owing to the gradual decomposition and consequent contraction and softening of the porphyry below, the ore lost its support, it broke off in plates corresponding to the porphyry-strata. It is not unlikely that its natural jointed structure has predisposed it to that effect.

Fig. 14 represents a cross-section through the backbone-vein in

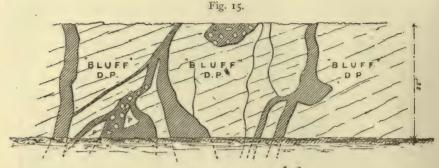
Fig. 14.



cut A, the working-level of which is about 60 feet below the summit of the Iron Mountain. The vein is here divided in two branches, from 12 to 18 feet thick each. These branches join above, enclosing a mass of broken ore and porphyry, mixed with quartz and apatite. This mass seems to be the product of destruction of numerous smaller ore-veins in porphyry, formerly existing in this same place or close by. The porphyry may have been broken up by the crystallization of the ore in its seams, and the ore by contraction and by the decomposition of the porphyry. All the pieces of ore have sharp edges and corners. All the porphyry is more or less decomposed.

P is slightly decomposed, but yet pretty hard, porphyry, passing into the loose, clayish "bluff" above it. The "bluff" on the southeast side of the vein is all a loose, soft clay of a yellow color. The surface of the big vein is wavy and very smooth. The small oreveins which cross the "bluff" in all directions vary from one-fourth to 3 inches in thickness, a few near P are up to 12 inches thick. The "bluff" does not contain any ore outside of these veins.

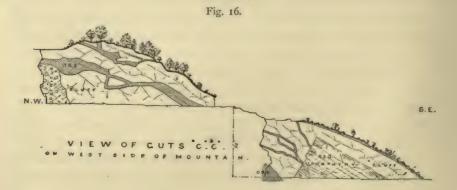
Fig. 15 gives a section of a characteristic part of cut B. This section shows in what a varied and often peculiar manner the smaller ore-veins cross the mass of the "bluff," enclosing larger and smaller pieces and blocks of decomposed porphyry, frequently changed into soft clay. Most of these veins strike about eastwest in this place. The ore has the same mineralogical character as that of the backbone-vein. It contains perhaps a little more



VIEW OF CUT .B.

quartz and more apatite, or crystalloid holes formerly filled by apatite. These holes are generally sitting on the walls of the veins. The ore resembles the surface-ore closely. The smaller the veins are, the more impurities they contain in proportion to the quantity of ore. The "bluff" is here very plainly a decomposed porphyry, sometimes imperfectly decomposed, in which case it is composed of a bluish-gray or bluish-brown matrix, enclosing white, decomposed feldspar-crystals. It seems from this appearance that a large part if not the whole of the "bluff" in this cut is derived from the normal porphyry.

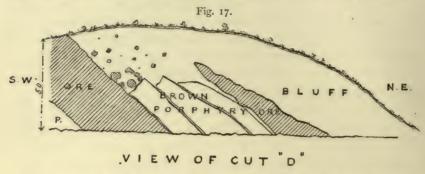
Fig. 16 represents a section through both the cuts C C, which lie about in one vertical plane, on the south-western slope of the Iron Mountain. The upper cut is about 60 feet, the lower one 120 feet, below the summit. The upper cut shows another characteristic section of veins running through the "bluff." What is marked as "detritus" is composed of blocks and pieces of colored clays



(decomposed porphyries), mixed with ore in pieces, the latter also somewhat softened by partial decomposition. Irregular holes and cracks in this mass are filled with yellow and red loam. This must be a part of a hole or crack which was opened after the complete solidification of both the porphyry and the ore, and was filled with broken porphyry and ore before the decomposition of the former. The section of the lower cut likewise presents some very peculiar features, indicating disturbances on a larger scale. An immense block, at least 35 feet in thickness, of thickly-stratified, red porphyry, lies here in the "bluff" in an inclined position, abruptly cutting off the ore-veins. This flesh-red porphyry is hard, and entirely fresh in its fracture and color, and encloses numerous oreseams, one-quarter to 3 inches thick. This block was evidently exposed to the influence of the solution that deposited the ore, but not exposed to those influences which produced the decomposition of the rest of the porphyry. To explain this satisfactorily, we must suppose that this decomposition was not effected by atmospheric influence exclusively, but that it was effected partly, or at least prepared and begun, by acid solutions, and that this block would have occupied an isolated or elevated position, while this, perhaps, preparatory action took place, and that it could not then be reached by the solutions. Afterward this block may have fallen over on the slope of the hill, sunk into the "bluff" gradually, cutting off the ore-veins, and may finally have been covered by a layer of bluff-clay washed down from the hill. The somewhat irregular character of the "bluff" overlying it seems to support this theory; otherwise the strange position of this block could only be explained by assuming that it was raised from below after the decomposition of the overlying porphyry.

However this may be, the presence of this intact, red porphyry with ore-seams, proves that the geological action by which the veins were filled was not the same as that which decomposed the porphyry, but that the latter took place much later than the former. For it might be supposed that the decomposition of the porphyry and the deposition of the ore had been effected, if not by the same solution, at least by a simultaneous chemical action. Another fact, however, besides the above, speaks against these suppositions. We find on Iron Mountain no instance of a total or partial replacement of porphyry by iron-ore, which replacement would have been the

unavoidable consequence of such a simultaneous chemical action, as I propose to show in my description of the Pilot Knob deposit.



In Fig. 17, which represents a section of cut D on the north-east slope of Iron Mountain, we find the brown or normal porphyry in a disturbed position, similar to that of the red porphyry in Fig. 16. We also see a cross-section of the north-east part of the backbone-vein, which is here in one solid mass, about 30 feet thick, and inclined toward the north under an angle of about 50 degrees.

P is a part of an immense solid mass of hard and intact normal porphyry, underlying the backbone-vein and being in close contact with it. This porphyry frequently contains specks and thin seams of green chlorite (?). Thin seams of ore also occur, though very rarely. The porphyry overlying the vein, in three flat pieces of a rather variable thickness from 3 to 10 feet, is very nearly of the same description. These three flats of porphyry, however, are not in close contact either with the backbone-vein, or with each other, or with the flat mass of ore that overlies them. All these masses lie loose over each other, being in contact at certain points only, while separated by clay-seams or empty spaces at other points. The "bluff" contains here no continuous ore-veins, but only single pieces of ore in such positions and so distributed as indicated in the above illustration.

This locality, as well as the whole north-eastern portion of the Iron Mountain, has evidently been subjected to great disturbances long after the formation of the ore and after the decomposition of the porphyry. The fact that the backbone-vein has not, so far, been struck by the cut F, which is situated on the line of its strike, and the presence of the above-mentioned conglomerates a little farther east, support this view, besides the appearance of cut D.

In the neighborhood of cut K, on the hill north-east of Iron Mountain, we find both the red and the normal porphyries. According to Dr. Litton's description (see Second Geol. Report, 1855) of a well bored near the furnaces, porphyries and large masses of ore exist there to a depth of more than 150 feet, overlaid by some magnesian limestone and sandstone. This shows that the porphyries are pre-Silurian, which fact is verified by numerous observations made in other localities.

When we look over all that has been said about the Iron Mountain, the geological history of this deposit naturally presents itself as follows:—

The whole Iron Mountain was composed originally of porphyries, which also filled the valley east and south of it.

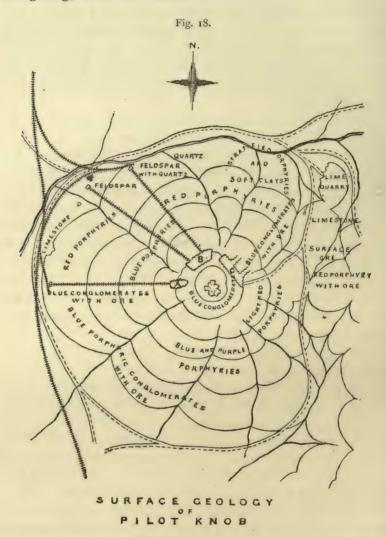
A great portion of these porphyries, especially on the north-west side, were of the red, the others of the brown or normal variety. These porphyries, either from the effects of contraction or from other causes, contained numerous large and small fissures. These fissures were kept filled with constantly renewed chalybeate waters for a very long period, during which these waters, through various chemical and physical influences, deposited the oxides of iron, which they contained in solution. The oxides of iron thus deposited were undoubtedly at first loose and soft, and mixed with water, but became denser and harder and less watery as their mass increased.

As the fissures were gradually filled, the access of the solutions became more difficult and more scarce, and was finally stopped. Then the ore dried in the veins, undergoing thereby a small contraction, which cracked and broke most of the veins without displacing their disconnected parts. After this had been done, the porphyry was acted on by atmospheric or other waters, probably containing carbonic acid, which decomposed the porphyry, removing the alkalies, and leaving a silicious clay. By this process these porphyric masses became so soft that rain and flood waters washed them off readily, the consequence of which was that, simultaneous with the erosion of the valleys, the cracked and disjointed ore-veins lost their support, and fell to the ground in single bowlders and pieces, thus forming the beds of surface-ore which now cover the slopes of the hill, and which fill a part of the now-eroded valleys.

Pilot Knob.—T. 34, R. 4, E., Sec. 29, Iron County.—For the

dimensions and the external appearance of Pilot Knob, I refer to Dr. A. Litton's description, given on page 79 of the "Second Annual Report of the Geological Survey of Missouri."

I give hereby, in Fig. 18, a topographical plan, showing the surface-geology of the Pilot Knob:—



Although the surface-geology does not always give perfectly reliable indications regarding the interior geological structure of a mountain, it generally allows us to draw certain valuable conclusions. The ore on Pilot Knob is not in veins, but forms a regular

bed in the porphyry. The top of the Pilot Knob, according to our sketch, is composed of "blue conglomerates." These consist of a dark, bluish-gray, porphyric matrix, enclosing large and small, but mostly sharp-edged, pieces of a light-gray, or reddish-gray to reddish-brown, porphyry. No distinct feldspar-crystals are visible in it. But the blue matrix contains numerous small, almost microscopic, crystals of iron-ore, more or less equally distributed through its mass. These conglomerates are all strongly magnetic with polarity. They have frequently a distinct though wavy stratification.

They form large groups of rocks on the summit, and compose the upper part of the mountain itself, directly overlying the ore-bed, into which they pass by degrees, becoming more and more impregnated with ore and mixing with ore irregularly above the bed.

They also lose their conglomeratic character in this direction, and a few feet above the ore-bed constitute a uniform bluish-gray porphyry, strongly impregnated with ore, and containing thin layers of a fine conglomerate.

Their maximum thickness, measured to the top of the rocks, may be estimated at an average of about 100 feet; that of the ore-bed at about 40 feet. Immediately below the ore-bed we find the same uniform bluish-gray porphyry, which directly overlies it, also mixed with small ore-crystals, although in a less number. These "blue porphyries" are likewise found on the surface for some distance down the slope of the hill, as is indicated on the above sketch. All of them contain a little ore in very minute, isolated crystals, not often visible to the naked eye.

A B and C represent mining excavations or cuts made in the orebed itself, but cutting also through the porphyry and conglomerate above it.

Lower down on the slopes of the hill, especially on the northern and eastern sides, we find the red porphyries cropping out in masses of such a size and position as to leave very little doubt that a large portion if not the whole of the middle and lower parts of the hill must be composed of them. They are mostly hard and unaltered, and correspond in their appearance to the general description given above, containing but rarely admixtures of feldsparcrystals or of quartz.

Such is the distribution of the rocks on Pilot Knob. We notice, however, two streaks of "blue conglomerates with ore," running

down the hill, one on the north-east and one on the west side. These conglomerates, with specks and impregnations of crystalline ore, are similar to those found on the summit. They are all broken up, in pieces and bowlders, partly decomposed, and are evidently washed down from above. This seems to have taken place on quite a large scale on the west and south-west sides, where these broken and either partly or entirely decomposed masses are spread in considerable thickness over a large area, and reach down to the foot of the hill. The ascent of the Pilot Knob is much less steep there than on the northern and eastern slopes. These loose masses are to a great extent altered into white or yellow clay. Several shafts have been sunk into them to a depth of near 70 feet without reaching the solid rock. We find another, though smaller accumulation of broken and decomposed materials along the foot of the hill on the north side. There we find also feldspathic rock, feldspar mixed with quartz, quartz without admixture, and pieces of stratified porphyry.

In a few places at the foot of the north-eastern and north-western slopes we find deposits of magnesian limestone, apparently overlying the porphyry.

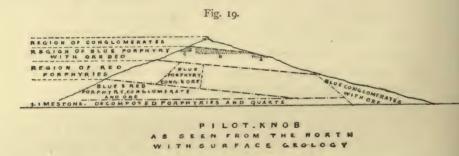


Fig. 19, giving a side-elevation of the Pilot Knob from the north side, with the surface-geology, will convey a clearer idea of the distribution of rocks over the hill, as described. It shows the red porphyries spreading over the lower two-thirds of the hill, but partly covered by decomposed materials and by loose, broken porphyries and conglomerates. Higher up it shows the region of the blue porphyries, with the mining-cuts A B and C, which indicate the position of the ore-bed. It finally shows the conglomerates on the summit.

It must be remarked that this sketch does not represent a sec-

tion, but a view of the hill. It is, however, probable that it would represent a pretty correct section through the Pilot Knob from east to west, if we would mark the whole mass below the "region of blue porphyry" as "red porphyry," leaving off the mention of the various detrital materials, which probably only cover the surface to a certain depth.

The ore-bed, and the stratified conglomerates above it, dip toward the south-west from 13 to 22 degrees, according to Professor Pumpelly, who also found the strike to be S. 50° E.

I will add here some sections through the ore-bed, as they appear in the three mining-cuts A B and C on Pilot Knob, to illustrate more fully the mode of occurrence of the ore in this deposit. The ore itself has been described in section B of this report.

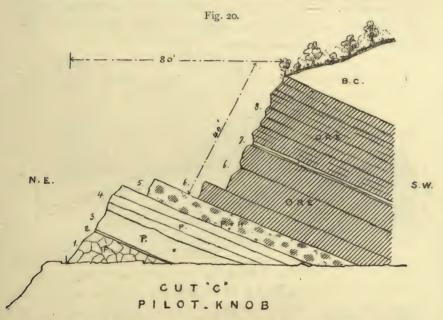


Fig. 20 is a cross-section through the highest and most eastern cut C made in the upper outcrop of the ore-bed. This section shows the following consecutive strata:—

1. Broken-up porphyric mass, without stratification, containing numerous irregular fissures. Its color is bluish gray. It encloses flakes of green serpentine, and small masses of white clay and of half-decomposed red porphyry. The thickness, as far as uncovered, is 15 feet.

- 2. Seam of hard, blue porphyry. 3 inches thick.
- 3. Compact porphyric mass, partly bluish gray, partly reddish brown, containing some serpentine in fissures. 10 feet. Between 3 and 4 is a large empty fissure, parallel to the strata, evidently produced by a sliding motion of layer 4 over layer 3.
- 4. Bluish-gray porphyry, distinctly stratified. Strata parallel to those of the ore. 8 feet.
- 5. Dark-blue, nearly black porphyry, mixed with ore irregularly. 4 feet.
 - 6. Good-looking though silicious specular ore. 16 feet.
- 7. Is composed of three parts, namely, a 1-foot stratum of hard, red porphyry without ore, a 1-2 feet stratum of ore, interstratified with red porphyry in very thin layers, and a 1-inch seam of light-gray clay-slate.
- 8. Stratified ore, good strata alternating with others which are intimately mixed with fine quartz. 22 feet.

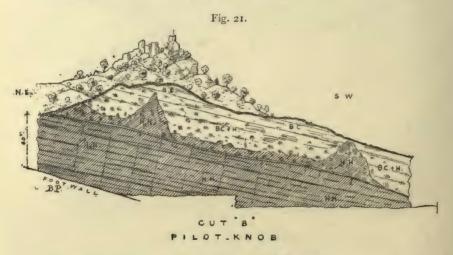


Fig. 21 represents the main cut B, which is about 420 feet above the valley west of the knob, and 160 feet below the highest point. This figure shows the foot-wall, consisting of a slightly ferriferous, blue porphyry; the main ore-bed (H H) nearly 40 feet thick, consisting of somewhat silicious specular ore, stratified more or less distinctly; a seam of light-gray clay-slate, varying from 6 to 18 inches in thickness, and containing no particles of ore whatever; another irregular layer of ore (H H) above the seam, mixing gradually with and passing into the blue conglomerate (B C).

The ore in the central part of the cut below the slate-seam is considerably softer than that either east or west of it. Analyses of these various ores have been given in section B.

If the slate-seam in cut B, Fig. 11, is the same as that in cut C, Fig. 10, as it appears to be, the thickness of the ore below the seam diminishes considerably toward the south-east, while the thickness of the ore above the seam increases in the same proportion, leaving the total thickness of the ore the same. The absence of ore in the slate-seam makes it probable that this seam is of later origin than the ore, being, perhaps, produced by a slide of the upper part of the bed over the lower part, whereby, through the irregularities of the surfaces, an empty fissure was left in places, similar to the empty fissure between layers 3 and 4 in Fig. 10, and was gradually filled up with a fine clay washed into it by surface-waters.

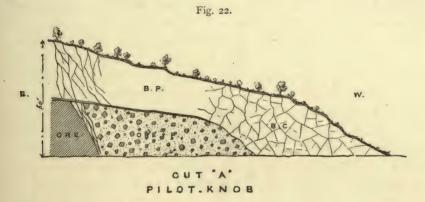


Fig. 22 is a section exposed by the mining-cut A, made on the lowest and western outcrop of the ore-bed, as may be seen on Figs. 8 and 9. We have here a disturbed locality before us. A portion of the ore-bed may have been underwashed here, broken off, and fallen down on the slope of the hill, the débris being then either carried away by floods or buried under the porphyric detritus. We see, therefore, in this cut an abrupt end of the ore-bed below the slate-seam. Above the ore and the seam, and close to the ore below the seam, is a nearly vertical streak of broken porphyric masses with some ore, permeated nearly vertically by numerous fissures mostly filled with red, yellow, and white clay or loam, washed into them from above. The porphyric parts are to a great extent converted into green steatite. Many pieces are thus altered

on the outside, while the inside is yet tolerably fresh porphyry. The broken ore is poor and silicious, similar to the less pure ore above the slate-seam in cut B. This whole irregular and mixed mass was evidently produced by a fall, and subjected, during a long period, to the influence of surface-waters. Adjoining it, below the seam, we find a breccia of ore, imbedded in more or less fine porphyric detritus, above the seam a mass of blue, ferriferous porphyry (B P), which is solid where the slate-seam is in its natural position, but broken up into a blue conglomerate (B C) where the slate-seam makes a sudden turn downward, indicating another break-down, produced by an underwashing of the stratum of blue porphyry.

The cut A, according to this description, presents two disturbances or falls, which have occurred at different times. The one affected the ore-bed and all the overlying strata; the second, of smaller extent, broke up a thick layer of porphyry only. The slate-seam runs across the first fall undisturbed, and must, therefore, have come into existence some time after the occurrence of the fall. As the latter produced ore-breccia, it must have occurred after the formation of the ore, from which it follows that the slate-seam is much younger than the ore-bed. The softness of the clay-slate that fills the seam corroborates this statement. The second fall in cut A breaks the slate-seam, and must therefore have taken place after the formation of the latter, and a long time after the occurrence of the first fall, and after the formation of the ore-bed.

As to the extent of the Pilot Knob deposit, we find, by throwing a look on the map, Fig. 8, that as far as opened at present, by the three cuts, A, B, and C, it seems to cover a triangular area, measuring 1,000 feet along the base, from east to west, and 600 feet in the height of the triangle. But the fact that the ore-bed dips in the south-western direction, at nearly the same angle as the surface of the ground, leaves the possibility of its extension over a considerable distance in that direction.

In regard to the geological and chemical action which may have created the ore-deposit on Pilot Knob, we must recall the introductory remarks on the formation of the specular-ore deposits of this region in general. Referring to this deposit specially, I fully agree with Professor Pumpelly, who, by more detailed and more thorough investigations, has come exactly to the same conclusion,

namely, that it has been formed by a gradual replacement of stratified porphyry by ore, effected by solutions similar to those which deposited the ore in the Iron Mountain and in the other places.

The stratification of the ore-bed, and of the impregnated and half-metamorphosed porphyries overlying it, is very plain and regular.

Stratified porphyries are found at the foot of the north-eastern slope of the knob, and very extensively in a large district east of it; but C, Fig. 20, shows a layer of red porphyry interstratified with ore.

The general appearance of the ore-bed, especially as shown in cut B, Fig. 21, as well as the appearance of numerous single specimens, and the partly impure and silicious character of the ore, naturally suggest the idea of an impregnation and gradual replacement of porphyry by ore, besides other circumstances mentioned in my introductory remarks.

Furthermore, such a replacement is not only possible, but it must be expected, at least partially, under certain circumstances.

When a solution of sulphate or chloride of iron, containing also carbonic acid, remains during a long time in contact with porphyry, the carbonic acid will decompose the porphyry, combining with its alkalies and dissolving them. The alkaline carbonates will almost simultaneously precipitate oxides of iron from the solution, and these oxides will fill the pores produced in the porphyry by the removal of the alkali. The silica is thereby set free, and will perhaps also to a small extent be dissolved, but the greater part of it will remain mixed with the ore, and make the latter silicious. The removal of the silicate of alumina contained in the porphyry is more difficult to explain. It has been shown, however, by Bischof, in his *Lehrb. d. chem. und phys. Geol.*, ed. 1863, vol. i. p. 86, that silicates of alumina can be decomposed by bicarbonate of iron in solution, and removed in the form of a soluble bisilicate of alumina and iron.

Considering, in our present case, that the solutions producing the transformation did not act perhaps on silicate of alumina previously existing as a free substance, but that they may be supposed to have acted on silicate of alumina which was undergoing at the same time its separation from the alkalies, and therefore, quasi in

statu nascendi, various other modes of decomposition or solution may be thought of.

This is not, however, the place to discuss such questions. My only purpose in touching them was to show the possibility of a chemical action, by which the formation of the Pilot Knob deposit by replacement may have, and undoubtedly has, taken place.

Another question regarding the formation of this deposit, is, whether the solutions from which the ore was precipitated were conveyed to this locality through fissures, and whether, therefore, veins of iron-ore will be likely to exist below or in close proximity to the ore-bed. It is very probable that the solutions were conveyed through either small or large fissures; but from this it does not follow necessarily that these fissures must contain deposits of ore; for the deposition of ores or of other substances, in fissures or cavities, is not alone dependent on the presence of the cavities and of the solutions which contain these substances. It is also dependent on various other circumstances, as temperature, unlimited or limited access of air, presence of other mineral solutions, facility of renewal of the matters to be deposed, rapidity of motion, and others. It is therefore not to be expected, with any degree of certainty, that cavities through which chalybeate waters flow, or were flowing, should contain deposits of iron-ore. Also a temporary deposition, and subsequent re-dissolution under altered circumstances, is possible and frequently met with.

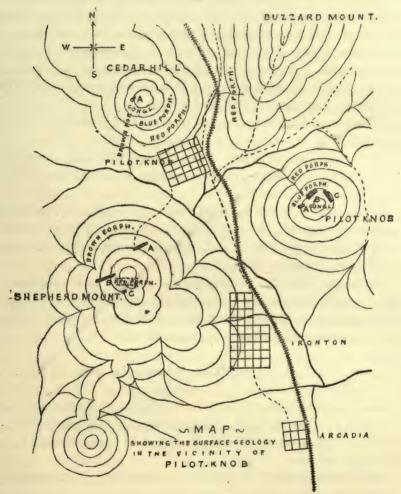
Thus, the existence of ore-veins below or close to the Pilot Knob bed is possible, but by no means certain.

Shepherd Mountain, T. 34, R. 4, E., Sec. 31, N. E. 1/4, Iron County.—This hill is named after Professor Forrest Shepherd, of St. Louis, who made the first investigations regarding the valuable ore-deposits it contains.

Dr. A. Litton, Professor at Washington University, St. Louis, has given an excellent description of the Shepherd Mountain, in Professor Swallow's second annual report of the Geological Survey of Missouri, to which description but little has to be added, because the mining operations which have been carried on there since that report was published, seventeen years ago, have proved the correctness of Dr. Litton's views, in nearly all their details. Although Shepherd Mountain is but little over one mile distant from the Pilot Knob, its ore-deposits are of an entirely different

character, being unquestionably veins, which, if they do not cut through the hill in its whole width, certainly extend over considerable distances. I give, in Fig. 23, a small map showing the relative position of the Pilot Knob, Shepherd Mountain, and Cedar Hill, also indicating the surface-geology of that district.

Fig. 23.



Shepherd Mountain is principally composed of normal porphyry, of a pretty uniform brown color, and containing either transparent or red and opaque feldspar-crystals, evenly distributed. This porphyry sometimes contains seams or small irregular masses of red

porphyry, which look as if produced by changes caused by infiltrations of some kind. The brown porphyry is in places beautifully banded, but never stratified. Some of it, on the upper part of the hill, is magnetic with distinct polarity, and is then found to be intimately mixed with single, microscopic particles of ore, absolutely invisible to the naked eye. The porphyry close to the ore-veins is considerably decomposed, quite soft, and mixed with large masses and veins of red and white clay. These clayish masses are sometimes impregnated with ore, or permeated by small veins of ore, which veins also penetrate more or less into the solid and intact porphyry. On the summit of the hill, south east of cut B, we find a zone of red porphyry, about 50 feet wide, running across the hill from east to west, as indicated in Fig. 13.

Three places have been opened on Shepherd Mountain, marked respectively A, B, and C on the map.

Cut A has disclosed a vein of specular ore, with some magnetite, 10 to 20 feet wide, nearly vertical, striking N. 62° E. (true). This cut is about 500 feet long.

Cut B on the west side has opened a length of about 1,000 feet of a quite similar vein, striking N. 68° E. (true). This vein, however, contains very little, if any, true magnetite, but is almost exclusively composed of specular ore.

The southern cut, C, is hardly opened enough to see its character. It looks thus far like a vein striking pretty nearly in the same direction as the other two veins, and containing a coarsely crystalline, specular ore, crossed by numerous thin seams of red porphyry.

It will be noticed that the strike of these veins is directed toward the Pilot Knob.

These deposits have undoubtedly originated in the same manner as those on Iron Mountain; but the surrounding porphyry has been very little altered, while most of that on Iron Mountain is entirely decomposed.

Smaller Deposits of Specular Ore in Porphyry, Cedar Hill, T. 34, R. 4, E., Sec. 30, Iron County.—The position of Cedar Hill can be seen on the map, Fig. 13. This map shows that the surface-geology, on the south-eastern slope of Cedar Hill, is very similar to that of Pilot Knob, the red porphyry at the foot passing into bluish and conglomeratic porphyries toward the summit. In the place marked A, however,

where a mine has been opened in summer, 1872, the porphyry is red again. This red porphyry, however, takes a bluish color in the proximity of the ore, and the walls of the ore-deposits are in some places composed of a blue, porphyric conglomerate, similar to that on Pilot Knob. This seems to indicate that the blue color, and perhaps also the conglomeratic to spherulitic structure, might be due to infiltrations of the same ferriferous solutions which deposited the ore.

The Cedar Hill is not sufficiently opened to allow a judgment regarding its ore-deposits. In September, 1872, about two acres of ground were freed from the soil, and showed two parallel, vein-like segregations, I to 4 feet in width, and about 100 feet in length, cropping out on the surface. Their strike was about 60–70° north-west. A shaft 10 feet deep was sunk on one of them, but showed it to be of an irregular section, so as to leave some uncertainty regarding its continuation in depth. These small segregations, surrounded by hard and solid porphyry, are not always workable in themselves, but they may be offshoots of some larger deposit, to the discovery of which they might lead.

Occurrences of specular ore were observed in some other localities in that vicinity, namely, in T. 34, R. 4, E., Sec. 18-19, 17-20, and T. 34, R. 4, E., Sec. 19, S. W. 1/4.

One and one-half miles east of Pilot Knob, in T. 34, R. 4, E., Sec. 28, N. ½, coarsely crystalline specular ore is found in seams through a red, banded, and stratified rock, of the appearance and fracture of jasper, overlying a series of stratified porphyries. The succession of strata which presents itself there, with a south-western dip, is from south-west to north-east, as follows:—

Slates of red, banded porphyry.

Stratified quartz-porphyry.

Slates of red porphyry.

Green porphyry.

Banded jasper.

Jasper, with specular ore.

Buford Hill, T. 35, R. 3, E., Sec. 26, two miles west of Iron Mountain, in Iron County, is a pretty high and steep hill, consisting principally of brown porphyry, occasionally mixed with a brown, jasper-like, feldspathic rock, and with light-red porphyry in some places. These three kinds of rocks seem to be mixed there

irregularly. Some micaceous iron-ore, mixed with quartz, is found in a depression between the hill itself and a low spur on its north-eastern side. Red porphyry occurs at a short distance south of it. The quartzeous ore is spread in large and small pieces over the ground, and extends along the slope of the hill over a distance of more than one hundred feet, in a line striking north-east to south-west. A little lower down on the same slope a streak of greenstone may be observed, running about parallel to the direction of the ore. This greenstone, which occurs in loose fragments, is fine grained, of a dark-green color, with white specks. It has the appearance of a diorite. It contains numerous black, microscopic crystals, of a metallic lustre, probably peroxide of iron.

The same kind of ore, accompanied by the same kind of greenstone, is said to exist in some places on the west side of the hill. Nowhere, however, are plain indications of the presence of larger masses of ore.

Buford Mountain, N. E. 1/4 Sec. 24, T. 33, R. 3, E., Iron County, contains a bed of manganiferous specular ore in decomposed porphyry, apparently of a not inconsiderable extent.

Big Bogg Mountain, S. E. 1/4 Sec. 13, T. 33, R. 3, E., the Russell No. 1 bank, E. 1/2 Sec. 3, T. 33, R. 3, E., and the Shut-in bank, N. 1/2 Sec. 2, T. 33, R. 4, E., all in Iron County, are also deposits of specular ore in porphyry, and have been mentioned and described by Dr. A. Litton in the second geological report.

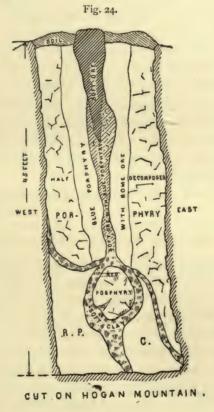
Lewis Mountain, S. ½ Sec. 6, T. 33, R. 4, E., Iron County, one and one-half miles south-west of Arcadia, is a small porphyry-hill, at the foot of which magnesian limestone is deposited in nearly horizontal strata. The slopes of the hill are covered with red and yellow clay and loam, 30 to 40 feet thick, mixed with half-decomposed pieces of porphyry. Above this is 2 to 4 feet of soil, mixed with rounded pieces of hard, specular ore, some pure, some mixed irregularly with brown porphyry in such a way as to indicate a gradual replacement of porphyry by ore. Indeed, in some specimens the brown porphyry, containing feldspar-crystals and brown quartz-grains, passes imperceptibly into specular ore, whereby the feldspar-crystals disappear, while the brown grains of quartz are preserved, and lie in exactly the same manner in the ore as in the porphyry.

An irregular vein, varying from I to 5 feet in thickness,

strikes across the hill about N. 75° W. This vein contains in some places specular ore, while in other places, where the vein is thinner, it is filled with red loam. This fact seems to show that the vein, which was originally of a more limited extent, has been opened more and extended by the crystallization of the ore, or by freezing water, and that the cracks thereby produced were afterward filled with loam washed into them from above. Some red, blue, and conglomeratic porphyries occur near the ore, and also small accumulations of micaceous ore-crystals, accompanied by quartz.

Cuthbertson bank, north-west quarter Sec. 19, T. 33, R. 4, E., and Ackhurst bank, south-west quarter Sec. 18, T. 33, R. 4, E., are deposits of manganiferous specular ores and magnetites and manganese-ores.

Hogan Mountain, south-east quarter Sec. 14, T. 33, R. 3, E.,



Iron County, contains irregular pockets of mostly soft, coarsely crystalline or micaceous, specular ore, in a peculiar, light-red, granular quartz-porphyry, which in the immediate vicinity of the ore is decomposed, and assumes a blue color when mixed with the ore.

Fig. 24 represents a section of such a pocket as opened by a cut. The massive, red porphyry on both sides of the cut is of the above description, and is covered by a dry soil, containing many pieces of broken porphyry. On both sides of the pocket we find the red porphyry half decomposed to a thickness of several feet. We further find, in immediate contact with the ore, a blue porphyry, mixed with some ore. The pocket itself is filled in its lower part with a blue porphyry conglomerate, strongly impregnated and mixed with ore, and in its upper part with pure, soft, crystalline, specular ore. This pocket, getting smaller below, runs out into a fissure, which splits in several branches, enclosing a large fragment of half-decomposed red porphyry, and a mass (C) of a gray conglomerate, mixed with pieces of a jasper-like porphyry and with quartz. The fissures are filled with a soft, loose, chloritic clay, undoubtedly a product of the decomposition of adjacent porphyries. R P is a dark, reddishbrown, very hard and massive porphyry.

It seems probable, from this section, that the solution which has deposited the ore has come from the side, or from above, out of fissures in porphyries which have since been destroyed and washed away, and that the cracks below the pocket were opened either by the crystallization of the ore or by the freezing of water, and afterward filled with fine clay.

There are several small deposits of a similar character on Hogan Mountain.

b. DEPOSITS OF SPECULAR ORE IN SANDSTONE.

The valuable and, in part, extensive deposits of specular ore in sandstone, the ores of which have been described in Chapter III. of the present report, occur in the eastern part of central Missouri, south of the Missouri River, especially in Crawford, Phelps, and Dent Counties, and constitute, together with the two following categories c. and d. of ore-banks, that iron-ore region in the State which in Chapter II. I have called the "Central Region." It has been mentioned there that many of these deposits have undergone great disturbances in their position. I shall describe under the present

head, b., only such deposits of specular ore in sandstone as are either entirely undisturbed, or which have been broken in place by being underwashed, or by a contraction of underlying strata or of their own mass, without subsequent removal of any great portion of their mass. Those deposits which were broken and separated into two or several parts, the single parts being shifted or moved to a greater or less distance, I shall describe under the two following heads, c. and d.

All these deposits belong to the Lower Silurian formation, and more especially to those strata which have been designated and described by Prof. G. C. Swallow, in the Second Annual Report of the Geological Survey of Missouri, page 125, as "Second Sandstone." Wherever I have been able to trace distinctly the geological position of these specular-ore banks, I have found them to be associated with this second sandstone, which has its place above the third and below the second "Magnesian Limestone." As the second sandstone is represented more or less extensively in the whole central part of southern Missouri, a great portion of which has been very little investigated as yet, we may hope that numerous other deposits will yet be discovered in the State, besides those to be described hereafter, and besides all those to be mentioned in our ore-bank list in Chapter V.

These deposits of specular ore have generally a lenticular shape, with either circular or elliptic outlines. They are frequently found in an inclined position, in which case they usually dip with the slope of the hill. Sometimes the ore is cut off abruptly at the outskirts, by nearly vertical walls, consisting of nearly vertical layers of clay, chert, and sandstone. In this case, these deposits appear like large, round, somewhat lenticular pockets in the sandstone, clad with layers of clay and chert, and filled with specular ore, which is often more or less altered into soft, red hematite. The thickness of these deposits is in the average about one fifth to onesixth of their average diameter. The ore is directly surrounded and underlaid by formerly continuous, but now broken and disconnected, strata of green or gray chert or flint, sometimes mixed with a fine, silicious, white clay, or with red loam. Below these chertlayers we find alternating strata of chert, sandstone, and of chertbreccia cemented by sandstone, sometimes continuous, but mostly broken. Below these are the regular strata of the second sandstone, running parallel with the above, and forming a circular or elliptic depression, in which the deposit lies. At the outskirts of this depression, where the sandstone strata suddenly change their nearly horizontal position, to curve downward and to run beneath the ore-deposit, the upper strata are frequently broken off, and form an annular outcrop round the deposit.

All these various rocks surrounding and underlying the specularore deposits, have in some cases their original and natural color and appearance, while in other cases they are ferruginous, or colored and impregnated by iron-ore, whereby the sandstone turns brown or black, and glittering with numerous fine ore-crystals throughout its mass, while the chert is colored green or red, and the clay or loam is transformed into a reddish-brown, sometimes pretty hard, ferruginous clay-rock.

In the Meramec bank we find regular layers of clay, chert, and sandstone, not only under but also above the ore-deposit. In most other places the ore is either covered with loose detritus, broken chert, and soil, or else it lies bare, in which latter case the specular ore is often changed into limonite near the surface.

The above description gives us the following two series of successive layers of rocks lying above, in, and under the specular-ore deposits in sandstone:—

Sandy and Cherty Soil.
 Sandstone with solid chert-layers. Same, impregnated with iron-ore. Sandstone, loam, and chert, broken and mixed. Solid chert. Broken chert and clay. Soft Red Hematite. Hard Specular Ore.
 White clay or red loam.
 Clay and broken chert.
 Broken sandstone and chert with layers of solid chert.
 Second, Lower Silurian, Sandstone.

Not all these strata are equally and invariably represented in all the deposits.

All the beds of rock thus associated with the ore-deposits seem to be in the same relative position in which they have originated, but to be in part broken, in part half-destroyed and altered.

It seems that these specular-ore deposits were originally formed in a lenticular shape, and imbedded in or on a sandstone containing layers of chert, and that they were afterward partially or wholly underwashed, some of the softer sandstone being thus removed, while the harder cherty parts and layers remained. In consequence of this action, a slight shifting of the whole mass of ore may have taken place, which somewhat crushed and mixed some of the underlying materials, and brought the deposit in a more or less inclined position. In some instances large caves, which are so common in all limestones, and which are undoubtedly formed by the dissolving action of acid waters, may have existed in the Third Magnesian Limestone, below the ore-deposits, and may have caused either a gradual or a sudden sinking, without which the origin of the pockets with almost vertical walls, in which such deposits are sometimes found, cannot easily be explained. The original lenticular masses of ore may have been formed either by deposition from chalybeate waters in depressions on the surface of the sandstone, and afterward covered by other strata, and condensed and altered by pressure and higher temperature, or else they may have come into existence by a gradual replacement of lenticular limestonedeposits formed in the above-described manner in the sandstone. While the first supposition appears as the more simple and as the more natural and intelligible one, the second one is supported by the two following facts:-

Irregular rounded masses of a very dense and hard orange-yellow, subcrystalline limestone, interspersed with pretty large gray crystals of carbonate of iron, are sometimes met with in the midst of the ore, and passing into the latter.

Mr. G. C. Broadhead found, in the lower coal-measures of Henry County, a stratum of limestone from 3 to 6 inches thick, which is partly converted into red hematite, containing no carbonic acid. This transformation has taken place from both the upper and the lower surfaces of the stratum, and has reached a thickness of three-quarters to one inch, from either surface, while the interior consists yet of the unaltered gray limestone.

Another fact, however, which speaks in a certain measure against the second of the above two suppositions, is this, that lenticular deposits of limestone have nowhere been observed in the Silurian sandstone of Missouri. I will not attempt to decide whether any such deposits have existed and have all been metamorphosed into ore-deposits, or whether the supposition is incorrect.

These undisturbed or slightly-disturbed deposits may be recognized by the following external characteristics:—

- 1. They mostly occupy a high topographical position, lying on summits of hills or of ridges if undisturbed, or on the upper part of slopes close to the summits if somewhat disturbed.
- 2. They have a circular or elliptical outline, inside of which the ground is covered all over with surface-ore of various sizes, partly specular, partly limonite, more or less rounded and smoothed, especially on the upper side, from being exposed to rain and storm. The larger the size of these pieces and bowlders on the surface, the more confidently a good deposit may be expected.
- 3. They are surrounded by annular outcrops of solid or broken red clay-rock, chert-breccia, black or brown impregnated sandstone, and finally of yellow or white sandstone. These annular sandstone outcrops are frequently very conspicuous.
- 4. The slope of the hill shows also, outside of these annular outcrops, streaks of smaller and more rounded surface-ore, evidently washed down from the original deposit, the main body of which always lies inside the sandstone outcrops. The surface-ore is sometimes spread over the whole hill. In other localities it is concentrated in depressions and ravines, the soil and sandstone being continually washed away, while the heavy ore is left and concentrated. In other not unfrequent instances the surface-ore on the slopes covers swellings of the ground, encompassed or cut by two or more diverging ravines. In such cases the surface-ore lies generally pretty thick and close, and thus protects the underlying softer materials from being washed away as rapidly as the less protected portions of the same slope. It is obvious that these swellings have been produced by such an unequal protection of the ground. When opened by shafts, these swellings are then found to consist of loose, sandy detritus, with little or no ore, overlying the solid sandstone. The extent of the surface-ore, however thick and close it may lie, when outside of the annular outcrops of sandstone, is therefore no proof of a corresponding extent of the deposit.

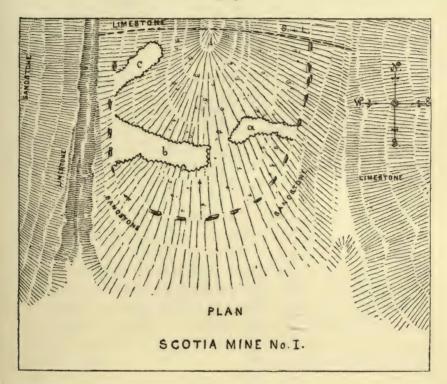
I will now proceed with the special description of a number of ore-banks belonging to this category of undisturbed or slightly-disturbed deposits of specular ore in sandstone.

Scotia No. 1, Sec. 1, E. 1/2 S. E. 1/4, T. 38, R. 3, W., Crawford County.

This deposit lies in a low sandstone-hill, which forms a spur on the southern end of a higher limestone-hill, and is separated by two

narrow valleys from the surrounding high hills, composed of Third Magnesian Limestone at the base, and of Second Sandstone in the upper part.

Fig. 25.

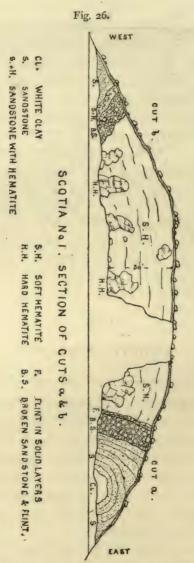


The Second Sandstone on which the ore is bedded occupies a much lower level than the Second Sandstone which caps the surrounding hills. It has undoubtedly sunk down gradually into its present location, which is at the level of the limestone, by which it is surrounded on three sides. The whole ore-bank sank down with the sandstone, and by its weight may have kept the latter in place, and protected it against destruction by the waters which effected the erosion of the valleys on both sides.

S and L, on Fig. 25, mean that pieces of sandstone and limestone are found there, mixed, on the surface. Distinct outcrops and openings of both these kinds of rocks are found on the other two hills, east and west of the ore-hill.

The broken line across the ore-hill shows the probable limit of

the ore. The annular outcrop of sandstone round the ore is also indicated. The surface-ore extends over an area about 180 feet wide and over 200 feet long. a, b, c, are cuts made for the purpose of mining the ore.



The foregoing, Fig. 26, represents a section through the cuts a and b, showing the interior structure of the upper part of the ore-deposit.

We see here nearly corresponding strata on both sides of the ore, all dipping toward and apparently under the ore. There is the Second Sandstone (S), yellowish-white, dipping about 40° on the west side, and considerably more on the east side. Next to this sandstone is, on the east side, a stratum of breccia of green, red, and yellow chert, mixed with pieces of sandstone, and cemented by red loam 8 feet thick (BS). This same stratum, with the same thickness, is represented on the west side; but it is there separated from the white sandstone by a stratum, 5 feet thick, of Sandstone impregnated with fine, greasy ore (S and H), probably produced by an accidental and local infiltration into the regular Second Then follows, on the east side, a stratum (F), 2 feet thick, of green and brown flint or chert, in solid and nearly unbroken but very irregular layers. These strata will undoubtedly meet below the level of the present section, and thus form a pocket, in which the ore is placed.

The ore itself consists of large, irregular masses of hard, blue, specular hematite (H H), getting more soft and light-colored outside, and passing into the soft, red hematite (S H), which surrounds them, and which constitutes the greater part of the bank, as far as opened at present. The soft, red ore is mostly greasy to the touch. It contains sometimes streaks of broken chert and of clay, and is in its upper part mixed with streaks and irregular masses of yellow ochre. All this ore must have been formerly one solid mass of specular ore, which was broken, and gradually softened, and subjected to such transformations as I have described in section B.

Scotia Bank No. 2, S.E. 1/4, Sec. 28, T. 39, R. 2, W., Crawford County.

This seems to belong to this category of ore-banks, although its exterior characteristics are not very plain, and although it is not sufficiently opened to allow an exact judgment regarding its character. To judge from the surrounding hills and from the surfacerocks, the hill seems to be composed of limestone capped by sandstone, or else of sandstone exclusively, and to contain a considerable mass of ore, situated above the sandstone on the summit of the hill. The ore is thickly covered by detritus on the north-east side, while it is but a few feet below the surface on the south-west side, near the summit. Its presence there has been proved both by a shaft and by a ditch, which are from 80 to 90 feet apart.

At the foot of the hill, which is about 150 feet below the summit, a tunnel was made, and struck immediately under the soil, the soft, red ore enclosing pieces and bowlders of specular ore, and numerous broken stalactites of specular, partly converted into red, ore. This ore, several feet thick, dips with the slope, and is underlaid conformably by a layer of green chert, 2 feet thick, which itself lies on a mass of broken chert and sandstone, mixed with clay and loam. All these materials that were struck by the tunnel are undoubtedly detached parts of the main ore-bank on the summit.

Cherry Valley, No. 1, E. ½ S. W. ¼, Sec. 4, T. 37, R. 3, W., Crawford County, 6 miles east of Steelville.

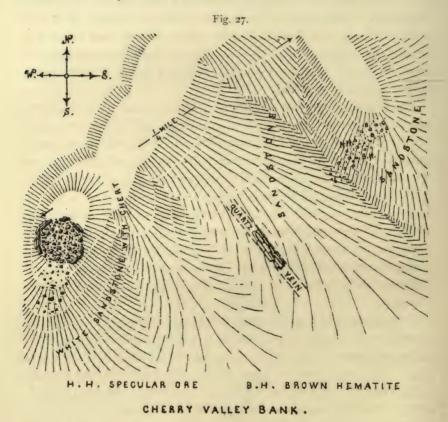


Fig. 27 gives a topographical sketch of the two Cherry Valley banks, neither of which is as yet opened. Nevertheless, the western or No. 1 bank will readily be recognized as a very distinct and characteristic example of a nearly undisturbed deposit of specular

ore in sandstone. The lower part of the hills in that region is composed of Third Magnesian Limestone, the upper part of Second Sandstone. On the summit we observe an annular outcrop, several feet thick, of white and yellow Second Sandstone, having in part the appearance of a vitreous quartzite, and dipping toward the centre, but so steep that the strata are in most places nearly in a vertical position. Inside of this outcrop of light-colored sandstone, and placed conformably to it, is an annular outcrop of a sandstone colored or impregnated by oxides of iron.

The circular space inside of these outcrops, 150 feet in diameter, is entirely covered with ore, the numerous large bowlders consisting principally of specular ore, while most of the smaller pieces are altered partly into limonite, partly into soft, red hematite. This space marks the position of the regular deposit, and a pocket of considerable depth, filled with ore, will certainly be disclosed here by future mining operations.

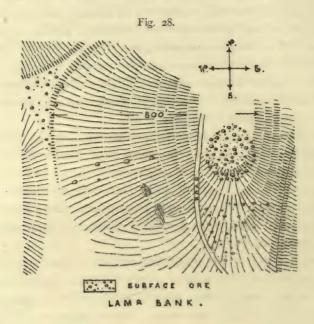
There is a gap in the sandstone outcrop on the south side, and there the surface-ore is spread in considerable quantity down the slope, outside the outcrops, in a streak 50 to 60 feet wide and about 200 feet long. The greater part of this surface-ore is changed into limonite. An extension of the underground deposit in this direction cannot, however, be expected.

The eastern or No. 2 Cherry Valley bank, which is sketched in Fig 27, does not show the exterior characteristics of an undisturbed bank, although it contains very large and very numerous bowlders of specular ore and of limonite on the surface, and although very large masses of ore will undoubtedly be found there underground, especially in the upper part of the hill. But it is, from its present appearance, a disturbed deposit, belonging to the third category C, of which I shall speak hereafter.

Lamb Bank, Sec. 35, T. 36, R. 6, W., Phelps County.—This bank is situated in the "Upper Meramec" district, on the dividing ridge between the Benton Creek Valley and the Norman Hollow, at the head of the western branch of Benton Creek.

The main part of the bank is situated close to the highest point, and is nearly round, 150 to 200 feet in diameter. No distinct annular outcrops are perceptible, however, and the limit of the body of massive ore can therefore not be determined with great accuracy. But an annular streak of ferruginous clay-rock and of

chert-breccia can be traced nearly all round the bank, from the loose pieces lying on the surface.



A horizontal outcrop of white sandstone is found half-way down the western slope, where also large bowlders of specular ore occur, as well as in the western ravine, which is about 130 feet below the bank.

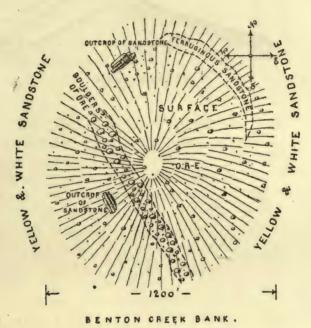
The low, triangular slope south of the circular bank is covered with small and rounded surface-ore, between the road and the little ravine on the east side. This ore was undoubtedly washed down from the main deposit.

Benton Creek Bank, Sec. 32, T. 36, R. 5, W., Crawford County.—On Benton Creek, in the "Upper Meramec" district.

The sketch (Fig. 29) shows a large hill, 150 to 200 feet high, covered with surface-ore, which is partly specular, partly limonite, partly strongly-impregnated sandstone. The surface-geology indicates yellow and white sandstone on the lower half of the hill, cropping out in several places on the slopes and dipping in each case toward the centre of the hill. On the north-west side we find a ferruginous or impregnated sandstone on the surface, extending in a curved streak round the hill. All this points toward the existence

of a large ore-deposit inside these sandstones. A peculiar feature of this bank is a straight zone of very large bowlders of specular ore running across the hill from north-west to south-east. This bank is somewhat disturbed, but it has preserved its circular character. The mass of ore will probably be found broken, but not scattered to a great extent.

Fig. 29.

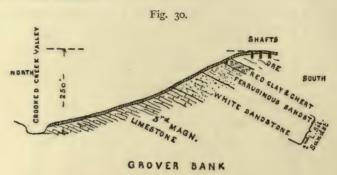


This is one of the largest ore-banks in central Missouri, judging from its appearance and dimensions.

Grover Bank, S. W. ¼, Sec. 2, & N. W. ¼, Sec. 11, T. 35, R. 4, W., Crawford County.—This ore-bank is situated in the "Upper Meramec" district, on the top of a high ridge, with pretty steep slopes, cut by numerous ravines, which descend gradually through lower ranges of hills into the broad valley of Crooked Creek.

The ore does not lie thick, either on the slopes or on the hill. It is more concentrated in the ravines. Fig. 30 presents an elevation, showing the various rocks met with in going from the Crooked

Creek valley up to the bank, namely, the Third Magnesian Limestone, the Second Sandstone, which becomes ferruginous near the bank, above this a thin streak of red clay with chert, and finally the ore on the summit.

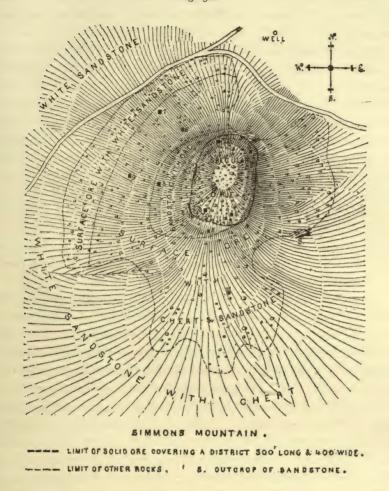


This succession of rocks and the situation of the bank seems to warrant the presence of a good ore-deposit, although the surface-ore is not very copious. Six small shafts have been dug on the top of the hill, five of which were too near the outcrop of the ore, and therefore, after cutting through 5 to 7 feet of soft red and of specular ore, struck either the underlying white clay or the chert-breccia or the impregnated sandstone. The sixth shaft was made nearer the central part of the summit, and struck soft, red hematite immediately below the soil, together with bowlders of specular ore up to one foot in diameter. This shaft was brought down six feet only in the ore, and then discontinued. The presence of a large amount of soft ore in this locality proves that the bank has been broken up and somewhat disturbed. But it is not likely that a considerable part of the original mass of ore should have been washed away.

Simmons Mountain, N. W. 1/4, Sec. 24, T. 34, R. 6, W., Dent County, 1/2 mile south-west of Salem.—This is one of the largest, if not the largest deposit of specular ore in the central ore-region. It received its name from its original owner, Mr. C. C. Simmons, of St. Louis.

Fig. 32 is a view of the Simmons Mountain, which is a nearly isolated hill about ninety feet high, above the plateau south of Salem, on which it is situated, and covering over thirty acres of ground.

Fig. 32.



The main body of the hill seems to be composed of Second Sandstone, which is found in pieces on the surface, and has been uncovered by a digging at the foot of the north-western slope, close to the road. The sandstone on the surface is mixed with pieces of chert on the southern and south-western sides, near the base. Higher up it is mixed with specular surface-ore, which extends over a very large district, increasing in frequency and size toward the summit.

Some of the surface-ore on the slopes is altered into a fine and pure limonite (brown hematite), but most of it is specular. The

latter occurs in bowlders, several feet in diameter. The following topographical sketch will give a better idea of the surface-geology.

We here notice, in addition to the occurrences just described, an elliptic district, about 400 feet wide and 500 feet long, enclosing the summit, and being very thickly covered with surface-ore. This is the position and extent of the original deposit. As may be seen on the sketch, it is surrounded by outcrops of sandstone (S), which are especially distinct on the north and west sides, and are ferruginous in several places. On the east side some outcrops of sandstone are found lower down the slope. The dip of the sandstone cannot now be distinctly recognized, but this rock will undoubtedly be found to form a large elliptic pocket, filled with ore. Inside of the upper sandstone outcrops, the surface-bowlders are of enormous size, evidently outcrops of an immense body of massive ore. Wherever the soil is removed between these bowlders, ore is found immediately below it.

Outside of this district, the surface-ore, although very large in places and very plentiful, must be considered as being broken off from the main deposit and thrown or washed down the hill. This ore may have been at first imbedded in large masses of detritus of sandstone which was broken off simultaneously with the ore. Afterward this ore was concentrated on the surface by the slow but unavoidable and merciless action of rain-water, which mechanically destroyed and removed the light sandy materials surrounding and underlying the ore, while the ore itself, being too heavy to be carried off by such action, remained in place. This outside surface-ore is therefore not indicative of the existence of large bodies of ore below it.

These views have been fully verified by a number of shafts which have lately been sunk on the Simmons Mountain, and which on our sketch are marked by the numbers 1 to 9. The shafts 5, 6, 7, 8, and 9, which are outside the elliptic district, disclosed 15 to 25 feet of loose, sandy detritus, and finally struck the solid sandstone. Shafts 5 and 6, which are the nearest to the deposit, met with more clayish materials, and streaks and masses of white clay and chert, which are so frequently found in close proximity to such deposits. The shafts 1, 2, 3, and 4, although sunk quite near the limits of the deposit, but inside of them, went through 25 to 30 feet of solid,

pure, specular ore, without reaching the foot-walls. I was lately informed that, since my last visit, one of these shafts has struck the clay at a depth of a little less than 30 feet. This is not at all astonishing, on account of the proximity of the shafts to the limits of the pocket. The fact that none of these shafts has reached the clay at a less depth, proves that the walls of the pocket are nearly vertical, and points to a great thickness of the ore in the central portion of the deposit. At the foot of the Simmons Mountain, and north of it, a well has been sunk, which is marked in Fig. 32. This well is over 60 feet deep. It passed through—

8—10 feet of soil and loose, sandy material.

6-7 feet of sandstone in broken layers.

15-18 feet of red, sandy loam.

6 feet of chert, in thick, broken layers.

6-8 feet of red, sandy loam.

3—4 feet of chert, in broken layers.

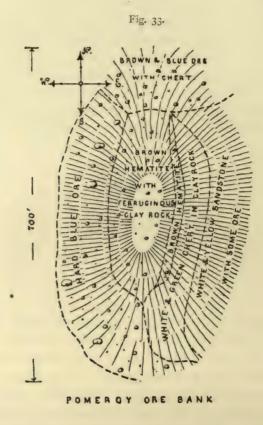
14 feet of chert, mixed with clay.

All the materials just mentioned seem to be remnants of destroyed sandstones which must have formerly surrounded the Simmons Mountain. The harder cherty strata have been evidently less subject to destruction, and have therefore been left in place and concentrated, while the greater part of the softer sandstone was destroyed and carried off by the waters and floods, which effected the erosion of the valleys in that region, besides washing away whole strata of rocks. The body of the Simmons Mountain was in a great measure protected against this action by the size and weight of its ore-deposit, which seems to be nearly undisturbed, resembling in this respect the Cherry Valley No. I and the Lamb banks above described.

Pomeroy Bank, Sec. 10, T. 34, R. 6 W., Dent County, three miles north-west of Salem.

This quite extensive ore-bank was evidently underwashed on the west side, and broken and turned or moved in that direction, as can be easily perceived from the study of the surface-geology on our sketch. The mass of the ore does not seem, however, to be much scattered, nor to have been removed to any considerable ex-

tent; so that we may rank this bank with the disturbed deposits (b).



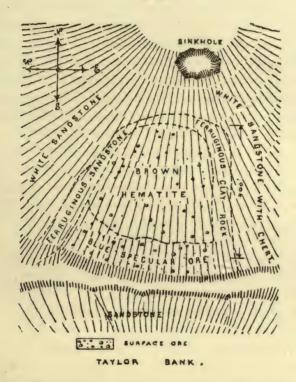
The top of the hill is about 120 feet above the eastern valley; but the surrounding hills are mostly higher than the Pomeroy hill.

In throwing a look on Fig. 33, we find that the principal mass of the surface-ore, although in large quantities and sizes, is here not situated on the summit, but on the western slope, where indeed, besides the cherty soil, hardly anything else but ore is seen on the surface. The summit is occupied by ferruginous clay-rock and pieces of ore altered into limonite. On the eastern slope we have a zone of the well-known breccia of white and green chert, cemented by clay-rock, and lower down the ordinary, white or yellow, Second Sandstone. No regular outcrops are to be seen; but

the succession of rocks from the east to the west, shows that a considerable mass of ore must exist in the western and central parts of the hill.

Taylor Bank, S. W. 1/4 of S. W. 1/4, Sec. 12, T. 34, R. 7, W., Dent County, eight miles north-west of Salem.

Fig. 34.

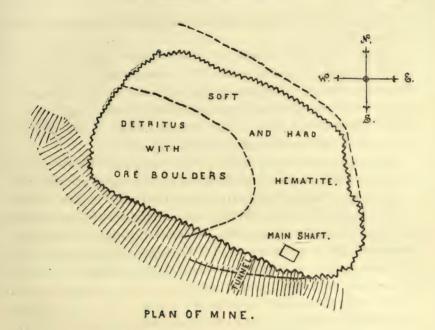


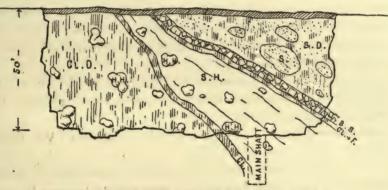
On this topographical and geological sketch of the Taylor bank we meet with circumstances very similar to those just described. We find about the same succession of rocks, and the whole orebank situated on the slope, the specular ore occupying the foot of the hill. The surface-ore extends over an area about 400 feet square. The main body of the ore will probably be found in the upper part of the semicircular space, which is surrounded by a zone of ferruginous rocks. This bank has been undoubtedly underwashed on the south side, and disturbed in its position. North

of the bank, near the summit of the hill, is one of the round sink-holes which so frequently occur on hills composed of Second Sandstone and Third Limestone. They are, perhaps, caused by the existence of large cavities in the limestone, which have caused a sinking of the overlying sandstone. This sink-hole does not at present seem to be in any connection with the ore-deposit.

Iron Ridge No. 1, N. E. 1/4, Sec. 29, T. 39, R. 5, W., Crawford County.—I give here a plan and section of the Iron Ridge mine. The ore-deposit seems to be of a lenticular shape, but curved horizontally, while dipping at an angle of about 40 degrees toward the east. The specular ore is all broken, and to a great extent altered into soft, red hematite, in which the remains of the specular ore are imbedded as half-converted bowlders. The deposit is entirely surrounded by loose materials, and has undergone considerable disturbances; but at the time when these took place the ore-bank must have been in a solid and intact state, because its limits are sharp and well marked, and the main body of the ore, although broken up interiorly, has not been separated into several smaller bodies, nor scattered about, as far as can be seen at present. The succession of rocks may be observed as follows:—

- I. Cherty and sandy soil. I-3 feet.
- 2. Clayish and sandy detritus, white, yellow, and light red, enclosing pieces of chert and chert-breccia in sandstone, and sometimes masses of soft sandstone. 40 feet.
- 3. Very hard breccia of sandstone, cemented by quartz. 2-4 feet.
- 4. Broken chert, imbedded in red and yellow clay or loam. I-3 feet.
- 5. Soft, red ore, partly greasy, enclosing bowlders of hard specular ore, exteriorly converted into red ore to a greater or less extent. These bowlders form about one-third of the whole mass, and grow larger with the depth, being apparently 5 to 8 feet in diameter at the bottom of the main shaft. The average thickness of the deposit, as far as now opened, is about 25 feet.





SECTION OF DEPOSIT BEFORE BEING WORKED

CL.D. CTAYISH DETRITUS

S.H. BOFT HEMATITE

CL. YELLOW PLASTIC CLAY

CL.4F, CHERT IN RED & YELLOW CLAY

B.S. BROKEN SANDSTONE

S.D. SANDY DETRITUS

S. SANDSTONE

LIMIT OF CUT

LIMIT OF ORE

IRON RIDGE ORE BANK & MINE .

- 6. Yellow, plastic clay, sometimes with pieces of specular ore. I-3 feet.
- 7. Red loam and white or yellow, sandy clay, irregularly mixed, enclosing large bowlders of decomposed specular ore. These bowlders are soft enough to be broken by picks and sledges, and present in their fracture a variegated appearance, red, brown, yellow, and black ore being mixed together, and containing specks of white clay and in some places seams of quartz.

The original geological position of the Iron Ridge deposit cannot be safely determined from its immediate surroundings, all the regular geological strata in that district being thickly covered by sandy and cherty detritus, undoubtedly produced by a very extensive and complete destruction of sandstones with chert-layers. A bore-hole, a few hundred feet south-west of the ore-bank, on the same ridge, went through 65 feet of this loose and irregular formation, without striking the solid rock. Several wells were sunk in various places in the valley to a depth of 45 feet. There, also, drifted masses were found to a depth of 12 to 15 feet, consisting of broken chert, of rounded pieces of sandstone, and of sand. Below this more solid, yet not quite undisturbed, rocks were reached, consisting of alternate strata of sandstone and more or less broken chert. Of these strata, a thickness of 30 feet was pierced, without obtaining water in desirable quantity.

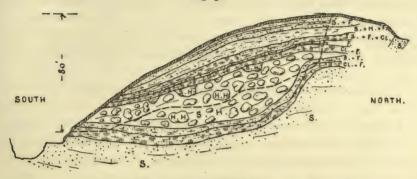
The ore in the Iron Ridge deposit seems to extend to a considerable depth. The main shaft is now over 50 feet deep, and has not reached the end of the deposit.

Meramec Bank, N. W. 1/4, Sec. 1, T. 37, R. 6, W., Phelps County, 7 miles south of St. James.

The Meramec bank, a section of which is given in Fig. 36, is a lenticular deposit of a nearly circular outline, lying in clay and chert beds, in the Second Lower Silurian sandstone. Its inclined position, as well as the broken condition of the ore, indicate former disturbances, probably caused by a partial destruction and removal of the underlying sandstone, especially on the south side, where a deep ravine or narrow valley has been eroded, perhaps by the same waters that may have underwashed the ore-deposit.

In this valley, through which a road leads down to the Meramec Iron Works, outcrops and bluffs of the Second Sandstone are exposed. The dip of the sandstone is very irregular, mostly, however, 10 to

Fig. 36.



SECTION OF MERAMEC MINE

5.	SECOND S	SANDSTONE	C6.+F.	CLAY WITH CHERT
5 . + F.	SANDSTONE	WITH FLINT	F.	SOLID FLINT
8.+ F. + CL	SANDSTONE	WITH FLINT&LOAM	8. H.	SOFT HEMATITE
.S. + H. + F.	SANDSTONE	WITH ORE & CHERT	H. H.	HARD HEMATITE
		- LIMIT OF	CUT	

20 degrees to the north-west, about in the direction of the orebank. Lower down, in the vicinity of the iron works, the Third Magnesian Limestone is exposed with a dip of 10 to 15 degrees north-north-west. A large spring, discharging about 10,000 cubic feet of water per minute, of a temperature of 58° F., uniform all the year round, comes out of this limestone, proving that it must contain enormous cavities and spacious subterranean channels. The limestone is thickly bedded and contains numerous chert concretions. The upper strata are more irregular, and enclose layers and masses of sandstone. The Second Sandstone in the vicinity of the ore-bank occurs on the north and west sides in distinct outcrops dipping toward the ore. This sandstone is often thinly bedded, and shows sometimes a wavy striation on the surface of the layers, similar to that which is frequently produced on loose river-sand by the waves of a shallow water.

The succession of strata in a section through the ore-deposit, as represented by Fig. 36, seems to be the following:—

- 1. Second Lower Silurian sandstone.
- 2. Broken chert, imbedded in red, sandy clay. 12 feet.
- 3. Chert-breccia in sandstone. 4 feet.
- 4. Soft, red hematite with many bowlders of specular ore, more or less altered on the outside; also, in places, irregular, large

masses of hard, yellow limestone, with seams and specks of carbonate of iron; also, soft, greasy "paint-ore," red to dark purple, the latter very pure in pockets on the surface of the deposit; also, streaks of soft, yellow ochre. The thickness of the deposit varies from 5 to 40 feet.

- 5. White clay, mixed with broken chert. 0 to 5 feet.
- 6. Layers of solid chert. I to 2 feet.
- 7. Broken sandstone and chert, mixed with loam. I to 10 feet.
- 8. Impure and uneven sandy rock, impregnated with oxides of iron and containing layers of broken chert. 5 to 20 feet.
- 9. Alternate layers of sandstone and of massive chert. 5 to 20 feet.
- 10. Dry, sandy soil, containing small, rounded particles of specular ore. 1/2 to 2 feet.

All the strata above the ore are very irregular in their position and thickness. Most of them can, however, be traced across the whole mining-cut. That portion of the section, Fig. 36, which is below the limit (— — —) of the cut, is imaginary, and has been added merely to give a clearer picture. The ore might, perhaps, in the central part of the deposit, extend deeper into the sandstone than is indicated in that section.

James Bank and Moselle No. 9, S. ½ of S. E. ¼, Sec. 29, T. 38, R. 6, W., Phelps County, 2 miles south of St. James.—These two banks are situated close together, on a low ridge, on the plateau of St. James, between the Dry Fork and the Bourbeuse Rivers.

The formation in that district is Third Magnesian Limestone, capped by Second Sandstone. This may be observed along the Dry Fork River, south of St. James. The two ore-deposits mentioned are in the sandstone. The sketch on next page, Fig. 37, shows their relative position.

Both these banks are nearly worked out; the best and richest ores are all taken out, and the walls and the bottoms of the deposits are laid bare. These places, therefore, offer a good opportunity for the geologist to study the character of these pocket-like deposits in the sandstone.

The ore formerly contained in these pockets was mostly soft, red, and in part greasy hematite, enclosing large bowlders of specular ore. The ore filled a nearly circular depression in the sandstone,



with pretty steep walls. The ore was 10 to 15 feet higher in the centre than at the circumference. The James bank had about 35 feet average height, and a diameter of over 200 feet, and furnished about 30,000 tons of ore.

The Moselle bank was 20 feet thick in the average, and 150 feet in diameter, and may have contained 12 to 15,000 tons of ore, some of which is yet in place.

The James bank lies at the edge of the northern slope of the hill, and dips slightly north. The Moselle bank lies at the edge of the western slope, and dips west. These last observations, made on two banks so similar in every respect, and in so close proximity to each other, go far to prove that the dip of such deposits follows the slope of the hill, and that both were produced by the same cause, namely, by erosion.

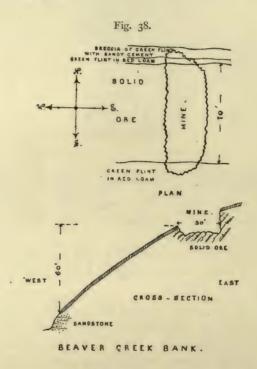
The walls of these two, now empty, banks, consist of a mixture of green and white broken chert, with yellow and red clay, partly soft, partly indurated.

The face of the walls is rather uneven, the ore reaching in places into the chert in irregular masses. The limits between the chert and the ore are, however, well marked. The bottom consists of a white or gray broken chert, mixed with white clay.

A shaft has been sunk into this mass, in the centre of the James bank, 22 feet deep, without reaching the solid rock. In the lower part of this shaft, the clay turned dark gray, green and black, and

was mixed with iron pyrites, in small concretions or as a fine crystalline grit.

Beaver Creek Bank, S. ½, Sec. 33, T. 37, R. 8, W., Phelps County, 5 miles south-west of Rolla.—I give here a plan and an elevation of the Beaver Creek bank, as far as it was opened and known in summer 1872.



It lies on the summit of a high ridge, near the head of Beaver Creek. An excavation has been made into the ore, 70 feet long, 30 feet wide, and 16 feet deep. The hill seems to consist of sandstone, which crops out on the slope about 60 feet below the mine, as seen in the above elevation.

The ore seems to be pretty solid, and in its greater part specular, but slightly altered or softened. The above plan shows that the mass of ore extends about 70 feet from north to south, being cut off on both sides by nearly vertical layers of green chert imbedded in red loam. Next to this a layer of chert-breccia may be observed on the northern wall. The extent of the deposit in other directions cannot as yet be estimated, because the ground surrounding the

bank is covered by soil, without any plain surface-indications of either rocks or ore.

Other deposits, which probably belong in this category, are the Craig bank, in the "Upper Meramec" district; the Wiggins and the Ziegler banks, in the Salem district; and the Mont Rouge, Mocassin Bend, and Hancock banks, in the specular-ore district, on the Middle Gasconade River, and in Miller County. The exact location, with a few particulars, of these banks is given in the general ore-bank list, section D.

c. DISTURBED DEPOSITS OF SPECULAR ORE.

The specular-ore deposits, of which I intend to speak under this head, were originally such as described under b. They were, however, not only broken by contraction, or by underwashing, or by more violent geological actions, but they were also divided into two or more large portions, which portions were separated from each other by the removal of one, or of more than one, or of all of them, from their original position. We may, accordingly, distinguish two kinds of such "disturbed deposits," namely:—

- I. Masses of specular ore which have been removed from their original position and deposited elsewhere, in a more or less irregular manner, and
- 2. Remaining portions of original deposits, from which other portions have been separated and removed.

Such disturbances must have taken place in some instances slowly and gradually, in other instances with more rapidity and violence, which difference of action must have exercised a marked influence on the condition in which the various deposits are found at present. The more rapid and violent this action was, or the greater the distance over which a certain mass of ore has been shifted, the more will the present ore-bank be broken up, and the less of those more solid rocks with which it was originally associated will adhere to it, as chert, sandstone, breccia, and the more directly will it be imbedded in loose detrital materials.

Some of the deposits, which I shall describe or mention in this category, are not yet sufficiently opened to give a final decision regarding their character. They may prove to be broken-off parts of larger deposits situated in close proximity, and may lead to the discovery of the latter.

FRANKLIN COUNTY ORE-DISTRICT.

Thurmond Bank, N. ½ N. W. ¼, Sec. 19, T. 41, R. 1, W., Franklin County.—This bank is situated 2 miles north of Stanton, in a rather rough country, with steep, high hills, separated by narrow valleys and ravines. The soil is mixed with, and in places covered, by broken white chert. No outcrops of regular geological strata are perceptible, nor any surface-rocks, which might give a clue to determine the formation. The Thurmond bank is as yet but imperfectly opened. The surface-indications consist of a number of large pieces of limonite, and of some small, sharp fragments of a very hard and silicious specular ore. They are scattered over a surface about 50 feet wide and 200 feet long, over a slight swelling of the ground extending down the slope of a moderately steep hill.

A shaft was sunk here a number of years ago, in a vain attempt to find copper-ores. It is said that this shaft, which is yet open to a considerable depth, passed through 37 feet of red iron-ore. Some heaps of soft, red, somewhat clayish hematite, mixed with pieces of soft "paint-ore," are seen at the mouth of the shaft. All appearances indicate that this is a greatly disturbed and dislocated deposit.

Old Copper Hill, E. ½ N. E. ¼, Sec. 23, T. 40, R. 2, W., Crawford County.

This bank is not opened. It has externally a great resemblance to an undisturbed bank, as which it would have to be considered if the surface-ore was larger, less rounded, and more concentrated on the summit of the hill. As it is, the bank has more the appearance of being the remainder of a disturbed deposit, large parts of which would have been removed. A circumstance which is very strange, and which also points to a disturbance, is, that fragments of white sandstone, in part sharp-edged, are found together with the surface-ore on the summit, while the upper part of the hill generally seems to consist of a dark-colored and ferruginous sandstone.

The hill is pretty steep, and nearly isolated. The surface-ore is specular, in some places pure; in others mixed with sand, and passing into a strongly-impregnated sandstone. The pieces are all rounded, none over head-size, most under fist-size.

The two shafts, indicated on the annexed sketch, were sunk to a

Fig. 39.

GILKERSON'S FORD ON GRAND RIVER, HENRY CO.



depth of perhaps 20 or 30 feet, in loose sand and clay, mixed with pieces of white sandstone. They did not reach any solid strata of rock.

STEELVILLE ORE-DISTRICT.

Cherry Valley No. 2, W. ½ S. E. ¼, Sec. 4, T. 37, R. 3, W., Crawford County.

This bank is represented in Fig. 27, and has been mentioned and characterized in connection with the description of the Cherry Valley No. 1 bank, from which it is only ½ mile distant. It consists of a streak of large and copious, specular and brown surface-ore, about 20 feet wide and say 200 feet long, extending down the south-western slope of a hill into a ravine, and a short distance up the opposite slope. This bank has the appearance of a disturbed

though undoubtedly very valuable deposit. Ore is found in less quantity in several other places on the surrounding hills.

Steelville No. 1, E. ½ S. W. ¼, Sec. 5, T. 37, R. 4, W., Crawford County, 2 miles west of Steelville.



The above sketch gives a section through this bank, which is opened by a large mining-cut. As Fig. 40 shows, this bank represents a typical example of a disturbed deposit of the first kind. We see here an irregular mass (S H) of soft, red hematite, with bowlders of specular ore lying at the foot of a hill, imbedded in loose materials, as white clay (Cl), clay mixed with broken chert (Cl+F), broken chert and sandstone mixed (F+S), red, sandy loam (R Cl), and fine sandstone-detritus with some broken chert (S D). The red loam encloses large bowlders (S) of a fine-grained, yellow, very hard sand-rock. The position of all these materials, including the ore, makes it evident that they must have slid down the hill, some simultaneously, others at various times, and must have been thrown there one over the other, in irregular layers.

Some of the loose materials round the ore are undoubtedly products of the destruction of cherty sandstone-strata, in which the deposit originally lay. It will be noticed that the excavation made by the miners has nearly reached the solid sandstone which seems to compose the hill and which is likely to cut off the ore. There are, however, indications of specular and red ores in other places, which make it probable that other loose masses of ore have been thrown down at the foot of this hill and buried under the detritus.

ORE-DISTRICT ON THE UPPER MERAMEC RIVER AND ITS TRIBU-TARIES.

Winkler Bank, S. ½, Sec. 14, T. 36, R. 6, W., Phelps County.
—This bank is situated on the plateau between West Benton Creek and Norman Hollow, and spreads over three flat hills, lying in a north-south line, somewhat curved toward the east.

The south-eastern slope of the most northern of the three hills is covered with good and large surface-ore, mixed with some broken chert. The central hill shows scarcely any ore on the surface, but frequently pieces of sandstone. The southern hill is very wide and flat, and bears on its western slope a very extensive streak of surface-ore, about 1,200 feet long and 100 to 400 feet wide. Most of this ore is rounded off, and not very large, and looks as if it had been drifted. A number of pieces, however, reach and exceed head-size. The ore at the south end is very hard and silicious, that at the north end is purer and softer.

This bank is untouched, and its exterior appearance does not convey an exact idea of its character. It is not unlikely that the ore on the northern hill forms a separate deposit from that on the southern hill, and that the latter deposit has been more disturbed and broken, and the ore scattered over a larger surface. From all appearances the Winkler bank seems to contain considerable quantities of specular ore.

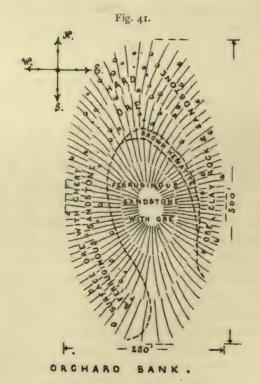
Arnold Bank, S. E. 1/4, Sec. 4, T. 35, R. 5, W., Dent County.—This bank is not yet thoroughly opened. But it presents a very similar appearance to that of the Steelville No. 1, and is undoubtedly a deposit which was formerly imbedded in sandstone, and fell or slid down to the foot of the hill simultaneously with the erosion of the ravine, near which it lies. The hill itself is sandstone, which has been struck by a shaft sunk 12 feet deep near the summit of the hill, about 50 feet above the ravine. The ore is principally spread over a swelling of the ground, reaching from the ravine about 40 feet up the slope, in a width of 40 to 50 feet. On this ground large bowlders and smaller pieces of surface-ore, mostly rounded, are found, together with pieces of white, yellow, and red sandstone, containing thin seams of specular ore. Also pieces of broken chert, and of an impregnated or ferruginous sandstone, are quite frequent. Numerous bowlders of ore are deposed

in the ravine. The hill on the other side of the ravine is likewise sandstone. A ditch made at the foot of the hill on which the ore is found struck red clay, mixed with paint-ore and with bowlders of specular ore. A second shaft, sunk 12 feet deep into the slope, on a place about 40 feet above the ravine, passed through red, sandy clay, mixed with pieces of ferruginous and of white sandstone, and with bowlders of specular ore. The whole slope is evidently thickly covered with sandy detritus, enclosing irregular and unevenly distributed masses of broken ore.

Other banks of this district, which are likely to belong in this category, are the N. G. Clark No. 2, C. C. Cook, Arthur, and St. L., S. and L. R. B. banks. Their location, etc., is given in the orebank list, in Chapter V.

SALEM ORE-DISTRICT.

Orchard Bank, E. ½ S. E. ¼ S. E. ¼, Sec. 13, T. 34, R. 6, W., Dent County, close to Salem.

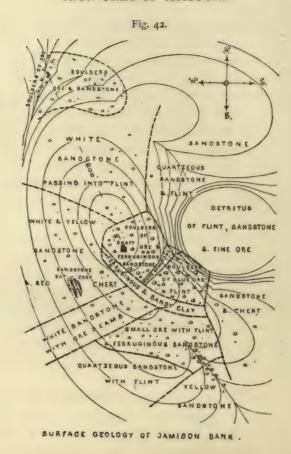


This bank is remarkable for the large development of sandstone, colored and impregnated by oxides of iron, which seems to compose the greater part of the body of the hill, as well as for the unusual relative position of ore and sandstone, the former occupying here an annular space round the latter. These facts, together with the flatness of the hill, the height of which is only about 30 feet, prove that this deposit has been greatly disturbed. It seems likely that the ore lay originally on the impregnated sandstone, and that both occupied a much higher level than they do now. A large part of the ore was broken into pieces varying from a pea to headsize. Another large part of it has undoubtedly been carried off. How much of the original mass of the ore is left in the hill, can only be ascertained by practical work. It may be observed in this locality, as in several others, that the white sandstone gradually passes into the yellow and into the ferruginous sandstone. argillaceous or calcareous cement that surrounds the single sandgrains is changed into red clay and into reddish-brown iron-ore. In other places, the original cement is replaced by amorphous quartz, so that the sandstone takes the appearance of a quartzite, which itself in places loses its grainy structure, passing into a solid flint or chert.

It also seems that under certain circumstances the sand-grains, when enclosed in a quartzous or ferruginous cement, have been dissolved and removed, leaving a mere skeleton of a former sand-stone, with a cellular structure. The cells and irregular holes of such masses have sometimes been filled up again, either partly or wholly, by a transparent quartz of a dark appearance, or by yellow jasper.

Jamison Bank, S. W. 1/4, Sec. 1, T. 33, R. 6, W., Dent County, 3 miles south of Salem, on the vast plateau dividing the waters of the Meramec from those of the Current River.

This bank occupies the highest point on a rather flat, semicircular hill, which lies round a nearly circular depression (sink-hole?), apparently filled with fine detritus of chert, sandstone, and specular ore. The surface-geology, as given in Fig. 42, is very irregular, but nevertheless seems to be grouped in a general way round that part of the summit and eastern slope over which the largest and most copious surface-ore is spread. The ore is specular, in part



pure, in part mixed with quartz. Some bowlders are 2 to 3 feet in diameter.

The principal surface-ore district is separated from the ordinary light-colored sandstone by a zone of ferruginous and clayish materials.

Another smaller district, with rounded surface-ore, is seen about 600 feet to the north-west, another in a small ravine to the south. Both are probably drifted outliers of the main deposit, which lay originally at a higher level, above the present top of the hill, perhaps a little north of it.

I view this bank in a similar light as the Orchard bank, and consider it as containing the remnants, perhaps pretty large, yet incalculable, of a former lenticular deposit in sandstone, which has been broken, and partly destroyed and removed.

The shaft marked on Fig. 42 was 10 feet deep, end September, 1872, and had not struck any solid rock, but stood in a red, sandy loam, with bowlders of sandstone and of specular ore.

Other banks in the Salem district, and in Shannon County, which may be supposed to belong in the category of more or less disturbed deposits, are the Barksdale, Merriam, Shannon, and Current River banks. The locations, etc., are given in Chapter V.

ST. JAMES ORE-DISTRICT.

Thornton Bank, N. E. 1/4 Sec. 33, T. 38, R. 6, W., Phelps County.—The situation and appearance of this bank may be seen from Fig. 43.

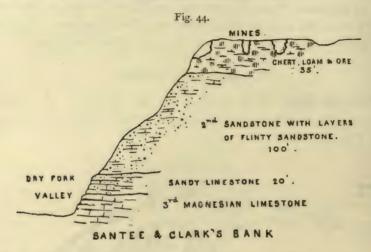


THORNTON PANK.

The ore is soft red, with small pieces of hard specular. No large bowlders have as yet been found in it. The character and position of the ore is such that it must be considered as a bed-like or a lenticular deposit, which has been brought into its present vertical position by some exterior disturbance, and then broken and decomposed. The deposit can so far be traced over a small space only, and is perhaps a removed portion of some larger bank. ore seems to be associated with the rocks in which it originally lay, or at least with their detritus. As the succession of these rocks is the same on both sides, the supposition suggests itself that the corresponding strata might come together below the ore, and thus constitute a pocket, which is crushed sidewise, in the direction from S.E. to N.W. The hill is Second Sandstone.

Santee and Clark's Bank, S. W. 1/4, Sec. 33, T. 38, R. 6, W. Phelps County.—This bank lies on a high bluff of Third Magnesian Limestone and Second Sandstone, on the east side of Dry Fork River.

Several small openings have disclosed irregular masses of red and brown ore, imbedded in layers of chert and loam.

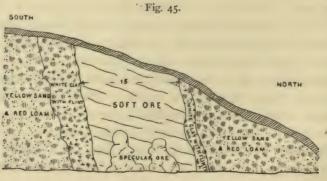


The character of this bank is very indistinct and doubtful. The present digging may lead to some larger, disturbed bank, or it may disclose a drifted deposit. The materials which surround the ore are evidently of a detritic nature, and not now in the place where they were formed.

Another deposit which might belong here is the South Mountain (see Chapter V.).

ROLLA ORE-DISTRICT.

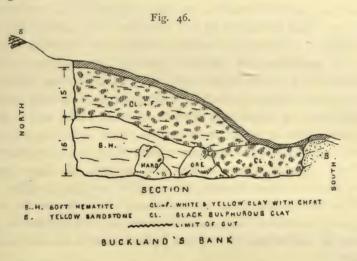
Kelly Bank No. 1, E. ½, Sec. 18, T. 36, R. 8, W., Phelps County.



KELLY BANK . Nº 1.

This bank is situated near the summit of a hill composed of Second Sandstone in its lower part, while no solid rock can be seen higher up. The bank itself has a decided resemblance to the Thornton bank, above described, but it seems to be more extensive, and contains larger masses of hard, specular ore. It differs, besides, by the detritic character of the associated rocks. There is next to the ore, on each side, a thick layer (1–3 feet) of white clay mixed with broken chert, and outside of this a mass of yellow sand and red loam, irregularly mixed, and free from chert. When opened further, this deposit may be found to be a large fragment of a disrupted-layer deposit, or else an original ore-pocket, which has been pressed and crushed sidewise. The layers of clay and chert are evidently in their original position relative to the ore, but they are broken and mixed. It is doubtful whether this statement could also be extended to the surrounding mass of sand and loam.

Buckland Bank, S. ½, Sec. 20, T. 37, R. 8, W., Phelps County.—This bank lies at the foot of a sandstone hill, in the crossing of two ravines.



I give here a section of the mining-cut made in it, which presents a very plain instance of a disturbed specular-ore bank. This deposit seems to be actually overturned; the ore, which in regular deposits lies above the clay and chert (Cl and F), is here covered by the detritus of these materials. On the south side of the cut we find a mass of black, tenacious clay (Cl) mixed with fragments of a half-triturated,

dark-gray clay-slate, and with pieces of pyrites, and impregnated with sulphate of iron. I have mentioned a somewhat similar sulphurous mass as having been found in the James bank, near St. James, below the ore and below the chert and clay beds that underlie it. It is therefore probable that also in the Buckland bank this mass lay below the clay and chert and ore formerly, but that the ore-deposit was underwashed with the erosion of the ravines in which it is situated, and broken and overturned.

The proximity of the regular sandstone on all sides, as marked on the sketch, shows that this deposit cannot extend horizontally much over the limits of the present cut; but it may extend somewhat in the depth.

Another bank in this district, which I shall mention in Chapter V. as Moselle No. 10 bank, seems also to belong in this category of disturbed deposits.

d. DRIFTED DEPOSITS OF SPECULAR ORE.

In the general introduction to this Chapter (IV.), I have given the reasons which induce me to add a category of "drifted deposits," although it is somewhat doubtful whether such deposits really exist. I understand by "drifted deposits," accumulations of loose fragments of destroyed or half-destroyed ore-banks, which fragments have been carried off by water over considerable distances, either alone or mixed with detritus of other rocks, and again deposited, either in more or less regular beds or strata, alternating with layers of other broken and triturated rocks, or irregularly distributed through large masses of such detritus. We have therefore two kinds of drifted deposits, the stratified and the irregular.

The detritus which accompanies such deposits always consists of sand, sandstone, chert, and red loam, of such a character as to leave no doubt that the original deposits were in the Silurian sandstone.

Specular-ore banks, having the exterior habitus of drifted deposits, are very numerous in the central ore-district of Missouri. I intend to describe a few of them in the following lines, but as none of them is sufficiently opened as yet to allow a clear insight into its interior composition, I must leave to future mining operations to decide whether any, and how many, of these banks, really are

what they externally look to be, namely, "drifted deposits," or whether, on the contrary, the bowlders and pieces of ore visible at present are only outliers of either intact or disturbed deposits, which now lie hidden in the ground.

FRANKLIN COUNTY ORE-DISTRICT.

Blanton Specular Bank, N. ½ S. E. ¼, Sec. 29, T. 40, R. I, W., Washington County.—Some rounded surface-ore, mostly small, is found on three flat spurs of a low ridge. Strata of solid sandstone crop out at the foot of these spurs, dipping slightly south-west. The spurs point about north.

A hole, dug 15 feet deep on the top of the most eastern spur, passed through drifted, sandy detritus, with little ore, and then struck a layer of chert.

This bank consists, according to these observations, of a low sandstone-hill, thickly covered with detritus, through which single pieces of specular ore are unequally distributed. The presence of larger and workable masses of ore is not impossible, but is nowhere plainly indicated. The ore itself is of good quality.

Primrose Hill, S. W. ¼ N.W. ¼, Sec. 32, T. 40, R. I, W., Washington County.—This bank is, as far as opened at present, of a similar character as the Blanton specular bank. But the prospects are here better. The surface-ore, mostly small and rounded, occurs on the inner side of a high horseshoe-shaped ridge, enclosing a deep ravine. The spurs ending the curved ridge point north-east. Pieces of a hard sandstone with quartz-cement, and of ordinary soft sandstone, are also found on the surface.

The ridge was investigated by three shafts, one on the northern slope of the western spur, the two others on the inner slope of the central and highest portion of the horseshoe. Neither of these shafts has reached the solid rock as yet. The two upper shafts are forty feet deep, in fine, sandy detritus, mixed with streaks and irregular masses of soft, red hematite, and of broken stalactites of half-decomposed specular ore, sometimes cemented by soft sand-stone.

It is not unlikely that workable masses of ore will yet be met with in this vicinity.

STEELVILLE ORE-DISTRICT.

The Scotia district contains one bank that may belong here, namely, the Bleeding Hill. The Steelville district contains the N. G. Clark No. 1, the Knox, the Sea and Marsh, and the Ferguson banks, all in Crawford County. The last-named bank is the most worked, and therefore the most interesting of them.

Ferguson Bank, Sec. 21, T. 37, R. 4, W., Crawford County.

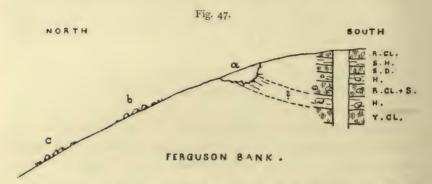


Fig. 47 gives an elevation of this bank, as it now appears. A flat northern hill-slope shows, in several places marked a, b, c, horizontal zones of larger and smaller specular ore on the surface. These zones are in some places very distinct, in others less so. They are from four to eight feet wide, measured down the slope. They seem to run across the slope, and to terminate on either side, in a ravine. The ravine on the western side is the deepest, and contains irregular accumulations of rounded ore. A shaft sunk, near the highest point of the slope, to a depth of twenty-two feet, passed through

- 6 feet of soil and red loam,
- 2 feet of soft, red hematite,
- 2 feet of red and yellow sandy clay,
- 2 feet of soft, red hematite, with pieces of specular ore,
- 2 feet of red clay, with pieces of sandstone and some chert,
- 8 feet of large bowlders of specular ore, imbedded in soft, red hematite,

below which a layer of light-yellow, clayish ochre was struck. All these materials seemed to be in layers of irregular thickness, generally dipping into the hill.

An opening made at the point marked a, has cut through a 6-feet layer of white sand and clay with bowlders of white sandstone dipping along the slope. Below this, layers of red clay with small ore were struck, dipping into the hill; below these, and dipping in the same direction, a 16-inch layer of broken chert with sandy clay, and 2 inches of fat, white clay; finally, large bowlders of specular ore, softened, and altered into red ore on the outside.

The above description would indicate that a considerable portion of this hill might be composed of alternate layers of broken ore and of detritus of rocks.

ORE-DISTRICT ON THE UPPER MERAMEC.

Smith Banks, Sec. 26, T. 36, R. 6, W., Phelps County.—The three Smith banks are situated on three very flat slopes or swellings of the ground, all pointing south, and lying about on an east-west line, within a distance of one half-mile. The two western banks, No. 1 and No. 2, are very near together, and have a very similar appearance. In both of them good specular ore, in very numerous, rounded pieces, is spread over a flat, triangular slope, encompassed by two small converging ravines, or water-runs, which unite at the lowest and southern end of the bank. The upper, wider, and most northern part of the slope, which forms the base of the triangle, passes into a plateau.

This triangular space, over which the ore extends, is about 250 feet wide at the base and 300 feet long in the western or No. I Smith bank; and it is 250 feet wide at the base and 600 feet long in the central or No. 2 Smith bank. The ore of the latter rarely exceeds fist-size, while that of the former is generally somewhat larger, and sometimes reaches head-size. Pieces of broken chert and sandstone are found with the ore. The No. I bank extends 12 to 20 feet over its western ravine and up the opposite slope, where the ore, however, has a somewhat different character, being mostly stalactitic, or "pipe-ore."

Three holes were dug, 8 to 10 feet deep, on various points of the Smith bank No. 1. They passed through loose masses of broken white sandstone, sand, broken chert, white clay, and red loam, all mixed irregularly, and containing in places some soft, red

hematite and some rounded specular ore, the latter principally concentrated in the soil or near the surface. I have explained the process by which such a concentration is effected, in my description of the Simmons Mountain, in division b of this chapter.

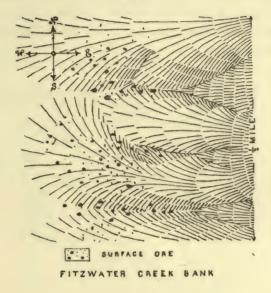
There is no doubt that the now visible portions of the two western Smith banks have the character of irregular, drifted deposits. They may contain, occasionally, larger and workable accumulations of ore; but no one can tell whether or where they exist.

The Smith bank No. 3 has a different appearance. On the upper part of a flat slope a circular depression of sandstone is perceptible, having a diameter of about 50 feet, and being marked by annular outcrops. Inside of these outcrops is a small accumulation of specular ore, in rounded pieces, from nut- to head-size. This description would indicate the presence of an undisturbed deposit of the category b; but the scarcity of the surface-ore, the small size of most of it, the light color of the surrounding sandstone, the absence of ferruginous materials, the very slight dip of the sandstone-outcrops, and the small diameter of the circular space they enclose, all this together makes me believe that this bank was formed by a slight depression of the sandstone, in which depression some drifted ore has found a resting-place.

Fitzwater Bank, Secs. 33 and 34, T. 35, R. 4, W., Dent County.—This bank occupies a pretty high position, being about four hundred feet above the Fitzwater Creek. It lies on the western slope of a ridge, which is composed of Third Magnesian Limestone, capped by Second Sandstone. The ore seems to overlie the latter. A sandy soil, mixed with fine chert, and with pieces and larger masses of chert, either porous or dense, covers the surface of the hills.

The ore-bank is as yet untouched. Fig. 48 is a topographical sketch of this locality, showing the manner of distribution of the surface-ore over one large central spur, and over the adjacent slopes of two spurs, north and south of the central one. The best indications extend about fifteen hundred feet north and south, and about eight hundred feet east and west. Most of the ore is below the size of the fist, and rounded off at the corners and edges. In some places, however, it reaches and exceeds the size of the head. This is especially the case in the ravines, where the most considerable accumulations are found. Pieces of yellow sandstone, and

Fig. 48.

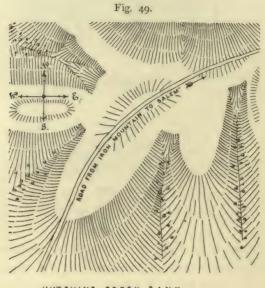


near the northern ravine, also, some pieces of chert-conglomerate, cemented by yellow sandstone, occur with the surface-ore. No ferruginous rocks have been observed. The ore seems to be most abundant at a certain level along the slopes, which level is about eight feet below the highest point on the two northern spurs, which are flat and low. The southern spur, a part of which only is visible on the sketch, is considerably higher, and the level of the most abundant surface-ore is there much farther below the highest point. These observations indicate that there might exist a thick and extensive layer of drifted ore, running nearly horizontally through all three spurs, and covered by a mixed detritus of sandstone, chert, and ore. It is, however, obvious that such a conclusion cannot be drawn with any degree of certainty. The character of this bank is not indicated with sufficient clearness, by its external appearance, to make a reliable judgment possible.

Other banks in this district, which may be supposed to be drifted deposits, are the Santee, the Anderson, the Blackwell, the Reuben Smith, and the Carson banks, the location, etc., of which will be given in Chapter V.

SALEM ORE-DISTRICT.

Hutchins Creek Bank, Sec. 15, T. 34, R. 4, W., Dent County, 9 miles east of Salem.



HUTCHINS CREEK BANK .

The topographical sketch, Fig. 49, shows that this bank has much resemblance to the Fitzwater bank, and that its character is even less pronounced than that of the latter. Good specular ore, from nut-size to one foot diameter, mostly rounded, is found in four ravines on the west and south sides of a high ridge. Quite scarce and only very small ore is seen on the spurs between the ravines. The ore in the ravines does not reach a higher level than about 30 feet below the top of the ridge. The hills are covered with soil and chert. Large and small pieces of white sandstone are met with on the lower part of the slopes. I could not find any ferruginous rocks. The district represented in Fig. 40, and containing the four ravines in which the ore is principally concentrated, measures over one-quarter of a mile in each direction.

Practical opening and working only can decide whether the ore in the ravines has come from a coherent deposit existing in the ridge, or whether it is derived from a drifted deposit, and has been concentrated in the ravines by the gradual erosion of the latter. Other deposits in this district are the Huzzah, the Pittsburgh, the Norris, the Hayes, and the Orchard & Young banks. For locations, etc., see Chapter V.

IRON RIDGE DISTRICT.

Iron Ridge No. 2, Sec. 33, T. 39, R. 5, W., Crawford County.— A pretty extensive tract of slightly undulating ground, ½ mile north of Iron Ridge Station, on the Atl. and Pac. R. R., contains in many places indications of specular ore, and occasionally shows large bowlders of good surface-ore. A number of ditches were made to investigate this tract, and disclosed irregular accumulations, mostly of small extent, of rounded ore with red clay, of white clay with pieces of chert, and of impregnated sandstone. This locality has decidedly the appearance of an irregular, drifted deposit.

Other banks, supposed to be of a drifted nature, are, in this district, J. P. Card & Co.'s, Senator Buckland's, the Dorey, the Isabella, and N. G. Clark & Co.'s banks; in the St. James district the Thompson, the A. C. L. No. 1, the Railroad Nos. 1, 2, 3, the Lenox, the Hall, and the Seaton banks. See the ore-bank list, Chapter V.

ROLLA DISTRICT.

Kelly Bank No. 2, N. E. 1/4, Sec. 21, T. 37, R. 8, W., Phelps County.—Two openings have been made, one on the summit and another on the eastern slope of a hill apparently composed of sand-stone covered by cherty soil.

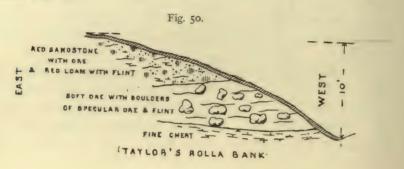
The lower opening shows a double succession of layers of clay, of broken chert, and of broken, ferruginous sandstone, dipping 45° south-east, and below this an irregular mass of soft hematite, red and brown, containing thin veins and small pockets filled with crystalline carbonates of iron. These carbonates are also found as cement of broken chert. On the north side of the cut is a large mass of a loose, gray rock, probably triturated calcareous sandstone. This rock contains single crystals of iron pyrites, and also veins of carbonate of iron. The latter is evidently formed after these masses were brought into their present irregular position.

The upper opening shows a bed of bowlders of limonite, some

1--2 feet in diameter, imbedded in red loam, without any chert. This limonite is in its general appearance more like that found in limestone, and unlike the limonite formed by alteration of specular ore.

The Kelly No. 2 bank is evidently a locality which has undergone several strong disturbances at various epochs. The materials met with in the lower cut especially may have been broken up, drifted, deposited, and after that once more disturbed and broken.

Taylor's Rolla Bank, S. W. 1/4, Sec. 15, T. 37, R. 8, W., Phelps County.—This bank is situated a short distance from the Kelly bank No. 2. It is less irregular in its formation, and has more distinctly the character of a drifted deposit, as may be seen from the section, Fig. 50.



The bottom of the cut consists of a bed of finely-broken chert. Above this is a layer, 3 feet thick, of soft, red hematite, in part clayish, and full of seams, specks, and irregular masses of spathic ironore (carbonate of iron), and enclosing bowlders and pieces of specular ore and of chert. Above this are 5 feet of alternate layers of red, somewhat ferruginous sandstone, and of red loam with broken chert. A cherty soil covers the slope.

Other perhaps drifted deposits are, in this district, the Hyer, the Cold Spring, the Coleman, the Piney Creek, the Baird, the Hudgeons, the Camp Creek, and the Railroad No. 4 banks; in the Gasconade district, the Frost, the A. C. L. No. 2, the Railroad No. 5, the Morgan, and the James pipe-ore banks; on the lower Osage River, the Wimar Creek, the Belans Creek, and the Linn Creek banks.

e. STRATA OF RED HEMATITE.

The red hematites of the carboniferous formation, which hematites I have mentioned in Chapter I. and described in Chapter II., do not occur as deposits with definite limits, lying as independent and foreign developments between or across the regular stratified or unstratified geological rocks; but, unlike all other deposits of ironore in Missouri, they form and compose in themselves regular geological strata.

These strata of red hematite, although always in the carboniferous system, do not seem, however, to occupy the same geological horizon in all the localities where they are found. While occurring in the so-called Ferruginous Sandstone of the subcarboniferous formation in Callaway and Cooper Counties, and on the Upper Osage River (in St. Clair and in the south-eastern corner of Henry County), the strata discovered near Calhoun, in Henry County, lies, according to Mr. G. C. Broadhead's investigations, in the Lower Coalmeasures.

The development of red hematite in the Ferruginous Sandstone seems frequently to extend over large areas. The sandstone in such districts becomes more and more impregnated with iron, contains more and more nodules and layers of pure ore, and finally entire strata of sandstone, varying in thickness from a few inches to three and more feet, are replaced by the ore.

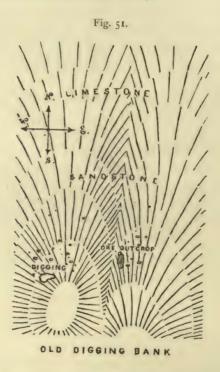
None of the deposits of this kind are as yet sufficiently opened and worked to allow the geologist to decide whether this ore was formed directly after and on the surface of the underlying sandstone, or whether it was infiltrated afterward, gradually removing and replacing either beds of limestone in the sandstone, or beds of the sandstone itself, which happened to be more soluble than other layers, or more liable to be attacked and altered by the chalybeate solution, at the temperature then existing.

CALLAWAY COUNTY DISTRICT.

Old Digging, Sec. 22, T. 45, R. 10, W., Callaway County.

Fig. 51 is a sketch of this locality. The lower part of the hills seems to be composed of subcarboniferous limestone, the upper of ferruginous sandstone. Large and small fragments of chert are found all over the surface of the ground. The ore has been dis-

covered in two places, near the top of the hill, on both sides of the ravine. On the western hill a hole was dug a number of years ago, and it is said that many tons of ore were taken out of it and were worked in a charcoal-hearth in the valley.

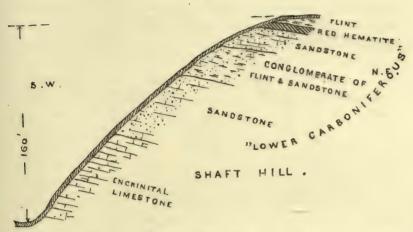


On the east side of the ravine, and rather close to it, an outcrop is perceptible, consisting of a 5-inch stratum of solid, pure, red hematite. The place is not opened, and the total thickness of the ore cannot be seen.

As the ore in both these outcrops, east and west of the ravine, seems to be in place, it is probable that a stratum of ore extends from the one to the other, and perhaps through the whole hill.

Shaft Hill, N. W. 1/4, Sec. 4, T. 45, R. 10, W., Callaway County.—This hill, of which I give a geological section in Fig. 52, is one of the most hopeful localities, and one whose structure is most clearly seen, in this ore-district. It is now being opened and worked. The annexed section was made from the indications perceptible on the surface. The strata seems to dip slightly north-east. The limestones, sandstones, and conglomerates are exposed in





several places, forming high bluffs. The ore and the strata overlying it are not exposed, and have to be judged from the pieces found on the surface. Fragments of ore are found at a certain level all round the hill. A stratum of ore undoubtedly runs through the hill, and its thickness may, from the surface-indications, be estimated as varying from one to three feet. In some places, however, it seems to reach a thickness of five feet. The hill is nearly round, and has, at the level of the ore-bed, a diameter of about 800 feet.

The ore occurs sometimes in nodules or lenticular concretions, composed of several concentric layers, and apparently imbedded in loose sand; sometimes in thin layers, alternating with layers of loose sand; sometimes as thick, massive strata.

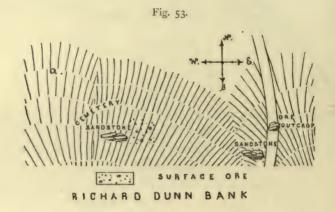
A shaft was dug, ten years ago, on the eastern slope of the hill, near a deep ravine, at a level considerably below that of the regular ore-bed. The shaft went eight feet deep through sand and broken ore and chert. Larger masses of ore have lately been discovered there. This part of the Shaft Hill deposit has probably been displaced, and is only the remainder of a portion of the regular ore-stratum which was underwashed and partly destroyed by the erosion of the rayine.

Raph Dunn Bank, S. E. 1/4 Sec. 32, T. 46, R. 10, W., Callaway County.—This bank is situated close to the Shaft Hill bank, and separated from the latter by a deep ravine. The ore-deposit must have been originally a continuation of that on Shaft Hill, to judge from the similarity of their position. Outcrops of the ore can be

observed on the east side of the hill, near the top. Toward the west and north the hill passes into a plateau. There is nothing to indicate how far the ore extends into this plateau.

Bloomfield Bank, W. ½ Sec. 32, T. 46, R. 10, W., Callaway County.—Good evidences of stratified red hematite occur on both sides of a little valley, adjacent to the village of New Bloomfield.

Richard Dunn Bank, Sec. 21, T. 46, R. 10, W. Callaway County.

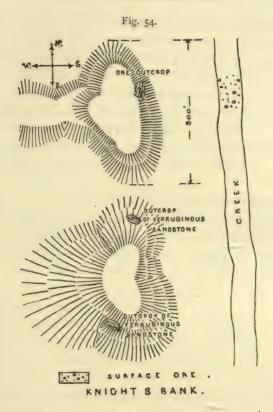


Strata of red hematite are perceptible three miles north of New Bloomfield, on the road to Fulton. The ore crops out in the road for a distance of about twelve feet down the slope. Sandstone is seen both above and below the ore. One-quarter of a mile west, on the same slope and level, stratified ore has been found immediately below the soil, in digging graves in a cemetery. A connection between those two points cannot be traced at present.

Knight Bank, Sec. 2, T. 46, R. 10, W., Callaway County.

A fine outcrop of dense and fine-grained hematite is seen on the eastern slope of the northern low hill, as represented in the sketch, Fig. 54. The ore is over two feet thick, and dips north about twenty degrees. It can be seen only in two places, about twenty feet apart. But the ore seems to be in place, and may therefore be expected to run through the hill. Due east of this hill, small and large pieces and plates of ore are found loose in the bed of the Middle Auxvasse Creek, as indicated on the sketch.

The southern low hill in Fig. 54 shows, at both its northern and southern slopes, outcrops of a ferruginous sandstone, overlaid by



thin seams of red ore. These indications are, however, not sufficient to warrant the presence of workable ore in this southern hill.

UPPER OSAGE DISTRICT.

Brown Bank, Sec. 23, T. 40, R. 24, W., Henry County.—The Brown bank is situated on the dividing ridge between Osage and Grand Rivers. This ridge consists of subcarboniferous rocks. Red, earthy hematite, partly changed into brown and yellow limonite, is found on the surface over a very large area, associated with ferruginous sandstone. The bank is not opened, and the thickness and extent cannot be estimated with any degree of certainty from the present appearances.

Gover Bank, Sec. 16, T. 39, R. 24, St. Clair County.—Large and small fragments of ferruginous sandstone, frequently very rich in iron, together with some brown and red hematite, are spread over

a zone several hundred feet wide, and about one-fourth mile long, across a limestone ridge. Smeltable ore is not now seen in large quantities on the surface; but all the sandstone is so strongly impregnated with oxide of iron, as to give hope that a larger deposit of ore may be discovered in this locality.

Collins Bank, Sec. 23, T. 39, R. 25, W., St. Clair County.—An outcrop of red, earthy hematite, partly somewhat argillaceous, extends over a distance of 200 feet, along a ravine at the foot of a steep slope, on which no rocks are perceptible besides broken chert above the soil.

Marmaduke Bank, Sec. 23, T. 39, R. 25, W., St. Clair County.— Fragments of earthy, red hematite, partly altered into a yellowish-brown, porous limonite, are found on the surface on the summit of a ridge, over an area measuring about 600 feet across, and 400 feet along the ridge. Some of the ore is sandy, and passes into a regular ferruginous sandstone in places. No rock is perceptible on the ridge. The soil is covered with broken flint, which is mixed with the fragments of ore. Most of the ore is good and the fragments large and sharp-edged, indicating the presence of a stratified deposit in the sandstone.

Other banks, of a similar character to those just described, seem to be, the Black Fork and the Lamine banks, both in Cooper County, and the Parkes bank, near Calhoun, in Henry County. The ore of the last-named bank is in the coal-measures, as mentioned above. For the location, etc. of these banks, see Chapter V.

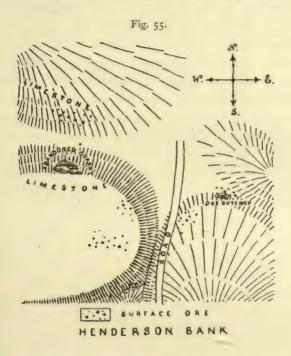
f. DISTURBED OR DRIFTED DEPOSITS OF RED HEMATITE.

I will describe under this head a few either drifted, or at least greatly disturbed, deposits of originally stratified red hematite. Although having at present the appearance of such half-destroyed deposits, they may lead to the discovery of coherent banks when they are more closely investigated.

Murphy's Hill, Sec. 15, T. 45, R. 10, W., Callaway County.— This locality is situated a short distance east of the "Old Digging" bank, which has been above described. No ore is here to be seen in place, but large, somewhat rounded pieces and plates of red ore are found in two ravines, and were evidently washed down from the hill. The hill itself seems to be principally composed of sandstone. Large masses of limestone are, however, projecting

from the lower part of the slopes, apparently between the sandstone.

Henderson Bank, Sec. 12, T. 45, R. 11, W., Callaway County.



The sketch, Fig. 55, shows the occurrences of ore to be observed at the Henderson bank, namely, loose and rounded surface-ore in several places on the two hills, west of the road; loose surface-ore along the road, on the northern slope of the eastern hill; a small and indistinct outcrop of stratified ore at the foot of this hill, near the ravine; and finally, loose surface-ore in the ravine.

The two western hills are composed of encrinital limestone, which is laid bare in several places, and seems to reach the summits, and to be covered only by soil. This soil is thick and copious on the plateau on the northern hill, and is there used for agricultural purposes. Loose ore is sometimes thrown up by the plough on this plateau. The hills are about 40 feet high above the creek. The hill east of the road is thickly covered with a fine sandy soil and by vegetation, and does not show any evidences of rocks. The ore-outcrop at the foot of this hill dips slightly north-west. The ore is a dark-red, fine-grained hematite in thin layers, and is asso-

ciated with layers of chert. The exposure extends, however, over a few feet only, and is therefore too small to allow a reliable judgment regarding its character.

All appearances at the Henderson bank seem to indicate that the encrinital limestone which composes the hills was formerly covered by sandstone-strata, containing strata of ore, and that this sandstone was destroyed and carried off, together with portions of the ore, while other portions of the latter were left, though in a dilapidated condition.

g. DEPOSITS OF LIMONITE ON LIMESTONE.

The distribution of the limonites over the State has been described in Chapter II. All undisturbed limonite deposits are found on limestone. The deposits along the Mississippi lie partly on the Upper Silurian shales and limestones, partly on the Second Magnesian Limestone, according to Shumard's reports. All the other deposits of limonite in the eastern ore-region, as well as those in Franklin County and in the central region, seem to lie on the Third Magnesian Limestone, as far as their position could be ascertained. The same geological position is occupied by the limonites on the Lower Osage and some of those on the Middle Osage, while those in the western parts of Camden and Morgan Counties, and those in Benton County, are on the Second Magnesian, and those on the Upper Osage on the lower carboniferous limestones.

The ore occurs neither in veins, nor in beds, nor as strata, nor in lenticular or other pockets of well-defined limits and regular shape. It is deposited in irregular cracks, pockets, and cavities, either on or near the surface of the various limestones. These cavities have sometimes very large dimensions, in depth as well as width, as will be seen from the following descriptions. In other instances they are quite small; but wherever they exist, they are not single, but a larger number of them is generally found together on a comparatively small space. They are also mostly near the present surface of the ground, and not covered by rock-deposits. The underlying limestone, especially the Third Magnesian, is often sandy, sometimes so much so as to be readily taken for a calcareous sandstone. The rock is more sandy and more loose where it is in immediate contact with the deposit, showing that it has been exposed to dissolving agencies.

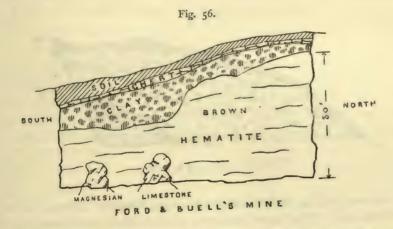
The ore is occasionally mixed with broken chert. In some localities, where the banks are sufficiently opened to make observations possible on this point, the ore in the upper part of the deposits is considerably harder and denser and richer than in the lower part, where it is more inclined to be light, porous, ochrey, and clayish.

This fact, and the invariably stalactitic structure of the ore, are proofs that the solutions from which the ore was deposited have been infiltrated from above. The chemical influence of the carbonate of lime has undoubtedly contributed, in no small extent, to precipitate and deposit the iron in the form of hydrated oxide.

SOUTH-EASTERN LIMONITE-DISTRICT.

The limonite deposits along the Mississippi River are described in Dr. B. T. Shumard's reports on St. Genevieve, Perry, and Cape Girardeau Counties.

Ford Bank, T. 33, R. 7, E., one-half mile from Cornwall Station, on the eastern branch of the Iron Mountain Railroad, in Madison County.—This bank is opened and mined, and is one of the largest and most coherent limonite deposits. The ore-indications on the surface extend about 1,500 feet along a low, flat hill, to a width of about 500 feet. The bank is mined in two different cuts: Fig. 56 represents a section through one of them.



The limestone, which evidently here underlies the ore, is not much uncovered, but is only visible in single, large, rounded masses

of irregular shape. This limestone is so sandy and loose on its surface that it has there the appearance of a soft, calcareous sandstone.

The ore lies immediately above it. It is a limonite, which in this deposit is softer and less distinctly stalactitic than in most other limonite banks. It consists of irregularly-mixed masses of yellow and reddish-brown, porous ores, and of somewhat harder, dark-brown, but generally porous limonite, frequently in botryoidal and mammillary forms. The softer ore passes occasionally into a pure yellow ochre. The best, purest, and hardest ore is in the upper part of the deposit; the softer and ochrey ore is found more in the lower part. But all the ores are mingled irregularly, without any perceptible law or rule, and without any sign of stratification. They are in some places clayish, and contain seams of brown and red clay. The thickness of the ore is very variable and irregular. It is in places only 10 feet and less, and reaches in other places 30 feet.

Above the ore is an irregular layer of reddish-brown clay, fine, pure, and pretty uniform in color, so as to be used as paint. This layer varies in thickness from a half-foot to 15 feet. Above this clay is a layer of broken chert, 2 to 3 inches thick, and above this 1 to 5 feet of soil, enclosing broken chert and surface-ore.

Deal Bank, Sec. 2, T. 31, R. 8, E., Bollinger County.—This bank has been opened somewhat, and presents the following aspect:—



We see in Fig. 57 four successive, very irregular layers, sloping with the hill. The lowest is a mass of solid, chocolate-brown limonite, taking occasionally a bluish color. It is in part mixed with yellow ochre, in part with white or yellow, fine or coarse, broken chert.

The layer above the ore is red clay, with broken veins of ore which enclose broken chert.

Above this is a yellow, sandy clay, mixed with fine chert, and interstratified with layers of this chert.

Above this is a cherty soil, with bowlders of good, hard, and dense limonite.

It may be hoped that the ore in this deposit will turn purer toward the bottom. This can be ascertained only by actual prosecution of the work, because none of the numerous localities where limonite occurs, mixed with chert in this manner, have as yet been fully opened, so as to give a basis for general conclusions in this respect. It seems, however, not unlikely that the chert has come into the ore from above, through the same apertures in which the solutions came, and that the chert has been retained in the upper ore, and that the lower portion of it, therefore, will contain less of it.

Irondale Banks.—Several limonite banks exist in the vicinity of Irondale, in Washington County. The ore is there deposited in numerous small, irregular pockets on the surface of the Third Magnesian Limestone. Larger coherent masses are rarely found.

Other undisturbed limonite banks in the eastern ore-district are the Russell No. 2, McLaughlin, Singer, Dinger, Lindsey, and Love banks in Iron County, the Jessie Lutz and Francis banks in Bollinger County, the Buffum bank in Reynolds County, and the Clarkson, Silvy, and Crane banks in Wayne County. The location, etc. of these banks will be given in Chapter IV.

FRANKLIN COUNTY ORE-DISTRICT.

Moselle Banks.—The various limonite banks in the vicinity of Moselle Furnace seem to have the same general character as the above-mentioned Irondale banks. Some of them, however, contain larger masses of ore, as, for instance, the

Bowlen Bank, N. W. 1/4 Sec. 5, T. 41, R. 2, E., Franklin County.

We have here the following succession of rocks, beginning with the lowest:—

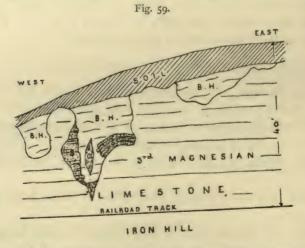
- I. Solid and uniform mass of pure, hard, chocolate-brown limonite, porous, with small, equally distributed pores (B H).
- 2. Clayish limonite, with irregular masses of yellow ochre, soft and friable, and easily crushed into a fine, dry, yellow dust (Cl+H).



- 3. Red loam, with green and gray broken chert (R Cl+F).
- 4. Sandstone, colored and impregnated with oxides of iron, in disturbed and broken layers (S).
 - 5. Dry soil, with some chert.

As no limestone has as yet been struck, the ore will certainly extend to a greater depth, and may prove to form a considerable coherent deposit.

Iron Hill, Sec. 17, T. 42, R. 1, E., Franklin County.—The Iron Hill deposit seems to consist of numerous smaller cracks and cavities on the surface of the Third Magnesian Limestone, which cavities are in part or wholly filled with brown limonite and with yellow ochre. Some of these cavities have been cut through by the railroad-line, three miles west of Moselle Station. The following illustration, Fig. 59, gives a section of one of these:—



The Third Magnesian Limestone is here thickly stratified and very sandy, especially in the lower layers, and near the irregular depressions and cavities. The latter all start from the surface and

reach more or less deep into the body of the limestone. In the deepest of these cavities, in Fig. 59, we find deposited a loose, coarse-grained and ferruginous, thinly-stratified sandstone, which has afterward been broken up again and partly destroyed, perhaps simultaneously with the opening of the crack in which the cavity terminates at its lower extremity. The point of this crack is filled with white clay and with broken, white chert (Cl+F). All the rest of the cavity is nearly filled with limonite (BH), in irregular, botryoidal, and stalactitic forms, mixed with yellow ochre and some The lower part is mostly ochre; the higher portions are harder, and form one coherent, porous mass of limonite, in places mixed with heavy-spar. The thickest and least porous forms of the limonite enclose sometimes a core of pyrites. An oblong space, now filled with an indurated red clay (Cl), exists in the centre of the lower part of the cavity, and seems to indicate that the ore has been formed gradually from the walls of the cavity toward the centre, as well as from the top toward the bottom. There can be no doubt that the infiltration has taken place from above.

Blanton Limonite Bank, S. ½ S. W. ¼ Sec. 29, T. 40, R. I, W., Washington County, on the southern slope of the Blanton Hills.



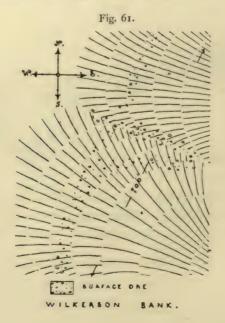
This bank is not opened. The ore seems to lie on the limestone and beneath the sandstone. The surface-ore occurs in pieces and large bowlders, and can be traced about 150 feet down the slope and 60 feet along the slope.

CENTRAL ORE-REGION.

Steelville No. 2 Bank, E. ½ S. W. ¼, Sec. 5, T. 37, R. 4, W., Crawford County.—This is a limonite bank, situated but a few hundred feet north of the Steelville No. 1 bank of specular and red ore.

This limonite bank presents a very fine show of large surface-ore on the eastern slope of a sandstone hill, near its foot. A brown, impregnated sandstone is found above the ore on the same slope, passing into a white sandstone, which forms the summit. It is not now to be seen what rock underlies the ore, because the latter descends to the foot of the hill and into the valley. The surface, over which the ore is spread, is about 400 feet long and 30 to 40 feet wide. A narrow belt of breccia of gray and green chert, cemented by an indurated clay, encircles the ore above, and separates it from the sandstone. Some soft, red ore has been found close to the bank north of it. These facts would indicate that this might be a transformed specular-ore deposit; but the absence of all specular ore and the mineralogical character of the limonite make it more probable that it will prove to be an original limonite deposit on the Third Magnesian Limestone.

Wilkerson Bank, Sec. 34, T. 36, R. 4, W., Crawford County.



As Fig. 61 shows, this bank exhibits a considerable quantity of surface-ore, in bowlders from one-half to two feet in diameter, lying in a curved line along the north-western slope of two hills, 15 to 20 feet below the summit, and being also concentrated in the ravine which separates the two hills. No distinct outcrop of rock can be seen. Pieces of chert are mixed with the surface-ore, and occasionally some sandstone on the southern, and single pieces of limestone on the northern, hill.

Other banks of limonite on limestone exist on Crooked Creek, and in several localities along the Gasconade River, also in numerous places in the southern part of the State. Many of these banks will be found in the list in Chapter V. of this report.

ORE-REGION ON THE OSAGE RIVER.

The most important limonite region in Missouri is on the Osage River. The banks on the Lower Osage, in Miller County, seem to be mostly disturbed, and do not therefore belong in our category g; but very numerous banks of this character exist on the Middle and Upper Osage Rivers, some of which I will now describe.

Furnace Bank, on Boulinger Creek, Sec. 4, T. 39, R. 18, W., Camden County.

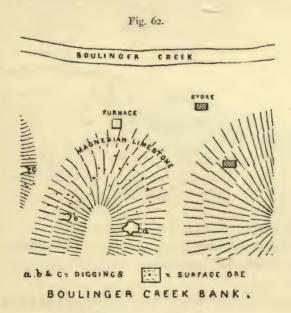
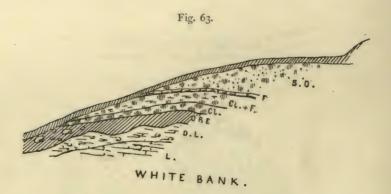


Fig. 62 is a topographical sketch, on a very small scale, showing the extent of the surface-ore, the three openings, a, b, c, made into the deposit by the miner, and the position of the blast-furnace, which has been erected near the foot of the main hill, to smelt the ore. The distance between a and c is about one-quarter of a mile. The openings a and b, out of which a considerable quantity of good limonite has been taken already, have shown that the ore lies on the irregular surface of the Third Magnesian Limestone, which composes the main body of the hill. This limestone is very sandy near the ore, and has there the appearance of a loose, calcareous sandstone. The ore seems to form in some places a layer of irregular thickness on the limestone, and, besides, to fill all the pockets and cavities on the latter. One such cavity, which has been struck in the opening a, has been mined to a depth of 12 feet in the solid and pure ore without reaching the bottom.

White Bank, S. E. 1/4 Sec. 7, T. 39, R. 18, W., Camden County.—The White bank, which has been opened by a tunnel and a shaft, has a great resemblance in its general character to the Furnace bank. The layer of ore on the limestone continues here pretty steadily some distance into the hill; but larger cavities filled with ore have not been met with, so far. The following section, Fig. 63, will give an idea of the position of the ore as seen at present:—



L is the regular Third Magnesian Limestone.

D L is a layer of decomposed limestone, presenting the appearance of a loose, calcareous sandstone; thickness varies from 2 to 30 inches.

Ore. Above this rock is a bed of limonite, I to 4 feet thick, soft

and earthy, enclosing irregular masses of hard, solid ore of more or less stalactitic structure.

- Cl. Dark-red to brown, strongly ferruginous clay or loam, ½ to 2 feet.
- Cl+F. White and green clay in thin and irregular layers, with sand and chert, 1 to 3 feet.
 - F. Layer of white chert, I to 3 inches.
- S. D. Layers of yellow sand and variegated clays and loams, with more or less broken strata of sandstone.

Palm Bank, on Dry Creek, N. W. 1/4 Sec. 12, T. 40, R. 19, W., Morgan County.



We have here a distinct outcrop of ore in the ravine upon the western slope of the hill, close to the foot. The ore is a limonite of good quality, about 4 feet in thickness, and seems to be in place. Around the outcrop, within a radius of 30 feet, is a large amount of surface-ore, which extends in smaller quantities to a distance of 50 or 60 feet up the slope. The soil on the hill is mixed with chert. No other rocks are visible.

Wigwam Bank, Sec. 10, T. 40, R. 19, W., Morgan County.— The ore is a limonite, which is very largely mixed with chert, so much so as to form a breccia in some cases. There are, however, portions of it which are pure. It is found on the western slope of a cherty hill, the lower part of which seems to consist of a sandy, magnesian limestone. The ore extends about 1,000 feet along the slope and 60 feet vertically. Some sandstone is found on the surface of the upper part of the hill, a short distance from the ore and apparently above it.

Cout's Bank, on Flat Rock Branch, Sec. 14, T. 40, R. 19, W., Morgan County.—The ore lies on the east slope of a hill in a zone about 30 feet wide, extending 150 feet down the hill-side. It is a limonite, massive, but frequently mixed with fine, broken chert. A large amount of broken chert is seen on the surface, but there is no rock exposed.

Walker Bank, Sec. 36, T. 41, R. 20, W., Benton County, is situated at the top of a high, cherty hill, over which single pieces of ore are widely scattered. The outcrop of ore is circular, about 20 feet in diameter, and consists of large bowlders, some of which are several feet in diameter. The ore is a limonite of good quality.

Gun Bank, Sec. 33, T. 40, R. 20, W., Benton County.—Here a large amount of surface-ore is scattered for a distance of 50 feet vertically and 500 feet along the northern slope of a low, flat hill. Two test-pits have been sunk and numerous drill-holes, all of which struck the ore at a depth from 4 to 6 feet below the surface.

The ore is a good-quality limonite, and the bank is one of the most promising in this region.

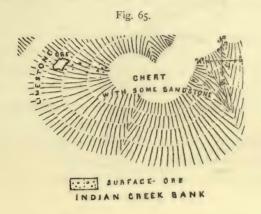
Richwoods Bank, Secs. 3 and 4, T. 39, R. 22, W., Benton County.—Here the ore lies upon the western slope, in a belt about 30 feet wide, and extending some 200 feet up the hill.

Above the ore is a yellow sandstone. The rock below is covered by soil, but at the foot of the same hill, a few hundred yards distant, is an outcrop of limestone, probably the Third Magnesian.

Indian Creek Bank, Sec. 26, T. 42, R. 21, W., Benton County.—The hill on which the ore is found is about 100 feet high. At the base, and extending probably 20 feet vertically, is a horizontal limestone, probably the Third Magnesian.

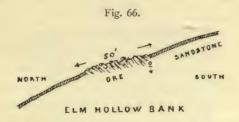
Above this, on the western slope, the surface is covered with chert and pieces of limonite-ore. At one place is a large bowlder of many tons' weight. It is partly formed of fine pipe-ore broken, and the pieces cemented again by ore. Other pipe-ore is mixed with the soil near by. Some surface-ore, though scarce, is found higher up on the southern slope, and on the top of the hill.

Sandstone probably forms the top of the hill above the limestone; the ore will be found, not forming a coherent deposit, but only larger or smaller accumulations in the cavities on the limestone.



Elm Hollow Bank, Sec. 36, T. 41, R. 22, W., Benton County.

—The ore at this place lies upon the northern slope of a hill, over a surface perhaps 50 by 100 feet. It consists of numerous pieces, varying in size, not rounded. Chert is scattered over the whole



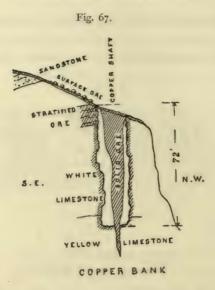
hill, and near the top, above the ore, occur pieces of sandstone. There is a probability of this being a coherent deposit. In some of the larger pieces of the ore a core of pyrites was found.

Copper Bank, Sec. 27, T. 39, R. 24, W., St. Clair County.

This bank is upon the north-west slope of a hill into which a shaft has been sunk to a depth of 72 feet, in search of copper.

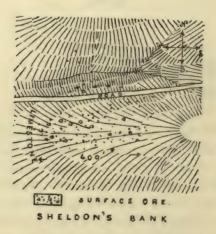
The shaft is in a whitish limestone, probably the encrinital, and follows a crevice which is filled with a soft, earthy limonite. The course of this crevice is north-east and south-west.

At the mouth of the shaft, stratified ore appears several feet thick, and above this is an outcrop of ferruginous sandstone.



Sheldon Bank, Sec. 8, T. 38, R. 24, W., St. Clair County.— This hill seems to consist of a coarse-grained, semicrystalline, gray limestone, filled with encrinites, which crop out over the slope to a height of 80 feet. Higher up pieces of ferruginous sandstone are

Fig. 68.



found scattered. The principal part of the ore is in bowlders, from one to three feet diameter, lying in the midst of large limestone outcrops. It is also found in smaller pieces higher up in the region of the sandstone. On the lower part of the slope the ore is solid limonite, somewhat argillaceous and ochrey, and inclined to stalactitic forms.

Higher up the hill it becomes more sandy. The hill north of this is 150 feet high, of the same geological formation, and shows some surface-ore on its southern slope.

This bank is one of the most promising in this region.

Greenwell Bank, Sec. 15, T. 39, R. 25, W., St. Clair County.

—The ore occurs scattered 70 feet along and 40 feet down the slope of a low, flat hill, which is covered with broken flint and large bowlders of crystalline, gray limestone. Part of the ore is hard and solid, and part is argillaceous.

Besides these already described, the Laclede, the Big and Little Manqua, the Carl, Newman, Turkey Creek, Brown, Gover, Collins banks, and many others in this Osage region, belong to this class of ore-deposits.

h. DISTURBED OR DRIFTED DEPOSITS OF LIMONITE.

Some of the limonite deposits seem to have been disturbed from their position, others partly destroyed, broken up, and re-deposited. We find, therefore, some apparently on the sandstone, others as drifted ore imbedded in the soil, subsoil, or in other destroyed and drifted materials. Few of these deposits are, however, as yet sufficiently opened to allow a reliable judgment regarding their character.

Matthews Mountain, Sec. 3, T. 32, R. 6, E., Madison County.—The prevalent rock here is limestone, with cherty soil, but porphyry-hills are abundant. The porphyry is brown or reddish brown, with crystals of a transparent feldspar. Matthew Mountain is a high, steep hill, seemingly composed of this porphyry, although the rock-mass is covered with a thick layer of porphyritic detritus, with large, sharp-cornered pieces of porphyry. These are mixed with pieces and large, rounded bowlders of dense limonite. Near by, some test-pits have been sunk to a depth of 5 to 10 feet, in which large bowlders of ore were found mixed with the detritus, but the solid rock was not reached.

As no deposits of limonite are known to occur in porphyry, it has probably in this case been drifted into the detritus, from some limestone which has been eroded or washed away.

Poblick's Bank, Secs. 23 and 24, T. 32, R. 8, E., Bollinger County.—The ore is a hard limonite, occurring imbedded with white flint and pieces of sandstone in the soil and subsoil on the southern slope of a high and steep hill. No limestone is seen in this vicinity. The geology of this locality is the same as at the Gilman, Turkey Hill, and Murdoch banks, hereafter to be described. The surface seems to consist of irregular, unconformable, and indistinctly stratified red and yellow clays, mixed with sand and chert, and with pieces of sandstone originating from destroyed strata. This formation, which has evidently a secondary character, and looks as if deposited by greatly agitated waters, probably rests upon the top of the magnesian limestone, which can be seen directly, in a few localities, only in some of the deepest ravines. The soil covering this formation is itself clayish, but mixed with fine and coarse chert as well as with large pieces of flint, and often with pieces of limonite. This surface-ore seems in some places to have no connection with coherent deposits, while in others it has led to the discovery of more substantial deposits of irregular shape.

Gilman Bank, N. W. 1/4 Sec. 1, T. 31, R. 8, E., Bollinger County.—At this place the soil is covered by, and encloses, a large quantity of white chert, mixed with fragments of a white sandstone. The Third Magnesian Limestone is seen cropping out in the bed of Crooked Creek, a half-mile from the ore-bank.

The mass of ore is of lenticular shape, and lies imbedded in yellow clay, mixed with fine, white chert. The ore is broken and fractured into pieces and blocks, which have, however, sharp corners and edges, and are so disposed as to indicate that they must have been at one time parts of the same coherent deposit, and have been broken apart by irregular shrinkage, or by movements of the surrounding masses.

Beds and layers of solid flint, which occur in the ore and clay, are also broken and fractured in a similar manner.

Turkey Hill, N. W. 1/4 Sec. 32, T. 31, R. 10, E., Bollinger County.—The ore here is found at the foot of a steep hill near the bottom of a ravine. It consists of bowlders, large and small, of a generally pure, dark-colored limonite, which are imbedded and irregularly distributed in a yellow, marly clay, containing much white cheft and flint.

A tunnel has been run about 20 feet into the hill, through clay containing ore-bowlders, without reaching the solid rock.

Murdoch Bank, Sec. 16, T. 30, R. 9, E., Bollinger County.—The ore of this bank is found scattered thinly over a surface of 2 to 3 acres on the top of a flat hill, and extending somewhat down the western slope. Near the top, the ore is impure, being mixed with a breccia of flint, the ore serving as the cement. On the slope it grows purer, but scarcer. There has been considerable prospecting done here, by scattered diggings and a shaft, but without disclosing anything but clay and chert.

Other banks, probably belonging to this class, in the eastern ore-region, are the Collins, Leeds Hill, Creder, Orth & Livering, Baker, and Lutz.

FRANKLIN COUNTY ORE-DISTRICT.

Stanton Hill Bank, S. W. ¼ of N. E. ¼, Sec. 36, T. 41, R. 2, W., Franklin County, is a circular depression of about 50 feet diameter, in a dark-colored sandstone which crops out all round, and toward the centre grows very ferruginous, where it has almost the appearance of crystalline, specular ore. The only pure ore found is a limonite. The lines of the strata in the sandstone are very much distorted, and its whole appearance indicates a disturbance from its original position.

OSAGE RIVER ORE-REGION.

In the Osage River ore-region there are very few deposits of class h, but of these the Carpenter and Grissom banks are the most important.

Carpenter Bank, Sec. 12, T. 46, R. 21, W., Benton County.

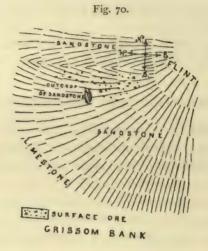


- L. Sandy limestone.
- CL. Ferruginous, reddish-brown loam.
 - C L+F. Clayish sand, with one distinct half-inch layer of fine, white chert.

The ore covers the surface for but a small area, of perhaps 20 feet diameter, but is found scattered in less quantity in various other places on the same flat hill.

There is no ore in the strata of the cut; it is only found in the soil covering the strata, having been drifted there; but it may possibly be only an outlier of some deposit higher up the hill.

Grissom Bank, on Turkey Creek, Sec. 28, T. 46, R. 21, W., Benton County.



The ore lies upon the western slope of a hill 150 feet high. At the foot, limestone is seen in position to a height of about 16 feet. Above this the mass of the hill seems to be of sandstone, covered with pieces of sharp-cornered flint. The ore is in large pieces of irregular shape, very porous, partly of stalactitic fracture. Some of it is sulphurous. A singular feature of this bank is the appearance of the ore in large quantity and exclusively above the sandstone outcrops. From this feature some disturbance was inferred. The Sample and Tuscumbia banks are the only others of class h in this region having come to my knowledge.

CHAPTER VI.

THE IRON-ORES OF MISSOURI.

BY ADOLF SCHMIDT, PH. D.

D. List of Deposits of Iron-ore in Missouri.—Explanation of the Signs used in the List of Deposits of Iron-ore in Missouri.

THE following list contains all the deposits of iron-ore in Missouri which have come to my knowledge. As this list has been made principally for commercial and industrial purposes, the deposits were arranged according to their position along the various routes of transport, railroads or navigable rivers, over which the ores would have to be carried to their respective markets, or to those places where they may be used directly in the manufacture of iron.

They were arranged as follows:—

Deposits	along th	ne Mississippi River	Nos.	I to	6
"	"	Iron Mountain R. R	"	7 to	56
"	"	Atlantic & Pacific R. R. east of			
		Cuba	"	57 to	76
"	"	St. Louis, Salem & Little Rock			
		R. R	"	77 to	125
"	"	Atlantic & Pacific R. R. west			
		of Cuba	"	126 to	181
"	"	Missouri Pacific R. R	",	182 to	198
4.6	**	Osage River	"	199 to	273
"	in othe	er parts of the State	44	274 to	278

The following columns are used in the list:-

Column 1. Consecutive Numbers.

Column 2. Name of deposits, or "banks." These names are

taken either from the present or former owners, or from the lessees, or from people who live in the neighborhood, or from creeks, towns, or counties, or from other objects or circumstances having some connection with these ore-banks.

Column 3. Location of deposits, giving the township, range, and section. These were mostly obtained from the owners or lessees of the banks to which they refer.

Column 4. Counties in which the banks are situated.

Column 5. Names of the owners or lessees, or both.

Column 6. Probable character of deposit. The signs used in this column refer to the division of iron-ore banks, as given and explained under I. C. of the present report, in brief thus:—

- a. Deposits of specular ore in porphyry.
- b. Deposits of specular ore in sandstone.
- c. Disturbed deposits of specular ore.
- d. Drifted deposits of specular ore.
- e. Strata of red hematite.
- f. Disturbed or drifted deposits of red hematite.
- g. Deposits of limonite on limestone.
- h. Disturbed or drifted deposits of limonite.

Ore-banks which were not visited by members of the Survey have no sign in the 6th column.

Column 7. Probable size of deposit. As mentioned under I. A. of the present report, I have divided the various ore-banks in five sizes, according to the number of tons of workable ore they are supposed to contain, from their appearance and condition in summer, 1872. These sizes are:—

- Estimated at less than 20,000 tons.
 20,000 to 100,000 "
- 3. " 100,000 to 500,000 "
- 4. 500,000 to 2,000,000 "
- 5. " more than 2,000,000 "

Column 8. Character of the ore. This column is required, because the 6th column cannot be filled for all the banks, and because many banks, though having, for instance, the general character of specular-ore deposits, contain also red hematites or limonites, produced in the course of time by altering influences.

Column 9. Distances from the nearest railroads or navigable rivers. I have considered this column indispensable, because the immediate industrial importance of the various banks depends in a great measure on their accessibility, and on their distances from the routes of transport. This column is less important for some deposits which are directly connected with iron-works.

EASTERN ORE-REGION OF MISSOURI, -ORE-DISTRICT ALONG THE MISSISSIPPI RIVER.

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FRANKLIN COUNTY ORE-DISTRICT.

No. Name. Location. County. Owners or Lessees. Chark & Cha
1. 40, K. 1, W. 3 cc. 32, S. Washington Hamilton Iron Co. d. I Hematites. A &P.R.R. rom

SCOTIA ORE-DISTRICT.

					:
	74 Scotia, No. 1 T. 38, R. 3, W., Sec. 1, E. 4 Crawford Scotia Iron Co b. 3 Spec. & red	o m.	76 Scotia, No. 2 T. 39, R. 2, W., Sec. 28, S.		, 9 m.
	4	K.K.	K. K.,	8	K. K.,
	8	S F.	κr.	6	κr.
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	cotia			cotia	
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	scotia,	3leedir	scotia,		
	74 S	75 Bleeding Hill T. 38, R. 2, W., Sec. 4, S.	2 94		_

CENTRAL ORE-REGION OF MISSOURI.—STEELVILLE DISTRICT.

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77 N. G. Clark, No. 1.	-
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S. L., S. & L. R. R. R. 8 m.	Specul'r ore S. L., S. & L. R. R.	bec. ore & Limonite S. L., S. & L. R. R. R.		matites. S. L., S. & L. R. R. R., o m	R. R., o m R. specular he-marities S. T., S. & L. R.	R. R., ‡ m ecular & red hema- tites S. & L. R. R. R., ‡ m
<u> </u>	Specul'r ore S.	Spec. ore & Limonite S.	Specular &	matites. S.	Red & specular hematites S.	Specular & red hematitiesS.
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V., Sec. 26, E. Crawford Samuel Knox	Thos. James' estate, Card, Lewis & Co., Lessees	. Thos. James' estate, Card, Lewis & Co., Lessees	Missouri Iron Co., (Crawford, Scott & others)	Missouri Iron Co., (Crawford, Scott	V., Sec. 5, S. 3 Crawford. Lea & Marsh	. Grover & Crawford
Crawford	Crawford	c. 4, W. Crawford	Sec. 5, E. ‡ Crawford	Crawford	Crawford	Crawford
T. 38, R. 3, W., Sec. 26, E.	Cherry Valley, No. 1 T. 37, R. 3, W., Sec. 4, E. 4 Crawford	Cherry Valley, No. 2 T. 37, R. 3, W., Sec. 4, W.	T. 37, R. 4, W., Sec. 5, E. ‡	Steelville, No. 2 T. 37, R. 4, W., Sec. 5, N. E. ‡	Lea and Marsh bank. T. 37, R. 4, W., Sec. 5, S. \(\frac{1}{2} \)	Ferguson bank T. 37, R. 4, W., Sec. 21, N. 4 S. E. ‡, & S. E. ‡ S. W. ‡ Crawford Grover & Crawford. d.
78 Knox bank	Cherry Valley, No. 1	Cherry Valley, No. 2	Steelville, No. 1.		Lea and Marsh bank.	Ferguson bank
78	79	8	81	82	83	\$4

ORE-DISTRICT ON THE UPPER MERAMEC AND ITS TRIBUTARIES.

Character of Distance from the By whom re-	S. L., S. & L. R.	Specul'r ore S. L., S. & L. R.			Specul'r ore S.	Specul'r ore S.		S. L., S. & L. R. S. I., S. & L. R. R. R., ½ m.
Probable size of Deposit.		63	H	J=4	=			4
Probable character of Deposit.	:	ပံ	ď	ਰੰ	rj	p.		· :
Owners or Lessees,	Santee & Wagner, of Rolla	D. Carson, A. & W. James	Missouri Iron Co d.	Missouri Iron Co	Missouri Iron Co	D. Carson, A. & W. James	Thos. James's estate, Card, Lewis & Co.,	N. G. Clark
County.	Phelps	. Phelps	Phelps	Phelps	Phelps	Phelps	Crawford	Crawford
Location.	T. 37, R. 6, W., Sec. 36 Phelps	Winkler bank T. 36, R. 6, W., Sec. 14, S. 4. Phelps	I. T. 36, R. 6, W., Sec. 26, S. W. ‡ S. W. ‡ Phelps.	2. T. 36, R. 6, W., Sec. 26, S. E. ‡.	3. T. 36, R. 6, W., Sec. 26, S. F. 4.	T. 36, R. 6, W., Sec. 35, N. E. ‡	T. 36, R. 5, W., Sec. 32, W. ‡ S. E. ‡, & E. ‡ S. W. ‡ Crawford.	N. G. Clark, No. 2. T. 36, R. 5, W., Sec. 34, W
Name.	Santee bank	Winkler bank	Smith bank, No.	Smith bank, No.	89 Smith bank, No. 3	Lamb bank	Benton Creek	N. G. Clark, No. 2.
No.	85	98	87	88	89	96	16	92

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S. L., S. & L. R.	L, S. & L. R.	L, S. & L. R. R. R. R. R. R. R. R.	S. L., S. & L. R.	L, S. & L, R.	L., S. & L. R.	L, S. & L. R.	L., S. & L. R.	L., S. & L. R. R. R. IA m. Shumard.	L., S. & L. R. R. R., 6 m
<u> </u>	Speculr ore S. L., S. & L. R.	Specul'r ore S. L., S. & L. R.	S	Limonite., S. L., S. & L, R.	Limonite S. L., S. & L. R.	w.	Limonite S. L., S. & L. R. R. R. 12 m	L'monite. S. L., S. & L. R. R. R. 14 m.	Specul'r ore S. L., S. & L. R. R., 6 m
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-:	<u>:</u>	:	· · · ·	<u>ئ</u> .		* :	•	:	
C. C. Cook	W. James & Twit-	Revold Bros., St.	St. Louis, Salem & Little Rock R. R. Co	Wilkerson, on Crooked Creek.	Judge Wm. Key	Clark, Wallace & Anderson	•		Grover & Raney
Crawford		, S. Crawford	N. 4 Crawford	•	W. F.N. Crawford	N. Crawford	Crawford	Crawford	
T. 36, R. 5, W., Sec. 34, S. Crawford C. C. Cook	T. 36, R. 5, W., Sec. 24, S. E. ‡ N. W. ‡ Crawford.	T. 36, R. 4, W., Sec. 18, S. E. ‡	S. L., S. & L.R. bank T. 36, R. 5, W., Sec. 24, N. E. ‡ S. E. ‡, & S. W. ‡ N. E. ‡	T. 36, R. 4, W., Sec. 34, W. 4, N. E. 4, & S. E. 4, N. W. 4	T. 36, R. 4, W., Sec. 27, W. \$\frac{1}{2}\ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \	T. 36, R. 4, W., Sec. 26, N. W. ‡ S. E. ‡	T. 36, R. 3, W., Sec. 15 Crawford	T. 36, R. 3, W., Sec. 36, N. 4. Crawford.	T. 35, R. 4, W., Sec. 2, S. Crawford
93 C. C. Cook bank	94 Craig bank	Arthur bank	S. L., S. & L.R. bank	Wilkerson bank	Key bank	Anderson bank	Dry Creek bank.	bank	Grover bank
93	94	95	96	16	86	66	00	IOI	102

ORE-DISTRICT ON THE UPPER MERAMEC AND ITS TRIBUTARIES.

nearest railroad or by whom re- navigable river,			L., S. & L. R. R. R. I m Shumard		• • • • • • • • • • • • • • • • • • •			÷	
or By	:	:	Shirt			: :			: :
m throad o	Specul'r ore S. L., S. & L. R. R. R. R. 14 m	Specul'r ore S. L., S. & L. R. R. R. 2 m	Specul'r ore S. L., S. & L. R.	I.R.	Specul'r ore S. L., S. & L. R.	S. L., S. & L. R R. R., 16 m.		Speculr ore S. L., S. & L. R.	Specul'r ore S. L., S. & L. R. R. R. 11 m
istance from nearest railroad navigable river.	L., S. & L. R. R., 14 m.	L., S. & L. R. R., 2 m.	S. S.	. ×. ×.	. S.	K. K., 18 m. L., S. & L. R. R., 16 m.		S. S.	F. K., 10 m. L., S. & L. R. R., 11 m.
Distance from nearest railroa navigable river	F. I.	F. 1.	E.L.	, L	1,	K., 1		Η,	7 1. X
H	ore	ore	ore S	Specul'r ore S. L., S. & L.	ore S	:		ore	ore
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Deposit		: 	.S.		Sp	:		S	SF
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s or L	i Iro	ennet		n Jan ames,	arsor c Iron	Carso	T.	3enne tsbur	ames,
Owners or Lessees.	Missouri Iron Co	Graff, Bennett & Co., of Pittsburgh		William James, of St. James	David Carson, of Marramec Iron Works	David Carson	SALEM DISTRICT.	Graff, Bennett & Co. of Pittsburgh	Wm. James, of St. James
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ty.							SALE		
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	. S. :	Sec. 12, S. 4 S. W. 4. Dent		T. 35, R. 4, W., Sec. 33 & 34. Dent.	3, E. Crawford.	T. 35, R. 3, W., Sec. 14 & 16. Crawford		T. 34, R. 3, W., Sec. 9, S.E. ‡ N. W. ‡, & N. E. ‡ S. W. ‡. Dent.	bank T. 34, R. 4, W., Sec. 15 Dent.
	T. 35, R. 5, W., Sec. 4, S. E. ‡	Sec. I	T. 35, R. 5, W., Sec. 32, S. W. 4.	c. 33	sec. I	c. 14		2.9.	ec. I
Location.	```	:	×., ×	V., Se	×.	V., Se		N. E.	W., S
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Name.	ank.	I ban	bank	r ban	Smith	anks.		ank.	Cree
4	Arnold bank,	ckwei	- ban	Fitzwater bank.	upen	rson b		ızzah	ıtchin
·	3 Arr	ro4 Blackwell bank.	10	Fit	107 Reuben Smith bank. T. 35, R. 3, W., Sec. 13, E. 4 S. E. 4	108 Carson banks		109 Huzzah bank	IIO Hutchins Creek
No.	103	IO	105	901	IO	10		10	II

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Isabelle Furnace Co. of Pittsburgh	Card & Lewis, of St. Louis	Card & Lewis.	Card & Lewis	C. C. Simmons, of St. Louis, & Missouri Iron Co	Graff, Bennett & Co. of Pittsburgh	Wm. James, of St. James	Taylor & Clark	Zane, of St. Louis	Missouri Iron Co	Mrs. Ziegler	Cable & Co. of Ohio.	Widow Merriam, of Cleveland, Ohio
	T. 34, R. 5, W., Sec. 12 Dent	T. 34, R. 5, W., Sec. 20, S. E. ‡ N. E. ‡.	T. 34, R. 6, W., Sec. 13, E. 4 S. E. 4 S. E. 4.	Dent		T. 34, R. 6, W., Sec. 10 Dent	T. 34, R. 7, W., Sec. 12, S. W. ‡ S. W. ‡.	N. Dent.	T.33, R. 6, W., Sec. I, S. W. Pont	T. 33, R. 5, W., Sec. 2 Dent	T. 33, R. 4, W., Sec. 26, E. ‡ Dent	T. 32, R. 7, W., Sec. 28, S. 4 Dent
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banl			nk.	Ioun	Youn	ank.	-74	ank.	nk.	k.	oank.	erriam bank, Pigeon Creek.
urgh	bank	bank	rd ba	ons N	rd &	oy b	r ban	ns b	on ba	r ban	dale	am eon (
ittsb	Norris bank	Hayes bank	Orchard bank	immi)rcha	Pomeroy bank	Taylor bank	Wiggi	Jamison bank.	Ziegler bank	Barksdale bank	Merri Pig
III Pittsburgh bank T. 34, R. 5, W., Sec. 13 Dent	112 N	113 I	114	115 Simmons Mountain	116 Orchard & Young b'lk T. 34, R. 6, W., Sec. 27, N. 4 Dent.	117 F	T 811	119 Wiggins bank	120 J	121 2	122 I	123 Merriam bank, Pigeon Creek

SALEM DISTRICT.

Distance from the By whom re- navigable river.			•			•		
By			:	:		•	: :	: :
the	X : X :	_	'n.	n.	nch	nch	. ii	i ii
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fro t rai ble ri	8 4 8 4 8 6 8 6 8 8 8 8 8 8 8 8 8 8 8 8		R. F	R. F	idge	dge	R. I.	R.F
istance from nearest railroad navigable river.	S. L., S. & L. R. R. R., 24 m S. L., S. & L. R. R. R., 24 m		A.& P. R.R., 2 m.	- L 2	on Ridge branc	n Ri	& P. R.R., I m	. P.
Dist ne na	S. S.	_	A.8	A.8	Irol	Irol	A.8	A.8 A.6
Character of Distance from nearest railroa navigable river.	S. L., S. & L. R. R. 24 m Specul'r ore S. L., S. & L. R. R. 24 m		:	A.& P. R.R., I m.	Spec. & red hematites Iron Ridge branch	Specul'r ore Iron Ridge branch	Specul'r ore A.& P. R.R., I m.	Specul'r ore A.& P. R.R., 1 m. Red & spec. hematites A.& P. R.R., 12 m.
ore.	cul',				c. &	cul',	cul'1	Specul'r ore Red & spec. hematites
	Spe		:	:	Spe	Spe	Spe	Spe Red
Probable size of Deposit.			:	:	m	H	H	
Probable character of Deposit.					0	d.	j.	
in a	St. : St.		:	:		ron	:	0 : :
Owners or Lessees.	W. James, of St. James of St. James of St. James		J. P. Card & Co	Senator Buckland		W. A. Dorey, of Iron Ridge	. Iron Ridge Co	Isabella Furnace Co., of Pittsburgh
or Le	es,		d &	uck	o e	rey,	e C	nrng Durg
ners	James	ICT.	Car	or E	Ridg	Do.	Ridg	la F
Own	Jan Jan	STR	P.	enat	uo	Z.A. Rid	on]	abel of I
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	T. 31, R. 6, W., Sec. 13 Shannon Ik. T. 31, R. 6, W., Sec. 16 Shannon		T. 38, R. 5, W., Sec. 13, S.	T. 38, R. 5, W., Sec. 12, N. E. 4. E. 4. R. E. 4. Sec. 20, N. Crawford.	E. 4.	T. 39, R. 5, W., Sec. 29 Crawford	FiF	W. ‡
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Name.	ban Rive	rd	Bu	. e	0	ınk.	ge,	- bank
Z ·	non	S	nk.	bank		y ba	Rid	ban
	Shannon bank Current River ba	5	bankSenator Buckland's	bank.		Dorey bank	Iron Ridge, No. 2., T. 39, R. 5, W., Sec. 33 Crawford.	
No.	124 Shannon bank T. 31, R. 6, W., Sec. 13 Shannon Shannon T. 31, R. 6, W., Sec. 16 Shannon		bank	128		129 1	130 I	132

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Crawford..... N. G. Clark & Co.....

..... A. & P. R.R., 5 m.

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	Dunn A. & P. R. R., 5 m. Limonite. A. & P. R. R., 4 m. Shumard.	Spec. & red		pec. & red	hematites A. & F. K. K., 2 m.		hematites A.& F. K.K., 3 m. oec. & red	hematites A. & P. R. R., 3 m. J A. & P. R. R., 2 m.	A.& P. R.R., 4 m. A.& P. R.R., 8 m.	A.& P. R.R., 6m.	A.&P. R.R., 8 m. A.&P.R.R., 12 m.
P	A.& P. A.& P.	d & P. D	- P	es A. & F.	es A. & P. F	p;	es A.& F.	A.& P.	A.& P. A.& P.	. A.& P.	A.& P. A.&P.F
Spec. & red	Limonite	Spec. & re	Brown&red	Spec. & red	hematit	Spec. & red	hematites Spec. & red	hematit			
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ps Thos. James's estate. b.	Th. James & E	os Lett, of St. James c.	A. James & Dunn	DS Th. James & Co		of Youngstown, O.	ps Thos, James's estate. b.	Agricultural College.		ps Lenox	ps Hall.
Phel	Phel Phel	Phel	Phel	Phel	Pheli		Phel	Phel	Phel Phel	Phel	Phel Phel
134 Meramec bank T. 37, R. 6, W., Sec. I, N. W. 4	Thompson bank T. 37, R. 6, W., Sec. 17, N. 4. Phelps	7 Information bank 1. 38, K. 6, W., Sec. 33, N. E. ‡	Santee & Clark's b'k T. 38, R. 6, W., Sec. 33, S. Whelps	South Mountain T. 38, R. 6, W., Sec. 23 Phelps.	140 Moselle, No. 9 T. 38, R. 6, W., Sec. 29, S. Phelns	,	141 James bank T. 38, R. 6, W., Sec. 29, S. 4 S. E. 4	A. C. L., No. 1 b'k. T. 37, R. 7, W., Sec. 1, W. 4. 3 Railroad b'k, No. 1. T. 37, R. 6, W., Sec. 6	Railroad b'k, No. 2. T. 37, R. 6, W., Sec. 8 Phelps. 5 Railroad b'k, No. 3. T. 37, R. 6, W., Sec. 32 Phelps. 1 Fanor bath v. S. 6. 32 Phelps.	7 Hall bank T 25 R 6 W See F N	
134	135	137	138	139	140		141	142	144	147	148

ROLLA DISTRICT.

By whom re-						
Character of Distance from the By whom re- Ore.	A.&P.R.R., 12m. A.&P.R.R., 12m. A.&P.R.R., 10m.	hematites A. & P. R. R., 7 m. pec. & red hematites A. & P. R. R., 6 m.	A.& P. R.R., 6 m. A.& P. R.R., 2 m. A.& P. R.R., 1 m. A.& P. R.R., 1 m. Specul'r ore A.& P. R.R., 4 m. Snem'r, ore A.& P. R.R., 4 m.	pec. & red hematites A. & P. R. R., ‡ m.	ed& br'wn hemat. & carb'nates A.& P. R.R., 2 m.	Red & spec. hematites A.&P. R.R., 14 m. A.& P. R.R., 3 m. A.& P. R.R., 6 m.
	S. Cara	hematites Spec. & red hematites	Specul'r ore	Spec. & red hematites	~	~ ::
Probable size of Deposit.		. H		. H	н	н ::
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Owners or Lessees.	Dr. Hyer. V. James. Coleman J.H. Brown & Co. of	Major Kelly, of Rolla	W. James, of St. James Baird Hudgeons Thomas Jones' estate b. Thomas James & Co.	Senator Buckland, of St. Louis; W. James, Lessee	Major Kelly, of Rolla	Taylor, of St. Louis. Love A. & P. R. R. Co
County.	Phelps Phelps Phelps Phelps	Phelps	Phelps Phelps Phelps Phelps	Phelps	Phelps	
Location.	T 36, R. 7, W., Sec. 26, W. ½ Phelps. T. 36, R. 7, W., Sec. 29, S. ½ Phelps. T. 36, R. 8, W., Sec. 34, N. ½ Phelps. T. 36, R. 8, W., Sec. 20, N. E. ‡.	I T. 36, R. 8, W., Sec. 18, E. 4 Phelps	T. 36, R. 9, W., Sec. 23, N. 4] Phelps. T. 37, R. 9, W., Sec. 35, E. 4 Phelps. T. 37, R. 9, W., Sec. 36, N. 4 Phelps. T. 37, R. 8, W., Sec. 33, S. 4 Phelps. T. 7, R. 8, W. Sec. 32, Phelps.	T. 37, R. 8, W., Sec. 29, S. 4, 77, R. 8, W. Sec. 20, S. 4	Taylor's Rolla bank. T. 37, R. 8, W., Sec. 15, S.	162 Camp Creek bank T. 38, R. 8, W., Sec. 32, N. 4 Phelps. 163 Railroad bank, No. 4 T. 38, R. 8, W., Sec. 12, E. 4 Phelps.
Name.	Hyer bank Cold Spring bank. Coleman bank Moselle, No. 10	Kelly bank No.	154 Piney Creek bank 155 Baird bank 156 Hudgeons bank 157 Beaver Creek bank		Taylor's Rolla bank.	Camp Creek bank Railroad bank, No. 4,
No.	149 150 151 152	153	154 155 156 157	159	191	162

ORE-DISTRICT ON THE MIDDLE GASCONADE RIVER.

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164 Frost bank T. 37, R. 10, W., Sec. 13, E. 4 Phelps.	165 A. C. L. bank, No. 2 T. 37, R. 10, W., Sec. 23 Phelps	Railroa	Bee Cre	Morgan		Waynes		Mocassi		171 Tames Pipe bank. T. 27. R. 12, W., Sec. 21 Pulaski.		172 Hancock bank T. 28, R. 12, W., Sec. 14 Miller.	P.		174 Bear Creek bank T. 36, R. 14, W., Sec. 25 Laclede .
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WHITE RIVER DISTRICT.

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30, R. 19, W., S. 29, R. 24, W., Set 27, R. 23, W., Set 27, R.	27, R. 23, W., Sec. 27, R. 24, W., Sec. 27, R. 24, W., Sec. 25, R. 25, W., Se
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Mark bank T. 29, R. 19, W., S. bank T. 29, R. 24, W., Seilson Creek bank. T. 27, R. 23, W.,	— bank
St. Mark bank T. 39, R. 19, W., S. L. bank T. 29, R. 24, W., Se. Wilson Creek bank. T. 27, R. 23, W.,	178 — bank. T. 27, R. 24, W., Sec. 19, W. ‡ Christian. 179 — bank. T. 27, R. 24, W., Sec. 14 & 15 Christian. 180 Smart bank. T. 27, R. 24, W., Sec. 24, E. ½ Christian. 181 Railroad bank No. 6. T. 25, R. 25, W., Sec. 2, N. Barry. Atl. & Pac. R.R. Co. Limonite. A. & P. R.R., 3 m. Swallow. Atl. & Pac. R.R. Co. Limonite. A. & P. R.R., 3 m. Swallow. W. ‡. Sec. 2, N. Barry.

ORE-DISTRICT ON THE LOWER GASCONADE RIVER.

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ORE-DISTRICT ON THE LOWER GASCONADE RIVER,

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Swallow.		Limonite Osage river, ‡ m Broadhead. Limonite(& Osage river, 4 m.	Meek.		
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ORE-DISTRICT ON THE LOWER OSAGE RIVER.

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ORE-DISTRICT ON THE UPPER OSAGE RIVER.

By whom reported.	
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Owners or Lessess.	R. H. Melton. Greenwell.
County.	Henry St. Clair.
Location.	T. 40, R. 24, W., Sec. 23 Henry. T. 39, R. 24, W., Sec. 16 St. Clair T. 39, R. 24, W., Sec. 29 St. Clair T. 39, R. 24, W., Sec. 28 St. Clair T. 39, R. 24, W., Sec. 27 St. Clair T. 38, R. 24, W., Sec. 27 St. Clair T. 38, R. 25, W., Sec. 15 St. Clair T. 39, R. 25, W., Sec. 15 St. Clair T. 39, R. 25, W., Sec. 23 St. Clair T. 39, R. 25, W., Sec. 23 St. Clair
Name.	Brown bank. Gover bank. Bradshaw bank. Copper bank. Greenwell bank. Greenwell bank. Marmaduke bank.
No.	265 266 267 271 273 273

OZARK COUNTY DISTRICT.

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., Sec. 36 Ozar	., Sec. 34 Ozar	., Sec. 29 Oza	., Sec. 23 Oza	a .
W., Sec. 36 Ozar	W., Sec. 34 Ozar	W., Sec. 29 Oza	W., Sec. 23 Ozz	W.
12, W., Sec. 36 Ozar	15, W., Sec. 34 Ozar	12, W., Sec. 29 Oza	12, W., Sec. 23 Oza	II, W.
R. 12, W., Sec. 36 Ozar	R. 15, W., Sec. 34 Ozar	R. 12, W., Sec. 29 Oza	R. 12, W., Sec. 23 Ozz	R. 11, W.
2, R. 12, W., Sec. 36 Ozar	2, R. 15, W., Sec. 34 Ozar	3, R. 12, W., Sec. 29 Oza	3, R. 12, W., Sec. 23 Oza	6, R. 11, W.
22, R. 12, W., Sec. 36 Ozar	22, R. 15, W., Sec. 34 Ozar	23, R. 12, W., Sec. 29 Oza	23, R. 12, W., Sec. 23 Oza	26, R. 11, W.
T. 22, R. 12, W., Sec. 36 Ozar	T. 22, R. 15, W., Sec. 34 Ozar	T. 23, R. 12, W., Sec. 29 Oza	T. 23, R. 12, W., Sec. 23 Ozz	T. 26, R. 11, W.
T. 22, R. 12, W., Sec. 36 Ozar	T. 22, R. 15, W., Sec. 34 Ozar	T. 23, R. 12, W., Sec. 29 Oza	T. 23, R. 12, W., Sec. 23 Oza	T. 26, R. 11, W.
T. 22, R. 12, W., Sec. 36 Ozar	T. 22, R. 15, W., Sec. 34 Ozar	T. 23, R. 12, W., Sec. 29 Oza	T. 23, R. 12, W., Sec. 23 Oza	T. 26, R. 11, W.
T. 22, R. 12, W., Sec. 36 Ozar	T. 22, R. 15, W., Sec. 34 Ozar	T. 23, R. 12, W., Sec. 29 Oza	T. 23, R. 12, W., Sec. 23 Ozz	E T. 26, R. 11, W.
T. 22, R. 12, W., Sec. 36 Ozar	T. 22, R. 15, W., Sec. 34 Ozar	T. 23, R. 12, W., Sec. 29 Oza	T. 23, R. 12, W., Sec. 23 Ozz	mk T. 26, R. 11, W.
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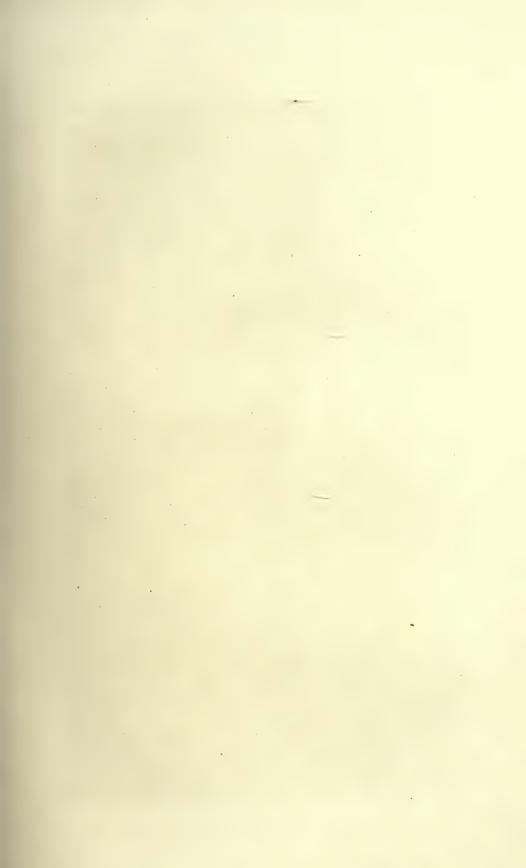


Photo by Childs.

JACKSON MINE Looking West in Nº1 Pit.

J. Bien, lith.

PART II.

IRON-BEARING ROCKS (ECONOMIC)

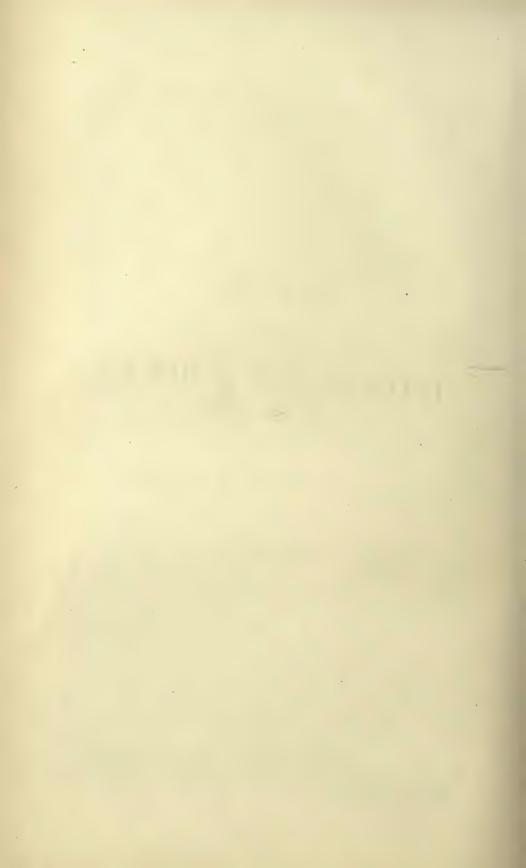
OF THE

UPPER PENINSULA OF MICHIGAN.

BY

T. B. BROOKS.

Memorandum.—The Appendices referred to in this part are found in Vol. II. of Michigan Geological Survey, 1873.



ATLAS PLATES

REFERRED TO IN PART II.

- 1. Map of the Upper Peninsula.*
- 2. Map of the Central Portion of the Upper Peninsula.
- 3. Map of the Marquette Iron Region.
- 4. Map of the Menominee Iron Region.
- 5. Map of Iron Mines at Negaunee.
- 6. Map of Republic Mountain and vicinity.
- 7. Map of the Champion Mine.
- 8. Map of the Washington and Edwards Mines.
- 9. Map of the Lake Superior and Barnum Mines.
- 10. Map of the New York Mine.
- 11, Magneto-geologic Chart of Republic Mountain.
- 12. Statistical Tables of the Iron Ore Trade and Iron Mines of the Upper Peninsula.
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^{*} This map refers also to Parts II. and III.



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INTRODUCTION.

It is customary to preface Geological Reports with a history of the surveys on which they are based; in this case, however, it will be impossible to give more than a brief sketch, without omitting some part of the report itself, the limits of the book, for the publication of which funds were provided, having already been considerably exceeded.

The first survey of the State by Dr. Houghton, which was discontinued on account of his death by drowning in Lake Superior in 1845, is noticed in the first chapter in connection with the discovery of iron ore. The present survey was inaugurated by act of the Legislature in 1869, which appropriated \$8,000 per year for the work, one-half of which went to the Upper Peninsula. This amount was again divided equally between the Iron and Copper Regions, which gave \$2,000 per year for each to cover all expenses, including salaries, supplies, instruments, travelling, etc. To the \$8,000 aggregate for four years from this source, the Geological Board added \$1,000 for chemical work, making \$9,000 in all received by me from the State for the survey of the Iron Region. In addition to this sum I have expended about \$2,000 of my own means, and have not received any compensation for my services.

This small sum would have been inadequate to have accomplished anything worthy of the importance of the work undertaken, had not several corporations and individuals generously come to my relief; indeed on this source of help I counted largely in undertaking the work, and made it an express condition in the arrangement that I should be permitted to avail myself of all the assistance of this kind I could obtain, and also that during the progress of the work I should be free to continue the practice of a profession from which I was sure to obtain further facts bearing on the objects of the survey.

The companies which have contributed valuable data in their

possession, or have instituted special surveys at my suggestions, with the view of furthering the object of the survey, are:—The Marquette, Houghton & Ontonagon Railway, The Portage Lake & Lake Superior Ship Canal Co., The Republic, Washington, Lake Superior, Champion, New York, Spurr Mountain, Iron Cliff, Cannon and Magnetic Iron Companies. E. Breitung bore a part of the expense of making Map No. V., and John Fritz, A. Pardee, and Daniel J. Morrell, of Pennsylvania, S. P. Ely, of Marquette, and A. B. Meeker, of Chicago, contributed generously to the chemical fund, the results of the analyses being given in Chapter X.

The law of 1869 established a Board of Survey, consisting of H. P. Baldwin, Governor; W. J. Baxter, President of the Board of Education, O. Hosford, Superintendent of Public Instruction, with power to select the Geologists, disburse the money appropriated, and perform other necessary duties. Prof. A. Winchell was made Director, who approved the plan for the survey of the Iron Region which I submitted to him, and which is contained in the following letter:

LETTER OF INSTRUCTIONS,

Referred to in Agreement with T. B. BROOKS, dated Negaunee, Mich., June 5th, 1869.

- " To Major T. B. BROOKS, Assistant of the Geological Survey of Michigan.
- "SIR:—You are hereby authorized and requested to make a Survey of the Marquette Iron District, and to draw up a report on the same, substantially in accordance with the following suggestions:
- "I. By the Marquette Iron District is meant the region embracing all the deposits of iron ore extending from the shore of Lake Superior on the east, through Townships 46, 47, and 48 north, as far as Range 31 west, inclusive, being the region which for the present finds its outlet by railroads through Marquette and Escanaba.
 - "Your report on this district would appropriately furnish-
 - " 2. A historical sketch of discovery in the Iron Region of Lake Superior.
- "3. A physiographical sketch of the Marquette Iron District; general topography, hydrography, timber, soil, climate, etc.
- "4. The general geological structure of the district (not entering into details, nor theoretical discussions); identification of iron range stratification; outline description of the rocks; general description of the ores of iron occurring in the district.
 - "5. The mines in general; their distribution and grouping.
- "6. Special notices of the mines and mining locations of the District; local structural geology, topography, mineralogical specialties of the ores,

- "7. Discovery of ores; geological principles applicable; the use of instruments.
- "8. The working of Iron Mines; methods in use here and elsewhere in analogous regions; advantages of each; machinery.
- "9. The manufacture of iron and steel; special adaptations of the different varieties of ore in the District; the use of charcoal and mineral coal; resources of charcoal in Michigan; manufacture of charcoal; fluxes; location of furnaces; construction and operation of furnaces.
 - "10. Transportation of iron ores, and of iron; market; prices.
 - "II. Commercial statistics of iron ores, and of iron.
- "In the discussion of the above topics, it is intended that you make such reference to other iron regions as may be necessary to thorough treatment and illustration of the general subject.
- "It is not intended to lay down any stringent rules for your procedure, but only to furnish a general conception of the ground to be worked over. It is desired to produce as complete a manual as possible of information relating to the finding, extracting, transporting, and smelting of the iron ores of the Lake Superior Region, and it is believed that your own experience and the suggestions which may occur to you in the progress of the work will render it proper to deviate from the letter of the foregoing programme, according to the dictates of your own judgment. Specimens are to be collected according to the requirements and provisions of the law of 1869.
- "In the prosecution of your field work, it is obvious that you cannot with the money at your disposal enter into detailed and complete examinations of individual properties, but it will promote the interests of the general work, if proprietors can be induced to defray the expenses of such detailed surveys beyond the limits to which you may be able officially to prosecute them; and it is evident that the interests of proprietors, no less than those of the State, will be promoted by committing such detailed surveys to your direction.

"The report, with the requisite maps, plans, and other illustrations, is to be ready for publication by the 31st day of December, 1870.

"(Signed) A. WINCHELL,

"Director Geological Survey,
"Ann Arbor, Mich."

On the completion of this survey of the Marquette Region, the Board decided to extend the work over the Menominee Region as well as further West before publishing, thus embracing all the known iron-fields of the Upper Peninsula. Professor Winchell having resigned in 1871, this part of the work was done under the direction of the Board.

Prof. R. Pumpelly has been engaged, with interruptions, in the Copper Region during the same period I have been at work in the Iron (see his Report, Part II.), and in the spring of 1871 Dr. C. Rominger commenced work on the Palæozoic rocks; his Report on the Silurian rocks of the Upper Peninsula is contained in Part III. of this volume.

The sum appropriated (\$20,000) for publishing 2,000 copies of the three reports, with Atlas of maps, enabled the Board to contract for no more than a 500 octavo-page volume, which at the time was deemed sufficient space. I have been generously allowed more than one-half this space, but find that it was not sufficient to contain the material which I had accumulated, and which it seemed to me could be advantageously embodied in the proposed report. It was for some time a question with me, whether I should attempt to consider all the points named in the above scheme (giving each its relative space), which plan would have excluded a large amount of valuable material, or whether I should only attempt to treat each subject in order, as fully as my material would admit and its importance seemed to demand, without attempting at this time to cover the whole ground.

I choose the latter plan, and have in consequence been obliged to entirely omit all consideration of the important subjects of the location, construction and operation of furnaces; of fuels, fluxes, and ore mixtures; of the resources and manufacture of charcoal in Michigan,* as well as the consideration of the question of steel manufacture. The question of the transportation of ore and iron, of markets and prices, was also forced out for want of space. A proper treatment of these subjects would fill a volume.

I trust those gentlemen, who have favored me with lengthy and carefully prepared replies to my numerous inquiries on these excluded subjects, will feel that no injustice has been done them in withholding their papers, until they can be properly presented.

The following named gentlemen are well acquainted with their respective localities on the Lower Peninsula, and are prepared to give information regarding the timber, etc., which is in many instances unsurpassed:

Joseph Dame, North Port.
E. E. Benedict, Manistee.
E. B. Mills, Mayville.
Geo. N. Smith, Bear river.
W. H. C. Mitchell, East Traverse bay.
Leroy Warren, Pentwater.
J. S. Dixon, Charlevoix.

O. W. HART, Torch lake.
A. G. BUTLER, Frankfort.
JAMES LEE, Bingham.
W. H. HURLBURT, South Haven.
DENNIS T. DOWNING, Little Traverse.
DELOS L. FILER, Ludington.
WILLIAM H. FREY, West Olive.

^{*} The subject of the resources of Michigan in Charcoal and the location of charcoal furnaces both on the Upper and Lower Peninsulas has been carefully worked up and illustrated by Timber Maps, but there is unfortunately no means provided for their publication.

It may be questioned, whether with the purely practical object I have had in view in preparing this report, and the limited space, that so large a place should be given to the subject of Lithology, so ably treated by Mr. Julien, in App. A, Vol. II. The reasons which led to this were my own inability to properly treat this subject, its great relative importance in the study of rocks devoid of fossils, but above all I had collected and catalogued during seven years a more complete suite of specimens from the Azoic of the Upper Peninsula, than had before been got together, which collection I believed worthy the study and paper referred to, and which I saw no better way of utilizing to the public, than as has been done It is open to question whether Mr. Julien's paper should not have been published through some scientific channel, rather than in an industrial report, where it will stand nearly alone as a contribution to science.

Grouping Iron Deposits.—It has been found convenient in this report to disregard such political divisions as counties and towns in designating localities, and to employ instead, either the precise and simple method of U.S. linear surveyors, which can be readily understood by an inspection of Maps II., III, and IV. of Atlas; or, by the use of what may be termed the mineral or industrial geography of the Upper Peninsula, by which it is conveniently divided into regions, districts, groups, etc., which, although not sharply defined, may be considered at present to have the following boundaries: The Marquette Iron Region (see Map III., Table XIII., and Chap. IV.) embraces all the developed iron mines of the Upper Peninsula, the ores of which now find their outlets via Marquette, L'Anse and Escanaba by the Marquette, Houghton and Ontonagon and Chicago and Northwestern railroads. This again is subdivided into the (1) Negaunee, (2) Michigamme, (3) Escanaba, and (4) L'Anse districts. These divisions may be conveniently carried still further by a subdivision of the Negaunee District into the Cascade Range, Negaunee Hematite Mines, Ishpeming Group, New England and Saginaw Range; and of the Michigamme District into the Washington, Champion, Spurr and Magnetic Ranges, and Republic Mountain Basin. The S. C. Smith is the only worked mine in the Escanaba District, and no ore has yet been shipped from the L'Anse District or Range. The Menominee Iron Region (see Map IV. and Chap. IV.), which as yet has sent no ore to market,

is divided into (1) The North Belt in south part of T. 42, (2) The South Belt in Ts. 39 and 40, and (3) The Paint River District. The Lake Gogebic and Montreal River Region or Range (Chap. VI.) is so little known that it may be questionable whether it should have a place in this economic grouping; it embraces the country between Lake Gogebic and the west boundary of Michigan, and is 100 miles west of the Marquette Region.

It but remains for me to express my obligations and gratitude to the many gentlemen who have contributed in various ways to the objects of this survey, to officially acknowledge their services and to thank them cordially for myself and on behalf of the Board for what they have done.

To S. P. Ely, of Marquette, the survey is more deeply indebted than to any other person; indeed, I would not have undertaken the work except from assurance of his support, which has been constant and generous from the beginning. To Messrs. H. B. and F. L. Tuttle, of Cleveland, Ohio, I am indebted for a considerable amount of the material embodied on Statistical Tables XII, and XIII. of Atlas, much of which I believe it would have been impossible for me to have procured, except through them; App. J, Vol. II., contains a letter from H. B. Tuttle, who has always, with great promptness and care, answered my various inquiries. To Major Fayette Brown, Cleveland, the survey is indebted for a most valuable paper on the amount of air required by charcoal furnaces and the mode of applying it, based on his experience with the Jackson Co.'s furnaces at Fayette, the almost unparalleled success of which gives his statements great value. S. L. Smith, on the part of the Marquette, Houghton & Ontonagon railroad, placed all the results of that company's explorations, made under my direction, at the disposal of the survey. J. J. Hagerman, Milwaukee, furnished a statement regarding the working of Lake Superior and Iron Ridge, Wisconsin, ores with anthracite and coke, and the successful use of the metal in making rails. John L. Agnew has furnished drawings of the new charcoal furnace, superintended by him at Escanaba, 50 feet high and 12 feet bosh, the largest, so far as I know, in the world. M. R. Hunt, Depere, Wis., has given full details of a remarkable long and successful blast of the First National Iron Co.'s furnaces.

The Historical chapter has been made far more complete and reliable than would otherwise have been possible through the contribution of facts and documents by Messrs. William and John Burt, Messrs. Everett, White, Harlow, Hewitt, and Ely, of Marquette; also by Messrs. Jacob Houghton and Charles T. Harvey. This chapter was rewritten by Charles D. Lawton.

I am indebted to so many persons for the facts embodied in the chapter on Mining, that I can only mention W. E. Dickinson, J. C. Morse, William Sedgwick, A. Kidder, Peter Pascoe, George and Eugene St. Clair, and D. H. Merritt, of Marquette county, and Prof. R. Akerman, of Stockholm, Sweden.

C. H. V. Cavis, S. H. Selden, and George P. Cummings, civil engineers, have greatly aided in the work by their personal efforts in procuring information which is embodied in the maps. The valuable explorations of C. E. and Frank Brotherton, and of A. M. Brotherton, deceased, made for the C. & N. W. and M. H. & O. roads, has been to a large extent placed at my disposal by the officers of these companies.

The nature of the valuable scientific aid given to this work by Alexis A. Julien, Prof. R. Pumpelly, Dr. T. S. Hunt, Prof. George J. Brush, Dr. H. Credner, and Charles E. Wright, are explained in the text of chapters III., V., VI., and in Appendices A, B, and C, Vol. II.

Edwin Harrison, of St. Louis, has given me full and detailed statements regarding the working of his Irondale furnace, which has one of the best records ever made by a charcoal furnace. Robert Wood has prepared most of the manuscript for the press, and, with Mr. Bien, will take care of the publication and indexing.

The survey is indebted to the University of Michigan, Ann Arbor, for the use of rooms without charge, and for the same courtesy (most cordially extended) to the School of Mines, Columbia College, New York, on which institution the survey had no claim.

The Marquette, Houghton & Ontonagon, Chicago & Northwestern, Michigan Central, the Great Western, and Grand Trunk railways have in every instance, when requested, granted passes to persons connected with the survey.

To the gentlemen and companies above named, as well as to Messrs. J. N. Armstrong, American Iron & Steel Association, S. C. Baldwin, William H. Barnum, J. B. Britton, C. M. Boss, J. R. Case, Mr. Childs, Girard Iron Company, C. H. Hall, A. Heberlein, Alexander L. Holley, E. C. Hungerford, Prof. Hayden, Gilbert D.

Johnson, F. B. Jenny, Prof. J. P. Leslie, J. S. Lane, A. W. Maitland, David Morgan, Capt. H. Merry, F. W. Noble, Charles H. Pease, New York Mine, J. R. Orthey, Freiburg Royal School of Mines, James M. Safford, Samuel Thomas, J. M. Wilkinson, H. N. Walker, Walter Williams, Capt. R. D. Weston (deceased), Washington Mine, Dr. White, and Charles R. Westbrook, who have in various ways promoted my work, I am under great obligations. Without their aid this report could not have been prepared. I have forwarded to the Board of Survey a full list of their names and addresses, with the request to furnish each with a copy of this Report and accompanying Atlas.

CHAPTER I.

HISTORICAL SKETCH OF DISCOVERY AND DEVELOPMENT.*

Note, -Statistical Tables XII. and XIII. of Atlas contain many facts relevant to this subject, which could not well be incorporated in the text.

MINERAL explorations along the south shore of Lake Superior began at a very early period, and the existence of copper was made known to the world as long ago as 1636, by La Garde, in a book published in Paris. During the subsequent portion of the 17th century frequent mention is made in the "Relations" of the Jesuit Fathers of the finding of this metal.

These Relations † extend from 1632 to 1672, and are made up of the reports or simple narratives of these humble but zealous missionaries, scattered as they were all over the region of the great Lakes, then controlled by the French Government, and are necessarily of inestimable value to the historian and archæologist; and also contain much that is highly interesting to the geologist, as indicating the early discoveries of minerals and the knowledge of their localities and uses, possessed by the natives. In illustration of the allusions to copper found in these reports, we quote simply from one, Claude Allouez, who seems to have been a man of intelligence, as well as one of the most persevering and deserving of these early missionaries. He first visited Lake Superior in 1666, and makes mention of a large mass of copper to be seen near the shore of the lake, with its top rising above the water, giving an opportunity for those who passed that way to cut pieces from it. The writer says, this "rock" has disappeared, having become buried, as he opines, beneath the sands, through the action of the waves. He also states that pieces of copper weighing from 10 to 20 lbs. are frequently found among the savages, who esteem them

^{*} C. D. Lawton, Esq., rendered much assistance in the preparation of this chapter.

[†] These valuable documents have been republished by the Canadian Government.

as domestic gods, and hold them in superstitious awe, preserving them, in some instances, time out of mind, among their most precious articles.

In 1672, a map was published in Paris of this region, which was made by these early Jesuits, and on which is represented 1,600 miles of coast and many islands, with what may be considered remarkable accuracy.*

In 1689, Baron La Houtan, in a book relating to travels in Canada, mentions that "upon Lake Superior we find copper mines, the metal of which is fine and plentiful; there being not a seventh part base from the ore."

In 1721, P. De Charlevoix described the native copper deposits, and the superstitions which the Indians had in regard to them, in considerable detail. The occurrence of native copper being so frequent, the wonder of the early voyageurs was naturally excited, being increased also by vague rumors (gathered from the savages) of the existence of gold, silver, and diamonds.

In 1765, Captain Jonathan Carver visited Lake Superior, and in his account dwelt so largely on the abundance of native copper, that a copper company was formed in England in 1771, which actually began mining operations on the Ontonagon river, under the direction of Mr. Alexander Henry, who seems to have been a better historian than miner; for he gives a detailed account of the winding-up of his operations in 1772 and concludes, as the result of his unsuccessful experiment in mining, that the country must be cultivated and peopled before the copper can be profitably mined.

In 1819, Mr. H. R. Schoolcraft accompanied as mineralogist and geologist a government exploring expedition along the south shore of Lake Superior, having for its object the investigation of the copper mines.

In 1823 another government expedition, under charge of Major Long, passed along the north shore of the Lake, having come from the northwest; and mention is made of their having observed copper boulders in the region of the headwaters of the Mississippi.

Steps had been taken with a view to an exploration of this region

^{*} A fac-simile of this map, and much other interesting matter relating to the early history of the copper region, may be found in Foster and Whitney's Report, Exec. Doc., 1850, Part I.

during the Presidency of John Adams, but nothing was ever The work of systematic, scientific exploration of the Upper Peninsula of Michigan was first undertaken by Dr. Douglas Houghton, the earliest State Geologist. Dr. Houghton had commenced his examination of this region in 1831, and in his first annual report to the Legislature in 1841 presented the results of his labors up to that period in so able a manner, that the attention of the world became directed to the Northern Peninsula with greatly increased interest. In 1840, Dr. Houghton wrote to the Hon. A. S. Porter, under date December 26th, regarding the mineral wealth of the south shore of Lake Superior: "Ores of zinc, iron and manganese occur in the vicinity of the shore, but I doubt whether either of these, unless it be zinc and iron, is in sufficient abundance to prove of much importance. Ores of copper are much more abundant than either of those before mentioned, and a sufficient examination of them has been made to satisfy me that they may be made to yield an abundant supply of the metal,"

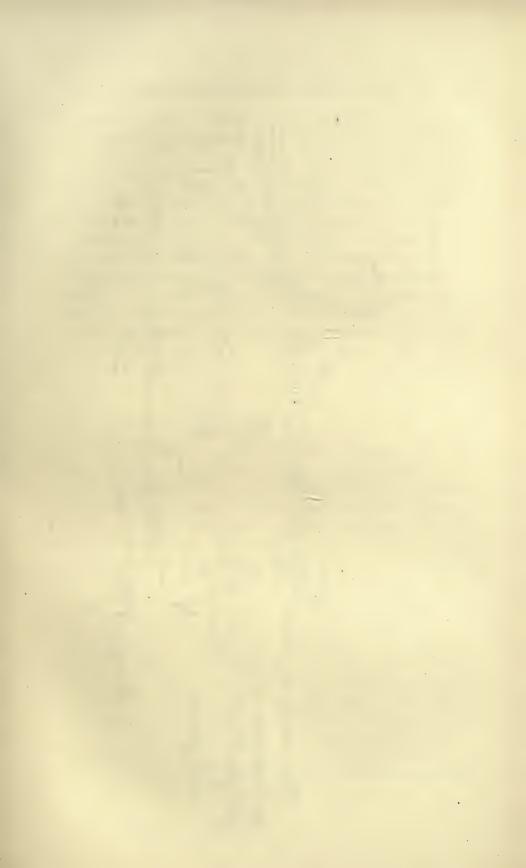
In his Geological Report of 1841 Dr. Houghton says: "Although hematite ore is abundantly disseminated through all the rocks of the metamorphic group, it does not appear in sufficient quantity at any one point that has been examined to be of practical importance." At this date Dr. Houghton had traversed the south shore of Lake Superior five times, in a small-boat or canoe, on geological investigations. It is therefore probable that up to 1841 no Indian traditions worthy of credence, in regard to large deposits of iron ore, had come to his knowledge. As there are, so far as known, no considerable outcrops of iron ore, which come nearer than seven miles to the shore of the Lake, it is plain that investigations, based on observations taken along the shore only, could have determined no more than its probable existence, which is plainly indicated in the extracts given. Dr. Houghton was not aware of the existence of iron ore in quantity, until the return of Mr. Burt's party of survevors to Detroit in the fall of 1844, his examinations in the interior of the country having been confined to the Copper Region. Attention at that early period was entirely directed to searching for ores of more value than iron, and it is worthy of remark, that the Jackson and Cleveland Iron Companies, which were the first two organized, were formed to mine copper, silver, and gold.

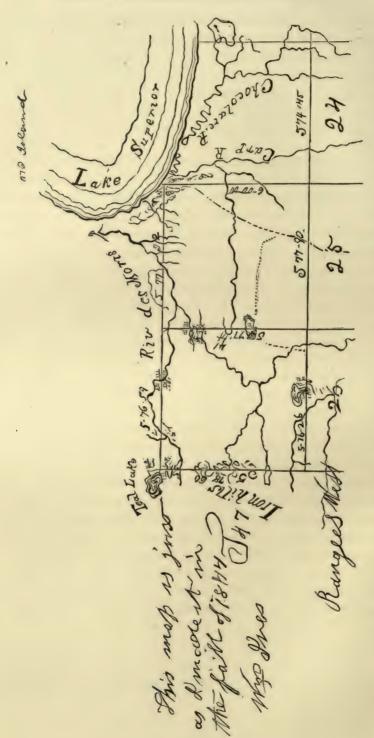
The remarkably rapid development of the mineral resources of

the Upper Peninsula is largely due, among other causes, to the fact that the United States Linear Surveyors were required to combine geological and topographical observations with their surveys. The use of Burt's solar compass, which permits of rapid and precise observations of local variations (so important in the economic survey of a primitive iron region), served greatly to enhance the value of the results, by making known the position of rocks containing magnetic ore.

The honesty, skill and enthusiasm with which the field-work was executed resulted in the collection of a large amount of geological data, which at the completion of the survey would have left little to be done save the final report, in which the master-mind should classify, group, and harmonize the facts, and thereby develop nature's law from the mass of material collected. Dr. Houghton's untimely death by drowning in Lake Superior, while in the midst of his labors, prevented him from performing the crowning work. Any one familiar with the geology of the Upper Peninsula, who will peruse the manuscript notes * left by Dr. Houghton, will be convinced that his views regarding the geology of the older rocks were far in advance of his time, and such only as geologists years afterward arrived at, and those which are but now, thirty years after he recorded them, universally accepted (see Appendix E, Vol. II.). A brief statement of the origin of a work from which such important results have accrued will be given. In 1843 the financial troubles of the State of Michigan arising out of the "Five Million Loan," as it was called, were of such a character as to cause the Legislature to withhold the annual appropriation for the Geological Survey, which then had been for several years in successful operation under the direction of Dr. Houghton. oughly interested in his scientific work, and believing that the best interests of the State and the cause of science demanded the continuance of the survey, Dr. Houghton asked from the General Government the aid which his own State felt unable to grant, and succeeded in obtaining, in the appropriation for the Public Surveys of

^{*} These manuscript notes are now in the University Library at Ann Arbor, having been presented to that Institution by Dr. Houghton's widow. Dr. Houghton, it will be remembered, was at the time of his death a Professor in the University of Michigan as well as State Geologist.





the Upper Peninsula of Michigan, an additional allowance per mile to cover the cost of the geological work. In order to expedite the work and insure the best scientific results from the adoption of his plan, Dr. Houghton himself took the contract from the Government for completing the surveys on the Upper Peninsula, which had been previously begun in 1840 under the direction of the Hon. William A. Burt, United States Deputy Surveyor. In the spring of 1844 Dr. Houghton commenced operations under his contract, the field-work being in charge of Mr. Burt, who received in compensation therefor the allowance granted by the Government. It is proper to add that Mr. Burt entered with deep interest into Dr. Houghton's plans and had, during his survey in the Lower Peninsula, collected for him many specimens and important geological information not required by his instructions.

In 1844 Mr. Burt, with a party consisting of William Ives, compassman, Jacob Houghton, barometer man, H. Mellen, R. S. Mellen, James King, and two Indians named John Taylor and Bonney* was engaged in establishing Township lines and making geological observations, as previously described.

On the 19th September, while running the east line of Town. 47 north, Range 27 west (the great iron Township), they observed by means of the solar compass remarkable variations in the direction of the needle, amounting to 87° from the normal. (See Appendix D, Vol. II.) Ascribing this phenomenon to iron ore, they sought for and found it in the ledges or outcrops at several points. Specimens + were collected and named by Mr. Burt and Dr. Houghton (See Appendix D, Vol. II.), and were described by them in their official returns; the fact of the great variation and large amount of ore being also especially commented upon. (See Appendix and official notes in Land Office, Washington, Lansing and Marquette.) A map made by Mr. Ives at the time, a fac-simile of which is given in Pl. I., has written along this line the words "Iron Hills." As the Tackson range is not magnetic at this locality, and does not outcrop on the line in question, it is not probable it was seen, but instead one or more of the ranges of flag or soft hematite ore further south.

^{*} Bonney's real name was Michael Doner; himself and Taylor are now dead.

[†] Mr. R. S. Mellen has still in his possession a piece of the ore found that day, which he brought away with him.

In the month of June following, Dr. Houghton and Mr. Burt, with their party, were engaged in subdividing the Township above mentioned (Town. 47 north, Range 26 west), when the former made a personal examination in reference to iron ore, especially at the corners of Sections 29, 30, 31, 32 (see Appendix D, Vol. II.), now known as the Cascade mines, and remarked to Jacob Houghton and others, who were members of the party, that it would some day be very valuable and the basis of an active industry.

It thus appears that the U. S. surveyors, in the fall of 1844, officially established the fact, that iron ore in considerable quantities existed in the Upper Peninsula of Michigan. It is also undoubtedly true, that Indians had previously observed the ore and were acquainted with locations of it, without, however, being able to identify it.

The Jackson Co.—The manner in which this, the earliest developed, and one of the most important of the iron properties on Lake Superior, was discovered (although the enterprise was not mainly undertaken with a view of finding iron), is reliably set forth in the following letter, written by P. M. Everett, now of Marquette, to Captain G. D. Johnson, now of the Lake Superior mine. The letter is dated at Jackson, Mich., Nov. 10th, 1845, and is as follows:—

"I left here on the 23d of July last and was gone until the 24th of October. I had considerable difficulty in getting any one to join me in the enterprise, but I at last succeeded in forming a company of thirteen. I was appointed treasurer and agent to explore and make locations, for which last purpose we had secured seven permits from the Secretary of War. I took four men with me from Jackson and hired a guide at the Sault, where I bought a boat and coasted up the lake to Copper Harbor, which is over 300 miles from Sault Ste. Marie. We made several locations, one of which we called iron at the time. It is a mountain of solid iron ore, 150 feet high. The ore looks as bright as a bar of iron just broken. Since coming home we have had some of it smelted, and find that it produces iron and something resembling gold-some say it is gold and copper. Our location is one mile square, and we shall send a company of men up in the Spring to begin operations; our company is called the 'Jackson Mining Co.'"

The actual discovery of the Jackson location was made by S. T.

Carr and E. S. Rockwell, members of Everett's party, who were guided to the locality by an Indian chief, named Manjekijik.*

The superstition of the savage not allowing him to approach the spot, Mr. Carr continued the search alone, resulting in the discovery of the outcrop, which he describes as indicated in Mr. Everett's letter. Previous to the discovery he was led to suppose from the Indians' description, that he would find silver, lead, copper or some other metal more precious than iron, as it was represented and found to be "bright and shiny."

July 23d, 1845, articles of association of the Jackson Iron Company were executed at Jackson, Mich., and by these articles Abram V. Berry was appointed the first *President*, Frederick W. Kirtland, *Secretary*, Philo M. Everett, *Treasurer*, and George W. Carr and William A. Ernst, *Trustees*.

Mr. Berry gives the following account of the early history of his company, in a letter dated at Jackson, Mich., Oct. 21st, 1870:—

"In the summer of 1845, an association was formed in this city, then a village, for the purpose of exploring the mineral region on the south shore of Lake Superior. The company consisted of P. M. Everett, James Ganson, S. T. Carr, G. W. Carr, F. W. Carr, E. W. Rockwell, F. W. Kirtland, W. H. Munroe, A. W. Ernst, F. Farrand, of Jackson, and S. A. Hastings, of Detroit (John Watkins, of Detroit, was interested with Hastings). Eleven individuals of the association procured permits from the War Department to locate one square mile each of mineral land on the south shore of Lake Superior. John Western, of Jackson, was then added to the

^{*} In reward for the service of the Indian on this occasion, the officers of the Jackson Company subsequently gave him a written stipulation, of which the following is a copy:—

[&]quot;RIVER DU MORT, LAKE SUPERIOR,

May 30, 1846.

This may certify that, in consideration of the services rendered by Manjekijik, a Chippeway Indian, in hunting ores of Location No. 593 of the Jackson Mining Company, that he is entitled to twelve undivided twenty one-hundredths part of the interest of said mining company in said location No. 593.

A. V. BERRY, Superintendent. F. W. KIRTLAND, Secratary."

This agreement on the part of the company was never fulfilled, and Manjekijik finally died in poverty; his relatives, now living in Marquette, are in the same miserable condition, without ever having received, as is averred by those who are cognizant of the facts, any compensation for the services mentioned.

company, making thirteen in all. In the fall of 1845 a company of explorers, consisting of S. T. Carr, P. M. Everett, W. H. Munroe, and E. S. Rockwell, visited Lake Superior, when what is now known as the Jackson location was secured by the permit granted to James Ganson, in the unsurveyed district, the section lines not having been run. The location was described by metes and bounds, commencing at a certain large pine-tree, the position of which was fixed by its course and distance from the corner of Teal lake. When the land was surveyed it was bought at \$2.50 per acre. * * *

"In the spring of 1846, another expedition was fitted out, consisting of F. W. Kirtland, E. S. Rockwell, W. H. Munroe and myself, members of the company and several other adventurers: the object being to make further examinations of the iron and to use the remaining permits, by entering other mineral land. * * * * I found our location much beyond what I had anticipated. After spending twelve days in the woods, exploring the surrounding country, including what was afterwards known as the Cleveland location and building what we called a house, we returned to the mouth of the Carp with 300 pounds of ore on our backs. We then divided; one party was left to keep possession of the location, another went farther up the Lake to use the remaining permits, while I returned to the Sault with the ore. It was my intention at this time to use another permit on the Cleveland location, but on arriving at the Sault I met Dr. Cassels, of Cleveland, agent of a Cleveland company, and having arranged with him that his company should pay a portion of the expense of keeping possession, making roads, etc., I discovered to him the whereabouts of the Cleveland location. He took my canoe, visited the location, and secured it by a permit. On arriving at Jackson we endeavored on two occasions to smelt the ore which I had brought down, in our common cupola furnaces, but failed entirely. In August of the same year, Mr. Olds, of Cucush Prairie, who owned a forge (in which he was making iron from bog ore), then undergoing repairs, succeeded in making a fine bar of iron from our ore in a blacksmith's fire, the first iron ever made from Lake Superior ore. In the winter of 1846-47 we began to get up at Jackson a bellows and other machinery for constructing a forge on the "Carp;" and in the summer of 1847 a company of men commenced building the same, and continued until March, 1848, when a freshet carried away the dam. * *

"-The association was then (1848) merged into an incorporated company, and by some means the pioneers in the enterprise are now all out."

In a book * on the mineral region of Lake Superior, with map by Jacob Houghton, Jr. and T. W. Bristol, published in 1846, only one iron company is mentioned—The Jackson. The description of the company's property is as follows:

Permit No. 593-somewhere in T. 46, N.-R. 27 or 28 W., while on Section 1 of T. 47, R. 27, Permit No. 158 is marked, which was granted to D. Hamilton, of Watervliet, New York. Section 3, same Township, embracing the New York mine, is covered by Permit No. 160, granted to T. Williams, of Newburg, N. Y. Section 10, same Township, embracing parts of the Cleveland and Lake Superior mines, was covered by Permit No. 177, granted to T. Ricket, of Copper Harbor.

In 1846 Fairchild Farrand explored the Jackson location and mined some ore. The company, under the superintendency of Wm. McNair, began, in 1847, the construction of a forge on the Carp river, three miles east of the mine, the first iron being made Feb. 10th, 1848, by forgeman, now Judge, A. N. Barney. Work was stopped in a few days by a freshet which carried away the dam. Mr. Everett came up in the summer of 1848, had the dam repaired and resumed the manufacture of blooms. The first iron made was sold to E. B. Ward, who employed it in the construction of the steamboat "Ocean." This forge was afterwards carried on under leases by B. F. Eaton, and later by the Clinton Iron Co., subsequently by Peter White and lastly by J. P. Pendill; it made but little iron and no money. The quality varied from the highest (as shown by the experiments of Major Wade, of the U. S. army) to indifferent, the trouble being a lack of uniformity in the blooms. The power was supplied by the Carp river, a dam 18 feet high having been constructed across the stream for this purpose. There were upon either side of the stone arch, and arranged opposite each other, four fires, from each of which a lump was taken every six hours, which was placed under the hammer and forged into blooms

^{*} This little volume, (afterwards revised by Mr. Houghton,) thus early issued, contains much interesting and valuable matter relating to the early discoveries and mining operations of Lake Superior, especially regarding the copper region.

four inches square and two feet in length; the daily product being about three tons, requiring two teams of six horses each to convey them over the intervening ten miles of horrible road to Marquette. These teams, when so fortunate as not to break down, on returning brought back supplies for the men and animals. The same difficulties attended the procuring of the supply of ore and charcoal. The power was also found to be insufficient, owing to a scant supply of water occurring at certain seasons of the year. These difficulties were too numerous and serious for the maintenance of the existence of the concern, and resulted in its abandonment in 1856.

On the 6th of June, 1848, a meeting was called to act on the question of the acceptance of an act of incorporation passed at the preceding session of the Legislature, and it was decided to incorporate the company under the act referred to. The organization was completed under the title of the Jackson Mining Co., of Jackson, Michigan-Fairchild Farrand, President, W. A. Ernst, Secretary, George Foot, Treasurer, F. W. Carr, F. W. Kirtland, Lewis Bascom, and John Western, Directors. The capital stock of the company, as also that of the New England Mining Co., organized at this time, was fixed at \$300,000, in shares of \$100 each; the purpose of each being the mining of copper as well as iron. April 2d, 1849, an amendment to the charter of the Jackson Mining Co., of Jackson, was obtained, when the title was changed to its present form—Jackson Iron Co. The first officers under this organization were Ezra Jones, President, Wm. A. Ernst, Secretary, John Watson, Treasurer, S. H. Kimball, James A. Dyer, and James Day, Directors.

In 1850, Mr. A. L. Crawford, proprietor of the iron works at Newcastle, Pa., took with him from Lake Superior about five tons of the Jackson ore, and there worked it up. Part of the ore having been made into blooms and rolled into bar-iron, was used for special purposes, and part used for lining in the puddling furnaces. The iron was found to be excellent. About the same time, General Curtis, of Sharon, Pa., proprietor of extensive iron-works at that place, came to Lake Superior to inspect the Jackson and Cleveland locations; his object being to secure an interest, with a view to a future supply of ore for his works, of a better quality than he then possessed. Failing to make an arrangement for the Cleveland, he bought up sufficient stock in the Jackson Co. to give

him a controlling interest in the management of its affairs; so that for some years the location was known as "Sharon."

It is proper to remark that General Curtis believed, as did also John Western before him, that, as soon as practicable, the best policy for Lake Superior iron mines to follow would be to sell their ore to the furnaces of Ohio, Pennsylvania, and elsewhere; and in 1852 about 70 tons of the company's ore were taken to Sharon, Pa., and there made into pig-iron in the "old Clay Furnace." There were frequent changes of officers and directors in the Jackson Co. up to 1860, and the history of the company was one of disappointment and financial embarrassment. 1860 and 1862 the gentlemen who now compose the Board of Directors came into office, and in 1862 the first dividend was made. The great demand for iron occasioned by the war caused the iron interests of Lake Superior, for the first time, to assume a very successful aspect. The first regular shipments of ore from the Jackson mine were made in 1856, which amounted to about 5,000 tons. Up to this time the different forges in the district had consumed about 25,000 tons of ore. (See Table, Pl. XII. of Atlas.) The Jackson mine, earliest discovered, and first opened and tested, became widely known from the outset, and has ever continued to remain the leading mine in the district. The important village of Negaunee, within whose corporate limits the Jackson mine is situated, dates its origin with the commencement of the company's o erations. As the Chicago and Northwestern and the Marquette, Houghton and Ontonagon railroads form a junction in Negaunee, facilities are thus afforded for shipments over either road—that is, by the way of Escanaba or Marquette. The "openings," or pits, are irregular and numerous, and extend from the west edge of the village of Negaunee west for three-quarters of a mile. The greater portion of the product finds its outlet through a tunnel, which enters the mines from the north side of the hill and is of sufficient size to admit railroad cars and small locomotive engines. From the main tunnel radiate several branches, which extend to, or are being extended to, the different stopes and shafts. The main shafts are supplied with ample steam-power for pumping and hoisting purposes. For details of workings, geological structure, etc., see accompanying maps, tables, and text.

The New England Mining Co. was, like the Jackson, incorporated by a special act of the Michigan Legislature passed in 1848. The purpose for which the organization was effected is stated as being the mining and smelting and manufacturing of ores and minerals in the State of Michigan, the language stating the company's objects being identical with that of the Jackson Company; the capital stock was placed at 300,000. It does not appear that anything noticeable was accomplished by this company, thus early organized. The charter came in 1855 into the possession of Capt. E. B. Ward, by whom it is now held.

The Marquette Iron Co.—In the summer of 1848, Mr. Edward Clark, of Worcester, Mass., was sent to Lake Superior by Boston parties, to look for copper, but at the Sault he fell in with Robert J. Graveraet, who induced him to stop at the Carp river and see the iron mines. The Jackson Company's forge was at work and had made a little iron. Clark, on his return to Worcester, carried with him a bloom and some ore from the Jackson Iron mountain, which, on being drawn into wire at a factory, proved excellent. Clark at once proceeded to form an association for the purpose of building a forge on the far-off shore of Lake Superior, assisted by Graveraet, who also appeared in Worcester at this time (having travelled from Marquette to Saginaw on snow-shoes); he succeeded in organizing a company, March 4th, 1849, consisting of E. B. Clark, W. A. Fisher, A. R. Harlow, of Worcester, Mass., and R. J. Graveraet, of Mackinaw; Clark and Graveraet putting in against the capital of the others leases of iron lands of which they claimed to have possession. These iron lands constitute what subsequently became known as the Lake Superior and Cleveland mines, and over which a long controversy arose as to which party should possess the land, and which was finally decided by the Interior Department at Washington in favor of what was known as the Cleveland Company. Mr. Harlow constructed and purchased the necessary machinery to the value of \$8,000, and in the spring of 1849 shipped it to Marquette, starting himself with his family on the 11th of June, and arriving in Marquette on the 6th of July thereafter. Graveraet had reached there on the 17th of May previous, taking with him a small party of men, among whom was Peter White, then a lad, but subsequently largely identified with numerous interests in the Iron

Region, and now President of the First National Bank of Marquette. The forge was completed, making the first bloom in just one year from the date of Mr. Harlow's arrival.

The Marquette Iron Co.'s works started with 10 fires, and used Cleveland and Lake Superior ores, mostly the former, making blooms exclusively, which were sold in Pittsburg at prices ranging from \$35 to \$50. The works were in operation somewhat irregularly until 1853, when the Marquette Company was merged into the Cleveland Company, under the auspices of which the forge continued in operation for a few months longer, and was finally destroyed by fire in 1854. Like all bloomeries started in Marquette County, it was from the first, financially, a failure. The cost of the plant was great, transportation difficult and expensive, and the price of iron during the entire period disproportionately low. There was no dock at Marquette, no canal at the Sault, scarcely a road in the country, no shop for repairs, no skilled labor but what was, together with all supplies, imported "from below," and no regular communication. During the summer of 1849 only three sailing vessels and five propellers arrived at Marquette. The stock of the Marquette Company was bought up by the Cleveland Company, and its property passed to the ownership of the latter.

In 1852 John Downey, Samuel Barney and others began the construction of a forge on the "Little Carp," but after having built some houses, constructed a wheel, etc., permanently abandoned the enterprise.

In 1849 and 1850 a whetstone quarry was opened in a bed of novaculite, near the outlet of Teal lake, and Messrs. Smith and Pratt established a factory, for the purpose of sawing these blocks, at the mouth of a small stream near the Marquette landing, and carried on a "thrifty business."

The Iron Mountain Railroad.—The question of transporting the rich ores of Marquette county to the coal of Ohio and Pennsylvania, being one that came to be seriously considered, it naturally suggested the necessity of a railroad from the mines (those near the present villages of Negaunee and Ishpeming) to Marquette bay. 1851 Messrs. Heman B. Ely and John Burt strongly advocated the enterprise, and in the following year Mr. Ely caused a survey to be made; at that period the entire population of Marquette county was

less than 150 persons. There being no general railroad law in the State at that time, the construction of the railroad was undertaken by Mr. Ely, assisted by his brothers George H. and Samuel P. Ely, of Rochester, New York, as an individual enterprise, he having previously made a contract with the Jackson and Cleveland Iron Mining Companies and Mr. John Burt, as the representative of other companies, for the transportation of their ores. This contract the two first-named iron companies subsequently attempted to break, and sought to defeat the railroad by constructing a plank-road in opposition to it, thus instituting a serious and embarrassing controversy, which continued until 1855, when all matter of dispute then pending between the Railroad Company, under charge of Mr. Ely, and the Plank-road Company, under charge of Mr. S. H. Kimball, were submitted to arbitration and settled to the satisfaction of both parties— Messrs. C. T. Harvey and Austin Burt being arbitrators. Immediately after the passage of the General Railroad Law of this State in 1855, the Messrs. Ely incorporated the railroad under the title of the Iron Mountain Railroad, and John Burt was first President. A year later the company was strengthened by the addition of Jos. S. Fay, Edwin Parsons, Lewis H. Morgan, and other capitalists; and in 1857 the road was completed and put in operation. Mr. H. B. Ely, to whose foresight and energy the origin and success of the enterprise was largely due, and to whom the interests of Lake Superior became otherwise greatly indebted, died in Marquette, in 1856, before the work upon which he had labored so intently was completed.

The death of his brother, and his own connection with the road, was the occasion of bringing to Marquette Mr. S. P. Ely, who is now more largely identified with the business management of many of the leading enterprises in the Iron Region than any person resident on "Lake Superior." The Iron Mountain Railroad became subsequently a part of the Bay de Noquette and Marquette Railroad, this becoming afterwards, by consolidation, the Marquette and Ontonagon Road, and still later, by further consolidation, a part of the through line of the Marquette, Houghton, and Ontonagon Railroad. The plank-road to which reference is here made was built by the Jackson and Cleveland Companies jointly, but was never used as a plank-road; longitudinal sleepers were laid down and covered with strap-rail, on which horse cars were run. The road was used for two seasons, and cost \$120,000, which

amount was practically sunk. The cost of transportation was nominally one dollar per ton; each team would make the round trip in a day, bringing four tons of ore. It is proper to add that the rates of transportation fixed by these H. B. Ely contracts, although afterward deemed by the iron companies much too liberal, were lower than any at which ore has ever been carried over the road; the present rates being more than double those agreed upon with Mr. Ely.

Among the most important enterprises early connected with the development of the Lake Superior iron interests was the construction of the Sault Ste. Marie Ship Canal. In the St. Mary's river or strait, connecting the waters of Lakes Superior and Huron, occurs, nearly opposite the village of Sault Ste. Marie, a rapid of about one mile in length, and about seventeen feet fall, forming a complete barrier to the communication between the lakes. Some years previous to the construction of the canal this barrier had been overcome partially, by the construction and use of a portage flatbar railroad, over which all articles of commerce between the lower lakes and Lake Superior were transported and reshipped in both directions. The important and growing interests of Lake Superior demanded more easy and effective means of commercial communication with the lower lakes. The matter being brought before the National Legislature, Congress granted to the State of Michigan, by Act approved Aug. 26th, 1852, 750,000 acres of land for the purpose of aiding in the construction and completion of a ship canal around the falls of Ste. Marie. On the 5th of February following, the State of Michigan, by an Act of its Legislature, accepted the grant of land above mentioned; and to further the objects thereof, authorized the Governor of the State to appoint Commissioners to let the contract for the construction of the canal, and to enter the lands authorized under the grant.

The Commissioners appointed under this legislative act entered into contract with Joseph P. Fairbanks, Erastus Corning and others for building the canal within two years from date thereof; the consideration being the U. S. Government grant of lands. This contract was soon after duly assigned to the Ste. Marie's Falls Ship Canal Co., which company had been organized in the city of New York on the 14th of May, 1853, under an Act of the Legisla-

ture of the State of New York, passed April 12th, immediately preceding. At the organization of the company, the following persons were chosen officers and directors of the company: Erastus Corning, President, J. W. Brooks, Vice-President, J. V. L. Pruyn, Treasurer and Secretary. Directors: Erastus Corning, J. W. Brooks, J. V. L. Pruyn, Jos. P. Fairbanks, John M. Forbes, John F. Seymour, and James F. Joy.

Subsequent to the passage of the grant by Congress, but previous to the acceptance thereof by the State of Michigan, Mr. Charles T. Harvey was authorized by Messrs. Fairbanks and Corning to cause a survey to be made, which he proceeded to do during the month of November, 1852, having secured the services of an experienced engineer from the Erie Canal, Mr. L. L. N. Davis. After the organization of the company, Mr. Harvey was appointed its general agent, and the supervision of the construction placed under his control.

Early in the season of 1853 Mr. Harvey, with 400 men, proceeded to the Sault, and on the 4th of June broke ground for the canal. The remoteness of the locality, and many other unfavorable circumstances, rendered the construction of a work of such magnitude exceedingly difficult, and necessitated at every step of the operations unusual care and energy in the management as well as heavy pecuniary expenditures. Mr. Harvey remained in control of the construction for one year, when he was relieved and placed in charge of the finance, and also appointed agent for the State to select lands under the grant in the Upper Peninsula. Mr. Harvey selected about 200,000 acres of land, 39,000 of which were taken in Marquette county, and were subsequently sold for \$500,000 cash, to the Iron Cliff Co. Among the copper land selected was the quarter section on which the Calumet and Hecla Company's mine is situated, and which was sold by the canal company for \$60,000, now worth, on the basis of late sales of stock, \$13,000,000. The 750,000 acres granted by the General Government were entered by the company as follows: on the Upper Peninsula, 262,283 acres of iron, copper, and timber land, and 487,717 acres of pine land in the Lower Peninsula. A land agency was established at Detroit for the purpose of locating the lands obtained through the grant.

During the summer of 1854 the difficulties necessarily attendant upon building the canal were very much enhanced by disease among the workmen; some 200 of whom died of the cholera, and among them was Mr. Ward, who had charge of the construction. Harvey was again placed in charge of the work, which, owing to the panic among the workmen, had become nearly suspended; but by the exercise of much skill and energy he succeeded in reorganizing the force, and pushing the work vigorously forward to final completion. On the 19th of April, 1855, the water was let into the canal, and in the following June the work was opened for public use, under the superintendency of Mr. John Burt.

The total cost of the construction of the canal, which includes also the expense attendant upon the selection of lands, as contained in the report of the company under date of January 1st, 1858, was \$999,802.46.

The State of New York, by act passed April 15th, 1858, granted a charter incorporating the "St. Mary's Canal Mineral Land Co." Under this act of incorporation, a company was duly organized, and to it was transferred the canal company's lands of the Upper Peninsula. It was soon found that the canal failed to meet the growing wants of the commerce of Lake Superior, owing to the variation in the general level of the Lake Superior becoming somewhat lower than when the canal was completed, thus making a variable difference in the depth of the canal of from one to one and one-half feet; and also that the General Government, by successive appropriations, has caused the channels through Lake George and the St. Clair Flats to be so widened and deepened, that vessels of far heavier tonnage than was originally anticipated could be employed. The Michigan State Legislature adopted a resolution in the session of 1869, offering to cede the canal to the U.S. Government; although Congress has not as yet formally accepted the offer made by the State, nevertheless, under its system of internal improvement, the General Government is now engaged in the enlargement of the canal. The width of the canal is to be increased to 300 feet, and its depth to 16 feet, the locks are to be double, 80 feet in width and 450 feet long. The amount of the government appropriations under which this improvement is being effected is in the aggregate \$800,000; and the work, when completed, will be fully adequate to the wants of commerce.

The report of superintendent Guy H. Carleton shows the following to be some of the principal exports and imports through the canal during 1871 and 1872:

	1871.	1872.
Flour, bbls	25,146	42,141
Pork, bbls	8,887	10,306
Beef, bbls	3,054	4,161
Bacon, lbs	163,763	242,475
Lards, lbs	283,141	213,394
Butter, lbs	519,545	559,137
Cheese, lbs	187,340	200,994
Tallow, lbs	104,354	106,170
Soap, boxes	21,799	18,205
Apples, bbls	18,359	20,025
Sugar, lbs	4,062,087	5,454,559
Tea, chests	3,864	7,980
Coffee, bags	5,228	7,815
Salt, bbls	36,199	42,690
Tobacco, lbs	258,179	321,836
Nails, kegs	29,843	34,984
Dried Fruit, lbs	115,366	73,230
Vegetables, bush	27,619	35,263
Lime, bbls	2,338	6,067
Window Glass, boxes	25,226	7,492
Cattle, head.	2,639	3,608
Horses and Mules	435	528
Hogs, head	1,625	1,567
Brick, M	1,225	9,067
Furniture, pieces	13,616	44,768
Machinery, tons	1,595	10,593
Engines	18	28
Boilers	17	34
Liquor, bbls	4,366	7,082
Malt, lbs	653,140	1,545,875
Coarse Grain, bush	283,503	444,875
Mdse., tons	23,245	38,215

The following are some of the principal exports from Lake Superior for 1871-72:—

	1871.	1872.
Mass Copper, tons	1,091	1,709
Ingot Copper, tons	7,666	8,547
Stamped Work Copper, tons	5,705	4,365
Iron Ore, tons	327,461	383,105
Pig Iron, tons	23,304	29,341
Fish, half bbls	26,041	14,529
Wheat, bush	1,376,705	567,134
Tallow, lbs	59,225	64,567
Flour, bbls	179,093	94,270
Barley, bush	25,320	898

	1871.	1872.
Silver Ore	464	306
Stone, building, tons	5,528	5,213
Potatoes, bush		636
Copper, manufactured, tons		395
Quartz, tons	****	591
Wool, tons	****	30

In 1853 the Lake Superior Iron Company, one of the three oldest companies in the district, was formed; articles of association were filed March 13th, capital stock \$300,000, in 12,000 shares of \$25 each. The capital stock was subsequently increased to \$500,000, which has all been returned to the stockholders in The incorporators were Heman B. Ely and Anson Gorton, of Marquette, 'Mich.; Samuel P. Ely, George H. Ely, and Alvah Strong, of Rochester, New York. The company commenced operations in 1857 on 120 acres of land in Sections 9 and 10, T. 47, R. 27, which was purchased of John Burt, being a part of the Briggs and Graveraet claim spoken of above under the Cleveland Company. Subsequent purchases enlarged the company's estate to 2,000 acres, its present dimensions. The company's principal openings are upon the land originally purchased. The first shipment of ore (4,658 tons) was made in 1858; since which the increase has been so great that its shipments now exceed those of any mine in the district, as will be seen by reference to the tables. This company have recently constructed, in Marquette, the Grace Furnace, which went into blast in December, 1872, using anthracite coal in the manufacture of pig-iron. The furnace is located on the shore of the bay, within the limits of the city, and is the first anthracite furnace built on Lake Superior. A map of the Lake Superior and Barnum mines accompanies this report.

The Eureka Iron Company was organized October 29th, 1853, with a capital stock of \$500,000 in 20,000 shares. The corporators were Eber B. Ward, Harmon De Graffe, Silas M. Kendrick, M. Tracy Howe, P. Thurber, Elijah Wilson, Thomas W. Lockwood, and Francis Choate, with office in Detroit. The organization was effected with a view of mining ore and of manufacturing charcoal pig-iron from Lake Superior ores; preparations were made to build a furnace in Marquette county, but the location was finally

changed and the furnace erected where now stands the flourishing city of Wyandotte, becoming the nucleus of the extensive iron works which have since grown up in that locality. The Eureka Company was also the first iron enterprise in which Captain E. B. Ward, subsequently so widely known as a successful iron master, became engaged. The company was formed by Philip Thurber, and a quarter section of land purchased near Marquette of Mr. A. R. Harlow, on which a few hundred tons of ore were mined; but it becoming evident that the ore did not exist in quantity, the work was abandoned. This land was subsequently sold back to Mr. Harlow for his shares of the company's stock, and is now known as Harlow's Mill.

The Cleveland Iron Mining Company filed articles of association March 29th, 1853; capital stock, \$500,000, in 20,000 shares. The corporators were John Outhwaite, Morgan L. Hewitt, S. Chamberlain, Samuel L. Mather, Isaac L. Hewitt, Henry F. Brayton, and E. M. Clark, with office in Cleveland, Ohio. The early history of this celebrated mine, one of the oldest and most important in the district, is referred to in connection with that of the Jackson Co.

Dr. J. Lang Cassels, of Cleveland, to whom reference is made in Mr. Berry's letter, visited Lake Superior in 1846, and took, as he expresses it, "squatter's possession" of a square mile for the Dead River Silver and Copper Mining Co. of Cleveland; the property here spoken of includes the mines of the present Cleveland Co. The Jackson Co. had previously taken possession of their lands, and Dr. Cassels obtained guidance thereto from an Indian, there being no white men in the region; the doctor went up from and returned to the Sault in a bark canoe. During the succeeding year, Cassels having left the country, the location was taken possession of by Messrs. Samuel Moody, John Mann and Dr. Edward Rogers, The two former claiming what became the Cleveland mine, and the latter what is known as the Lake Superior. When the Marquette Forge Co. was organized in Worcester, as previously described. Clark had authority from Mann and Moody to lease their location, and Graveraet had similar power from Rogers.

In this manner leases of these lands were put into the organization against \$20,000 cash capital, to be paid by Messrs. Harlow and

Fisher. Both the Cleveland Co. and Graveraet, representing Messrs. Moody and Mann, claimed priority of right to the land under a "pre-emptor's mining act." These conflicting claims went before the Department at Washington, where a decision was rendered, which gave the right of purchase to the Cleveland Co. The entries which the Cleveland Co. made did not cover the Lake Superior location, Graveraet still claiming it, in behalf of the Marquette Co., on the ground of the Rogers pre-emption. Previously Isaiah Briggs had been on the land, but, leaving it, Rogers had taken possession. Rogers lost his interest, however, by not being present at the Government sale of lands in November, 1850, and establishing his claim, having been detained by a storm on the lake while endeavoring to proceed to the Sault (where the land office was located) for that purpose. The location was purchased by John Burt, on the basis of the Briggs claim, he having agreed to lease an undivided onehalf interest to Graveraet, who was also present in behalf of the Rogers claim. This lease to Graveraet was assigned by him to the Marquette Co., passed with the company's other assets into the possession of the Cleveland Co., and was finally sold for \$30,000 to the Lake Superior Iron Co., that company having previously purchased the Briggs title.

The Cleveland association, although formed in 1849, did not do any business in Lake Superior until 1853; at that date the Cleveland and Marquette companies became finally merged by the former company purchasing (including 64 acres of land on which the forge was located) the assets of the latter, and the present Cleveland Iron Co. was formed. The Cleveland Co. continued to run the forge for about two years, until it was burned down. The company mined in 1854, 4,000 tons of ore, which was made into blooms at the different forges in the vicinity. In 1855 they shipped 1,449 tons of ore to the furnaces "below," thus preceding the Jackson Co. one year, and becoming the first to send out of the region any considerable amount of ore. The Jackson Co. had sent a few tons to the World's Fair in New York in 1853, and in 1852 some had been sent to Sharon, as before mentioned. The Cleveland Co. has also an ore dock at Marquette, entirely similar to the docks of the M. H. & O. R. R. Co., of which full descriptions and illustration are given.

On Nov. 8th, 1853, the Collins Iron Co. filed articles of associa-

tion, with a capital stock of \$500,000 in 20,000 shares. The corporators were Edward K. Collins, of New York, Solon Farnsworth, Edwin H. Thomson, Robert J. Graveraet, and Charles A. Trowbridge, with office in Detroit.

The company built a forge in 1854, and began to make blooms late in the fall of 1855; Robert J. Graveraet, Supt., and C. A. Trowbridge, Managing Director. E. K. Collins largely interested himself with a view of obtaining a superior quality of iron for the shafts of his ocean steamers. In 1858, about the time the Pioneer Furnace was completed, Mr. S. R. Gay, who had been engaged on that work, leased the Collins Forge and put up a cupola there in which he made some pig-iron. The company immediately thereafter constructed a blast-furnace under the direction of Mr. Gay. This furnace was completed and put in operation December 13th, 1858, with a single stack; all the necessary power being afforded by the Dead river, upon which the furnace is located.

On August 28th, 1854, the Peninsula Iron Co. filed articles of association, with a capital stock of \$500,000 in 20,000 shares. The corporators were Wm. A. Burt, Austin Burt, Wells Burt, John Burt, Heman B. Ely, Samuel P. Ely, and Geo. H. Ely; the two latter of Rochester, N. Y., the others of Michigan. Office of the company, Marquette, Mich. The company originally owned 800 acres of iron lands, which it sold in 1862 to the Lake Superior Iron Co., and determined on building a blast-furnace at Hamtramck, Detroit, Mich., which furnace was completed in February, 1863, and is still in successful operation. The company also operated a sawmill for a few years, which they built on the Carp river, a short distance from Marquette.

Oct. 11th, 1854, the articles of association of the Chicago and Lake Superior Iron Mining and Manufacturing Co. were filed. Capital stock, \$500,000, in 20,000 shares. The corporators were B. S. Morris, Isaac Shelby, Jr., Geo. Staley, Henry Frink, and Samuel S. Baker, all of Chicago, Ill.; and Solomon T. Carr and Fairchild Farrand, of Jackson, Mich. No permanent mining work was ever done by this company.

The Clinton Iron Co. Organized by forgemen from Clinton Co.,

New York, Jan. 20th, 1855. Capital stock; \$25,000. Corporators, Azel Lathrop, Jr., H. Butler, Chas. Parish, and Daniel Brittol.

The object for which the organization was effected was to lease and operate the Jackson Forge. The company being composed of workmen, who at the time were employed in that concern and were locally styled the "Mudchunk." The market price of blooms being much below the cost of their manufacture, they were enabled to operate the forge but a brief period, and having become hopelessly involved in indebtedness, the company permanently suspended.

The Forest Iron Co. filed articles of association, September 22d, 1855, with a capital stock of \$25,000 in 1,000 shares. The corporators were Matthew McConnell, Wm. G. Butler, Wm. G. Mc-Comber, M. L. Hewitt, and J. G. Butler. This company was organized for the purpose of putting up a bloom forge on Dead river, and the location became known as Forestville. McConnell, Butler and McComber commenced operations at this point as early as 1852 on their own private account, but becoming financially embarrassed, they sought relief by organizing a company as above indicated, who continued the manufacture of what was called half blooms, the production of which cost-them from \$180 to \$200 a ton. These selling in Pittsburg for \$35 to \$40, on six months' time, it naturally resulted in the ruin of the company.

To the original projectors of the Pioneer Iron Co. belongs the credit of having established the first blast-furnace on Lake Superior; previous to that all the iron manufactured had been made in bloomeries. Mr. C. T. Harvey was the mover of the scheme, and the originator and manager of the company. He induced capitalists (chiefly in New York) to embark in the enterprise, Mr. E. C. Hungerford of Chester, Conn., being chosen Secretary and Resident Treasurer. Although the business was unknown to a single man on Lake Superior, the most sanguine views prevailed from the outset, and a two-stack furnace was constructed near the Jackson mine.

The late S. R. Gay and L. D. Harvey, now Superintendent of the Northern Furnace, were the builders; the work being commenced June, 1857, and completed so as to make the first iron in February of the next year.

Much of the material, including two millions of brick, was brought from Detroit and had to be hauled 13 miles from Marquette by teams; the engines were made at the West Point Foundry. The original stock was \$125,000, in 5,000 shares; the articles were filed July 20th, 1857, the corporators being Moses A. Hoppock, Wm. Pearsall and Chas. T. Harvey. Most of the parties interested in the concern were totally ignorant of iron-making and as an instance illustrating the fact, it is related that one of the directors, during the period of construction, inquired when the furnace would be completed so that it might be sent up to Lake Superior; he supposing it was being made in Detroit. These unfavorable circumstances, combined with the financial depression of 1857, at which time the company were obliged to sell their iron for \$22, while the cost of its production was \$24 per ton, gave no return save anxiety and disappointment.

In the spring of 1860, the furnace was leased for four years to Mr. I. B. B. Case, he agreeing to deliver the pig-iron on board the vessels at Marquette for \$17.50 per ton, and paying all the expenses of its manufacture; the company furnishing the timber, standing, for the charcoal, and giving him the advantage of a contract with the Jackson Company for the ore, the royalty for which (\$1.00 per ton of iron) he paid. This price proved to be less than the iron could be made for. The furnace was burnt down August 9th, 1864; number two stack was at once rebuilt and put in operation in January following, by Mr. Case.

In 1865, Dr. J. C. McKenzie, then President of the Pioneer Iron Company, entered into negotiations with the Iron Cliff Company, which subsequently resulted, largely through the instrumentality of Major T. B. Brooks, Vice-President of the latter company, in an arrangement (ratified by the stockholders of both companies, March 10th, 1866) by which the Iron Cliff Company came into possession of the furnace, on consideration that it pay to its former proprietors one-third of the profits of the business. Soon after the two companies became practically one, through the purchase of the stock of the Pioneer by the Iron Cliff Company.

The Detroit Iron Mining Company filed articles 15th August, 1857. Capital, \$500,000, in 20,000 shares at \$25 each, with office in Detroit. Corporators were Patrick Tregent, Guy Foot, Joseph

P. Whittemore, John H. Harmon, John W. Strong, Oville B. Dibble, Nelson P. Stewart, Andrew T. McReynolds, Thornton T. Brodhead, Henry T. Stringham, Henry J. Buckley, Joseph L. Langley, of Detroit, and Edwin H. Thomson, of Flint. The company having ascertained, as they believed, that their lands did not contain sufficient ore for mining purposes, sold them to Mr. J. P. Pendill, and upon them is now built a portion of the village of Negaunee. The McComber mine, which lies at a short distance south of that village, is on this land.

The Excelsior Iron Company filed articles October 6th, 1857. Capital stock, \$100,000; 4,000 shares, at \$25 each. Corporators were: C. T. Harvey, Sarah V. E. Harvey, E. C. Hungerford, George P. Cummings, and Joseph Harvey, all of Marquette. This company did little but organize. It originated with Mr. C. T. Harvey, and some of the land which it owned has since proved to be valuable mining property, as it embraces the Barnum mine, now owned by the Iron Cliff Company; upon it is also situated a portion of the village of Ishpeming.

The Lake Superior Foundry Company filed articles of association July 14th, 1858. Capital stock (paid in), \$10,000; 400 shares, at \$25 each. Corporators: John Thorn, Isaac Maynard, Thomas Maynard, Nathan E. Platt, of Utica, N. Y., and Charles T. Harvey, of Marquette, Mich. This establishment, which was started in 1858, is now running on a much enlarged scale, under the name of the Iron Bay Foundry, D. H. Merritt, proprietor. The location is near the bay, within the city of Marquette.

The Grand Island Iron Company filed articles May 3d, 1859. Capital, \$400,000; 16,000 shares, at \$25 each; paid in, \$110,000. Corporators: Thomas Sparks, Henry W. Andrews, William Lippincott, John L. Newbold, John D. Taylor, John R. Wilmer, Samuel Pleasants, William M. Baird, Samuel J. Christian, L. de la Cuesta, William A. Rhodes, Charles Lennig, James C. Fisher, Samuel T. Fisher, Lewis Seal, Coleman Fisher, Henry Maule, William Gaul, J. T. Linnard, Howard Spencer, Caleb Jones, Charles W. Carrigan, of Philadelphia, and Devere Burr, of Washington, D. C., with office in Philadelphia. The property belonging to

this company, consisting of 3,000 acres of land, situated on Grand Island harbor, in Munising Township, was sold in 1867 to the Schoolcraft Iron Company, and their operations were confined to some minor improvements in the way of wharves, etc.

The Northern Iron Company filed articles May 16th, 1859. Capital stock, \$125,000, in 5,000 shares of \$25 each. Corporators: John C. Tucker, Moses A. Hoppock, of N. Y., and Charles T. Harvey, of Marquette, with office in Marquette. This company was formed through the efforts of C. T. Harvey, and constructed a blast-furnace at the mouth of the Chocolate river, 5 miles south of Marquette, with a view of making pig-iron with bituminous coal, being the first enterprise of this kind inaugurated in this region. After making about 1,000 tons of iron, the furnace was changed into and run as a charcoal furnace up to June, 1867; since which time it has not been working, and it is now being changed back into a bituminous coal furnace. This is the first charcoal furnace on the Upper Peninsula that has been permanently blown out.

1863.—The great financial prostration of 1857, combined with numerous causes which readily suggest themselves, naturally embarrassed and, in instances, extinguished the new and struggling enterprises of Lake Superior to the extent, that comparatively little was done in the manufacture of iron or the mining of ore up to the opening of 1863. During this interval of time no companies of importance filed articles of association in this region. Very early in the war, however, the greatly increased demand for iron which it occasioned, began to be felt over the country and finally extended its influence to Lake Superior, causing the revival of the languishing enterprises already started and the organization of many new ones. The abundance of ore, together with its surpassing richness in iron and freedom from deleterious substances, the facility with which it could be mined and the greatly improved means of transportation, were becoming generally known, and the strength and exceeding tenacity of the iron manufactured therefrom universally acknowl-Thus altogether there was opened to the Marquette region an outlook of prosperity, which it had not heretofore experienced, enabling its mining and iron manufacturing companies to assume a basis of more successful operation, and confidently to push forward their improvements.

The articles of association of the Teal Lake Co. were filed on the 7th of June, 1863, with a capital stock of \$500,000, in 20,000 shares, and an amount paid in of \$100,000. The corporators were George A. Fellows, John W. Wheelwright and Charles L. Wright, of New York, with office in New York. Beyond some explorations this company never did any work on Lake Superior, confining its operations chiefly to stock speculations, it being the only iron mining company organized in this region, whose stock was sold at the Brokers' Board in New York.

The articles of association of the Morgan Iron Co. were filed on the 1st of July, 1863, with a capital stock of \$50,000, in 2,000 shares, and \$26,000 paid in. Corporators were Joseph S. Fay, of Boston, Lewis H. Morgan, of New York, Harriet H. Ely, Samuel P. Ely, Ellen S. White and Cornelius Donkersley, of Marquette, with office in Marquette. The capital stock was subsequently increased to \$250,000, in 10,000 shares fully paid. The company own 20,000 acres of timber land. In 1863 they constructed the Morgan Furnace, eight miles west of Marquette on the M. H. and O. R. R., and the location has since become known as "Morgan." The furnace was put up under the supervision of Mr. C. Donkersley and has been successful. It went into blast Nov. 27th, 1863, making that year 337 tons of iron, and was the first furnace company in the region to pay a dividend to its stockholders. The extreme high price of iron, created by the war, enabled the company to realize, during the first ten months of the operation of the furnace, a dividend of 100 per cent. over and above the total outlay in its Having exhausted the fuel in the vicinity, the company constructed charcoal kilns upon their lands at a distance of nine miles north from the furnace, and provided for the transportation of the coal by building a wooden railway thereto. The kilns and railway were made in 1869, and most of the coal now used is prepared at these kilns.

In 1867 the Morgan Company built the Champion Furnace, which went into blast Dec. 4th of that year. This furnace is located at what is now Champion village, on the line of the M. H. and O. R. R., 31 miles west from Marquette. The ore used is mainly magnetic from the Champion mine, and the record of the furnace is one of gratifying success.

The articles of association of the Marquette Iron Co. were filed April 9th, 1864, with a capital of \$500,000, in 20,000 shares of \$25 each. Corporators: George Worthington, Truman P. Handy, Samuel L. Mather, N. B. Hurlbut, Richard C. Parsons, G. D. McMillen, John Outhwaite, of Cleveland, Ohio, and Charles I. Walker, of Detroit, Mich. This company was organized for the purpose of mining iron ore and owns 400 acres of land, lying contiguous to, and south of, the Cleveland mines, 240 acres of which was originally held by the latter company. Its stock is held by stockholders of the Cleveland Company. The year of its organization it shipped 3,922 tons of ore, and has been somewhat regularly in operation since that period.

The Magnetic Iron Co. was organized in 1864; the articles of the company were filed May the 6th of that year, with a capital stock of \$500,000 in 20,000 shares. Corporators: John C. McKenzie, Alex. Campbell, of Marquette, and Edwin Parsons, of New York. Office in Marquette, but now in Philadelphia, Pa. The property owned by this company consists of 520 acres of land on Section 20, T. 47, R. 30. A shaft 60 feet in depth has been sunk, and other explorations made to test the ore-deposit and the company expect to take out ore, as soon as a branch road is built to the mine.

The Chippewa mining property comprises Section 22, T. 47, R. 30, W., owned by J. S. Waterman, of Philadelphia, and S. S. Burt, of Marquette; considerable exploring has been done on the property and some fair ore found, but no mining done. This property lies on the east side of Michigamme river and opposite the Magnetic and Cannon properties.

The Phœnix Iron Co. filed its articles of association June 7th, 1864. Capital, \$500,000, in 20,000 shares, of which \$20,000 was paid in. The Corporators were Wm. C. Duncan, Henry J. Buckley and Simon Mandlebaum, of Detroit, with office in Detroit. No mining or manufacturing was ever done in the Marquette Region by this company.

Washington Iron Company filed its articles of association July 30th, 1864. Capital stock, \$500,000, in 20,000 shares, at \$25 per

share; amount paid in, \$100,000. The Corporators were Edward Breitung, I. B. B. Case and Samuel P. Ely, of Marquette, Joseph S. Fay, of Boston, and Edwin Parsons, of New York.

This company made its first shipments of ore (4,782 tons) in 1865, and has since been in active operation. The land owned by the company comprises 1,000 acres in the northeast part of T. 47, R. 29, which was purchased of Silas C. Smith, J. J. St. Clair, J. C. McKenzie, and Alexander Campbell, who derived their title from the United States Government. The mine is on the M. H. and O. railroad, at a distance by rail from Marquette of 27 miles. All the company's surplus earnings have been expended in making extensive improvements, of which an adit or tunnel, now over 1,100 feet long, constitutes the chief. Their plans and expenditures have been on an extensive scale, and contemplate operations for a long period to come. The details of the mine, shafts, adit and underground workings, together with the geological structure, are fully shown by the map of the Washington mine, accompanying this report.

The Bancroft Iron Co. filed its articles of association September 12th, 1864; capital stock being \$250,000 in 10,000 shares, of which \$100,000 was paid in. The Corporators were Wm. E. Dodge, of New York, Samuel L. Mather, John Outhwaite and Wm. L. Cutter, of Cleveland, Peter White and Samuel P. Ely, of Marquette, and Henry L. Fisher and L. S. McKnight, of Detroit, with office in Marquette.

The location of this company is the same as that of the Forest Iron Co., heretofore described; the property of the latter having been purchased by Mr. S. R. Gay, in 1860, he erected on the water-power employed by the old forge a blast-furnace, this being the second furnace he had built on Dead river, the one at Collinsville having been constructed by him the winter before.

Mr. Gay * having died in 1863, his furnace at Forestville passed to the ownership of the Bancroft Iron Co., who have since continued

^{*} It is a fact worthy of note, in connection with the services rendered by Mr. Gay, that he was the first among the iron men who visited Lake Superior to recognize the value of the hematite ores; while engaged in the construction of the Pioneer Furnace, he observed that the Jackson Co. were wasting their soft hematite in large quantities, they supposing it to be worthless. He at once called their attention to its value.

to operate it. The furnace is worked by Mr. L. Huillier on contract, the company paying him a certain price per ton for the iron delivered on the dock in Marquette.

The articles of The Iron Cliff Co. were filed September 15th, 1864, with a capital stock of \$1,000,000, in 40,000 shares at \$25 each. Corporators: William B. Ogden and John W. Foster, of Chicago, and Samuel J. Tilden, of New York. Office at Negaunee, Mich. This company in 1864 purchased of the St. Mary's Ship Canal and Mineral Land Co. the 38,000 acres of land which that company owned in Marquette county, Subsequently, as heretofore mentioned, the Iron Cliff Co. came into possession of the Pioneer Co.'s property, thus increasing its estate to over 40,000 acres. The company soon began the construction of a furnace near the Foster mine, which has never been completed. They own and are working the Barnum and the Foster mines, the latter of which was opened in the spring of 1865. The product is a soft hematite, which forms a good mixture with hard ores. This mine is situated on Secs. 22 and 23, T. 47, R. 27. The first shipment of ore therefrom was made in 1866, and the mine has since been continually worked.

The Barnum mine is situated on Sec. 9, T. 47, R. 27, connecting with the Lake Superior Co.'s principal opening. The first shipments of ore were made during 1868, the ore being specular and of excellent quality. The C. and N. W. R. R. has a branch running into the mine, over which shipments are made. The mine is supplied with pumping and hoisting machinery. The map of the Lake Superior mine, which will be found in the accompanying Atlas, embraces the Barnum mine.

On that portion of the estate purchased of the Excelsior Company, in addition to the Barnum, a deposit of specular ore has been found near the corner of Secs. 5, 6, 7, and 8, T. 47, R. 27, which promises well; a branch railroad has been surveyed to it. Besides those already mentioned the company have several other openings. One on Sec. 15, adjoining the Pittsburgh and Lake Angeline Co., opened during the past season, which gives a fine showing of hematite ore. The Cliff-Parsons, also opened during the past season, adjoins the Old Parsons, on Sec. 21, T. 47, R. 27.

Another opening is near the quarter-post between Secs. 17 and 18, T. 47, R. 26, from which ore was shipped during the season. A second opening is being made on this same line, at a point farther north, near the section corner. These openings belong to the Negaunee Hematite Group. In addition to their own mines the company are working the Pioneer opening of the Jackson mine on a lease. Near the Foster mine the company have in operation a sawmill, to which is attached shingle and lath mills.

In 1864 the Ogden and Tilden mines, situated on Secs. 13, 23, and 24, T. 47, R. 27, were extensively opened, and the branch road, which also extends to the Foster, built to them. The ores, however, proved of too low a percentage to sell in the then existing market, and the work was abandoned. The purchasers of the Iron Cliff estate also controlled the Chicago and Northwestern Railroad, and a short time previous to the purchase effected a consolidation with the Peninsula Road of Michigan, with a view to the future development of iron deposits on this extensive property, and the control of the railroad facilities for transporting the product of these and other mines to Lake Michigan.

The Iron Mountain Mining Co. filed its articles of association Nov. 1, 1864, paid in \$100,000. Corporators: Geo. E. Hall, of Cleveland, O., Richard Hays, Henry A. Laughlin, and Irwin B. Laughlin, of Pittsburgh, and Gilbert D. Johnson, of Ishpeming. The company own 320 acres of land, being the S. ½ of Sec. 14, T. 47, R. 27. The first shipments of ore were made in 1865, a branch of the C. and N. W. R. R. extending into the mine. All work at this mine has been discontinued, owing to the leanness and refractory nature of the ore, its yield being less than 50 per cent. of iron in the furnace. This mine has been recently leased to Messrs. Clark and Colwell, under whose auspices work will be resumed in the spring of 1873, with the view of finding hematite.

The Michigan Iron Co. filed its articles of association Dec. 30th, 1864. Capital stock, \$500,000, in 20,000 shares of \$25 each. *Corporators:* Henry J. Colwell, Andrew G. Clark and Samuel P. Ely, of Marquette, with office there.

This company own a large amount of woodlands, two furnaces and considerable other manufacturing property. The Michigan

The mining operations of this company are conducted in the southeast 1/4 of southeast 1/4, Sect. 3, T. 47, R. 27, being 16 miles west from Marquette and adjoining the Cleveland. The mine is worked under a lease from Mr. A. R. Harlow and the stock is all held by Mr. S. J. Tilden and Messrs. W. L. and F. W. Wetmore. Operations were commenced in the mine in 1864, during which year 8,000 tons of ore were shipped. The statement of its yearly product and other details will be found by reference to the tables in this work; the workings and geological structure are shown by a map. This company is identical with the New York and Boston Iron Mining Co., and also with the New York iron mine, incorporated March 31st, 1865; it soon after changed to the New York Iron Mining Co., as above described.

The Pittsburgh and Lake Angeline Iron Co. was incorporated Nov. 11th, 1865. Capital stock, \$500,000, in 20,000 shares of \$25 each. James Laughlin, President, T. Dwight Eels, Secretary and Treasurer. The company own 1,376 acres of land, situated in T. 47 and 48, R. 27 and 28, of the former Town., and R. 31 of the latter. They also hold a lease of about 300 acres, on which is located the Edwards mine. The company's mines consist of the Lake Angeline and Edwards; the Lake Angeline mine is situated on the south shore of Lake Angeline and on the line of the M. H. and O. and C. and N. W. R. Rs., 17 miles from Marquette and 66 miles from Escanaba, and produces both specular and hematite ore, the latter of first quality.

The Edwards mine lying contiguous to the Washington, is also on the line of the M. H. and O. R. R., distant from Marquette 28 miles, and produces only magnetic ore. Work was commenced in 1865, the first shipments being made in the following year. The mining is all conducted underground, the ore being raised to the surface through shafts and is the only mine in the Iron Region which has been exclusively worked in this way. The results of this company's operations are shown in the accompanying tables and the mine workings by maps and illustrations.

The Schoolcraft Iron Co. filed articles of association April 8th, 1866. Capital stock, \$500,000, in 20,000 shares of \$25 each. Paid in, \$250,000; the remaining 10,000 shares being held by the com-





JACKSON IRON CO'S. FURNACES AND COAL KILNS Fayette, Big Bay de Noquette, Mich.

Corporators: Hiram A. Burt, Peter White and H. R. Mather, of Marquette; office at Marquette, Michigan.

A furnace was constructed by this company at Munising, Schoolcraft county, on Grand Island bay, which went into blast in June. 1868, and was blown out in about six months thereafter. The furnace continued "in and out" of blast somewhat irregularly, until the company went into bankruptcy. In 1871 the furnace and other property, including 40,000 acres of hard wood land, which had belonged to them, passed into the hands of Peter White, Esq., by whom it was transferred to the Munising Iron Co., an organization effected for the purpose of owning and operating this estate, which is now being successfully done. Mr. Peter White, of Marquette, is managing director.

The Marquette and Pacific Rolling-Mill Co. filed its articles of association Oct. 1st, 1866. Capital stock, \$500,000, in 20,000 shares of \$25 each. The corporators were John Burt, Samuel P. Ely, Wm. Burt, Edward Breitung, Timothy T. Hurley, Cornelius Donkersley, W. L. Wetmore, Peter White and Alvin C. Burt, of Marquette. Office in Marquette, Mich.

The company has constructed at Marquette a bituminous blastfurnace, with rolling-mill connected therewith. The works are located near the lake shore, at a short distance south from the city. went into operation in the summer of 1871, and are connected with the M. H. and O. R. R. by a branch track. Upon their land at Negaunee, the company have opened a mine of manganiferous hematite ore, to which a side track has been extended, connecting it with both railroads; from this mine the company's furnace at Marquette is in part supplied. This rolling mill is the first erected on Lake Superior, and the furnace the first which has continually used bituminous coal. H. A. Burt is superintendent.

The Fayette Furnace was constructed and put in operation in December, 1867, the enterprise originating with Major Fayette Brown, general agent of The Jackson Iron Co. It is located at "Snail Shell Harbor," in Big Bay de Noquette, 20 miles east of Escanaba, and about it has grown up the beautiful village of Fayette. It is owned by the Jackson Iron Co., with general office in Cleveland, Ohio. The company own 16,000 acres of land, excellently well timbered with hard wood, and generally adapted to

agricultural purposes, the soil being of limestone formation. From the ledges of limestone, which exist in the immediate neighborhood, material for the necessary flux is obtained, as well as for the manufacture of all the lime used by the company. They possess a full complement of charcoal kilns, and a large portion of the necessary wood is purchased, the company preferring to save their own timber as long as possible. This wood is delivered by the parties of whom it is bought at the furnace, or along the line of the company's railroad, of which they have constructed for this purpose six miles, laid with T-rail, and operated with two small locomotive engines, it being the only furnace on the Upper Peninsula that operates a locomotive railway for the exclusive purpose of transporting fuel. The company have also a saw-mill, machine-shop, etc. The furnace, as originally started, consisted of a single stack, which is shown in the accompanying illustration. A second one was subsequently erected, and both stacks have since been in operation with results more favorable, than any other charcoal furnaces using Lake Supe-The extraordinary favorable working of these furnaces will be fully realized from the following statements, furnished from the company's reports: During the 73 days immediately preceding April 13th, 1872, there were made in the No. 1 stack an average of 27 tons per day, using 94 bushels of charcoal and 125 lbs. of limestone per ton, the ore being from the Jackson mine and yielding from 62 1/2 to 64 1/2 %. On August 4th following, the same stack again went into blast, making, during the first quarter, a period of 91 days, 2,258 tons of iron, an average of 27 8 tons per day, using by measure 92 bushels of charcoal per ton. No. 2 was also in blast during a portion of the same period with corresponding results. On December 14th No. 2 stack had produced, during the previous four weeks, an average of $26\frac{39}{100}$ tons per day, and on January 18th, 1873, had produced, during the previous five weeks, an average of 29,34 tons per day; the charge used during this time was 261/2 (called 30) bushels of charcoal, 1,000 lbs. of ore (1/2 soft and hard specular Jackson), 35 lbs. of limestone and 10 lbs. of clay.

These results require no comment relative to the efficiency of the management. The coal is of the best quality, kept dry under shelter, as is also the ore, which is crushed finer than is customary. The stacks are each 42 feet high inside and 9 feet 6 inches bosh;

4 feet 8 inches, and 5 feet 8 inches diameter, 3 feet below the top, and 4 feet and 5 feet at the top respectively. The hearths are 4 feet diameter battering from the bottom; the tuyeres, three in number, with 31/2 inch nozzle, are placed 40 inches above the bottom of the hearth. Two blowing engines are used, the cylinders respectively 36 and 48 inches in length, with diameter of 50 and 44 inches. The engines make from 24 to 28 revolutions per minute, and both of them are only run when the two stacks are in operation. The temperature of the hot blast averages in one about 600° and in the other 750°. Originally No. 2 stack had a five-foot cone, but did not make as much iron, nor as cheaply, as the other, until the cone was reduced in height to 4 feet 4 inches, since which time it has worked equally well with the other. The total product of these furnaces during 1871 and '72 was 19,117 tons, which were used as follows:

For	Bessemer Steel	17,465	tons.
"	Malleable Iron	88	44
66	Wheels	787	66
**	Foundry, etc	400	"
	Forge purposes	377	166

Genl. Agt., Major Fayette Brown, Cleveland, Ohio. Local Agt., C. L. Rhodes, Fayette, Mich. Founder, Jos. Harris, Fayette, Mich.

The Deer Lake Iron Company.—Articles of association were filed July 9th, 1868. Capital, \$75,000-3,000 shares at \$25 each. Corporators: George P. Cummings, of Marquette, Edward C. Hungerford, of Chester, Conn., Gardner Green, Caleb B. Rogers, Moses Pierce, Samuel B. Case, Theodore T. McCurdy, John E. Ward, James Lloyd Greene, James C. Colby (Ex'r), Daniel T. Gulliver, William R. Potter and Enoch F. Chapman, of Norwich, Conn.; Giles Blague, Jr., New York, Geo. Smith, New York, G. F. Ward, E. R. Ward, Old Saybrook, Conn., and James H. Mainwaring, of Montville, Conn., with office at Marquette, Mich.

This company organized for the purpose of smelting iron ore, and immediately constructed a furnace, which went into operation in Sept., 1868. This furnace, the smallest in the district, is located at Deer lake on the Carp river, two miles north from the village of Ishpeming on the M. H. and O. R. R., with which place it is

connected by a tram railway. The stack is 33 feet high and 7 feet 8 inches bosh, thus making it perhaps the smallest furnace which has been built in the United States during the past 7 years. Another peculiarity of this furnace is the comparatively enormous size of its hot-blast oven, to which is doubtless due in part the favorable results, which, considering its small size and peculiar management, the furnace has accomplished. The oven, on the Pleyer plan, contains 45 tons of metal, which is 50 per cent. more than that contained in the ovens of our largest charcoal furnaces; having twice the capacity of the Deer lake stack. The furnace is driven by water, employing an 18-inch turbine wheel under 35 feet head, thus leaving all the gas available for heating the blast, which is brought to an extremely high temperature. It runs but six days in the week, "banking up" Saturday night and starting again on Sunday night. Notwithstanding an arrangement necessarily disadvantageous to the greatest production, the furnace has averaged during several consecutive weeks II tons of pig-iron per day, using 110 bushels of charcoal to the ton, one-half of which is made from pine slabs,—the ore used being hard ore from the New York mine, averaging 66 per cent. The origin of this enterprise is due to Mr. E. C. Hungerford, who also determined its unusual size and the peculiar policy under which the furnace has been managed. Near the present one the company are now building a new iron shell furnace, 9 feet bosh.

The Cannon Iron Company.—Articles filed July, 1869. Capital, \$500,000; 20,000 shares, \$25 each. *Corporators:* Bernard A. Hoppes and Wm. H. Berry, of Philadelphia, and Samuel S. Burt, of Marquette, with office in Philadelphia. This company organized for the purpose of mining iron ore, but beyond making explorations on their lands with this view, nothing has as yet been done.

Bay Furnace Company.—Articles filed July 19th, 1869. Capital stock, \$150,000; 6,000 shares at \$25 each. *Corporators*: William Shea, of Munising, Mich., George Wagner, Jay C. Morse, Frank B. Spear and James Pickands, of Marquette, John Outhwaite, of Cleveland, and John P. Outhwaite, of Ishpeming, Mich., with office in Marquette.

This concern organized for the purpose of smelting iron ore, and

immediately proceeded to the construction of a blast-furnace for that purpose. This furnace was completed and went into operation on the 6th of March, 1870. It is located at Onota, in Schoolcraft county. on Grand Island bay, 40 miles from Marquette. But one stack was originally constructed; a second one, however, has since been erected and put in readiness for the blast. The ore used is from the Cleveland and McComber mines, received by the way of Marquette. This company own about 20,000 acres of land, mostly hard wood timber, from which the fuel for the furnace is obtained.

The Whetstone Iron Company.—Organized Aug. 20th, 1869. Capital stock, \$150,000, in 6,000 shares of \$25 each. Office at Marquette. This company have not commenced operations. Corporators were William Burts, Samuel Peck, A. A. Cole, Thomas O. Hampton, Clark Stratton, A. S. Harvey and A. G. Benedict.

Champion Iron Company.—Organized August 23d, 1869, with a capital stock of \$500,000, in 20,000 shares of \$25 each. Corporators: Joseph S. Fay, of Boston, Edwin Parsons, of New York, Thomas C. Foster, of Cambridge, Mass., and Samuel P. Ely and Peter White, of Marquette. The company own about 1,600 acres of land, but their mining operations are conducted on that portion of their land comprising the south half of Sec. 31, T. 48, R. 29, being 32 miles by railroad from Marquette. The ore is principally magnetic, though a large amount of slate ore is obtained. Champion mine is upon the south outcrop of the magnetic ore basin, which underlies Lake Michigamme, and near the village of Champion, about half a mile distant from the furnace of that name. The company are now working chiefly underground, as is fully shown in Map VII. of Atlas, where the geological structure and all other important details will also be found.

The Lake Superior Foundry Company filed their articles of association Sept. 2d; 1869, with a capital stock of \$50,000-2,000 shares at \$25 each. Corporators: Daniel H. Merritt, Lotan E. Osborn, Henry J. Colwell, William L. Wetmore, Jay C. Morse, Alfred Kidder, James Pickands and Thomas Fitzgerald, of Marquette, Mich.; Gilbert D. Johnson, Seymour Johnson, Harvey Diamond and Robert Nelson, of Ishpeming. The works (located at

Ishpeming) are quite extensive and adapted to general and particular foundry and machine work. (See Iron Bay Foundry, p. 33.)

Silas C. Smith Iron Company.—Articles of association filed Jan., 1870. Capital, \$500,000, in 20,000 shares at \$25 each. *Corporators:* Silas C. Smith, of Ashtabula, O., Oliver F. Forsyth and Wm. H. Lyons, of Flint, Mich., with office at Ashtabula, O.

The property of this company consists of 703 acres of land in Sections 18, 20, and 28, T. 45, R. 25, upon which have been made numerous openings, showing soft hematite ore in quantity, the main one being near the E. ½ post of Sect. 18. A tunnel is being driven into the deposit, of sufficient size for the admission of railway cars from a branch road five miles in length, which connects with the Chicago and Northwestern railroad. The ore at present is loaded into the cars from temporary docks, provided with pockets for that purpose. The principal stockholders are Silas C. Smith, the discoverer, General James Pierce, of Sharpsville, Pa., and Henry Fassett, of Ashtabula, O. The shipments of ore and other details will be seen by reference to the mining tables.

The Pittsburgh and Lake Superior Iron Co. filed articles of association June 28th, 1870. Capital stock, \$500,000, in 20,000 shares of \$25 each. Corporators: James McAuley, C. T. Spang, C. G. Hussy, Thos. M. Howe and James M. Cooper, of Pittsburgh; Sherman J. Bacon, of New York, Joseph G. Hussy, of Cleveland and W. M. Sinclair, of Philadelphia; with office at Pittsburgh, Pa. The company own 2,691 acres of land in Towns. 47 and 48, Ranges 25 and 26, their title to which was derived direct from the United States Government. Work was commenced on their property near the Cascade mines in Sept., 1872, houses, etc., were erected, a railroad side track built and a pit opened on Sec. 32, which is called the Hussy mine, and from which about 2,000 tons have been shipped.

The Republic Iron Co. was organized Oct. 20th, 1870. Capital stock, \$500,000, in 20,000 shares. Office in Marquette. Corporators: E. Breitung, S. P. Ely and Ed. Parsons. This company own 1,328 acres of land, being in part in Sections 6, 7, and 18, T. 46, R. 29, comprising what was formerly known as Smith mountain, which

is unquestionably one of the largest deposits of pure specular and magnetic ore on the Upper Peninsula, if not in the United States. The great extent and value of this deposit was observed and commented on by the early United States surveyors, when engaged in running the township lines in that locality in 1846. The property was explored and selected by Silas C. Smith, of Marquette, and entered in the name of Dr. James St. Clair, in 1854 and 1855. branch from the M. H. and O. R. R. has been constructed to the mine, over which the shipments of ore are now being made. See Tables, Plts. XII. and XIII. of Atlas. A complete map of this property, based upon careful surveys, exhibiting the topography, geological structure, magnetism and other important details, will be found in the Atlas accompanying this work, together with full descriptions.

The Cascade Iron Co. is an association of Pittsburgh men, owning 3,120 acres of land in Sections 19, 20, 29, 30, 31, and 25, T. 47, Ranges 26 and 27. These lands were entered by Waterman Palmer and purchased by the present company in 1869. An examination of the iron deposits in this locality was made by Dr. Douglas Houghton, in 1845, while engaged in running the interior section lines. (See Appendix D., Vol. II.)

The company's mines are provided with side tracks, connecting with a branch road of six miles in length to the C. and N. W. R. R. Mining operations commenced in 1871, and the openings (including the leased mines) are seven in number. There are other improvements, such as a saw-mill run by water, a store, sufficient number of dwellings, barns, repair-shop, etc. The expenditure which these improvements (including the branch railroad and side tracks) have necessitated has been very large, and future operations are contemplated upon a scale of considerable magnitude. (See Statistical Tables.)

The Cascade Company, under another organization, to wit, The Escanaba Iron Co., are constructing a blast-furnace at Escanaba, to consist of two stacks, one of which will go into operation in January, 1873; the height of stack, 56 feet; diameter of bosh, 12 feet. The entire structure is built in the most complete and substantial manner, and when finished, will probably not be surpassed, if equalled, in capacity, durability, or beauty, by any similar furnace in the United States. The principal owners are Joseph Kirkpatrick, William Bagaley, James Lyon, William Smith, Samuel Riddle and Samuel Hartman; Joseph Kirkpatrick, *President*, James Lyon, *Treasurer*, and John L. Agnew, *General Superintendent*.

The Emma Mine, one of the Cascade openings, is on the E. ½ of E. ½ of N. E. ¼, Sec. 31, and is being worked under a lease from the Cascade Company by an association of Pittsburgh gentlemen, who are represented at the mine by Mr. James E. Clark. They commenced shipping ore in 1872.

The Bagaley Mine, likewise one of the Cascade openings, is also worked under a lease from the Cascade Company, by Messrs. Wilcox & Bagaley, and its total product is about 6,000 tons.

The Gribben Iron Co., having a capital stock of \$500,000, in 20,000 shares of \$25 each, was organized 1872. The mining property comprises a lease on the S. E. ¼, Sec. 28, T. 47, R. 26, being on the Cascade range. Mining and exploring operations during the season have resulted in taking out considerable ore, some of which has been shipped for testing. The company have built a side track, which connects with the Cascade branch of the C. and N. W. R. R. Officers of the company are: W. C. McComber, President, C. H. Hopkins, Secretary, and James Mathews, Treasurer; all of Negaunee, Mich.

The Carr Iron Co. was also organized in the summer of 1872, with a capital stock of \$250,000. Its real estate comprises forty acres of land, situated on Sec. 33, T. 47, R. 26, being also in the Cascade range. The officers are Amos Root, *President*, Jackson, Mich.; E.W. Barber, *Secretary*, Jackson, Mich.; and W. H. Maynard, *Managing Director*, Marquette.

Negaunee Hematite Mines. A large number of new companies have recently been organized for the purpose of mining hematite ore in the vicinity of Negaunee. These new locations, which have been and are in process of being developed, are situated in Sections 6, 7, 8, and 18, T. 47, R. 26, and comprise what are known as the Mc-Comber, Grand Central, Rolling Mill, Himrod, Ada, Negaunee, Calhoun and Spurr, Green Bay, Allen, the Iron Cliff "Sec. 18," and other mines. The McComber mine, opened by William C. Mc-

Comber in 1870, is worked on a lease from J. P. Pendill, of Negaunee, at a royalty of fifty cents per ton for ore. The mine has been worked for the past three seasons, and in the spring of 1872 the lease was sold to parties interested in the Cleveland mine, who in July organized a company. The Rolling Mill mine, heretofore spoken of, is worked in part under a lease from A. L. Crawford. The company, however, own the greater portion of the land.

All these workings, except Sec. 18 and the McComber, are worked on leases from Edward Breitung, at 75 cents per ton royalty, he having leased from the owners, Messrs, Harvey and Reynolds, at 50 cents per ton royalty. Some of these pits have been worked during the past season, and nearly all of them are prepared for active operations during the coming year. Railroad side tracks are either completed, or in process of construction, to the several mines; dwellings and other improvements have been made. or are contemplated at each, and several of the locations bid fair to be the scene of active mining operations. The product is for the most part a soft hematite, containing usually from one to five per cent. of manganese, which renders the ore more easily worked in the furnace and is probably beneficial to the iron. The yield of metallic iron of the best of these ores is 50 per cent. and upwards, the average, however, is below that. See Map No. V. and Table Pl. XII. of Atlas.

Among the promising iron properties upon which work has been . commenced during the present season, and from which large shipments may be reasonably anticipated, are the Michigamme and Spurr Mountain mines, at both of which work has actively commenced; side tracks are being constructed at both places, connecting with the M. H. and O. R. R. The mines are situated upon the same magnetic range and are about two miles apart.

The property of the Spurr Mountain Co. (which company was organized in September last) comprises 160 acres of land, and the point at which mining operations have been commenced is at what is known as Spurr mountain. The preliminary work has uncovered the south side of a very large mass of magnetic ore of a great degree of purity; rising at the highest point to a height of 60 feet above the surface of the ground at the base of the hill. This remarkable outcrop of ore is situated (as will be seen by reference to the accompanying map) 900 feet east and 700 feet north from the west and south boundaries respectively of the company's property. It was first discovered to the public in 1868. The examinations which have been made, established beyond any reasonable doubt the presence of the ore in a very large quantity and of a uniform purity and quality. The natural facilities afforded at Spurr mountain for commencing mining operations are excellent, and with the exception of Republic mountain there is, so far as known, no other locality in Marquette county where occurs so large an exposure of pure ore, rising at so great an elevation above the general level and at which there is apparently so little preliminary work necessary.

This range has been explored to a considerable extent in either direction; westerly, across the east half of Sec. 23, owned by the M. H. and O. R. R. Co., the examinations show the presence of the ore, but to how great an extent the deposit exists future workings alone can determine; easterly, as is elsewhere more fully related, the range has been traced along the north side of Lake Michigamme for several miles. The officers of the Spurr Mountain Co. are: H. N. Walker, Esq., of Detroit, *Prest.*; Col. Freeman Norvell, *Supt.* and *Sec.* The distances from the mine to the ports of L'Anse and Marquette are respectively, by rail, about 24 and 39 miles.

The Michigamme Co. was organized in the winter of 1870-71, the organization being effected mainly by persons already largely identified with Lake Superior iron interests. The land owned by the company comprises 1,400 acres, situated on the north side of Lake Michigamme. Preliminary work was begun in the spring of 1872, and prosecuted during the summer. The point selected for the commencement of mining operations is near the shore of the lake, and upon each side of the line between Sections 19 and 20, the developments resulting from this work thus far being of the most promising character. Improvements, not previously indicated, consist of a large, substantial steam saw-mill, with other machinery attached thereto, an office, dwellings, etc. At a short distance south and west from this location the company have laid out a village plat, to be called "Michigamme," and which promises to be built up with considerable rapidity. The distance to L'Anse is about 26

miles, and to Marquette 37, by rail. The officers of the company are: William H. Barnum, of Lime Rock, Conn., Prest.; James Rood, of Chicago, Sec. and Treas.; and Jacob Houghton, Supt.

The Keystone Iron Co. also organized in the fall of 1872, with capital stock of \$500,000, in 20,000 shares of \$25 each. The property comprises the southeast 1/4 of southwest 1/4, Sec. 32, T. 48, R. 29, distant from Marquette, by rail, 29 miles, from Escanaba 77, and from L'Anse 35. The company are at work preparing for mining the ensuing season. A. P. Swineford, Marquette, General Agent.

A number of mining enterprises, comprising The Albion, Saginaw, Lake Superior Company's new openings, The New England, Winthrop, Shenango, and Parsons, in Secs. 19, 20, 21, 16, T. 47, R. 27, are situated east and west, parallel and contiguous ranges of specular and hematite ore, are all connected by branches with the M. H. and O. R. R., and soon to be with the C. and N. W. Road.

The Albion mine, opened in 1871 by the brothers St. Clair, who hold the property comprising the northeast 1/4 of the northwest 1/4, Sec. 19, on a lease from Messrs. E. Breitung and S. L. Smith, at a royalty of 75c. per ton; up to the present time but a small amount of ore has been mined. The opening is immediately west of the Saginaw mine and on the same ore belt.

The Saginaw Mine, situated on the northwest 1/4 of the northeast 1/4 of Sec. 19, T. 47, R. 27, was opened in 1872, and during the same season shipped (via M. H. and O. R. R.) 19,000 tons of specular ore. The mine was worked on a lease by Messrs. Maas, Lonstorf and Mitchell, of Negaunee, on a royalty of 50c. per ton for the ore. During the fall of 1872 the lessees sold out to parties representing the Cleveland Rolling Mill Co. for \$300,000, and immediately thereafter the Saginaw Mining Co. was organized with a capital stock of \$500,000 in 20,000 shares. A. B. Stone, of Cleveland, Prest., and A. G. Stone, of Cleveland, Sec. and Treas. A side track has been surveyed, to connect with the Chicago and N. W. Railroad, and the grading finished to the Winthrop mine. The land on which the Saginaw mine is located was purchased of the State of Michigan, with four other contiguous "40's" situated about the

centre of same section, seven years ago, by Messrs. Heater, Elison and Conrad; the latter having made the selections.

Between the Saginaw and New England mines, on Sec. 20, the Lake Superior Iron Co. have a very promising opening, from which a considerable shipment of specular slate ore was made in 1872.

The New England Mine, on same range, is situated on the east ½, northeast ¼, Sec. 20, T. 47, Range 27. The shipments from this mine commenced in 1866, and up to the present time about 60,000 tons of ore have been mined and shipped via Marquette. The property is mainly owned by Captain E. B. Ward, of Detroit, and the mining operations are conducted by H. G. Williams under a contract. The principal part of the product is a hematite ore. A very narrow bed of excellent specular slate ore was worked several years, but not proving sufficiently profitable, work was discontinued. The ore is chiefly consumed at the extensive works controlled by Capt. Ward at Chicago, Milwaukee, and Wyandotte.

Adjoining the New England is the Winthrop Mine, situated in the southwest ¼, Sec. 21, T. 47, R. 47, owned by A. B. Meeker and A. G. Clark, of Chicago, and H. J. Colwell, of Marquette, and opened in 1870 by Messrs. Richardson and Wood, who work the mine on contract. Up to the close of 1872 about 25,000 tons of ore have been shipped, and the indications are favorable for increased shipments during the coming year. The product is a hematite ore, one of the richest of the class in the district. A. B. Meeker, of Chicago, is *Prest.*, A. G. Clark, *Sec.* and *Treas.*, and H. G. Colwell, Clarksburgh, *Gen'l Agt*.

The Shenango Iron Co. was organized in September, 1872, with a capital stock of \$500,000, in 20,000 shares of \$25 each. The land worked by the company comprises the north-west ¼ of south-east ¼ of Sec. 21, T. 47, R. 27, and adjoins the Winthrop, the deposit being a continuation of that mine.

The officers are C. Donkersley, of Appleton, Wis., *Prest.*, and H. D. Smith, *Sec.* and *Treas.*; in addition to these, E. Decker, Charles Reis and George L. Hutchinson, constitute the Board of Directors. A small amount of ore was shipped during the fall of 1872, and the company are erecting machinery, including the sink-

ing of a shaft 60 feet in depth, with the view of doing considerable mining the coming season. The land is leased of the Williams Iron Co., who in turn lease of the Pittsburgh and Lake Angeline Co., who are the owners of land. The ore is mined by Messrs. Hurd and Orthey, part owners, on contract.

The Boston Mine, situated on the southwest 1/4 of the northeast 1/2 of Sec. 28, was organized in 1872, and a lease of the property above described secured by Messrs. Day, Anderson and others, with a view of mining operations. The lease of these parties is the same as that of the Shenango.

The Parsons, or "Old Parsons," mine is located between the New England and the Lake Superior Companies' opening on Section 16, northeast of the Winthrop. Several thousand tons of specular slate ore were shipped from each of these mines, but work has been discontinued.

The Kloman Iron Co. was organized in December, 1872, with a capital stock of \$500,000, in 20,000 shares. The corporators were Andrew Kloman, William Coleman, Thomas M. Carnegie, Jacob Houghton and T. B. Brooks. The company own 437 acres of land adjoining and northwest of the Republican mountain, being in part in Sec. 6, T. 46, R. 29, on the west side of the Michigamme river. The company have commenced mining on the continuation of the Republic mountain deposit and are building a short railroad to connect the mine with the Republic branch.

The Howell Hoppock Iron Mining Co. filed articles of association January 13th, 1873. Corporators: Lewis J. Day, Wm. R. Bourne, Wm. Rice, James S. Ward and Frank Austin. Office in Ishpeming, Mich. Organized to mine on the northwest 1/4 of northeast 1/4 of Sec. 28, T. 47, R. 27. Capital stock, \$500,000, in 20,000 shares.

The Watson Iron Co. filed articles of association January 16th, 1873, with capital stock fixed at \$500,000, in 20,000 shares of \$25 each. Corporators: C. J. Hussey, E. T. Daro, Thomas M. Howe, M. K. Moorhead, George F. McLeane, W. J. Moorhead, Charles F. Spang, John W. Chalfant, Campbell B. Herron and James W.

Brown, all of Pittsburgh, Pa., and James W. Watson, of Marquette county, Mich. The property of this company comprises the northwest ¼ of Sec. 32, T. 47, R. 26 and which constitutes \$325,000 of the capital stock. This ¼ section is a part of the estate of the Pittsburgh and Lake Superior Iron Co. and is on the Cascade range. Operations were commenced in September last by this latter company, of which mention has already been made under the Hussey mine.

In the **Menominee Iron Region** two companies, called respectively the Breen and Ingalls Iron Mining Companies, have been organized and are engaged in explorations, and in addition to the operations inaugurated by these companies, explorations are being made by private parties. The completion of the Peninsula railroad from Escanaba to Menominee, affording better promises for transportation, will stimulate operations of this character, which have heretofore been deferred from want of railroad communications.

The Breen Mining Co. owns 120 acres of land in Sec. 22, T. 39, R. 28, distant from Escanaba by proposed road 35 miles, from Menominee 55 miles and from Deer river 28 miles. The ore is chiefly flag, with some hematite. The property is being explored by Capt. E. B. Ward, J. J. Hagerman and J. W. Vandyke, who have an option of leasing or purchasing the mine. The officers are E. S. Ingalls, *Pres.*, T. B. Breen, *Sec.*, S. P. Saxton, *Treas.*, Thomas Breen, Bently Breen, and S. P. Saxton, *Directors*—all of Menominee, Mich.

The Ingalls Mining Co.'s property constituted 240 acres of land situated in Sections 8 and 9, T. 39, R. 29. The distance from Escanaba by proposed road is 44 miles and from Menominee 64 miles. The officers are E. S. Ingalls, *Pres.*, C. L. Ingalls, *Sec.*, and F. S. Mullburg, *Treas*.

An effort has been made to manufacture pig-iron by using peat as a fuel, but has not as yet proved in the requisite degree successful. A peat furnace was constructed at Ishpeming and went into operation early in the year 1872, but very soon went out of blast; subsequently it started again and made about 200 tons of iron and

again stopped, it being the intention to alter and enlarge the stack. the better, it is thought, to adapt it to the peculiarities of the fuel. The peat is prepared from a bed of the material which exists in proximity to the furnace.

The Ericson Manufacturing Co. was organized in April, 1872, to conduct general manufacturing operations, with a nominal capital of \$150,000. Corporators: Peter E. Ericson, John Carlson, A. J. Burt and Wm. Burt.

The company are operating a foundry and machine-shop, which they have built on Whetstone brook, within the city of Marquette. The machinery is driven by water-power.

Mr. Ino. Burt commenced, in September, 1872, the construction of a charcoal furnace, on the lake shore, at the mouth of the Carp river, south of Marquette. The stack is being built of stone, with a nine-foot bosh, and the whole is to be completed and put in operation in the spring of 1873. It is intended to supply the fuel from points along the lake shore, transporting it to the furnace in boats in the same manner that the wood for the Burt furnaces in Detroit is obtained, of which latter furnaces the one being built at the Carp will be a duplicate, and will be the first built on the Upper Peninsula based on this plan of obtaining fuel.

Very recently The Carp River Iron Co. has been organized, and own the furnace and about 500 acres of land at that point, including the water-power on the Carp, etc. The business office will be in Marquette.

SANDSTONES.

The Lake Superior sandstones are very carefully described by Dr. Rominger in his accompanying report, commencing with page 80, and the results of his observations, as therein described, are of great practical and scientific interest. There are two organized companies now engaged in quarrying and marketing sandstone within the limits of the city of Marquette, the locations being contiguous.

The Marquette Brown Stone Co. was organized in August, 1872, with a capital stock of \$500,000, in 20,000 shares. The corporators were Peter White, Wm. Burt, F. P. Wetmore, S. P. Ely,

Sidney Adams, J. H. Jacobs, H. R. Mather and Alfred Green. In addition to quarrying stone, the company's franchises include the mining and smelting of ore, etc. Office in Marquette, Mich.

This company's property was previously known as the Wolf Quarry, located on the farm formerly owned by J. P. Pendill, and has been worked for some time past, the stone being principally used in Chicago. It is of a uniform dark-brown color, free from pebbles and clay holes. It apparently exists in great quantity, and is readily quarried and transferred to vessels. Mr. Peter White is constructing in Marquette a fine business block with a variety of stone from this quarry, which is variegated and striped with different colors, giving to the building a unique and pleasing appearance.

The articles of association of **The Burt Free Stone Co.** were filed Oct. 3d, 1872. Capital stock \$500,000, in 20,000 shares of \$25 each. The corporators were John Burt, William Burt, Hiram A. Burt, A. Judson Burt and Wm. A. Burt. Office in Marquette.

This company have opened a quarry of sandstone adjoining the one described above and the deposit is similar, the stone being lighter colored.

Both companies are prepared to furnish stone in large quantities. For full description of the sandstone found in these quarries, see Dr. Rominger's report, pages 90 and 91.

In addition to the above, **The Lake Superior Stone Co.** has been more recently formed with the amount of capital stock and number of shares as the preceding. The company own and hold in lease about 296 acres of land, situated on the west side of Keweenaw bay and on the north side of Portage Entry. The stone outcrops horizontally in a bluff, which rises from the water of the bay and is thus readily accessible for removal from the bed to vessels.

It is intended to begin operations in the spring. The corporators are H. H. Stafford, V. B. Cochran, W. S. Dalliba, E. J. Mapes and A. Kidder. Office, Marquette, Mich. See Dr. Rominger's report, page 95.

The fine new Court-House at Milwaukee is built with sandstone obtained from Bass island, near Bayfield, on Lake Superior, at which point stones have been quarried for several years.

The quarry described by Dr. Rominger, page 89 of his report, is

now owned by Messrs. Winty and Mossinger, of Chicago, and Thomas Craig, of Marquette.

ROOFING SLATE.

There are three companies which were organized for the purpose of quarrying and selling roofing slate; but one of them, however, has actually commenced operations and is now at work on explorations.

The Huron Bay Iron and Slate Co. filed articles of association January 19th, 1872. Capital stock, \$500,000, in 20,000 shares. The corporators were Peter White, W. L. Wetmore, F. P. Wetmore, J. C. Morse, James Pickands, A. R. Harlow, M. H. Maynard, D. H. Ball, Wm. Burt, D. H. Merritt, Sidney Adams and H. R. Mather. Office, Marquette, Michigan. The company own 2,000 acres of land in T. 51, R. 31.

The Huron Bay Slate and Iron Co. was organized subsequently, with same capital stock and number of shares. The corporators are W. L. Wetmore, Peter White, M. H. Maynard, Wm. Burt, Thomas Brown, J. J. Williams, S. L. Smith, Alex. McDonald, John H. Knight, W. C. Wheeler, H. R. Mather, Jas. D. Reid, F. P. Wetmore and R. C. Wetmore. Office in Marquette. The company own 1,100 acres of land in T. 51, R. 31, and have commenced work near Slate river, about four miles south of Huron bay, on the northeast quarter of section 33 in the above town. The slate apparently exists in very large quantities.

The Stafford Slate Co., an association comprising H. H. Stafford, V. B. Cochran, E. J. Mapes, A. Kidder, J. M. Wilkinson, Wm. Burt A. J. Burt and W. S. Dalliba, own 1,900 acres in T. 51, R. 31. The operations of this company thus far consist in having cut out a road from L'Anse to their property on Section 27, in the above town, a distance of 15 miles.

The color of the slate found in T. 51, R. 31, is somewhat varied, the green, purple and gray are found on Sections 14, 15, and 16. South of this are found large deposits of black slate, extending several miles east and west, with an apparent thickness of several hundred feet, the cleavage planes dipping to the south.

SAW-MILLS.

The following saw-mills are now in operation, all of which, with the exception of the ones at Whitefish Point, at Onota and Fayette (the two former of which are in Schoolcraft county and the latter in Delta), are in Marquette county:

Name of Firm.	Location.
Decker and Steele	Eagle Mills.
Edward Fraser	Cherry Creek.
George Wagner	Laughing Whitefish Pt.
A. R. Harlow	
H. A. Stone	Bancroft.
Jackson Iron Co	Negaunee.
Iron Cliffs Co	
Mr. Jackson	Palmer Falls (Cascade).
Hartman and Connelly	Little Lake.
Cleveland Iron Co	Ishpeming.
Lake Superior Iron Co	
Deer Lake Iron Co	Deer Lake.
Michigan Iron Co	Clarksburg.
Michigamme Iron Co	Michigamme.
Edward Breitung	Republic Mt.
C. T. Harvey	Chocolate.
Bay Iron Co	Onota.

These mills produced in the aggregate, during the year 1872 (besides shingles, laths and a small amount of hard wood), thirteen and a half million feet of pine lumber, all of which, excepting the product of the three mills above designated, was, or will be, consumed in Marquette county. The total product during the coming year, if the winter is favorable, will be much greater, as most of these companies are preparing to get in a larger amount of logs. The Michigamme mill, which has a nominal capacity of 4,000,000 feet, has but recently started, and thus did not contribute to the total product of 1872.

COMPLETION OF THE RAILWAY SYSTEM.

Marquette, Houghton and Ontonagon R. R.

Among the most important events affecting the interests of this portion of our State, which transpired during the year 1872, was the extension of the C. and N. W. R. R. from Menominee to Escanaba, the consolidation of Marquette and Ontonagon Railroad with the Houghton and Ontonagon, and the completion of the line to L'Anse, thus making complete railroad communication from the head of Keweenaw bay to Chicago, a distance of 462 miles.

The development of the mineral resources of a country are so intimately blended with the improvement of its facilities for transportation, as to render it essential in considering the progress of the former, to give due credit to the latter. Iron ores having a low value per ton must be reached by rail or water before their value can be realized; differing in this particular from the ores of the precious metals, which will bear wagon or even pack-mule transportation. Especially is this true with reference to an isolated region like the Upper Peninsula, which is as yet a comparative wilderness, possessing but a small population, a rigorous climate, few thoroughfares and with a surface so rough and rocky in portions of its territory, as to render their construction a matter of much difficulty. It naturally follows, that the addition of two so important avenues of communication to the railroad facilities of the Peninsula becomes in a pre-eminent degree a matter of congratulation and importance. The history of the enterprise, which has thus resulted in the connection of the bays of Marquette and Keweenaw, is in brief as follows:

As has been previously related in speaking of the Peninsula road, the United States granted to the State of Michigan, by an act passed on the 3d of June, 1856, every alternate section of land for six sections in width, designated in odd numbers, to aid in constructing a railroad from Little Bay de Noquette to Marquette and thence to Ontonagon, and from the two last places to the Wisconsin State line. The State, by an act passed Feb. 14th, 1857, conferred this grant upon the Little Bay de Noquette and Ontonagon Railway

Co., and two other railroad corporations, all of which lines were required to be completed within ten years, a condition with which neither of the companies complied.

In 1863 the State conferred the forfeited franchises and grant previously given to the Marquette and Ontonagon Railway Co., upon the Marquette and Ontonagon Railroad Co., under certain conditions. Congress in 1864 extended the grant five years, in the subsequent year added four sections per mile thereto, and in 1868 fixed the time for a full compliance with the conditions of the grant until Dec. 31st, 1872. During the period of its existence, the company built twenty miles of main line of railroad, commencing near the Lake Superior mine at the terminus of what was formerly the Bay de Noquette road, and extending to a point on the south side of Lake Michigamme.

In 1870 the State decided that the company, by reason of its failure to complete any extension of their lines, had forfeited the greater portion of the grant. On the 24th of Jan., 1871, the Legislature confirmed the action taken by the State Board of Control during the month of April previous, which conferred the forfeited or unearned lands upon the Houghton and Ontonagon Railroad Company, a new organization, incorporated Jan. 15th, 1870, and of which the following Michigan men were among the principal stockholders: H. N. Walker, President, S. L. Smith, Chas. H. Palmer, Geo. Jerome and S. F. Seager. The conditions of the act of Congress required the completion of thirty miles of road before the close of the year 1872, which fortunately this company have succeeded in accomplishing. Jacob Houghton was chosen Chief Engineer; and having located the line from Champion to L'Anse during the winter, the construction was begun in the spring of 1871 at the L'Anse terminus, and on the 16th of Dec., 1872, the first train passed over the entire line to Marquette, sixty-four miles; the whole having been placed under one management by the consolidation of the two companies effected during the previous summer. The completion of the road to L'Anse, exclusive of innumerable other advantages, opens to market the products of several iron mines, among the most promising of the region.

In anticipation of future shipments of ore from L'Anse, the company have constructed at this terminus of the road an extensive dock, a full representation of which from careful drawings





Photo.by Childs.

ORE DOCK, MARQUETTE

M.H & O.Rd. (Looking East) Vessels loading- Break water.

is herewith presented.* They have also built, at this point, in a very substantial manner, a round-house, turn-table, machine-shop, etc.

The charter of the company and the grant of lands provide for the extension of the road to Ontonagon, and it is but reasonable to assume that the energy, which has characterized the prosecution of the enterprise thus far under its present efficient management, will result in the accomplishment of the work before the expiration of the time fixed by law. The length of the main line is 62 miles, of branches 20 miles and of sidings 18 miles, making 100 miles of road now constructed and in operation.

The dimensions and capacity of the company's railroad dock at Marquette, a representation of which is given in the accompanying view, are as follows: - Total length, 1,222 1/2 feet; working length, 720 feet; height above water, 38 feet, and width of top, 53 feet, on which are four tracks for cars. Whole number of pockets, situated on both sides, 136, of which 120 have a capacity of 55 tons each, and 16 (steamboat-pockets) of 100 tons each. From both sides 8 vessels can be loading at the same time, and 6,000 tons have been loaded in a single day. Three vessels arrived on Saturday, after 8 o'clock in the evening, and were loaded and gone early Sunday morning. Vessels with a capacity of 476 tons may be loaded in one hour and fifteen minutes; vessels of 683 tons, in one hour and thirty-five minutes; the average time is three hours. The average capacity of vessels is about 650 tons, ranging from 400 for the smallest to 1,100 for the largest. Total amount of ore shipped over the dock from May 12th, 1872, to the following Nov. 25th, 301,210 tons, of which 75,000 tons were taken by steam, and 225,-000 by sail-vessels; the estimated capacity of the dock, with a sufficient number of vessels to receive the ore, is 500,000 tons.

The working capacity is indicated by the amount of rolling stock, which at the opening of navigation, 1873, will consist of 1,600 orecars, 50 box and platform-cars, 7 passenger and baggage-cars and 28 locomotives. The present officers are: H. N. Walker, of Detroit, *President*, S. P. Ely, Marquette, *Vice-President*, Moses Taylor, New York, *Treasurer*, Freeman Norvell, Detroit, *Secretary*, Jacob Houghton, Michigamme, *Chief Engineer*.

Directors: H. N. Walker, Detroit, C. H. Palmer, Pontiac, S.

P. Ely, Marquette, John Steward, New York, Alexander Agassiz, Boston, S. L. Smith, Lansing, George Jerome, Detroit, Moses Taylor, New York, C. Francis Adams, Jr., Boston.

By the Peninsula division of the Chicago and Northwestern Railway the distance from Escanaba to Lake Angeline is 67 30 miles, and the branches completed and in course of construction, 37 90 miles; sidings, 15 90 miles; making a total length of track between these points of 121 10 miles.

The total amount of track between Escanaba and Menominee is $65_{\frac{70}{100}}$ miles, of which $2\frac{30}{100}$ are side-track, making a total amount of track between Menominee and Lake Angeline, inclusive of sidings and lurches, 186,80 miles.

Estimated amount of rolling stock, which will be necessary and available for the business of 1873, between Escanaba and Negaunee:

Number of locomotives	
other cars	100
For the estimated business between Escanaba and Men	ominee :
Number of locomotives	6
" cars (exclusive of ore-cars)	100
S. C. Baldwin, Div. Supt. Marvin Hughitt, Gen. Supt. C. & N. W. R. R.	

Statistics showing past production, with present condition and capacity of the mines and furnaces of the Upper Peninsula, might properly follow this historical sketch, thus bringing it to date and supplying facts, which could not well have been incorporated into the text. It was thought better, however, to arrange such information in tabular form, which has been done on Plates XII. and XIII. of Atlas, to which attention is here again called.

The Marquette Mining Journal, of Marquette, Mich., publishes an interesting yearly exhibit of the product and condition of the mines and furnaces.

In Appendix G, Vol. II., will be found statistics of population for the whole Upper Peninsula, from the United States Census for 1870.

CHAPTER II.

GEOLOGICAL SKETCH OF THE UPPER PENINSULA.

(Where to Explore.)

I. GEOGRAPHICAL DISTRIBUTION OF THE ROCK SYSTEMS.

In prospecting for valuable minerals the intelligent explorer should constantly observe several kinds of phenomena. If his search degenerates into a simple blind hunt for ore, he would deserve the success of a hunter who went into a gameless region, or who hunted for game whose habits he did not understand. The following general geological facts and laws will possess value to the explorer in enabling him to wisely select his field of labor and in prosecuting his work.

As all the sandstone suitable for building, which has yet been found in the Lake Superior region, belongs to a system of rocks named by geologists Lower Silurian, and all the workable deposits of iron ore have been found in another system called the Huronian, while all the copper and workable silver, in a third system appears known as the Copper-Bearing Rocks; and as no workable deposits of useful minerals have yet been found in the fourth and oldest system, the Laurentian or granitic rocks, it follows, that it is of the utmost importance to the explorer that he be acquainted with the boundaries of these several fields and not waste his energies on unproductive ground. I do not mean to assert that iron ore will not be found in the Silurian sandstones, for in St. Lawrence County, N. Y., and in the Maramec district, Missouri, valuable deposits of ore exist in rocks of this age. Large deposits of iron ore also occur in the Laurentian (granite) rocks of Canada and Northern New York, and again, the iron ores of Thunder bay are contained in rocks which the Canadian geologists declare to be the equivalents of our Copper series; but at this date it is a fact, that no workable deposits of iron ore have been found in the Upper Peninsula in rocks of these systems, and an explorer or miner would not be considered

wise, who should search for iron outside the Huronian limits. It is not only important that he be acquainted with the boundaries of the four great rock systems, but also with their leading characteristics. We will therefore first sketch in some detail the geographical distribution of these systems, as developed on the south shore of Lake Superior, beginning with the youngest and uppermost. The reader should have before him the map of the Upper Peninsula Pl. I. of the Atlas. The boundaries marked are not always exact, but embody the best information available and are not far wrong.

I. Lower Silurian.—The Lower Silurian system, the youngest or lowest division of the Palæozoic rocks represented on the Upper Peninsula, is made up of various sandstones and limestones which are fully described in Dr. Rominger's Report, Part III. The entire Peninsula, east of the meridian of Marquette, is underlaid by Silurian rocks and the "Copper range" is flanked by a Silurian flat on the south side, which separates it from the iron series, until the two, together with the South copper range, come together west of Lake Gogebic.

About two-thirds of the whole area of the Upper Peninsula, or 9,982 square miles, is underlaid by this system.

II. The Copper-bearing Rocks, corresponding with the upper copper-bearing rocks of the Canadian geologists, occupy a narrow belt on the northwestern edge of the Upper Peninsula. These rocks have less superficial extent than either of the other formations, underlying only about 1,186 square miles, or, say 7 per cent. of the whole surface. For descriptions of them see Prof. Pumpelly's Report, Part II.

III. The Iron-bearing Rocks, corresponding, it is assumed, with the Huronian system of Canada, consist of a series of extensively folded beds of diorite, quartzite, chloritic schists, clay and mica slates, and graphitic shales, among which are intercalated extensive beds of several varieties of iron ore. The same rocks occur on the east and north shores of Lake Superior, where they also contain iron. The Huronian area represented on the map equals about 1,992 square miles, or nearly one-eighth of the whole area of the Upper Peninsula.

IV. The Granitic Rocks, which so far have produced no useful minerals, and which are believed to be the equivalents of the Lau-

rentian of Canada, are represented as underlaying about 1,839 square miles, equal to 12 per cent. of the total area.

As our examinations in the southwestern part adjoining the Wisconsin line have not been thorough, there is considerable uncertainty regarding some of the lines dividing the Huronian and Laurentian rocks, and a portion of this region, equal to about 668 square miles, or 4 per cent. of the whole area, is left blank on the map.

While, as has been stated, it is not proven that iron ore may not exist in the other great systems in workable quantities, there is every reason to believe, that by far the greater part, if not all the workable deposits, are contained in the Huronian area above described. It must not, however, by any means be understood, that all of this area is iron-bearing. The several iron districts, which have been more or less explored, will be described in another place; they will be found to cover not more than about one-fifth part of the Huronian area, or, say one-fortieth of the whole area of the Upper Peninsula, and on less than one-half of this area have the ores been proven to have commercial value.

Recapitulation.

I. Lower Silurian area, about9,982	square miles.
II. Copper-bearing area, about	166
III. Huronian or Iron-bearing area, about1,992	.66
IV. Laurentian area, about	"
Unknown area, about 668	**
Total area of Upper Peninsula,	
exclusive of islands, about 15,667	. "

In a complete and systematically arranged geological sketch the lithology of the four systems would properly belong here, but what is written on this subject necessarily pertains almost entirely to the Huronian, the whole matter will therefore be considered in Chapter III., following, and in Appendices A, B and C, Vol. II.

II.—TOPOGRAPHY.

It is of importance to the prospecter to carefully observe the topography or form of the surface, for it is well known that useful minerals generally occur in corresponding topographical positions over considerable areas; again, the topography is the very best key to the nature of the underlying rocks, if these be concealed by earth, as is often the case. As the human physiognomy indicates the fundamental characteristics of the man, so the earth's physiognomy suggests the forces and materials lying beneath. It is safe to assert that within certain limits an experienced topographical geologist can, from a correct topographical map, judge of the nature of the rock underlying the surface represented; and conversely, from a geological map, he can predict the general form of the surface. In the same way, an experienced explorer does not hesitate to express an opinion as to whether he is on the "mineral range," from the form of the ground. We will now sketch in some detail the characteristic topography of the four great systems.

I. Silurian.—The prevailing surface characteristic of the Silurian region is a nearly level plain, underlaid by horizontal sandstones and limestones, often swampy and sometimes, where fire has destroyed the timber, a desert. The tame, flat, sandy and swampy country along the line of the Chicago and Northwestern Railroad, between Escanaba and Negaunee, is underlaid by Silurian rocks, but is far below the average in the value of its timber. Where rivers or water-courses have cut into these rocks, or waves wasted them, perpendicular bluffs are presented, which afford an excellent opportunity to explore and study the formation. The famous "Pictured Rocks" are bluffs of this character, from 50 to 200 feet high. From the top of these bluffs the country is flat, proving that they are the results of the action of water cutting its way into a horizontal plane, and are not, so to speak, built up and completed hills like those of the older rocks.

There is one apparent exception to this general flatness of the Silurian topography. Many of the highest hills and mountains in the Menominee iron region are capped with horizontal sandstone and limestone, which is never found in the valleys; the base, however, embracing the great mass of these elevations is always an old rock, and in the iron fields always Huronian. There is no doubt but that the sandstone once filled the valleys, extending in an unbroken bed of irregular thickness across the whole of the Menominee region, covering the older rocks, just as it now covers them further east. Since its formation it has here been mostly

eroded, but still caps the elevations as described. If it were all gone, the hills, made as they are, largely of highly inclined beds of quartzite, marble and ferruginous rocks, would remain, but with somewhat diminished heights.

Should the eastern part of the Upper Peninsula be elevated at any future time, so as to bring the underlying azoic rocks above the lake level, the Silurian rocks may there also become so eroded as to only cap the Huronian hills, as they now do in the region described. That the older rocks extend eastward under the Silurian, is, I suppose, a geological necessity, and is, I think, directly proven by the existence of local magnetic attractions in this Silurian area, which are undoubtedly due to the existence of beds of iron ore in the underlying Huronian. The explorer in the Menominee region finds these beds of sandstone much in his way, covering, as they do, in some instances, the ores.

Small lakes of clear water, with sandy bottoms but no outlets, are a characteristic feature of the Silurian area. The U.S. Survey maps represent about one-half of the whole surface of these rocks, which underlie the central and eastern portion of the Upper Peninsula, as swamp; the solid rock has often been found within a few feet of the surface in the swamp region. The western Silurian area being the prolongation of the Keweenaw Bay valley west, and embracing in part the Sturgeon, Ontonagon, Presquisle and Black rivers, has fewer lakes, much less swamp, and is more broken, than the eastern part already described.

Soft woods, including pine, are more prevalent on the Silurian rocks than on the older series; but on the other hand, some of the finest bodies of sugar-maple and beech found on the Upper Peninsula, are on these rocks. Beech has not, so far as I know, been found growing on the older rocks; whether this be due to climatic or soil influence has not been determined.*

The water divide, or height of land, of the central and east part of the Peninsula, is much nearer Lake Superior than Lake Michigan. It is an irregular line, approximately parallel with the shore of the lake, having an elevation where it crosses the Peninsula railroad of about 650 feet. See Map, Pl. I.

II. Copper-bearing Rocks.—There is probably no more striking

^{*} A timber map has been prepared, but could not be published for want of means.

topographical feature in Michigan, than the "Mineral" or Copper range, including Keweenaw Peninsula, of which it is the backbone. Ranges would better express the fact, for west of the Ontonagon river there are three; the Main or central Range which extends from Keweenaw Point far into Wisconsin, being flanked on the north by the Porcupine mountain range and on the south by the South copper range, each separated from the other by broad Silurian flats. The general trend of the three ranges is north, 60° east, and south 60° west, but they are not quite straight, as may be seen on the map. The ridge is broad, generally more than three miles, and the crest quite even, but is cut down to lake level at Portage lake, and further west is deeply eroded by the Fire steel, Flint steel, Ontonagon and other rivers. The surface of the ridge or plateau is from 500 to 600 feet high in the vicinity of Portage lake, and rises to a height of 884 feet at Mount Houghton, near Keweenaw Point. Between the Ontonagon river and Lake Gogebic the Central range attains, in isolated peaks, an elevation of 1,100 feet, and the Porcupine mountain range is over 900 feet high; the range is more broken towards the west, and in the vicinity of Rockland presents a series of oval mammillary hills with steep escarpments on the south side. This is also the character of the South copper range, between Lake Gogebic and Montreal river.

The iron range immediately south of the South copper range, and west of Gogebic, is lower, the hills having more gentle slopes; the range being in places obscured by low ground. As this is the only part of the Upper Peninsula, so far as I know, where the iron explorer may come in contact with copper rocks, it is important to observe the topographical differences above noted, especially as the copper traps in some places resemble the diorites or greenstones of the iron region. Lakes and swamps, so numerous in the iron and granite regions, are infrequent on the copper belt, as must follow from the form of the surface. The reason for the striking regularity in the leading topographical features of the copper range is to be found in the great uniformity in the strike and dip of the rocks, as is explained under Stratigraphy. The timber of the copper range is generally sugar-maple, is abundant and of excellent quality; very little pine or other soft wood occurs here.

III. Iron-bearing Rocks.—The topography of the Huronian rocks differs essentially from that of either the Silurian, or the copper

series. It is almost everywhere hilly and often mountainous, forming peaks higher than any in the copper range; but instead of a continuous range, or series of parallel ranges, it is rather a broad belt or irregular area of mountains, hills, swamps and lakes. It may be said, that the ruling topographical features, especially the mountains, have a general east and west trend, but there are numerous exceptions to this law; for example, the Michigamme river, from the lake to Republic mountain, runs northwest to southeast; and Michigamme lake itself has a north-south arm, nearly as long as the main lake, which runs east-west. The ridges west of Paint river, in T. 42, R. 33, run north-south, conforming with the bedding of the rocks.

Probably one of the most persistent ridges in the Marquette region is formed by the "lower quartzite," which outcrops on the shore of Lake Superior just south of Marquette, and rising rapidly from the lake it forms Mt. Mesnard on Sec. 34, T. 48, R. 25; from this peak it extends westerly, crossing the railroad at the Morgan furnace. then by way of the old Jackson Forge and along north side of Teal lake to south side of Deer lake, it holds its westerly course for a total aggregate distance of over 15 miles. The Chocolate and Morgan flux quarries and the Teal lake whetstone quarry are in this range. More persistent and conspicuous, and nearly as long, is the Greenstone ridge, which skirts the north side of the Michigamme and the Three lakes extending from the Bijiki river to the west end of the First lake, a distance of eleven miles:-points on this range are three hundred feet above Michigamme lake, which is 950 feet above Lake Superior. Summit mountain, one mile easterly from the Foster Mine, is one of the prominent landmarks of the region, looking as it does from an elevation of about 1,300 feet over the flat granite and Silurian region to the south. It forms one of a chain of hills which extend from the south end of Lake Fairbanks westerly for about 10 miles, but which form in no sense a ridge.

The mountains, or hill ranges, above described are exceptional in their regularity and continuity. Broken chains of irregular hills and short ridges of various sizes, separated by lakes and swamps, is the prevailing character; the highest hills are seldom over 300 feet above the low grounds at their base and about 1,300 feet above Lake Superior. Outcrops of rock, forming often perpendicular ledges of moderate height, are more numerous in the iron-bearing

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rocks, than in either of the systems described, except in the westerly part of the copper range. Although the relief of the surface is considerably modified by drift, it is generally plain that the strike, dip, and texture of the underlying rock has determined the general outline or contour; we should therefore expect that the great variation in these rocks, hereafter to be described, would produce this varied topography.

The topography of the Marquette region is very like the iron region of southern New York and northern New Jersey, except in its smaller elevations; a profile running north and south through the Jackson Mine, Marquette, would closely resemble a profile running northwest and southeast through the Sterling Mine, New York, platted say to half the scale.

Passing to the Menominee iron region, we find greater simplicity in the geological structure and a correspondingly less varied surface.

Obeying the influences of the great rock beds beneath, the elevations there have a tolerably uniform east-west trend and consequent parallelism. The south iron range, of which the Breen Mine is the east end so far as known, can be traced through a greater part of its course by a ridge, often bold, which crosses Town. 39, R. 29, and T. 40, R. 30, for a distance of over 15 miles, the bearing being west-northwest. The north iron range, about 12 miles from the other in the south part of Town. 42, Ranges 28, 29 and 30, is in places a prominent topographical feature. The capping of horizontal sandstones, which has already been mentioned as characterizing the Menominee hills, gives a somewhat more even character to the crest lines, and in places produces a strikingly different profile.

The Gogebic and Montreal river range, above referred to, is better marked by its running parallel with and lying south of the South copper range, than by any essential character of its own.

IV. Laurentian.—The surface of the granite country south of the Marquette region, at the same time the most extensive and best known, is not unlike that of the iron-bearing rocks on a much smaller scale. There are no mountains, the hills are lower, being usually mere knobs, seldom exceeding 50 feet in height; the ridges shorter and swamps more numerous. A coarse pitting of the surface, or promiscuous sprinkling of little hills, and low, short ridges may convey the idea. Sometimes the knobs range themselves in

lines constituting low ridges, with jagged crest line; these ridges, when near the Huronian rocks, are usually parallel with them; if they have any prevailing direction, it is east and west.

Perpendicular walls of granitic gneiss 15 to 40 feet in height sometimes face the ridges for several hundred feet in length, constituting the most regular topographical feature within the Laurentian area.

Small beaver meadows are common here as in the other rocks, and sometimes a succession of dams, one above the other, forms a long narrow meadow, which produces considerable quantities of wild hay.

This region was once heavily timbered, largely with pine, which has been prostrated by a hurricane, and since burned over several times. The soil, naturally light, has burned up and so washed away, as to expose the white-gray, pink and dark-green rocks in every direction, affording an unsurpassed opportunity to study this series; the boulders are very numerous and often of great size. The light colors of the rock, scarcity of vegetation and an abundance of standing trunks of dead trees give the landscape a peculiar aspect; but a second growth of poplar and wild cherry is rapidly changing this dismal character.

The fallen timber, swamps, steep bluffs and ledges, and numerous boulders, make travelling through the Laurentian area difficult and laborious in the highest degree. Florida swamps have denser vegetation and are much larger; sea-coast marshes often have more mud; the highlands of the Hudson present more formidable elevations, but, all in all, the writer believes it requires more physical exertion to travel 5 miles per day (all a man can accomplish with a pack) through Lake Superior granite windfall, than in any other region east of the Mississippi. The trees were prostrated by northwesterly winds, judging by the direction in which they lie; persons have travelled in a southeasterly direction on the trunks of fallen trees (mostly pine) for over a mile without once touching the ground.

III.—STRATIGRAPHY.

Scarcely second to the two classes of phenomena already mentioned is the observance of the rock masses, or strata, as to their direction or strike, and inclination or dip; the order of their superposition and thickness; but more important than either is to ascertain between what rocks the mineral sought for occurs. Useful minerals which occur in beds, like the iron ores of Lake Superior, will usually be overlaid and underlayed by rocks, having different characters and which maintain those characters for considerable distances. Next to finding the ore itself, it is desirable to find the hanging or footwall rock. Whoever identifies the upper quartzite in the Marquette region, or the upper marble in the Menominee region, has a sure key to the discovery of any ore that may exist in the vicinity.

With few exceptions, all the rocks in the region we are describing are stratified—that is, arranged in more or less regular beds or layers, which are sometimes horizontal, but usually highly inclined. This stratification or bedding is generally indicated by a difference in color of the several layers, oftentimes by a difference in the material itself, but occasionally the only difference is in the texture or size and arrangement of the minerals, making up the rock. Thus, rocks made of quartz, sand and pebbles, may vary from a fine sandstone to a coarse conglomerate. In general, a striped rock, whether the stripes be broad or narrow, plain or obscure, on fresh fracture or weathered surface, is a stratified rock. Usually rocks split easier on the bedding planes, than in any other direction; but the converse is true in the case of most clay slates and in some other rocks, which split more easily on their joints and cleavage planes, the direction of which seldom coincides with the bedding and is often at right angles with it. If a rock splits most easily along its striping, it is always safe to assume, the true bedding planes have been found. Such planes are supposed to have had their origin in the original deposition of the mud and sand, of which most rocks are made. Similar marks can be seen in excavations in sand and clay, which may be regarded as unconsolidated rocks. The cleavage and joint planes above indicated, which are always more regular in strike and dip, than the others, are supposed to have originated from pressure, subsequent to the formation of the rock.

The term plane, as used in describing bedding, must not be understood to signify a straight-line surface; on the contrary, they are usually curved planes, sometimes folding and doubling on each

other, so as to produce a very intricate structure. Not only do these plicatures take place on the small scale, as shown in hand specimens, but precisely similar folds exist in masses of rock, which may be hundreds of feet thick. The resulting curved strata take the name of troughs or basins, if the convexity is downward, the general term synclinal structure being applied to this form. Connecting the synclinal troughs and basins are anticlinal domes and saddles. The whole may be described as rolling or wave-like forms. Sometimes the power which produced the folds seemed greater than the rocks could bear, and cracks or breaks, and faults or throws, are the result, though these are not numerous in the Lake Superior region. Cracks so produced and filled with material, other than that constituting the adjacent rocks, are called dykes; or if the material be crystalline and metalliferous, veins. As iron ore in workable quantities does not occur in this form in this region, vein phenomena will not be considered here.

An examination of the four great rock systems will illustrate and prove the above remarks on stratification.

- I. Beginning, as before, with the uppermost or youngest, which is at the same time the softest and lightest rock, the *Silurian* brown and gray sandstones and limestones, so well exposed on the south shore of Lake Superior, we have a perfect illustration of the regular and horizontal bedding, without folds, faults, or dykes. An inspection of the Marquette quarry, or any of the numerous natural exposures, will convince any one that these rocks are but consolidated sandbanks.
- II. The Copper-bearing Rocks.—Some beds of this series are sandstones nearly or quite identical with the Silurian in appearance, but the great mass is made up of different varieties of copper trap, which are often amygdaloidal; interstratified are beds of a peculiar conglomerate. The stratification of these rocks, considered in large masses, is nearly as regular as the sandstones, and differs only in the fact that the layers are inclined, dipping northwest and north toward Lake Superior at a varying angle, which seems to be greatest on the south side of the range, and is there often vertical. It is least at Keweenaw Point, where it is as low as 23°.
- III. The Iron-bearing or Huronian Rocks are immediately beneath, and are exposed to the south of the copper rocks. This series are, on the average, heavier and harder, than either of the

others and folded to a far greater degree. The prevailing rock is a greenstone or diorite, in which, like the copper traps, the bedding is usually obscure; but the intercalated schists and slates which usually bear strong marks of stratification, make it usually not difficult to determine the dip of the beds at any point. This dip varies both in amount and direction, but is generally at a high angle, and is more apt to be to the north or south than in any other direction.

IV. Descending to the oldest or bottom rocks of the Lake Superior country, the granites and associated beds (*Laurentian*), we find the bedding indications still more obscure and often entirely wanting. Here there is, if possible, more irregularity in strike and dip, than in the Huronian.

IV.—BOULDERS (FLOAT ORE).

Fragments of iron ore which have been detached from the parent ledge and are found loose on the surface, or in the drift beneath, possess great interest to the explorer, and are among his most important helps and guides. The same remarks are applicable, but to a less extent, to boulders of other rocks. As a rule, in the iron region of Lake Superior, it is safe to assume, that when boulders of a particular variety of rock are abundant on the surface, a ledge of the same will be found in place very near—if not immediately under the boulders, then up hill from them, or perhaps a little to the north or east; the more angular or sharp-cornered the boulders, the nearer we would expect to find the ledge.

In the Menominee region it may almost be said, that this rule is invariable, as there seems to have been less movement of the drift material here than farther to the north.

In the Michigamme district a large amount of float ore is found some distance south of the iron range, part of the fragments being very large and containing at least 100 tons of ore. Sections 19, 29, and 30 of T. 48, R. 30, and Sections 25, 36, and 35 of T. 48, R. 31, contain many such boulders, which were probably derived from the Michigamme range. Considerable digging has been done at several of the larger boulders, which has failed to find the ore in place, and the magnetic attractions are of a character which

indicate detached boulders and not a continuous ledge. For mode of distinguishing boulders of magnetic ore, see chapter on use of the magnetic needle.

These Michigamme ore boulders are all found south of the iron range which produced them, and but few at a greater distance than two and one-half miles, most of them being much nearer. This southerly and westerly direction of the drift is, so far as I know, universal in the iron region of the Upper Peninsula, and it is fully confirmed by the direction of the drift scratches in the solid rock, which vary from north to east, averaging about northeast and southwest.

Therefore, if iron boulders be found in considerable abundance, the explorer may assume, especially if they are angular, that he has iron underneath the surface; if rounded or abraded, the ledge may be to the north or east. If the boulders be magnetic, the place of the ledge should be found, with comparative ease, by means of the needle; but if specular, it may be an expensive and difficult work. Soft hematite, from its nature, can never occur in the form of boulders, as it would weather into a reddish soil. Iron boulders are often met with in digging test-pits and shafts; in such instances, if near the ledge, I have generally found the ore in place very near; if considerably above it in the drift, the same rules would apply as to surface boulders.

Attention should be given to the character of boulders other than iron, which may be associated with it, or found where there is no iron. Occasional granite boulders occur everywhere in the Lake Superior iron region and have no economic significance. I have never seen an abundance of granite boulders, however, except over granitic rocks, and so far, these rocks have not produced workable deposits of iron.

Boulders of quartzite, diorite and slate usually accompany those of iron in the Marquette region, and marble boulders, as well as quartzite, are most significant in the Menominee region.

The above laws, regarding the occurrence of *iron boulders*, give the facts regarding their geographical distribution great importance in iron explorations. If, where there are iron boulders, we may confidently look for iron, then conversely, where there are none, we should not expect to find iron. I do not assert that every deposit of hard ore is marked by float or boulders, but, so far as the

facts have come to my knowledge, this is the case in the region under consideration.

Except in one or two instances, which have not been verified, I have heard of no iron boulders in the so-called silver-lead region, which extends north from the Marquette iron region to Lake Superior, which would lead one to believe, that merchantable hard ores will be found there. And except the L'Anse range in north part of T. 49, R. 33, this is true of the belt of country, west from the so-called silver-lead region. The region, without iron boulders, may be briefly described by saying, that it is bounded west and south by the line of the Peninsula division of the Chicago and Northwestern, and by the Marquette, Houghton, and Ontonagon railways. In other words, a person travelling by rail from Escanaba through Negaunee to L'Anse would have the region of iron boulders on the left, and the boulderless region on the right hand, or towards the lake.

Limiting their distribution still further, we may say, that iron boulders have only been found in quantity and quality, which would point toward economic importance in (1.) T. 45, R. 25, in the vicinity of the S. C. Smith mine, which is the most easterly locality in which they have been observed on the Upper Peninsula of Michigan; (2.) the Negaunee and Michigamme iron districts, extending in belts of irregular width from Negaunee west to the First lake in S. 17, T. 48, R. 31; (3.) the L'Anse iron range, in north part of T. 49, R. 33; (4.) south and southwest from Michigamme lake, embracing wholly or in part Towns. 44 to 47 north, and Ranges 39 to 32 west; (5.) the Menominee iron region, embracing wholly or in part Towns. 39 to 42 north, and from Range 28 west to the Menominee and Brulè rivers, but not west of Range 33; (6.) the Lake Gogebic and Montreal river iron belt, south of the South copper range.

Hunting for boulders is something like hunting game; when on the ground the best woodsman, the most active and observant will be the most successful, assuming, of course, that he knows at sight what he is looking for. (See chapter on Explorations.) I have found Indians good help in this kind of work, and believe that the incentive of a bonus in money for boulders or outcrops is often good policy. The best places in which to observe boulder phenomena is in the beds of rapid streams and under the roots of trees, the latter, probably, having been the most fruitful field. A

windfall is as good as five thousand dollars' worth of test-pits to the section.

With boulder phenomena may be classed the reddish or brownish earth, which comes from the disintegration of iron ore rocks of a hematitic character, and magnetic sand, which is very generally distributed, and which comes from the disintegration of magnetic ore. Such material may, for our purposes, be regarded as made up of minute boulders and the same remarks will apply, except that I should not expect to find red earth far removed from the ferruginous rock which produced it. Minute quantities of magnetic sand can be found almost everywhere in this region.

CHAPTER III.

I.ITHOLOGY.* (Mineral Composition and Classification of Rocks.)

In the preceding sketch the terms sandstone, limestone, conglomerate, trap, diorite, granite, etc., occur. It is evident that no satisfactory and useful progress can be made in geological fieldwork, which includes prospecting, until one has learned to recognize and name the more common varieties of rock. For this purpose we have to give attention to their mineral composition, that is, we must ascertain of what simple mineral or minerals the rock in question is chiefly made up and to observe, whether such minerals are angular, presenting bright facets (crystalline), or whether they are rounded like sand and gravel (fragmental). Not only must the prospector be able to recognize at sight the mineral he is seeking, but in case it is not exposed, which often happens, then those rocks, which are known to indicate its presence or absence. Experienced prospectors will not spend much time in looking for iron among granite rocks, nor in the copper traps, nor yet in the region of horizontal sandstones and limestones.

The mineral composition of rocks, by which they are identified, described and named, constitutes the science of Lithology, one of the most abstruse departments of Geology. A high authority on this subject has remarked:—"In all attempts to define and classify rocks, it should be borne in mind that they are not definite lithological species, but admixtures of two or more mineralogical species, and can only be arbitrarily defined and limited." When rocks present recognizable crystalline minerals, the task of describing and naming is comparatively easy; but when the constituent minerals are obscure, as is often the case in the rocks we are considering, the attempt to employ specific names, which shall define such vaguely compounded aggregates, will be exceedingly difficult.

^{*} The stratigraphical order of the rocks here considered will be found in the succeeding chapter.

The difficulty may be illustrated by supposing, were an attempt made, to give such name to a common brick, as will designate its composition and structure. Bricks are made in general of sand and clay, but several varieties of sand, and as many of clay, are employed in different localities, which, being mixed in various proportions and differently burned, give rise to a wide variation in composition and appearance and could not be expressed by a single word or term. In the case of rocks we have, of course, no previous knowledge of the numerous ingredients employed in their composition, by which the difficulty is greatly increased. It may seem at first sight, as if chemical analysis should form a reliable basis for rock nomenclature, but this is not the case. Van Cotta asserts, that a rock containing 72 silica, 11 alumina, 2.8 oxide of iron, 1 lime, 1.2 magnesia, 1.2 potash, 2 soda and 0.4 water, may be either a granite or a gneiss, protogine, granulite, quartz-porphyry, felsite, petrosilex, pitch-stone, trachyte-porphyry, obsidian, or pearlstone; and by giving a little range in the percentages of some of the constituents, half a dozen other rock names could be added. Here we have eleven different rocks, having precisely the same chemical composition, but widely different in physical character.

It must be borne in mind, in studying this subject, that the solid crust of the globe is almost entirely made up of ten or eleven simple *chemical elements*, which variously combined, according to the laws of chemistry, produce the few *minerals* which in turn, mechanically mixed, constitute ordinary *rocks*; hence we should expect, that the average chemical composition of a series of rocks, wherever found and of whatever character, would nearly agree.

The materials of the first formed rocks, whatever their origin, have been worked over and over by rains and waves and chemical forces, distributed over sea-bottoms, consolidated and elevated, to pass again through the same process by just such means, as are now at work in producing similar results.

The reader who may not be familiar with the physical characters and composition of the minerals—quartz, feldspar, hornblende, chlorite, talc, argillite, mica and the oxides of iron and manganese, which make up the great bulk of the rocks herein described, is advised to refer to some elementary work on geology or mineralogy.

Extensive rock formations are now generally named after the locality, where they were first thoroughly studied, or are best ex-

posed, and their minor beds and layers are often named according to their peculiar mineral composition, or with reference to their relative age, that is, order of superposition. The names Laurentian, Huronian and Silurian are geographical names of the first class. attempt will here be made to describe the lithological character of either the Copper bearing traps, conglomerates and sandstones, nor the Silurian sandstones and limestones; these will be fully treated by Prof. Pumpelly and Dr. Rominger, respectively. been and will hereafter be said of the geographical distribution and topographical and stratigraphical character of these rocks was considered necessary, to acquaint the prospector and explorer with those general principles of geology, which lie at the foundation of intelligent and successful work. Whoever would become thoroughly acquainted with these systems is referred to Parts II. and III. of this volume. A number of specimens from the Laurentian are described in Appendix A, Vol. II. (see descriptions 252 to 299); but they do not cover all the lithological families represented in that system.

In subdividing the Huronian or iron-bearing series, which we have particularly to study, the rocks have been grouped (1) lithologically, i.e., according to their mineral composition, and (2) stratigraphically, i.e., according to relative age. As this system was first described and named by the Canadian geologists, their names have been employed as far as possible in the body of this report; the identity in composition of many of our rocks with theirs, having been established by an examination of a large number of Marquette specimens by Dr. T. Sterry Hunt.

Alexis A. Julien, A.M., of the School of Mines, New York, has made careful studies, both in the field and laboratory, of a large number of specimens from the Lake Superior region, his results being in part given in Appendix A, Vol. II. As his paper was not obtained in time to modify this chapter and the geological descriptions which follow, in accordance with Mr. Julien's nomenclature and orthography, what follows may be regarded as an independent and popular presentation of this subject, which is scientifically and more fully treated in the Appendix, the practical needs of the explorer and miner being here chiefly considered.

The specimens examined by Mr. Julien are in part from the Marquette region; the L'Anse, Menominee, and Gogebic districts

are also well represented, thus embracing an area over 125 miles long and having an extreme width of 60 miles. The specimens described belong to a catalogued collection, numbering over 2,500 specimens, being probably the most complete suite of rocks from the Azoic of the Upper Peninsula yet collected. Those from the Montreal river and Gogebic district were collected by Prof. R. Pumpelly and myself, and are believed to be the first described from that region. Prof. Pumpelly took very full lithological notes in the field, but has not yet, so far as I know, made them public. Dr. H. Credner's publications are very full on the lithology of the Menominee region, he having spent two seasons in that field.

Appendix B, Vol. II., contains a list (named by Mr. Julien) of the specimens constituting the State collection, over thirty duplicate suites of which were collected and have been distributed among the incorporated colleges of Michigan and other leading institutions and cabinets, of this country and Europe.

Appendix C, Vol. II., contains a list of 76 specimens, number 1,001 to 1,076, determined by the microscope by Chas. E. Wright, under the direction of the Faculty of the School of Mines, Freiberg, Saxony. A suite of these rocks is at Freiberg and others in Michigan.

The several beds or layers of the Huronian system, as developed in the Marquette region, are numbered upwards from I. to XIX., always written in Roman numerals. These strata being particularly described as to thickness, geographical extent, etc., in following chapters, it need here only be said in general that I., II., III., IV. are composed of beds of silicious ferruginous schist, alternating with chloritic schists and diorites, the relations of which have not been fully made out; V. is a quartzite, sometimes containing marble and beds of argillite and novaculite; VI., VIII. and X. are silicious ferruginous schists; VII., IX. and XI. are dioritic rocks, varying much in character; XIII. is the bed which contains all the rich specular and magnetic ore, associated with mixed ore and magnesian schist; XIV. is a quartzite, often conglomeritic; XV. is argillite or clay slate; XVI. is uncertain, it contains some soft hematite; XVII. is anthophyllitic schist, containing iron and manganese; XVIII. is doubtful; XIX. is mica schist, containing staurolite, and alusite and garnets. This classification, it will be borne in mind, applies only to the Marquette region, the equivalency of the rocks of the Menominee and other regions not having been fully made out.

These beds appear to be metamorphosed sedimentary strata. having many folds or corrugations, thereby forming in the Marquette region an irregular trough or basin, which, commencing on the shore of Lake Superior, extends west more than forty miles. The upturned edges of these rocks are quite irregular in their trend and present numerous outcrops. While some of the beds present lithological characters so constant, that they can be identified wherever seen, others undergo great changes. Marble passes into quartzite, which in turn graduates into novaculite; diorites, almost porphyritic, are the equivalents of soft magnesian schists. In this fact is found the objection to designating beds by their lithological character, while to numbers or geographical names no such objection exists. The total thickness of the whole series in the Marquette region is least at Lake Superior, where only the lower beds exist, and greatest at Lake Michigamme, where the whole nineteen are apparently present, and may have an aggregate thickness of 5,000 feet.

Near the junction of the Huronian and Laurentian systems, in the Marquette region, are several varieties of gneissic rocks, composed in the main of crystalline feldspar, with glassy quartz and much chlorite. Intersecting these are beds of hornblendic schist, argillite and sometimes chloritic schist. These rocks are entirely beneath all of the iron beds, seem to contain no useful minerals or ores and are of uncertain age. No attempt is here made to describe or classify them.

The following description and classification has resulted from an examination of a large number of specimens of "ore and rock," collected with the view of embracing all varieties found in the iron-bearing series of the Marquette region, together with a study of the parent masses in the field, which latter is of great importance on account of the variations in composition of the same bed, to which attention has been directed.

The specific gravity of over five hundred specimens, weighing from 3,000 to 10,000 grains, was determined by a balance, which turned when loaded, by the addition of two grains. The magnetic properties were carefully examined and are given in part in the chapter on the magnetism of rocks. Most of the specimens examined were arranged into ten *lithological groups* (having no reference to age), which are designated in what follows by the first ten letters of the alphabet. When a specimen represented a very

small and unimportant layer, it was thrown out as exceptional and not important to the object of this report.

It must be constantly borne in mind, that the divisions between these ten lithological groups or families are not sharply marked; one passes into the other by insensible gradations, thus producing many intermediate varieties, which it was difficult, if not impossible, to classify or describe. The first family, A, will include all valuable iron ores, the remaining nine (B to J) will include "rocks." But as iron ore, in large masses, has all the geological characters of the associated rocks, the popular general classification of minerals into "ores" and "rocks" will be disregarded except as above mentioned. Except in a few instances, where Mr. Julien's collection was incomplete, all minute lithological descriptions have been omitted, for such, frequent reference will be made to his paper; and for the reason that he had not access to maps and sections, which gave the stratigraphical distribution of the various rocks, this part has been made quite full in that respect.

In a few instances reference is made to the full suite of Marquette rocks, numbered 6,000 to 6,222, deposited by me in the cabinet of the University of Michigan, at Ann Arbor.

A. IRON ORES.

(Occurring in formations X., XIII., XIII. and below V.)

Only such ores as are now employed in the manufacture of iron will be described under this head. They are in order of present supply, the (a) specular hematite or red specular ore, as this class is designated in the iron trade; (b) the magnetic; (c) the "mixed" or second-class ore, which may be either specular or magnetic; (d) the soft hematite, and (e) the flag ores. Another variety, the magnetic specular, might be added, which, as the name implies, is a mixture of the black and red oxides, which gives a purple streak. The local terms "hard," embracing both the magnetic and specular ores, and "soft," for the soft hematites, are convenient.

The commercial statistics, modes of mining, and composition will be considered under their proper heads,* attention being directed here chiefly to the mineralogical and physical character of each

^{*} See Chapters IX. and X., Plate XIII. of Atlas, and Appendix J, Vol. II.

ore. Under Woodcraft and Surface Explorations, Chapter VII., are given some brief practical rules for distinguishing iron ores, for the benefit of those, who know little or nothing of rocks.

All the specular, magnetic, and mixed ores, and a part of the soft hematites, are found in one formation; bed XIII. of my arrangement, which has its most easterly exposure near the Jackson mine and extends irregularly and indefinitely westward, embracing all the mines now producing rich hard ore.

It may be said of these ores in general, that they are essentially oxides of iron, with a few per cent. of silica added, and generally contain minute quantities of sulphur and phosphorus, but no titanium. Alumina in quantity not exceeding two and one-half per cent., with one-fourth as much manganese, is sometimes found, together with alkalies, which seldom aggregate over one and onehalf per cent. The soft hematites are in part hydrated sesquioxides, hence contain water and usually more silica, than the hard ores; traces of organic matter are sometimes found, and manganese is almost exclusively confined, to the soft ores. Many specimens of specular and magnetic ore have been analyzed, which gave ninetyeight per cent. of oxide of iron, the balance being nearly pure For numerous analyses of all the ores, see Chapter X., Appendix J, Vol. II., and Plate XIII. of Atlas. Weathering has no appreciable effect on the hard ores, except to crumble and cover with soil the more granular varieties. The exposed surfaces of the compact ores (by far the most prevalent variety) are of almost as high lustre as fresh fractures, and are often highly polished, showing no weathered coating like almost all other rocks. In the "mixed ores" the jasper bands are sometimes slightly elevated on the weathered surface, due to their greater hardness.

a. Red Specular Ores.—Miners divide these into slate and granular. The former resembles closely in its structure the soft greenish chloritic schists, commonly associated with it. The slabs, into which the slate ore easily splits, are not uniform in thickness like roofing-slate, but taper always in one and often in three ways, producing elongated pieces often resembling in form a short, stout, two-edged sword-blade, with surfaces as bright as polished steel, but striated and uneven. See Specimens 46, 47, 48, State Collection, Appendix B, Vol. II., and 1,050 Appendix C, Vol. II. Thin edges of such slates can be pulverized into a bright scaly powder by the finger-nail, and

occasionally the whole mass is too friable for economic handling. The magnet will generally lift one or two per cent. of the powdered ore, and occasionally one-fourth of the whole, in which case the streak is purple. These last, constituting magnetic slates, are more friable than the pure red specular slates, due in some way to the larger admixture of magnetite. See Specimen 49, State Collection, Appendix B, Vol. II.

The granular or massive specular ore shows no tendency to split in slabs, and is made up usually of minute crystalline grains, which are sometimes, however, so large that their octahedral form can be easily recognized without the aid of a lens; fine specimens of this variety occur at the Cleveland and New York Mines. Mineralogists apply the name martite to the red oxide of iron, when it has the crystalline form of the octahedron, which belongs to magnetic ore. See Specimens 2, 43, 44 and 45, State Collection, Appendix B, Vol. II. It is not improbable, that all of the granular specular ores under consideration may have once been magnetic and in some way have gained the two per cent. of oxygen necessary to change them from black to red oxides. See Dana's System of Mineralogy, 5th ed., p. 142.

The granular ore is generally firm in texture and never friable, like the granular magnetic. Some highly compacted varieties, which contain a little silica, are very hard, constituting the hardest rock to drill which the miner encounters. This variety is called the "fine-grained steely ore;" some specimens of it possess almost the highest specific gravity observed, 5.23, while the rich softer ores of the same class averaged about 4.85. See Spec. 45, State Collection, Appendix B, Vol. II.

From the examination of a considerable number of specimens of red ore, it was found that the magnet would usually lift an appreciable portion of the powder. In the case of one coarse-grained specimen of pure ore from the New York mine, one-third of the pulverized ore was removed by the magnet. Spec. 1060, App. C, Vol. II. The percentage of powder lifted by a magnet in twenty-one specimens, together with color of powder, is given in Table, App. H, Vol. II. Numerous specific-gravity determinations of this variety of ore will be found in App. B, Vol. II.

b. Magnetic Ore.—The description given above of the granular specular ore applies with equal force to this class, except that the

latter is more of granular and often friable, has the magnetic property and gives a black or purple powder instead of red. Sometimes the rich magnetites crumble easily into grains, like some Lake Champlain ores, to which the term "shot ore" is applied; again, it is very hard, as in Pit No. 8 of the Washington mine. See Specs. 39, 40, 41 and 42, State Coll., App. B, Vol. II. The compact tabular form so frequent in the magnetic ores of New Jersey and Southern New York is not common in the best ores of the Marquette region, nor are the latter ores as highly magnetic as the former, or at least good loadstones are not so common; the ore from the Magnetic mine (see Spec. 17, State Coll.) has most of this tabular character.

Typical slate ores occur with the magnetites, but they are of the character already described, that is, mixtures of the two oxides, the magnet not removing over one-fourth of the powder, while it takes all in the case of the granular variety. The specific gravity of the granular magnetic ores, as will be seen in Appendix . B, Vol. II., varied from 4.59 to 5.01, the average of many specimens being 4.81. Specs. 1,054 and 1,059 of Appendix C, Vol. II., are also varieties of this ore.

The following minerals and rocks are most commonly associated with hard ores: a soft grayish-green *chloritic schist*, which sometimes, owing to bad sorting, goes to market in sufficient quantity to perceptibly reduce the furnace yield. The magnesia it contains might tend to stiffen the slag, otherwise it can have no effect in the furnace further, than what is mentioned above. This rock is described under Group D. See Specs. 53, 54, and 55, State Coll., App. B, Vol. II.

Micaceous red oxide of iron often occurs in scales and bunches, particularly in proximity to jasper. It has been improperly called plumbago, but is in reality in no way related to it, being chemically pure oxide of iron, having the crystalline structure of mica. A soft whitish mineral, often called magnesia, and appearing not unlike flour, occurs occasionally in specular ore and frequently in "soft hematite." This substance is usually most abundant in the more jaspery varieties of specular ore; an examination by Prof. Brush determined it to be kaolinite, a hydrated silicate of alumina (clay) in minute crystalline scales. The presence of this clay in small quantity could not but help the working of the furnace, by

forming a more fusible slag, but it would of course diminish the yield of iron, if in quantity.

The needle and velvety forms of the mineral Göethite (a hydrated oxide of iron) are not uncommon at the Jackson mine, and "Grape ore" (botryoidal limonite), sometimes finely colored with yellow ochre, is found at several of the mines, but always in soft hematite. Fine specimens of crystallized quartz are rare, and no form of lime has been observed, although analyses show minute quantities. Bunches of iron pyrites are occasionally found, especially in the magnetic mines. At the Champion mine a thin layer containing this mineral occurs next the hanging wall, but it is easily separated from the ore, and is not sent to market. Hornblende, so generally present in the magnetic mines of New York, New Jersey and Sweden, is rare in the Marquette mines, of XII. and XIII.

c. Second-class Ore. -By far the most abundant, and commercially objectionable ingredient in the Marquette ores of all kinds, is the so-called jasper, a reddish ferruginous quartz, which is invariably found associated with the best ores, usually in thin seams or lamina conforming to the bedding, but sometimes in a form approaching a breccia. In the hard ores this impurity can usually be readily distinguished, but in the soft hematites it is often only found by analyses. As this rock possesses considerable scientific as well as commercial interest (the better varieties constituting the second-class ores), I will attempt to describe and illustrate it somewhat minutely. It consists of jasper, varying from bright red to dull reddish-brown, with occasional seams of white quartz, and usually pure specular or magnetic ore of high lustre. These materials are arranged in alternating lamina, varying in thickness up to one inch. These lamina are often highly contorted, zigzagging, and turning sometimes in opposite directions within a few inches. The jasper bands are in places broken up into little rectangular fragments, which are slightly thrown out of place, as it were, by tiny faults; the ore fills the break, so that the whole mass has the appearance of a breccia. There can be little doubt, but that the true breccia at the east end of the Jackson mine has this origin, and it would be interesting to consider whether this idea might not be extended to other conglomerates in the Huronian series. The contorted laminated structure, with the striking contrast of colors, is beautiful, and affords fine miniature examples of the anticlinal and

synclinal folding and faulting of large rock masses. Sometimes the lamina are very irregular and indistinct, and one or the other of the minerals greatly preponderates. When the jasper layers all thin out (as they usually do somewhere), the ore becomes first class. Some phases of this interesting rock, with descriptions, are given in Appendix K, Vol. II., Figures 19 to 29. See Specs. 36 and 37, State Coll., Appendix B, Vol. II.

The miners call this material "mixed ore;" and those varieties in which the jasper does not constitute over 20 per cent. of the whole, are sold as second-class ore, yielding about fifty per cent. in the furnace; for rail-heads and some other uses requiring a hard iron, the presence of silica in the ore is not objectionable. The quantity of "mixed ore" is greatly in excess of the pure ore, and it will some time undoubtedly have considerable commercial value. Its nature is such, as to admit of the ready mechanical separation of the pulverized ore from the jasper by jigging, a process now employed in separating ores in the Lake Champlain region. For fixing puddling furnaces, or for any branch of iron industry which may demand pulverized ore (as the Elerhausen process promised to), it is very probable that this method may advantageously be employed, and a cheap ore produced.

"Mixed ore" is seen in outcrops far oftener than the purer ores, the softer character of which has caused their erosion, whereby they had become covered with soil; but as the mixed ores are usually associated with the pure varieties, their outcrops possess great significance in prospecting. It is important in this connection not to confound the "flag ores," (e) to be described, which they sometimes closely resemble, with this variety. The quartz of the magnetic mixed ore is usually white, or lighter colored than the red mixed ore.

d. The soft hematites of the Marquette region differ entirely from the ores above described, and are closely related to the brown hematites of Eastern Pennsylvania and Connecticut. In color they are various shades of brown, red and yellow, earthy in form, and generally so slightly compacted, as to be easily mined with pick and shovel. They are invariably associated with, or rather occur in, a limonitic silicious schist, from which they seem to have been derived by decomposition and disintegration. These ores occur in two distinct formations, X. and XII., and probably in others, in irregular bunches or pockets, surrounded by the schist and passing by gra-

dations, often abrupt, into it. Scattered through the ore, and conforming in their positions with the original bedding of the rock, are fragments of the schist. When the ore shows stratification, which it often does not, it also conforms with the bedding of the schist. The specific gravity of the soft hematite ore varied from 3.50 to 3.81, the average of five specimens being 3.59, and specimens of the schist varied from 2.80 to 3.38. Strictly this schist should be described under the next group of rocks, B, to which it belongs, but its assumed parentage of the hematite ore, here considered, has led to the digression. See Specs. of soft hematite 1,067, 1,077, 1,079, and of schist 1,040, 1,065, and 1,069, Appendix C, Vol. II.; also, Specs. 25 and 26, State Coll., App. B., Vol. II.

The following analyses of the schist and ore, from the Foster mine, by Dr. C. F. Chandler, will help to make their relations better understood:—

	Schist.	Ore.
Sesquioxide of iron	44.33	79.49
Alumina	2.14	1.19
Oxide of Manganese	.16	.25
Lime	.36	.27
Magnesia	13	33
Silica	47.10	9.28
Phosphoric Acid	0.13	0.19
Sulphuric Acid	0.17	0.17
Water	5.19	8.74
	99.71	99.91
[Iron	31.03	55.64
Equivalent to Sulphur	.068	.068
Phosphorus.	.057	.083

It will be observed that the essential difference is in the amount of silica, of which the schist has over 47 per cent., while the ore has less than 10 per cent., and again the ore has 25 per cent. more metallic iron than the rock. The one would evidently be converted into the other, both as to its chemical and physical characters, by the abstraction of the greater part of its silica. It is not at all improbable, that this change may have been brought about by the alkaline waters of former thermal springs, such as are now producing similar results in other parts of the world. There seems to be very little sand or clay in this ore, and washing has not appeared to

improve its quality, as is the case with the eastern ores which it resembles. If the fragments of silicious rock, which are scattered through it, are carefully picked out by the miner, an ore uniform in character is obtained. Except the ever-present silica, there are only two minerals, which it is necessary to mention as being generally associated with this variety of ore. 1st. The white clay (kaolinite), above described, which is far more abundant in this ore than the hard ores; bunches as large as a hen's egg being sometimes seen. There can be no doubt but that the kindly working of the furnace usually obtained by using the best quality of this ore, is due in part to this clay as well as to the porous character of the ore. (Calcining the ore would expel the water, of which it contains from 2 to 9 per cent., and should also cause it to reduce more easily in the furnace.) The second and most important mineral to be mentioned is the oxide of manganese, usually if not always in the form of Pyrolusite; minute quantities of this metal, always less than one per cent., are sometimes found in the hard ores, but from I to 4 per cent. is constantly present in several of the hematite deposits, which is so important an element in their value, as to almost warrant the subdivision of the soft hematites into two classes, the manganiferous and non-manganiferous.

The recently developed hematite mines near Negaunee, belonging to formation X., contain most manganese; others contain little or none. Scarcely enough of the ore has been worked to determine its place in the market; but there can be no doubt, that when equally rich in metallic iron, the manganese would give this ore the advantage, as a mixture for the furnace, over the non-manganiferous varieties. See Spec. 25, State Coll., App. B, Vol. II.

The hematite ores now in the market, as a class, vary greatly in richness, from an average of not exceeding 40 per cent. of metallic iron for some deposits, to at least 55 per cent. in the case of others. This difference is in part brought out in Chapter X.

Passing from the Marquette region to the undeveloped districts, we find on the L'Anse range, at the Taylor mine, a large deposit of hematite of excellent quality. At the Breen mine, on the south belt of the Menominee region, is also a good "show" of hematite. Promising indications of this ore were also found between Lake Gogebic and Montreal river; all of these localities and their ores will be described hereafter.

e. The last variety of merchantable ore, to be described in this report and designated Flag, has been in use so short a time, that but little can be said of its metallurgical character. It corresponds more nearly with the second-class ores (c), than with either variety described, differing from it more in structure than in composition. The ores embraced under this head are abundant and have received various local names, which will be found significant and convenient, as lean ores, iron slates, magnetic slates and silicious ores. They have also been called "lower ores," in reference to their subordinate geological position, being older than the rich ores of formation XIII., already described. Flag ores are in reality only varieties of the ferruginous schists, constituting Group B, next to be described, which are sufficiently rich in iron, to possess market value. The percentage of metallic iron in these ores and the associated schists varies from say 5 to nearly 60, those above 50 now constituting a merchantable ore. The remaining material is generally silica. always silicious, but sometimes contains more or less chlorite, manganese, argillite, mica, garnet, or hornblende added. This ore is always flaggy in structure, the layers being occasionally thin enough, to warrant the application of the term slate. All forms of the oxide of iron can be observed, a mixture of the black and red prevailing. The hydrated oxide, producing limonitic silicious schist, has been described above, as the rock from which the soft hematite ore seems to have been derived, and an analysis is there given, to which nothing need be added here.

Stratigraphically these rocks are older than the ores described under a and b, and constitute at least four beds, X., VIII., VI., and below V., separated by diorites, chloritic schists, quartzites and argillites. Like the mixed ores (c) they are banded, but the marking is seldom bright and often obscure, produced by the interlamination of a dull reddish or whitish quartz, with dull *silicious* instead of *pure* ore. There are exceptions to this rule, but they are not numerous in this region. As this is a point of much importance to iron prospectors, it may be asserted, that when white or red quartz (jasper) is found banded with an ore which can be scratched with the knife, it is in all probability the "mixed ore," which accompanies the pure ores of bed XIII.; but if the quartz be dull and not sharply defined in its layers, and particularly if the knife marks the ore layers like a pencil, instead of cutting them, then we probably have

one of the flag-ore formations. It is difficult to say, whether the red or black oxides prevail in many flag ores; hence whether particular varieties should be described as hematitic or magnetic.

All ores and ferruginous rocks become more magnetic as they are followed west in the Marquette region, the maximum amount of magnetite occurring in the Michigamme district. The ferruginous schists of the Republic Mountain series are among the most highly magnetic rocks in the whole region. At the Ogden mine, Section 13, T. 47, R. 27, the abrupt transition of the hematitic into the magnetic variety can be plainly observed, by following the *strike* of the beds less than 200 feet. This transition probably often occurs in the same bed, and, of course, might occur still oftener in crossing the formations, that is, in passing from one bed to another.

Several varieties of *flag ore* will now be described, showing a wide range in lithological character, which we should not be warranted in grouping together in a strictly scientific classification; but our arrangement of rocks, as has been stated, is rather economic and for the use of practical men.

(1) A showy, granular, chloritic, specular ore was found in a small pocket-like mass at the north ¼ post of Sec. 26, T. 47, R. 26, at locality known as the Gillmore mine. A specimen having a specific gravity of 4.28 gave Dr. C. F. Chandler metallic iron 60.46, alumina 3.49, lime 0.60, magnesia 1.33, silica 7.05, sulphur 0.30, phosphoric acid 0.08, water and alkalies not determined 0.77.

A similar ore, but containing some magnetite and peculiar white glistening spots, which appear to be mica scales, is found at the Chippewa location, Sec. 22, T. 47, R. 30. A specimen of this gave Prof. A. B. Prescott metallic iron 53.17, and insoluble silicious matter 20.20. Neither of these varieties are flaggy. See Specs. 6,156 and 6,206, University of Michigan cabinet.

- (2) A specular slate ore, holding reddish specks on freshly fractured surfaces, is found at the Cascade location, bedded with layers of jasper, having the local significant name of "Bird's-eye Slate." A specimen of this gave J. B. Britton metallic iron 59.65, insoluble silicious matter 12.24, alumina 0.88, lime 0.14, magnesia 0.08, oxide of manganese 0.02, water 1.08, with traces of sulphur and phosphorus. See Spec. 6,190, University of Michigan cabinet, and Spec. 6, State Coll., App. B, Vol. II.
 - (3) South of the Cascade range is a flag ore, beautifully banded with

red jasper and silicious iron ore, closely resembling some of the mixed ores of Bed XIII. above described, and interesting on this account.

- (4) Northeast of the Cascade location, and near the centre of Sec. 29, T. 47, R. 26, is a granular slate ore showing on fresh fracture a peculiar fine reticulated appearance and indistinct octahedral forms. A specimen of this gave Mr. Britton 59.42 per cent. of metallic iron. See Spec. 6,191, University of Mich. cabinet. Since the foregoing was written, shipments of flag ore have been made from the Cascade mines (see Plate XII. of Atlas), and with it a considerable amount of a good quality of specular ore.
- (5) At the Tilden mine, while the prevailing ore is a 40 per cent. ordinary red flag ore, there are seams or layers of bright steely ore, very hard and heavy, which yield, according to analyses made by Dr. Draper, 62 per cent. metallic iron. This ore possesses particular interest from its close resemblance to the Pilot Knob ore, Mo.
- (6) While the most abundant ore at the Iron Mountain mine, Sec. 14, T. 47, R. 27, is much like the Tilden and Ogden ores already mentioned, there is a peculiar variety, containing manganese, which is also found on the hills south of Negaunee and on the lands of the Deer Lake Company, north of the New York mine. This ore is a very dark-colored silicious hematitic schist, containing on the average several per cent. of manganese, single specimens of which have proved to be nearly pure oxide of manganese. Some of this ore from Iron Mountain was tested in the furnace as a mixture, but was found to be silicious. The need of ferro-manganese in steelmaking would make ores of this character a legitimate object of exploration. An experienced iron-master recently expressed the opinion that a 30 per cent. iron ore, with 12 to 20 per cent. of manganese, would soon have commercial value. It is possible that such a variety may exist in some of the beds under consideration. The soft or hematitic variety of this ore has already been mentioned.
- (7) Passing from the Negaunee to the Michigamme district, we find two flag ores worth noticing. On the Magnetic Company's property, Sec. 20, T. 47, R. 30, is a large amount of a very compact, hard, heavy, highly magnetic ore, laminated with a greenish horn-blendic mineral, producing an unusual banded structure. A piece of one of the layers of ore gave Mr. Britton 56.78 metallic

iron, 19.44 insoluble silicious matter, less than one per cent. of alumina, lime and magnesia, and a trace of phosphorus. See Spec. 18, State Coll., App. B, Vol. II.; also Chapter X. Recent explorations have developed a workable deposit of this ore.

(8) Adjoining this property, to the southeast is Sec. 28, owned by the Cannon Iron Co., on the north side of which is a thin layer of micaceous specular ore, closely resembling that described above under A, but containing more silica. A specimen of this afforded Professor Prescott 55.12 metallic iron, 19.80 insoluble silicious matter, with traces of sulphur and phosphorus. This and the banded ore associated with it, has a closer resemblance to the slate and "mixed ore" of some of the old mines, than any place I have seen in the flag-ore series, to which it seems to me geologically to belong; its relation to the associated mica schist is interesting. See Group H below. The Chippewa ore, near the Cannon, has already been mentioned above in connection with the Gillmore.

The foregoing brief descriptions of several varieties of flag ore embrace all those, which have come under my notice in the Marquette region and give promise of having early commercial value.

As will be elsewhere (Chapter V.) more fully described, the hard ores found in the Menominee region up to October, 1872, are more nearly allied to flag ores than to either of the first-class ores of the Marquette region. Flag ores of a low grade have also been found in the L'Anse and Gogebic districts, as will be mentioned hereafter.

A very limited experience in working these ores, together with the little I have been able to learn from others, leads me to believe, that they require more limestone and coal and produce a harder metal, having comparatively little strength, but which is probably well adapted to making rail-heads. I think a large mixture of manganiferous hematite might help the working of a furnace consuming flag ore. Precisely the same remarks may be made of the second-class ores (c); indeed, these two classes are to all intents and purposes identical in their metallurgical character, and are only separated here because of their different geological occurrence. The second-class ores are, it will be remembered, simply inferior grades of the rich hard ores of XIII.

The flag ores have here received relatively far more attention, than their present commercial importance warrants, for the following reasons:—Ist, Their quantity, so far as can now be judged, is

greater by tenfold than the first-class hard ores, and for this reason they must, at some future time, constitute a large part of the total production of the region. 2d. Very serious disappointments and losses have occurred in the past, and are likely to be repeated in the future, from mistaking flag ore for first-class ore. This arises from the fact, that the better varieties of flag ore closely resemble the poorer varieties of the rich ore. So close is this resemblance, that the best judges of ore in the Marquette region have erred. It is doubtful, if the matter can be settled definitely, except by thorough explorations, aided by the well-known laws of the geological occurrence of the two ores, which will be more fully brought out in succeeding chapters.

It is not asserted that first-class hard ores may not be found associated with the flag ores, hence below and older than formation XIII.; but it is a fact, that over one million dollars have been sunk in such search, and excepting the West End mine of the Cascade range (if that is an exception), no workable deposit of strictly high grade hard ore has been found in the flag-ore series.

B. Ferruginous, Silicious, and Jaspery Schists.

(Occurring in formations XII., X., VIII., VI., and below V.)

The best general idea of the character of the rocks embraced here can be conveyed by saying, that they are identical with the flag ores last described, except in containing less iron and usually more silicious matter. On geological grounds, as has been remarked, the flag ores should be embraced under this head and described as a subclass, rich in iron. It remains therefore for me to mention briefly, a few of the remaining varieties of this series, which are so poor in iron as to render it highly improbable that they will ever possess value as ores: I design to embrace in this group Mr. Julien's quartz schist, silicious schist, and jasper schist, Appendix A, Vol. II. For minute lithological descriptions of numerous varieties see Specs. 154 to 173, App. A, Vol. II.

At Republic Mountain are three highly magnetic beds of silicious, chloritic and hornblendic schists, numbers VI., VIII., and X. See Map No. VI. of Atlas. The peculiar striping—whitish, greenish, brownish, and yellowish—exhibited in the large outcrops suggested the name "rag-carpet schist." A specimen made up of numerous

chippings of this rock gave 31 per cent. of metallic iron; this is believed to be above the average. Both the red and black oxides are present, and some of the layers hold an ore, which, if it could be separated, might yield 50 per cent.

South of the Washington mine these rocks contain the minimum amount of iron, a specimen of which gave Charles E. Wright less than 5 per cent. Garnets and anthophyllite, or mica, seem to replace the iron, producing a grayish and brownish schist, the mineralogical character of which is obscure. See Group I. The old Michigan mine ore, Section 18, T. 47, R. 28, seems to be a variety of this peculiar schist, but much more highly charged with metal, specimens of which, I should judge, would afford 30 to 40 per cent. of metallic iron.

Passing to the Negaunee district we find in the railroad cut at the northwest end of Lake Fairbanks a chloritic, magnetic, silicious schist of a brownish gray color, faintly banded and very hard; it is aphanitic in character, and shows no disposition to split on the planes of bedding. In the railroad cut near the centre of Section 8, one mile and a half southeast of Negaunee, is a soft variety of ferruginous rock, affording some good red chalk. The rock seems to be chloritic, layers of which are impregnated with red oxide of iron. A similar material was found in numerous test pits in the east part of Section 18, T. 47, R. 26. Recent explorations in this vicinity prove this rock to be associated with the Negaunee hematites, which are fully described in Chapter IV.*

One of the best characterized and abundant varieties of this group is the banded ferruginous jaspery schist, which constitutes in the Michigamme district the whole of formation XII., and is also abundant in parts of ore formation XIII. Such varieties of "mixed ore," as contain too little iron to give them commercial value (unfortunately the greater part), would be classed here. The full descriptions and illustrations already given of "mixed ore" under A, will make any further description unnecessary, for this is a similar rock with little or no iron. See Spec. 32, State Coll., App. B, Vol. II., and for several other varieties of this group see Specs. 1,026, 1,034, 1,061, and 1,064, Appendix C, Vol. II. The Felch mountain series contain a large amount of a similar rock.

^{*} It is questionable whether this rock should be classed under D or G.

C. DIORITES, DIORITIC SCHISTS AND RELATED ROCKS (Greenstones,)*

(constituting formations XI., IX., VII., and one or more beds below them.)

These obscurely bedded rocks, locally designated greenstones and sometimes traps, are co-extensive with the ferruginous rocks A and B, very abundant, outcropping throughout the Huronian region, and present much variety in appearance. They range in structure from very fine-grained or compact (almost aphanite) to coarsely granular and crystalline, being sometimes porphyritic in character. The color of the fresh fracture is from dull-light to dark or blackish green, the weathered surface being usually lighter and of a grayish green or brownish color, not unfrequently spotted or mottled, showing a dark-green, or black, lamellar mineral (hornblende), set in a whitish, and sometimes reddish, softer mineral (feldspar). rock is exceedingly tough, powdering under blows of the hammer rather than break. It can be scratched by the knife, giving a light gravish-green powder, and is fused without difficulty before the blow-pipe. On the one hand, it graduates into a heavier, tougher, blacker variety, which is unquestionably hornblende rock, with some feldspar, well shown at the Greenwood Furnace quarry, on Sec. 15, T. 47, R. 28. See Specs. 1,018 and 1,020, App. C. Vol. II. On the other hand, it passes into a softer, lighter colored rock of lower specific gravity, which, while it has the same streak, weathers similar to the true diorite, is eminently schistose in character, splitting easily, and appearing more like chloritic schist than any other rock. The Pioneer Furnace quarry at Negaunee contains this schist and several transition varieties, some of which approach the granular massive rock. See Specs. 1,001, 1,005, 1,006, and 1,015, App. C, Vol. II. On the north side of Lake Michigamme, and west, varieties occur having a true slaty structure in appearance, although not splitting easily. See Spec. 1,028, App. C, Vol. II.

At several points dioritic schists, semi-amygdaloidal in character, were observed, and in one instance the rock had a strong resemblance to a conglomerate. See Spec. 1,024, App. C, Vol. II.; and

^{*} See Dr. Houghton's Notes on Diorites, Appendix E, Vol. II.

Spec. 71, State Coll., App. B, Vol. II. It is of much practical importance to distinguish between the schist of this group and the true chloritic schist to be described under the next head, D, which is usually found associated with the pure ores of Bed XIII.*

At Republic mountain a dioritic schist graduates into black mica schist, and large garnets are there found in typical diorite. Iron pyrites are usually seen sprinkled through the rock, and epidote is sometimes observed. Dr. Hunt found chromium in two specimens. South of the Old Washington mine, in Bed XI., occurs a variety, which in places may almost be described as hornblendic schist; that in other parts of the same bed, near at hand, graduates into the above-described dioritic schist.

In the railroad cut at the foot of Moss Mt., west of Negaunee, is an exposure of soft dioritic schist, in which are imbedded rounded lumps of diorite, which, when broken, show a crystalline reddish feldspar. See Specs. 1,001 and 1,002, App. C, Vol. II. Spec. 77, App. B, Vol. II., is another beautiful and rare variety, in which the feldspar is red. On the south side of Sec. 9, T. 49, R. 33, is a heavy bed of coarse-grained friable diorite, which has in places disintegrated into sand. Mr. Julien regards this and the associated dioritic rocks of the L'Anse range as possessing such distinctive characteristic as to warrant him in describing them as a distinct variety. See Specs. 342 to 353, App. A, Vol. II. He also classes the well-known peculiar serpentine rock of Presque isle with the diorites. See Spec. 321, App. A, Vol. II, also App. E.

The magnet usually lifts less than one per cent. of a powdered diorite, but in one case it took nearly all, and the specimen attracted the needle. This piece was from the ridge south of the New England mine; it had the essential character of a compact, perhaps hornblendic diorite, but its magnetic property and very high specific gravity, 3.29, prove that it is exceptionally rich in iron. It will be shown below, that in addition to the magnetite, seventeen per cent. of metallic iron exists in some diorites in the form of combined protoxide, which does not attract the needle. The specific gravity of the typical rock varied from 2.84 to 2.96, the average of six specimens being 2.91. The hornblendic varieties ranged as high as 3.01, while the schistose variety fell as low as 2.70,

^{*} See Julien's remarks under Chloritic schist, App. A; Vol. II.

averaging 2.82. A garnetiferous specimen, from Smith Mountain, gave 3.02, while a peculiar variety from north of Greenwood Furnace, which appeared to be feldspathic in character, gave but 2.71. Numerous additional specific gravity determinations are given in App. B, Vol. II. The precise character of the constituent minerals of this rock is obscure. Mr. Julien has minutely described numerous varieties in App. A, Vol. II., Specs. 302 to 353.

The following analysis of a specimen from bed XI. is from Foster & Whitney's Report, Part 2d, p. 92. The specimen was from Sect. 10, T. 47, R. 27, on south side of the Cleveland and Lake Superior ore deposits:—

	OXYGEN.
Silica 4	6.3124.06
Alumina I	1.14 5.21
Protoxide of iron 2	21.69 4.82
Lime	
Soda	6.91 1.78
Water	4.44
Magnesia t	trace.
-	
10	00.17

From this it is deduced that the rock is a mixture of labradorite feldspar with hornblende or pyroxene. Regarding the presence of water, numerous analysis of similar rocks in Canada show the same result. See Geology of Canada, pages 469, 604, 605, and 612. Dr. Hunt expresses the opinion, that in the case of the Marquette diorites, the hornblendic mineral often becomes softened and hydrated, passing into a degenerate form more nearly allied to chlorite or delessite (in which water is an essential constituent), than to a true hornblende. This chloritic mineral is sometimes seen scattered through the body of the rock, and very often near the weathered surface.

The absence of *magnesia*, which is regarded as an essential ingredient of chlorite and delessite, and as very rarely absent from hornblende, as shown by the above analysis, deserves notice. Dr. Hunt remarks that the hornblendic element may very likely be the iron hornblende described by Dana, System of Mineralogy, 5th ed. p. 234, under the name grünerite. The unusually large amount of

iron shown by Whitney's analysis and the high specific gravity observed would favor this view. The conversion of this non-magnesian diorite into a magnesian schist (chloritic or delessitic) would require the introduction of the magnesian element under some law of pseudomorphism, the possibility of which is proven by chemical geology.

Magnesia is not, however, absent from all varieties of the diorite. A chromiferous specimen from near the centre of Sec. 36, T. 48, R. 28, was found by Dr. Hunt to be rich in magnesia, containing more of this element than of lime; the specimen was not a typical one, but showed a tendency to pass into a steatitic rock, which might be expected to contain magnesia. Until, however, the presence of magnesia in the schists and its absence from the diorites is proven by more analyses, it is not worth while to conjecture in the matter, and I here digress only to record a few facts, bearing on an interesting and unsettled question in chemical geology. In the absence of any additional light, we adopt the hypothesis that the Marquette "greenstones" are diorites, composed essentially of a non-magnesian iron hornblende and some feldspar other than orthoclase.

It is of great importance that the prospector should have a good practical acquaintance with this rock, for it is everywhere associated with iron ores in the Upper Peninsula. He should be able to recognize it at sight, to distinguish its varieties, and especially he must not confound the Huronian diorite with a similar rock, found in the Laurentian, nor with Copper trap. More than one piece of land has been bought for iron on the Laurentian area, because "greenstone" was found on it.

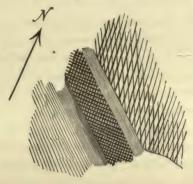
The bedding of these rocks is generally obscure, and in the granular varieties entirely wanting. It is usually only after a full study of the rock in mass, and after its relations with the under and overlaying beds are fully made out, that one becomes convinced, whatever its origin, it presents in mass precisely the same phenomenon as regards stratification, as do the accompanying schists and quartzites.

I have nowhere seen the granular diorites show more unmistakable evidence of bedding than on the small knob southwest of Bear Lake, Republic Mountain, shown in Fig. 1, scale $\frac{1}{50}$ th. The cross shading represents massive diorite, and the parallel shading a slaty silicious iron ore.

No reference is here made to the false stratification or joints, which are numerous and interesting, but which, unfortunately, for want of space, can receive no other attention here, than to warn the observer against mistaking *joint* planes for *bedding* planes, which is sometimes done, even by experienced observers.

This description, as has been stated, is intended to apply to the diorites of the iron-bearing or Huronian series, and more especially

Fig. 1.
Stratification of Diorite.



to the Marquette region; but a similar rock, as has been observed, occurs abundantly in dykes or veins, and probably in beds in the Laurentian rocks. A fine example of such a dyke can be seen penetrating a granitic gneiss, near the northeast corner of Sec. 7, T. 46, R. 29. At other points in the Laurentian area immense masses of a dioritic rock were observed, the stratigraphical relations of which to the gneiss and granites was not made out. The average specific gravity of the dyke diorite was 3.03. Mr. Julien describes some specimens of diorite from the Laurentian in App. A, Vol. II.

The following designated specimens, in addition to those already referred to, constitute a tolerably full collection of the more important varieties:—Granular diorites, 1,007, 1,008, 1,009, 1,010, 1,011, 1,012, 1,014, and 1,016; Dioritic schists, 1,001, 1,019, and 1,023 of App. C, Vol. II. The State Collection, App. B, Vol. II., also contains a large number of specimens of diorite of several varieties.

The distribution of this rock in the Huronian of the Upper Peninsula is interesting. It is far more abundant in the Marquette region and contiguous to the ore deposits, than elsewhere. The related rocks in the L'Anse region are abundant; but in the West iron dis-

trict, and on its prolongation into Wisconsin, where it forms the Penokie range, diorites are rare. In the Menominee region they seem to be replaced to a great extent by chloritic schists and horn-blendic schists, as described in Chapter V. Whether future explorations will prove that the best ores are always associated with the typical diorite, remains to be seen.

D. MAGNESIAN SCHISTS (mostly chloritic).

(See Mr. Julien's description, Specs. 179 to 188, App. A, Vol. II.)

Intercalated with the pure hard and mixed ores, at all the mines worked in formation XIII., are layers of a soft schistose rock, of some shade of gravish green, and often talcy in feeling. The Cleveland, Lake Superior and Champion mines are good localities for an examination of this rock. It is unquestionably a magnesian schist, varying from chloritic to talcose in character, and sometimes apparently containing a large percentage of argillite. In places, as at the Old Washington, its character is unmistakably talcose. Specimens obtained there held 4.2 per cent. of water, and had a specific gravity of 2.81, with light gravish-green color, and other characteristics of talcose schist. See Specs. 1,046, App. C, Vol. II. The corresponding schist at the Champion mine is also decidedly talcy. On the same magnetic range, but further west, at the Spurr Mountain, the equivalent schist is unmistakably chloritic. See Specs. 179 to 181, App. A, Vol. II. A rare variety of talc schist is represented by Spec. 74, App. B, Vol. II., obtained at the Grace furnace, Marquette.

In the Lake Superior and Barnum mines this rock is, in places, of a light green color, less soapy in feel, has a higher specific gravity and is of uncertain composition. See Spec. 55, State Coll., App. B, Vol. II. At this locality it has a marked cleavage structure, the planes of which trend east and west, and are nearly vertical, being distinct from its bedding, which latter is very obscure. Its structure bears a striking resemblance to that of the specular slate ores, noticed under A, even to the presence in both of minute octahedral crystals. Prof. Pumpelly has suggested, that one may be a pseudomorph after the other. In this connection it may be remarked, that no gradual transition of one into the other was observed, the division planes being in each instance sharply defined.

Specimen No, 1,043, App. C, Vol. II., from the Washington mine, is grayish, less schistose in structure than the last described variety, and gave up, when pulverized, one-third its bulk to the magnet. A similar massive variety from the same mine, which contained three per cent. of water, held black hard scales, which Prof. Brush decided had the character of ottrelite.

A reddish gray variety of this rock (see Spec. 6,164, University of Mich. Cabinet), holding grains of vitreous quartz, is from a heavy bed on the northeast side of the S. C. Smith soft hematite ore deposit, on Sections 17, 18, and 20, T. 45, R. 25.

South of the Edwards mine, at the Republic Mountain, and at other places in the ferruginous schists, occur bunches and thin irregular beds of a pure chlorite, often micaceous, which always contain garnets. See Spec. 6,097, University of Mich. Cabinet. This specimen shows, under the lens, minute elongated crystalline faces, closely resembling those seen in the diorite. Spec. 184, App. A, Vol. II., is garnetiferous. The "keal" or red chalk, found at several mines, is a variety of this schist impregnated with oxide of iron. See Spec. 6,183, University of Mich. Cabinet.

A very peculiar occurrence of this rock are the so-called "slate-dykes," which can be seen at the New England, Lake Superior and Jackson mines, but still better in the quartzite ridge, just north of the outlet of Teal lake. These dykes are often several feet in width, cut across the stratification, and are filled with a magnesian schist. If space permits, this subject will be more fully considered elsewhere. See Specs. 1,053, 1,068, App. C, Vol. II.

The Lower Quartzite bed V. often contains talc in bunches, small beds and disseminated, producing in places a talcy rock. The *novaculite* of that formation is due to the presence of talc and argillite. These rocks will, on account of their association, be more fully described in the Quartzite group.

It would be difficult for a skilled lithologist, and impossible for me, to draw the line between the chloritic schists here considered and the dioritic schists mentioned under Group C. So far I have chiefly noted occurrences of the magnesian schists, in formations XIII. and V., where they are not associated with true diorites. But at the Marquette quarries we find what may be called typical chloritic schists, bedded with granular diorites. See Specs. 182 and 183, App. A, Vol. II. At this locality the planes separating

the two kinds of rock are well defined; at others, which have been designated, the transition is gradual.

Along the north border of the Laurentian area, which lies south of Lake Gogebic (see Map I.), are numerous exposures of a chloritic schist (see Specs. 187 and 188, App. A, Vol. II.), which in places becomes massive and granular, a form designated "greenstone" by the United States Linear Surveyors, and so marked on their maps. See Specs. of Diorite, 309 and 212, App. A, Vol. II.

The specimens of Laurentian Gneiss, 275 and 299, App. A, Vol. II., contain chlorite as an essential ingredient, proving this mineral to be as widely disseminated in the Laurentian as Huronian. An examination of Prof. Pumpelly's very exhaustive chapters on the lithology of the copper-bearing rocks, will show chlorite to be of frequent occurrence in that system; demonstrating it to be next to feldspar and quartz, one of the most universally diffused minerals in the Azoic of the Upper Peninsula.

E. QUARTZITE—Conglomerates, Breccias, and Sandstones.

(Principal development in Formations V. and XIV. See Mr. Julien's descriptions, 126 to 140, and also 358 and 359, App. A, Vol. II.)

After diorite and the ferruginous schists, no rock is more abundant in the Marquette region, and none more frequently found in outcrops, than the different varieties of this group. Two extensive beds exist—XIV. lies immediately over the ore formation, and V. near the base of the series. The last appears to be the most persistent and wide-spread member of the Huronian system. It can be traced from the shore of Lake Superior, near Chocolate river, westward for 40 miles, and possesses unusually economic interest from its affording the marble, used to a limited extent as furnace flux, and the whetstone rock (novaculite), which was at one time quarried for market. This quartzite has also recently been successfully employed as lining for Bessemer converters.

The Upper Quartzite (XIV.) is co-extensive with the ore formation XIII.; it is seen as the hanging wall of the most easterly point, at which rich hard ore is mined, and overlays the most westerly deposit yet explored. Between these is a third bed, seen in the railroad cut

near the west end of Lake Fairbanks, the extent of which has not been made out. See Spec. 21, App. B, Vol. II.

At the west end of Lake Michigamme, near the centre of Sec. 25, T. 48, R. 31, is a large mass of quartzite, which appears to be a ledge, but if so, the bed is concealed to a greater extent than usual, for it has not been observed elsewhere. No. XVIII. is assigned for this quartzite, or for whatever rock may be found in the gap between Beds XVII. and XIX. The Cascade iron range is divided by a thin bed of quartzose rock, which varies from a quartzite to the coarsest conglomerate I have observed in the region, but which, like the two last-mentioned beds, seems to be local. At the Greenwood furnace is a heavy and persistent bed of quartzite, in which are intercalated layers of clay slate; its age has not been determined; it resembles the lower quartzite.

The extreme hardness of quartzite (the knife makes no impression on it, and it will readily scratch glass), and its general dissimilarity to the other members of the series, renders its recognition easy and much description unnecessary.

Vein quartz, occurring in bunches, seams and veins, in nearly all rocks, is not embraced in this description; nor are those slightly ferruginous quartz schists, already described in Group B, which a strictly scientific classification would place under this head. Ouartzite is seldom white, often light-gray, or dark-gray and sometimes reddish or greenish. The effect of weathering does not penetrate the rock beyond a mere film, dulling the lustre and color of a fresh fracture rather, than changing it; but the latter effect is sometimes produced in the impure varieties. Broken pieces often show grains of glassy quartz; and the arenaceous character is sometimes so plain, as to leave no doubt in the mind, that the rock is a metamorphosed sandstone or conglomerate (see Fig. 2). Again, the whole mass is compact, having much the appearance of vein-quartz. In structure it is usually massive, and the bedding obscure; but in places, as at the northeast corner of Teal lake, it is banded, presenting a flaggy structure, like the ferruginous schists. The mean specific gravity of a large number of specimens was 2.69. See App. B, Vol. II.

The foregoing description applies in general to all the beds; but as it is often of importance to the explorer to distinguish the Upper bed on account of its relation to the ore formation, a few points of difference will be noted. As has been remarked, the Lower bed is often calcareous, turning in places into a true marble, as at the Morgan Furnace; and the same formation is often talcy in character, containing in certain localities bunches and beds of a talcy material and in other places beds of argillite. An intimate mixture of these minerals with the quartzose material produces novaculite, which was formerly quarried just east of Teal Lake outlet. See Spec. 13, State Coll., App. B., Vol. II. Red oxide of iron in grains and small bunches, is not infrequent in the Lower bed, as can be seen in northeast quarter of Sec. 22, T. 47, R. 26.

So far I have seen neither marble, talc, nor novaculite in the Upper Quartzite, and only once, at the Lake Superior Mine, have I seen argillite associated with it. As this exception has much interest, it will be fully considered in another place. The Lower Quartzite is seldom conglomeritic, the upper one often so, and in places on the Spurr Mountain range it is a true conglomerate, containing pebbles of white and glassy quartz and jasper. See Specs. 115 to 118, App. A, Vol. II. At Republic Mountain large fragments of ferruginous schist are seen in the base of the Upper bed. Southwest of the Old Washington mine it is a coarse conglomeritic rock, which is in places schistose or slaty. See Spec. 122, App. A, Vol. II.

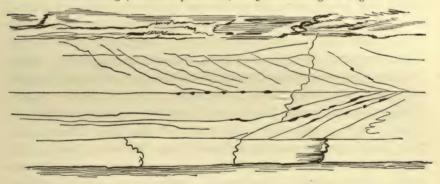
The matrix of this variety (See also Spec. 6,085, University of Mich. Cabinet) is a soft, micaceous, slaty material, containing fine grains of specular ore and holding pebbles of white quartz. The Upper bed overlying the east end of the Jackson, and that over the New York mine, also hold pebbles. Mica scales and epidote were found in the same bed at the Republic Mountain, and in places it had almost the appearance of fine-grained granite.

As if to leave in our minds no shadow of doubt, as to the sedimentary origin of this rock, nature has, in addition to the conglomerate on the Spurr Mountain range, given us a variety of the Upper Quartzite, which can only be described as a fine-grained, friable, banded sandstone. See Specs. 358 and 359, App. A, Vol. II. The alternations of magnetic sand with quartz sand, producing the stripes, is very interesting in connection with the origin of these ores. It is doubtful if any true breccias (conglomerates with angular pebbles) occur associated with the rocks here described, if at all in the region. The brecciated rocks, a variety of "mixed"

ore" found in formation XIII., is believed to have had the origin ascribed under Group A.

Specimens of University of Mich. Cabinet, Nos. 6,193, 6,084, 6,180, 6,211, 6,219, and 6,122 are from these quartzite beds. Specs. 8 to 14, State Coll., App. B, Vol. II., are from the Lower bed, and Specs. 50, 51, and 52, same Coll., are from the Upper. The extensive beds of quartzite, which occur in the Menominee region, will be fully considered in Chapter V. This rock is also of frequent occurrence in the L'Anse range and toward the Montreal river, as will appear in following Chapters. A beautiful example of false stratification, or discordant parallelism, was observed in this last-named region, as is shown by Fig. 2, sketched near the south quarter post of Sec. 10, T. 47, R. 45. It was a true granular quartzite, but showed deposition marks almost as plainly as a fresh-cut sandbank.

Fig. 2.
False bedding (discordant parallelism) of Quartzite—Gogebic Region.



F. MARBLE (Limestone and Dolomite).

(See Mr. Julien's descriptions, 101 to 113, App. A, Vol. II.)

The association of this rock with the Lower Quartzite, or rather the transition of the latter into marble, has been mentioned. This transition is seldom complete, the marble being always more or less silicious. As is usual in such cases, the change is gradual, producing all varieties, from calcareous quartzite to silicious marble. The prevailing colors are light gray, salmon and reddish. The purest varieties often present a sparry structure, with large lamellar facets like orthoclase feldspar, with which it is often confounded,

but from which it can readily be distinguished by its softness. Beds of argillite are invariably associated with the marble. See Fig. 19, App. E, Vol. II. Outcrops often present minute ribs or ridges of the more silicious layers, left by the weathering away of the purer marble.

The mean specific gravity of a large number of specimens averaged 2.82. See App. B, Vol. II. Pure marble has the same composition as pure limestone, of which it is simply a crystalline or highly altered form, that is, it is a carbonate of lime;—if carbonate of magnesia is present in considerable quantity, as is often the case on the Upper Peninsula, the rock becomes a *dolomite*. Marble is readily distinguished from its effervescing with acids, when pulverized.

Marquette marble has been considerably used as a blast furnace flux, for which purpose it only answers passably well, on account of the silica so generally present; silica, in the form of quartz, and jasper being always present in the *ores*, it is very desirable to have none in the *flux*, for it is to get rid of silica in the form of slag, that lime is used in the furnace. Large amounts of Kelly island limestone, which is quite pure, is now being imported. For building purposes, its hardness, variability in texture and the difficulty of securing large blocks, have so far prevented its use; beautifully variegated small blocks can, however, be easily procured. Specs. 6,198, 6,199, 6,200, University of Michigan Cabinet, are from the Morgan Furnace quarry, and Specs. 106 to 113, State Col., App. B, Vol. II., from the Chocolate quarry, just south of Marquette, all belonging to formation V., represent the chief varieties of this rock.

No marble has been observed in the L'Anse district, nor between Lake Gogebic and Montreal river, but it is one of the most abundant rocks in the Menominee region, where it occurs in a much purer form than in Marquette, usually more dolomitic. See Chapter V. and Specs. 102 and 103, App. A, Vol. II. Marble of similar quality is also abundant in the vicinity of Fence and Michigamme rivers, in Towns 44 and 45, R. 31. See Spec. 105, App. A, Vol. II.

G. ARGILLITE OR CLAY SLATES AND RELATED ROCKS.*

(Constitutes bed XV., and occurs in bed V. and elsewhere.)

It was previously mentioned under Groups E and F, that beds of clay-slate were sometimes interstratified with layers of quartzite and marble. Fine examples of this, in the case of both rocks, can be seen respectively at the Greenwood and Morgan furnaces. In addition to these, at least two distinct beds of argillite have been made out; one immediately beneath the ferruginous schist of formation X., to be seen in outcrop on the south shore of Teal lake, near west end, and in the railroad cut about one mile east of Negaunee. See Spec. 20, App. B, Vol. II. Another and far more extensive bed is XV., which forms the stratum next above the Upper Quartzite; boulders of this bed, which had the appearance of being near the parent ledge, were found in the railroad cutting, near the pockets at the Washington mine. At the Champion this formation is exposed in the branch railroad, and it is found at numerous points on the north shore of Lake Michigamme.

The prevailing color of this rock is usually dark brown or blackish, but where associated with the marble it is sometimes reddish. It has a true slaty cleavage, distinct from the bedding, but seldom splits in sufficiently large or regular slates to warrant us in supposing it may in places produce roofing slates, although experienced persons express the belief, that good slates will yet be found in the Marquette region. Black carbonaceous matter is often present in this slate, a preponderance of which produces the rock which will be described hereafter under J. A variety at the Greenwood furnace contains a large amount of iron-pyrites; and the first stack built of it had to be taken down, from the decomposition of this mineral. The slate in the branch railroad cut, at Champion, shows a slight tendency to be micaceous and holds garnets. See Spec. 56, App. B, Vol. II. Silicious bands often exist

^{*} Mr. Julien has in App. A, Vol. II., given the results of much study of these rocks, and has divided them into the true argillites and several other varieties possessing a different composition. See descriptions 189 to 225. As this difference cannot readily be made out by the unscientific, and as it is not important to the practical man, it will not here be attempted to separate these varieties.

in this rock, faintly marking its bedding at an angle with the cleavage, as can be seen in Spec. 20, App. B, Vol. II.

Overlaying the Lake Superior and Barnum ore deposits, hence occupying the place of the Upper Quartzite, is a greenish-gray schist, obscure in its composition, and somewhat like the magnesian schists D, but apparently of the same general character as this group. See Spec. 55, App. B, Vol. II. This rock may very properly be regarded as the connecting link between Groups D and G, which evidently graduate into each other, as did C and D. It is frequently stained reddish-brown along the seams and cracks, proving the presence of protoxide of iron, and shows in places beautiful dendritic delineations of manganese. This formation does not show the cleavage structure, so conspicuous in the schists of Group D, which are bedded with the pure ore at these mines. At the most westerly opening of the Lake Superior, thin beds of quartzite appear, indicating that the presence of argillite in this bed is probably only local. See Map No. IX.

An example of a magnesian schist (D) graduating into an argillaceous variety can be seen in the slate which overlies the specular ore of No. 1 pit, New England mine, which, by its high specific gravity (3.03), evidently contains considerable iron. Another ferruginous and probably chloritic variety occurs on N. W. 1/4 Sec. 31, T. 47, R. 25, where explorations for iron have been made by the Morgan Iron Co.

The average specific gravity of a number of typical specimens of argillite was 2.75. See App. B, Vol. II. The rocks above described are illustrated by Specimens 1,039, 1,072, and 1,036, App. C, Vol. II.

Beyond the limits of the Marquette region, we find in the recently explored Huron Bay district, particularly in the south part of T. 51, R. 31, the finest clay slates so far discovered in Michigan. Several competent experts have examined this district, and pronounced the slates of the best quality for roofing and other purposes, and in immense quantity. See Spec. 81, App. B, Vol. II. Companies are now at work in this district, the organization of which is given at the end of Chap. I. For an account of the clay-slates in the Menominee region, see Dr. H. Credner's papers (Leipsic).

This rock also occurs west of Lake Gogebic, as will be mentioned hereafter.

H. MICA-SCHIST.

(Formation XIX. contains the principal development of this rock. See Mr. Julien's description, No. 301, App. A, Vol. II.)

There appears to be but one extensive stratum of this rock, the character of which is unmistakable, which is at the same time the youngest and one of the thickest beds of the whole Huronian series. This formation, which I have numbered XIX., forms the surface rock along the south shore of Michigamme lake, among its islands, along the outlet for several miles, and westward from the lake through the southern parts of T. 48, Ranges 31 and 32, as shown on Map III. The rock is sometimes so silicious as to be rather a micaceous quartzite, but usually its true character is very plain. It frequently contains seams and bunches of white quartz, occasionally seams of black hornblende, and often holds numerous imperfect crystals of a delicately pink-colored, coarsely fibrous mineral, which Prof. Brush decided was andalusite, and brownish, smaller, and more perfect crystals of staurolite.

Andalusite and staurolite have not been observed elsewhere in the Marquette region in rocks of any age. Imperfect small reddish garnets are sometimes abundant, but they were not observed at the same places as the first-named minerals, and seemed to be nearer the base of the formation. The mica, which usually holds but little quartz, is of a brownish color on fresh fracture, weathering more grayish; its scales show a constant tendency to bend themselves around the imbedded crystals, like the fibres of wood around a knot. The projecting rounded crystals give the weathered rock a warty look, having somewhat the appearance of a conglomerate, as can be seen on the most southerly islands in Lake Michigamme. The specific gravity of this porphyritic mica-schist varied from 2.81 to 2.89, the mean being 2.84. See Specs. 1,031, App. C, and 61, App. B, Vol. II.

Descending in the series, the next mica-schist to be noticed is entirely different from the above, in being black, and decidedly dioritic in its affinities. It occurs in the upper part of diorite bed XI. at Republic Mountain. The deposit is not extensive, and its relations with the diorite indicate that it is a local variety, apparently graduating into dioritic schist.*

^{*} The local micaceous character of bed XV. has been noticed.

One other mica-schist, that associated with the Cannon ore on Sec. 28, T. 47, R. 30, deserves notice. This rock resembles XIX. only in the brownish color of its mica; it contains no crystals of other minerals, and is always quartzose, sometimes to the point of becoming a micaceous quartz-schist. The age of this rock has not been satisfactorily determined, but it is near the base of the series. The striking peculiarity of this variety is the fact, that in places the mica is replaced by micaceous specular iron ore, thereby becoming a specular schist, a rock very nearly related to the itaberite of some writers. The Cannon Iron Company's explorations, in which a fair specular slate ore has been found, are located in a highly ferruginous part of this bed. See Spec. 16, App. B, Vol. II. The relations of this rock with the lower quartzite of the North belt, Menominee Iron region, is fully discussed in another place.

I. ANTHOPHYLLITIC SCHIST.—(in bed XVII. and others.)

(See Mr. Julien's descriptions 174 to 178, App. A, Vol. II.)

Immediately below the great mica-schist bed, XIX., and probably separated from it by a stratum of quartzite, XVIII., is a well-defined stratum of a slightly magnetic rock, varying in color from brownish-black to dull slate on fresh fracture, and grayish to black-ish in outcrop. It often shows manganese,* and always a fibrous, light-brown mineral, which Prof. Brush, from the examination of some imperfect specimens, decided to be anthophyllite,† a variety of hornblende, and suggested the name here employed for this group.

Numerous outcrops of the rock occur along the north shore of Michigamme lake, and a fine development at the mouth of the Bi-ji-ki river, as well as at the Champion furnace, where layers rich in manganese occur. A specimen afforded Dr. C. F. Chandler 25.2 per cent. of metallic iron, and 4.37 per cent. of metallic manganese. See Specs. 58 and 59, App. B, Vol. II., and 178 App. A, Vol. II.

Below the ore formation XIII., at the Spurr Mountain, are layers of schist of a similar character, a specimen of which afforded Mr. Britton 45.21 metallic iron, 1.78 metallic manganese, 26.36 silica.

^{*} This variety resembles plumbago, and may contain carbon.

⁺ Prof. Dana now regards anthopholite as a distinct mineral.

A moderate increase in the percentage of iron and manganese therein found (which may very likely take place in some part of the bed) might render this rock a workable ore, particularly as the associated mineral is an easily fusible hornblende instead of the silica so common in the other ores. Ores containing 12 to 20 per cent. of manganese need not be rich in iron, to give them merchantable value.

Underlying this formation (XVII.), or perhaps forming its base, is a rock, numbered XVI., which at Champion and on Sec. 26, T. 48, R. 31, shows a tendency to pass into a *limonitic schist*, and may very likely afford workable soft hematite ore in some part of its course. The propriety of giving this rock, about which so little is known, a distinct stratigraphical designation, may be questioned; but its ferruginous character, pointing toward the possibility of commercial value, led to this course.

South of the Washington mine, and therefore stratigraphically below the ore formation,—for the whole dips north,—there is an obscure schistose rock of a gray color, weathering brown, and containing very little iron, often garnets, but made up chiefly of a light brownish fibrous mineral, which is probably anthophyllite, but which in places resembles mica. These rocks are extensive, stretching from the Champion mine eastward to the old Michigan mine. They are generally slightly magnetic, and unquestionably occupy the place of the silicious ferruginous schists of Group B. The diorites associated with them are also peculiar, the two sometimes resembling each other. This obscure series is well illustrated by Specimens 6,086 to 6,099, University of Mich. Collection. See also Specs. 174 and 175, App. A, Vol. II., and 27, App. B, Vol. II. Their affinities are apparently with this group.

J. CARBONACEOUS SHALE.

(See Mr. Julien's descriptions, 246 to 251, App. A, Vol. II.)

The presence of plumbago or graphite (a form of carbon) was noticed in the anthophyllitic schists, last described. Carbonaceous matter has also been observed in various clay-slates, as was noticed in describing the Argillite Group, and we could have placed this rock there as a variety of clay-slate, very rich in carbonaceous

matter. It is of a bluish-black color, but burns white before the blow-pipe, marks paper like a piece of charcoal, is soft and brittle, slaty in structure, and is the lightest rock yet found, having a specific gravity of but 2.06.

This rock has been found in the Marquette region only at two localities: 1. The S. C. Smith mine, T. 45, R. 25, where it seems to bound the iron-ore formation on the northeast. See Spec. 6,163, University of Mich. Collection.

(2.) On the south side of Sec. 9, T. 49, R. 33, along Plumbago brook, as will be fully described in the account of the L'Anse Iron range, is a large deposit of carbonaceous shale, a specimen of which gave Prof. Brush—carbon, 20.86; earthy matter, 77.78; moisture, 1.37. Another sample from same locality gave Mr. Britton—moisture and carbonaceous matter, 22.51; oxide of iron, 4.37; earthy matter, 73.12. See Spec. 64. App. B. Vol. II. These analyses prove the material to have no commercial value, but possess scientific interest as proving the existence of a large amount of carbon in the Huronian rocks. The equivalency of these shales with the members of the Marquette series has not been established; they are undoubtedly Huronian, and are, I suppose, younger than the ore formation XIII.

CHAPTER IV.

GEOLOGY OF THE MARQUETTE IRON REGION.

I. MICHIGAMME DISTRICT.

IN describing the geological structure of the Marquette Iron series, I shall begin with the Michigamme district, because its structure is simplest, the iron ranges easily followed on account of their magnetism, and because my explorations and surveys have there been more thorough than in either of the other districts.

The Champion mine, 33 miles west of Marquette, is at one of the most extensive, regular and typical deposits of ore in the whole region (see Map No. VII.). The strike is a few degrees south of west, and dip north at an angle of 68°. The extent and nature of the workings at the date of the survey may be seen by reference to the map. Up to this time the mine has produced an aggregate of 225,000 tons of magnetic and slate ore of first-quality. The general form of the ore mass is that of a huge irregular lens, or flattened cylinder-shaped mass, which thins out to the east and west to so narrow a width, as not to be workable. The easterly portion of the deposit is black, fine and coarse-grained magnetic ore; the westerly portion is specular slate ore, with a small admixture of magnetite. The local magnetic attractions are very strong and are fully considered in Chapter VIII. The position of the plane dividing the two varieties is approximately shown in the sketch of workings on Map No. VII. The whole mass here described is not, however, pure ore, as may be seen by inspecting plans of the first and second levels on the map. Minor irregular lens and pod-shaped masses of pure ore, "mixed ore" (banded ore and quartz), together with whitish and greenish magnesian schists, alternate like the muscles of an animal, forming, as a whole, a comparatively regular deposit. Overlying the ore on the north side is a hanging wall of gray quartzite, the thickness of which is considerable, but could not be accurately determined on account of the drift. Immediately south

of the ore, if it may not be regarded as a part of the ore formation, is a banded jaspery or quartzose rock, containing some iron. Next south, and underlying the whole ore formation, as may be seen by an outcrop near the east end of the mine, is a bed of diorite ("greenstone"); this rock in places becomes schistose and chloritic in character. South of the diorite is a silicious schist and then a swamp. The arrangement of these beds may be seen in geological section A—A," on the map, where they are numbered in Roman numerals X. to XIV., the latter designating the quartzite.

Following the Champion range east one mile, we arrive at the Keystone Company's mine,* where but little work has been done, and the arrangement of the rocks in consequence not so easily made out. A small bed of magnetic ore was opened at this locality two years ago, and what is said to be a large deposit of specular ore has but just been discovered on the same place. Five hundred feet north are a number of outcrops, indicating the presence of a heavy bed of conglomeritic schist, which holds masses of quartzite, varying in size from pebbles to others two feet by one thick, and even larger. It also contains flattish fragments of various schists and slates. Further north it passes into a brownish schist, containing pebbles of quartzite. This rock is believed to correspond with the overlying quartzite of the Champion, and is marked XIV. on the map and sections. North of this, and exposed in the railroad cut, is a micaceous slate, containing garnets, marked XV., and represented by Specimen 56, State Collection, App. B, Vol. II.

North and west of this locality, about one-fifth of a mile, are a number of test-pits, in many of which is exposed a soft, brownish, ferruginous rock, which affords hand specimens of soft hematite ore. This rock is marked XVI., and is represented in the State Collection by Specimen 57, App. B, Vol. II. Immediately south of the Keystone workings is a specular schist or conglomerate, in which flattened pebbles, or very uneven lamina of quartz, are contained between thin layers of micaceous specular ore. This formation is believed to be the equivalent of XII. of the Champion mine section, and is so numbered on the map.

West and south are numerous extensive outcrops of a brownish banded magnetic schist, marked X. on Section C—C", Map VII.

^{*} Late " Parsons Mine."

The arrangement and character of the rocks along the intermediate section, B—B,' will be sufficiently understood from the above descriptions and an inspection of the map. The other formations represented will be considered in another place.

At the Spurr and Michigamme mines we find rocks identical in their general character and sequence, although the order is reversed, this series being on the opposite side of the basin from the Champion. Projecting all the facts observed along the north shore of Michigamme Lake on one plane, which we will assume to pass north and south through the Spurr Mountain mine, the following Geological Section is easily made out:

Commencing at the most southerly and uppermost bed (the whole series dips to the south), we have, first, a comparatively soft, grayish and blackish flaggy rock, containing considerable iron, a little manganese and often made up largely of a hornblendic mineral, which occurs in needle-shaped crystals. Professor Brush calls this rock anthophyllitic schist. See Specimens 58 and 59, State Collection, App. B, Vol. II., and Chap. III.

This rock is numbered XVII. on geological section No. 9, map of the Marquette Iron region, which see. It is also well exposed at the mouth of the Bi-ji-ki river, in the railroad cut just east, at the Champion furnace, and at numerous projecting points along the north shore of the lake.

The next rock to the north, in descending order, (numbered XVI. on the map and section,) on account of its tendency to decomposition, has never been seen in outcrop; it is exposed by the explorations for ore, made on the north side of Sec. 26, T. 48, R. 31, and at the Champion; its character was indicated in describing the Champion series, and need not be repeated here. As will be seen, this rock has the same number in each section, and the two exposures are believed to belong to the same bed. It is not improbable that future investigations may prove it to be a variety of the ferruginous anthophyllitic schist XVII., already described, a point which was considered in Chapter III., Group I.

Next below is a dark-colored clay-slate, which also, on account of its softness, is seldom seen in outcrop. It is, however, exposed on the point in northeast part of Section 29, and at other places along the north shore of the lake. On the Spurr mountain, geological section No. 9, this formation is numbered XV., and is

believed to underlay the swamp and creek immediately south of the mountain which finds easterly prolongation in Black bay. As will be seen by reference to the Champion sections, this rock is regarded as the equivalent of the micaceous clay-slate XV., there described.

North of this clay-slate, and immediately overlying the ore at both the Spurr and Michigamme mines, is a quartzose rock numbered XIV., which is in places a hard conglomerate, and again, especially when in contact with the ore, a fine whitish sandstone. See Specimen 52, State Collection, App. B, Vol. II., and Julien's descriptions, Specs. 358 and 359, App. A, Vol. II. This rock is unquestionably the equivalent of the upper quartzite XIV. of the Champion section, which, on the whole, it closely resembles in its lithological character. See also Group E, Chapter III.

The prevailing variety of ore of the mines on this range is a fine-grained, somewhat friable, rich, blackish magnetite. See Specimens 40 and 41, State Collection, App. B, Vol. II., and also Iron Ores, Chap. III. There is also at the Michigamme mine a hard, fine-grained, steely magnetic ore, in considerable quantity. Analyses of these ores will be found in Chapter X. The surface indications, magnetic attractions, explorations and mining operations but just commenced, point unmistakably to large deposits of high grade magnetic ore at both localities.

The Spurr Mountain is an east and west ridge, the summit of which is 118 feet above Lake Michigamme and 75 feet above the creek, which passes south of it. This ridge terminates abruptly to the west near the centre of the northwest 1/4 of the southwest 1/4 of Sec. 24, T. 48, R. 31, where there is a natural exposure of merchantable ore 40 feet thick horizontally, being the largest outcrop of pure magnetic ore I ever saw. Mining operations, just begun, have demonstrated the thickness to be still greater, and the deposit to extend at least several hundred feet east and west, with a probability, based on magnetic attractions, of its extending much farther. The bold face, small amount of earth covering, softness of the ore, its apparent freedom from rock, convenience of the railroad and accessibility, present facilities for mining and shipping. which could not well be surpassed. The magnetic observations made at this locality, where the attractions were remarkably strong, are given with illustrative diagrams in the special chapter devoted

to that subject. It is easy by means of the dip compass, to follow this iron range two-thirds of the distance along the north side of Michigamme lake, and west-northwest from the Spurr to the First lake, an aggregate distance of over nine miles, as may be seen by the map of the Marquette Iron region, No. III. It must not by any means, however, be supposed, that here is a workable deposit of ore nine miles long; this has not been proven, but on the contrary, it has been proven that for a considerable portion of this distance the ore is not workable, having altogether too large an admixture of rock. Therefore, while it may be confidently asserted, that all of the rich hard ore which will be found in this vicinity, will be in or near the belt of magnetic attraction already described, it may be asserted with equal truth, that at least three-fourths of the whole length of this belt is barren ground, according to the present standard of merchantable ore. The law of the distribution of the rich "chimneys," "shoots," or "courses of ore," as they are designated in different mining regions, along a given iron range, has not been made out. The subject is more fully considered in Chapters VII. and IX.

Besides the deposits already described on this range, one other has to be mentioned, that on the east side of railroad Sec. 23, adjoining the Spurr on the west. The magnetic attractions here are remarkably strong, and explorations have revealed the existence of a small workable deposit of first-class magnetic ore. Whether this deposit connects with the Spurr or not, was not fully determined.

As has been remarked, both the granular and compact varieties of magnetic ore occur at the **Michigamme mine**. The explorations on this location, which were conducted by the writer, developed in a distance of 1,200 feet, east and west, seven places, where pure ore existed of a thickness of from seven to thirty-five feet, rendering it probable, that the ore deposit is continuous and workable for the whole of this distance. Mining operations, which have commenced at this location, confirm these results. Pure ore was found in place at two points on same range west, on Sec. 19 of the Michigamme Company's property, but not enough work was done to prove their extent. Eastward the ore can be traced by the magnetic needle into Michigamme lake, on the south side of Sec. 20.

There can be no doubt these deposits and the Champion belong

to the same horizon, being the opposite croppings of the synclinal basin, which passes under Michigamme lake; although the Champion deposit has not been traced westward, nor the Michigamme range eastward, to points where they come directly opposite each other. Whether the specular slate ore found so abundantly at the Champion will be found on the north side of the lake, remains to be seen. I see no reason why it should not; the explorations, so far, have been based entirely on magnetic attractions, and would therefore not be likely to result in finding specular ore.

Underlying the pure ore here, as at the Champion, is a ferruginous quartzose rock, which has an immense development on the Spurr-Michigamme range, where it is a well-characterized reddish quartz schist (jasper), containing thin layers of pure specular ore; these layers being occasionally thick enough to afford hand specimens. See Specimen 33, State Collection, App. B, Vol. II. A similar rock is found, as will be seen hereafter, at the Republic mountain, where it has the same relative position and number, XII.

Underlying this iron series we find, as at the Champion, a diorite (greenstone), but which here has a much greater development, forming a conspicuous ridge which borders the Michigamme and Three Lakes valley on the north, and which has already been described under Topography in Chapter II.

This greenstone ridge is separated from the granite region to the north by a valley about half a mile wide, which is underlaid by various schists and quartzites, about which little is known. Two are marked X. and V. on the Spurr-mountain section No. 9.

The most easterly developed mines in the Michigamme district are the Washington and Edwards, represented by map No. VIII. The general structure, which we are now considering, can be easiest made out at the Edwards and "old mine," which are adjacent, and about three-fourths of a mile west of the Washington mine proper. The general character and order of the ore and accompanying rocks at this locality is so similar to that of the mines already described, that a careful inspection of the map and accompanying sections leaves but little to be said. The Upper Quartzite XIV. is fully exposed in outcrop, as well as in the railroad cut, just west of the mines, where it is a coarse conglomerate, often schistose, as is shown by Specimen No. 51, State Collection, App. B, Vol. II.

The same formation is a compact gray quartzite at the Edwards mine, and at other points in the vicinity.

The ore formation XIII. affords at this group of mines all the varieties, already designated as being found at the Champion, Spurr, and Michigamme mines. Like the Champion, here are intercalated beds of magnesian schist, the arrangement of which are shown on the sections of workings given on the map already referred to, as well as in the plan of the Edwards mine, by A. Kidder, Plate XIX., Chap. IX., where the subject of detailed structure is more fully considered. One of these schists, of a decided talcy character, is represented by Specimen 54 of State Collection, App. B, Vol. II.

The underlying ferruginous quartzose rock, XII, has a large development south of the Edwards mine, and to it probably belongs the "red ore" of the old Washington. Southwest of the latter mine are large exposures of the peculiar conglomeritic specular schist, mentioned as occurring on the Keystone property, east of the Champion.

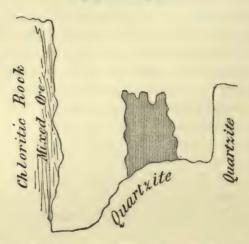
The dioritic formation, XI., is represented by a large outcrop of a greenish schistose rock, apparently chloritic, which can be seen immediately south of the old mine. Below this formation are alternating schists and diorites of different varieties, which are sufficiently well shown on the map and sections. One of the most interesting varieties is represented by Specimen No. 27, State Collection, App. B, Vol. II., procured 500 feet south of Pit No. 9, Washington mine.

The Washington mine proper presents some of the most complicated structural problems, to be found in the Marquette region, and I will not here either attempt their solution, or even advance the hypothesis which I have formed. Suffice it to say that, in general, the mine is a monoclinal deposit, dipping away from the St. Clair mountain (which term I apply to the high ground to the south) to the north and under the great swamp. The minor rolls, the peculiar faulting at the East Hill, and the trap dykes, would, if fully considered, occupy a chapter.

I cannot, however, pass to another mine, without noticing the singular manner in which the mass of ore, known as Anderson's cut, or Pit No. 1, is terminated in its downward course, as shown by Figs. 3 and 4. It will have been observed, that the usual form of ore masses is *lenticular*, *i.e.*, they generally terminate by wedging

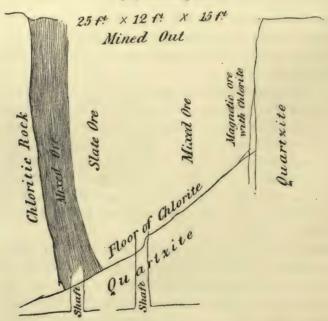
out more or less gradually each way. This exceptional mass, as will be seen, is obliquely and abruptly cut off, the bottom rock be-

Fig. 3.-Looking East.



ing a quartzite of the same kind, that bounds the deposit on the north, and there is no evidence of faulting on the plane of this floor,

Fig. 4.—Looking East.



or along the quartzite wall. An hypothesis to account for this phenomena, based on a sedimentary origin for these rocks, will readily suggest itself and need not be stated.

The Republic mountain and its prolongation on the Kloman lot, is the only remaining ore deposit of the class under consideration, which remains to be described in the Michigamme district. See map No. VI. The Magnetic mine group, embracing the Cannon and Chippewa locations, belong to a different geological horizon, produce different ores, and will be considered hereafter.

The immense mass of pure specular ore, which was naturally exposed near the centre of the north ½ of the southeast ¼ of Sec. 7, T. 46, R. 29, could leave no reasonable doubt in the mind of the experienced observer, that this deposit of ore was one of the largest; if not the largest, in the Marquette region. This outcrop, the extent of which is shown on the map of the Republic mountain, being there marked "pure specular ore," is, so far as I know, the largest outcrop of any equally rich ore, ever found in the United States.

The elevation of the ore, 120 to 150 feet above Michigamme river, gives an unsurpassed opportunity for mining operations, which began in the spring of 1872, and confirm, as far as they extend, the "surface show." Several other small outcrops of pure ore occur in the iron belt, one of the largest of which is near the centre of the Kloman mine lot, in southwest fractional 4 of Sec. 6, same Township.

The numerous outcrops of rock and ore at this mountain, the strong magnetism possessed by three of the beds, the remarkable uniformity in thickness of the several formations, and the bold topographical features presented, all of which were carefully surveyed and are faithfully represented and explained on the accompanying topographical, geological, and magnetic maps and charts (Plates VI. and XII. Atlas), leave but little more to be said in this place, regarding the general structure of the Republic mountain.

The lithological character of the rocks and ores will also be fully understood from the 14 specimens from this locality, which are embraced in the State Collection, App. B, Vol. II. The ten formations represented by colors on the map, as composing the Huronian series, will now be enumerated, commencing with the

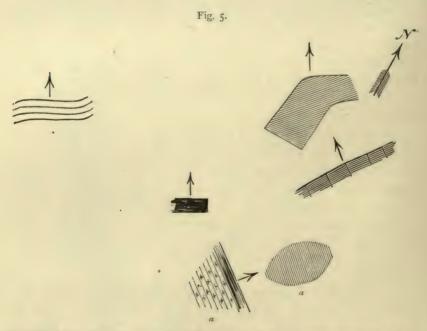
lowest, which reposes non-conformably on the Laurentian granites and gneisses.*

The lowest bed of the series will be numbered V., for reasons which will hereafter appear.

V. A quartzose rock, which is exposed at but a few points, and is best seen near 4,600 southwest and 6,200 southeast (see rectangular ordinates on map), from which locality Specimen 8, State Collection, App. B, Vol. II., was obtained.

VI. Is a magnetic, bright, banded, silicious and chloritic schist, containing considerable iron. See Specimen 15, State Collection, App. B, Vol. II., from near locality of Specimen 8. Very large exposures of this schist occur on the northeast side of the mountain, and southeast of Bear lake. The regular, various-colored stripes,

^{*} This sketch (6, 100 southeast and 4,700 southwest, Map VI.) represents outcrops of Huronian quartzites and schists dipping north-northwest, and the Laurentian gneisses, a a, dipping northeast, the latter being within 50 feet of the former. The actual contact is not seen, but the stratigraphical relations indicated, in connection with the wide difference in



their lithological character, leaves no doubt in my mind of the non-conformability of the two systems, the Huronian being the youngest. This non-conformability can also be observed on the L'Anse Range. See page 156.

which this formation, as well as VIII. and X. displays, strongly suggests a rag carpet. The greenish layers are apparently chloritic, the whitish and grayish are quartz, and the brown and dark gray are silicious layers of the red and black oxides of iron. Some of these lamina are quite pure iron ore, and the whole mass may contain from 15 to 30 per cent. of metallic iron. The magnetic power displayed by these schists is remarkable, as will be seen by inspecting the charts and explanatory text already referred to.

VII. Is a diorite of the general character of those, so fully described by Mr. Julien in App. A, Vol. II., as will be seen by reference to Specimen No. 18, State Collection, App. B, Vol. II.

VIII. This magnetic silicious schist in its lithological character differs in no essential particular from No. VI., already described. See Specimen No. 19, State Collection, App. B, Vol. II. This formation is noticeably thin, not exceeding 40 or 50 feet, the other beds being from three to five times this thickness, as can be seen on the map.

IX. Is a Diorite similar to VII. See Specimen No. 22, State Collection, App. B, Vol. II.

X. A magnetic silicious schist similar to VIII. and VI., but containing in places more iron, as at 5,600 southeast and 2,500 southwest, from which locality Specimen 23, State Collection, was obtained. This, it will be observed, is a fair specimen of magnetic flag ore, containing probably 45 per cent. of metallic iron.

XI. This formation is made up of a coarse-grained diorite, in which a light grayish and reddish feldspar is a conspicuous ingredient, as may be seen on the Kloman lot, as well as at the knob southwest of Bear lake, from which Specimen No. 29, State Collection, App. B, Vol. II., was obtained.

A schistose variety, containing considerable black mica, occurs in the same formation, at 3,400 southwest and 5,300 southeast, where Specimen No. 30, State Collection, was obtained, although it does not truly represent the prevailing variety at this locality.

XII. This is a reddish quartz or jasper schist, containing thin lamina of specular ore, and very similar to the corresponding formation of the Spurr mountain series already described, as will be seen by an examination of Specimen 32, State Collection, App. B, Vol. II.

XIII. We have now reached the iron-ore formation, the principal

outcrops in which have been enumerated. Four varieties of material chiefly make up this formation, which in the order of apparent quantity are as follows:

a. A banded rock made up of alternating layers of red quartz or jasper and specular ore, designated by the miners as "mixed ore," the richer varieties of which are now shipped as second-class ore. See Specimens 36 and 37, State Collection, App. B, Vol. II. The contorted and plicated lamina of this rock, brought out by the alternating bright red and steely bands, and which could be but poorly illustrated in Figs. 19 to 29, App. K. Vol. II., are very beautiful, being often contorted and plicated in a striking manner. See Iron Ores, Chapter III. It may be remarked in passing, that such contortions in the constituent lamina of rock formations generally indicate the presence of great folds in the whole formation, as is plainly the case at this locality.

On the southwest side of the basin, at points in the ore formation marked "specular conglomerate" on the map, occurs a true schistose conglomerate, in which pebbles, chiefly quartz, are bedded in a matrix of silicious ore. On the supposition that this rock may be a secondary form of the laminated or mixed ore, and from a desire not to multiply subdivisions in this connection, it will at present receive no further consideration.

- b. Next to the mixed ore in quantity, so far as can be judged by what can be seen, is the pure *specular* ore. See Specimen 46, State Collection, App. B, Vol. II. The specific gravity of these specimens varied from 5.09 to 5.56, the average of four being 5.24, or greater than that of any other ore in the region, which should indicate a somewhat greater richness in metallic iron; whether furnace work will confirm this, remains to be seen.
- c. The next in supposed order of quantity is a rich, black, magnetic ore, similar to the Spurr and Champion ores, but much coarser in its grain. See Specimen 39, State Collection, App. B, Vol. II.
- d. Dividing the specular ore below, from the magnetic ore above, can be seen, in cut No. 1, Republic mine, a bed several feet in thickness of a magnesian schist similar to that previously mentioned, as being found in the Washington and Champion mines. See Specimen 53, State Collection, App. B, Vol. II.

XIV. The Upper Quartzite at Republic mountain is a gray massive rock, sometimes banded, and, near the contact with the iron,

sometimes conglomeritic, containing large and small flattened fragments of flaggy ore. The prevailing variety is represented by Specimen No. 50, State Collection, App. B, Vol. II.

XV. Near the south point of Smith Bay is a considerable outcrop of what appears to be a dioritic schist, not unlike Specimen 31, State Coll., containing mica and garnets. It has some resemblance, as will be seen by the description, to the micaceous clay-slate of corresponding number of the Champion section, Specimen 56, App. B, Vol. II.

The horse-shoe form of the surface rocks, as indicated by outcrops, which is so conspicuous a feature on the map, taken in connection with the dip of the strata, as indicated by the arrows and geological section, leave no doubt whatever as to the structure of Republic mountain. It is evidently the south-east end of a synclinal trough with Smith's Bay in the centre, under which, at an unknown depth, all the rocks represented would be found and in the same order. The conjectural division plane, dividing the quartzite and ore (see section), may be regarded as hypothetical, only as to its position, which of course can finally be determined by boring.

It will be observed, that where the northeast side of the horse-shoe crosses the river, there is an offset of about 250 feet to the right, and that where the southwest arm of the shoe should cross the river, but very little appearance of Huronian rocks can be discovered on the west side, the Laurentian rocks to a great extent taking their place. These facts can be best explained, by supposing a fault to follow the line of this portion of the river, the east being the down side. On this supposition the Huronian rocks on the west side would have been eroded to a much greater extent than on the east, leaving as a consequence the narrow and incompleted series, shown on a section through the Kloman mine.

The proximity of the Champion ore deposit to the Laurentian, it being only about 400 feet distant, while at the Keystone (three-fourths of a mile east) the distance is three or four times as far, leaving room for a greatly increased thickness of vertical brownish banded magnetic schist (see Map VII.), can be best explained, by supposing a *fault*, similar to that just described, but having a direction nearly at right angles; that is, east by south.

These two instances are the best established cases of faults on a large scale, that have come under my notice, in the whole region.

Calling to mind the series of rocks, which have been described as occurring at the Spurr, Michigamme, Champion, Keystone, Edwards, Washington, and Republic mines, we are irresistibly led to the conclusion, that they are equivalents of each other, belong to the same series, and are of the same age. This hypothesis has already been introduced and carried through the descriptions by the corresponding numbers, which have been attached to equivalent formations in each section; it will no longer be regarded as an hypothesis, but accepted as a demonstrated theory. The Republic mountain section, it will be seen, is most complete for the rocks immediately below the iron, and the Spurr mountain section for those above. The latter embraces one formation of great extent and interest, which was not described, viz.: -XIX., which is made to include the several varieties of mica schists, so extensively developed on the south shore and among the islands of Lake Michigamme. This schist is often very silicious, and, in places, contains numerous crystals of garnets, and alusite and staurolite. See Specimen 61, State Collection, App. B, Vol. II., and Group H, Chapter III.

Near the centre of Sec. 25, at the west end of the lake, is a large mass, probably a ledge, of light-gray quartzite, which may fill in part at least, what appears to be a blank between the anthophyllitic schist XVII. and the mica schist XIX., just described. The number XVIII. is provisionally attached to this quartzite.

We have now described fifteen members of the Huronian series, from V. to XIX., both inclusive. This mica schist is the youngest member of the series, so far as my observations extend, to be found on the Upper Peninsula. It is proper to remark, however, that equivalency, member for member, of the Marquette rocks with the L'Anse, Gogebic and Menominee series, has not been established; they are all Huronian, and it is doubtful if any are younger than XIX.

With regard to the strata below V., there is less certainty as to their order and equivalency. I believe, that the iron ore and associated rocks, to be seen at the **Magnetic**, Cannon, and Chippewa locations, belong here. They are in any event the equivalents of each other, and are very near the base of the Huronian series. See Geological Section, No. 10, map of the Marquette iron region,

which extends from the Cannon to the Chippewa. At the latter location is a considerable deposit of ferruginous, silicious schist, or lean flag ore, in which occurs, in what I understand to be an irregular pocket-like mass, a peculiar specular ore of fair percentage, greenish-gray color, and containing numerous bright facets, which resemble scales of mica. This is in comparatively low, wet ground, and the extent of the deposit has not been determined. It resembles the Gilmore ore at north side, Sec. 26, T. 47, R. 26, Cascade range, the two being unlike any other ores in the region.

About 100 tons of 55 per cent. ore was taken from the latter location several years since, but work was not continued. The Gilmore deposit, as well as the Chippewa, is nearly in contact with the Laurentian.

At the Cannon location is a banded jaspery rock, holding thin layers of specular ore, which bears a striking resemblance to the rock of formation XII., and even to some varieties of "mixed ore." See Specimen 16, State Collection, App. B, Vol. II. A seam, several inches thick, of pure specular ore, was found here, but did not enlarge on being followed downward. The remarkable characteristic of this schist is the fact, that on following the range northwest and southeast, mica replaces the ore, and we have a micaceous quartz schist, or mica schist depending on the quantity of the latter mineral. These facts, already noticed, possess interest in their bearing on the nature of the Felch mountain ore deposit of the Menominee region, hereafter to be considered.

By far the most promising mine of this group, so far as existing explorations reveal, is the **Magnetic**, in south ½ of northwest ¼ of Sec. 20, T. 47, R. 30. The existence of a workable deposit of magnetic ore of medium richness has been proven. This ore, although highly magnetic, differs entirely in its character from those already described, as will be seen by inspecting Specimen No. 17 of State Collection, App. B, Vol. II. It is very hard, exceedingly fine-grained, and breaks into cubic or tabular pieces. Its structure is more like the flag ores than the first-class magnetites. It should yield about 55 per cent. in the furnace, although none has as yet been worked. The gangue is largely actinolite, instead of the more common quartz, which will help the reduction of the ore.

The relative geological position of this ore is shown in the accompanying north and south section, in connection with Map No. III.,

already referred to. As to the age of the series represented, I have but little doubt on account of their proximity to the Laurentian, and on lithological grounds, that they are the equivalents of the lowest rocks of the Republic mountain series, and are probably older than the lower quartzite V.

Fig. 6.

Geological Section (looking west).

Magnetic Mine. Sec. 20, T. 47, R. 30.



Level of water in Lake Michigamme 950 feet above L. S.

SCALE OF FEET.

A. Granite. B. Micaceous Quartz Schist. C. Quartzite and Quartz Schist.

D. Banded Magnetic Schist (ore). E. Greenstone or Diorite. F. Dioritic Schist.

B, C, D are undoubtedly the equivalents of the specular and micaceous schists of the Cannon series.

The line of magnetic attraction, running southwest and south, and finally south by east from the Magnetic mine, which has been traced to Sec. 9, T. 45, R. 30, is one of the longest and most persistent belts of attraction in the whole Lake Superior region. The maps of the United States Linear Surveyors mark its position very plainly, as is shown in the chapter on the Magnetism on Rocks, Plate v. Comparatively little exploration has been made on this range; but I see no reason why deposits of the character and equal in value to the magnetic, may not be found along it.

A large amount of very poor ore, and a small amount of very good ore, has been found in south part of Sec. 7 and the north part of 18, T. 47, R. 28; and quite recently a workable deposit of first-class specular ore is reported to have been found there, the locality being known as the **Michigan Mine.** Specimen No. 2, State Collection, App. B, Vol. II., is from this deposit.

Clarksburg, Geological Section No. 6, map of Marquette iron region, records the leading facts to be observed in this vicinity. The Roman numerals marked on the several formations express

their relative ages correctly; whether they also express the equivalency of these rocks with the Washington and other series previously described, I am not quite certain. Specimen No. 3, of State Collection, from formation marked III., possesses lithological interest, as being a Huronian rock allied to the Laurentian gneisses.

2. NEGAUNEE DISTRICT.

Following the same principle here that guided us in describing the mines of the Michigamme district—that is, beginning with those simplest in geological structure—we find on the Saginaw and New England range of mines (being the most westerly of this district), a structure almost identical with that of the Champion and Spurr mines. Referring to Geological Section No. 4, map of Marquette iron region, the rocks in the vicinity of the New England mine are represented as follows:—The ore formation XIII. is made up, as at the Republic mountain, of "mixed ore" (banded ore and jasper), magnesian schist and pure specular slate ore; magnetic ore being absent here, as in all the mines of this district. The quantity of specular slate ore at this mine is, so far as known, small; the small lens-shaped mass, that was formerly worked, having been abandoned.

Overlying the ore formation is the Upper Quartzite, XIV., dipping at a low angle to the north, as may be seen just north of the Parsons mine. This quartzite again comes to the surface about half a mile north, in a flat synclinal, where it again dips north and does not rise until we reach the New Excelsior mine, owned by the Iron Cliff Co., which is shown on the section.*

Returning to the **New England mine**, we find between the ore XII. and the quartzite XIV., a mass of specular conglomerate, somewhat similar to that described as existing at the Republic mountain, where it was regarded as belonging to the ore formation. The fact that it overlays the pure ore at this locality, and has lithological affinities with some of the conglomeritic varieties of the Upper quartzite, leads me to doubt in which formation it should be included. I incline to the view, that it belongs in XIV.

^{*} This general section was constructed more than a year before ore was found at this locality, but it has not been found necessary to make any changes in it.

Formation XII., underlying the ore, is here widely different in its lithological character and economic value from the corresponding formation of the Michigamme district, where, it will be remembered, it was a valueless reddish quartz schist, containing thin lamina of iron. If we suppose tepid, alkaline waters to have permeated this formation, and to have dissolved out the greater portion of the silicious matter, leaving the iron oxide in a hydrated earthy condition, we would have the essential character exhibited by this formation as developed on the New England-Saginaw range, and as will be afterward seen at the Lake Superior mine. This is not offered so much as an hypothesis to account for the difference, as to illustrate the facts observed. The prevailing variety of rock in this formation is a brownish silicious schist, containing a considerable amount of iron (Specimen 26, State Collection, App. B., Vol. II.). Scattered through this formation are here and there large and small pockets of soft earthy hematite ore, having usually the most irregular forms, that can possibly be conceived. This subject was discussed under iron ores, Chapter III. Specimens 34 and 35, State Collection, are ores of this class.

The Winthrop and Shenango mines are in this formation, and are producing hematite ores as rich as any now worked in the district, and excepting perhaps the Lake Superior and McComber, richer than any other of this class, as indicated by analyses, Chapter X.

Underlying this hematite formation is a diorite, XI., similar in its general character to the rock, having a corresponding number in the Michigamme district; below this and south, are various ferruginous schists and diorites, corresponding in a general way with the Michigamme series, but which have not been carefully examined in the vicinity of the New England mine. Recent explorations afford opportunities for study, which did not exist when this section was made.

The series at the **Saginaw** and intermediate mines, as well as further west, is so near an exact duplicate of what has been given above, as to require no further mention than to state, that the deposits of specular ore are larger than at the New England, which has been mentioned as being rather small for profitable working. There has been too little work done at these new mines, to determine the extent of the deposits, but I see no reason to suppose that any of those now worked will prove very large. The fact that Sec. 16,

the Parsons and New England mines, have produced specular ores and have been abandoned, is significant. No doubt, considerable amounts of first-class ore will be taken out on this range at a profit. The only question is, whether they will continue to produce such ore in quantity for a series of years, at a fair cost for mining.

This range of ore has been traced westerly into the northeast ½ of Sec. 24, T. 47, R. 28; west of this the drift becomes very deep and the ore range is lost. A shaft 67 feet through the sand in this vicinity found no ledge. Whether there is any stratigraphical connection between this ore formation and the Washington, six miles distant west by north, is not determined. So far as is now known, it is economically a blank in the Marquette iron belt. Work now in progress at the new Michigan mine, already noticed, may throw light on this interesting and important question. It is not at all improbable, that the Negaunee and Michigamme districts may be independent ore basins, in which case the intervening rocks, which are all Huronian, would consist of the lower members of the series, that is below XIII. Even should this be the case, valuable hematite and flag ores may be found in this now barren district.

The new **Excelsior Mine**, previously mentioned and shown on the New England section, is near the southeast corner of Sec. 6, T. 47, R. 27, and is, as will be seen, the opposite cropping of the basin. There is so much drift between these ranges, that not much can be said definitely about the nature of the intervening rocks; but it seems probable that we have here a great basin, underlaid by ore at an unknown depth, and that the New England and Excelsior deposits are related to each other in the same way, as it was assumed are the Champion and Michigamme deposits. This could be cheaply tested, and possibly an important discovery of ore made, by a drill-hole through the quartzite, near the railroad on the west side of Section 16. All efforts to find an extension of the Excelsior deposit east and west have so far failed.

Returning to the New England range and following it eastward, we find that near the south 1/4 post of Section 16, it bends suddenly to the northeast, making its way diagonally across this section to the Lake Angeline Mine, which produces specular ore, having such admixture of jasper, as to cause it to rank intermediate in the market between first and second class ores. Whether the deposit worked at this mine belongs to bed XII. or XIII., I have not determined,

the ore partaking somewhat the character of each. The overlying rocks on the north are covered by the waters of Lake Angeline.

To the south is a high ridge of diorite, XI., on the south side of which is an extensive deposit of soft hematite, owned and worked in part by the Lake Angeline and Iron Cliff Companies.

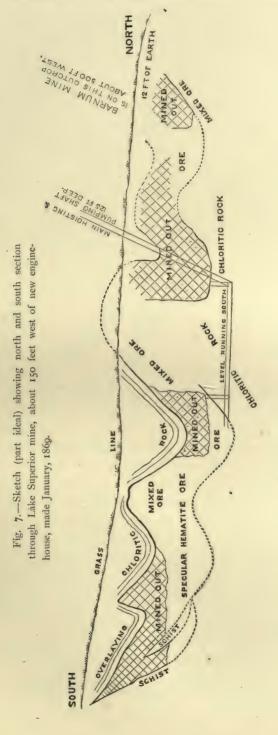
I suppose this hematite to belong to formation X., and therefore of the same age as the Negaunee and Foster hematites, which will be fully described below. It will be borne in mind, that the hematite ores on the Saginaw range occur in formation XII.

Without attempting to point out at present the structural relations of the Lake Angeline and Lake Superior ore deposits, we will pass at once to a consideration of the latter mine, one of the most extensive, productive and geologically interesting in the Marquette region.

The accompanying map, No. IX., representing the **Lake Superior** specular and hematite workings, together with the **Barnum mine**, is intended to give the geological facts to be observed in considerable detail, as well as the condition of the workings in 1870. The structure of the east half of this mine is more complicated, than that of any other in the district, and some questions connected with it remain unsolved.

Regarding for the present the west half of the mine only, we find presented on a small scale about the same structural phenomena, which is so prominent a feature in the Republic mountain rocks. The basin, or trough, in this case, however, abruptly narrows up, the sides and bottom being as it were gathered in, as if to be tied, at a point just south of the engine-house; to the west the outcropping edges of the basin diverge rapidly, and its bottom sinks into the earth in the same degree. If we suppose the frustrum of a hollow cone, lying with its axis horizontal and its small end towards the east, to be cut in two by a horizontal plane, representing the surface of the ground, the lower half will represent my conception of the form of the Lake Superior-Barnum ore basin. Conceive now this cone to be made of sheet-lead, and to be considerably bent and dented, and the illustration will be still more applicable.

A study and comparison of sections D—D', C—C', B—B', and A—A', in connection with the plan of the mine (Map IX.) will, I think, render it plain that this conception of the structure is in accordance with the facts; although the minor folds and faults con



siderably obscure and confuse the general structural question. Of course, it is not absolutely proven, that the Barnum deposit dipping south, and the continuation of the main Lake Superior deposit, now worked in Pit No. 25, which dips north, are opposite croppings of the same bed, and that the intervening space is underlaid by the ore formation, and that, therefore, if work continue long enough they will eventually connect under ground; but certainly all the facts point to this conclusion. The importance of this theory in

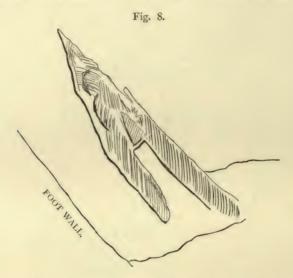


Fig. 8, represents on a large scale the south or left-hand end of the section represented in Fig. 7, and brings out the peculiar form of the "horse" of magnesian schist, which is shaded, the ore being white.

its bearing on explorations for ore, mining and valuing ore deposits, is very apparent. It shows, that such formations are not vein or dyke-line deposits, but true stratified beds, like the rocks by which they are enclosed. Their structure is therefore essentially the same as the coal, limestone, sandstone, and slate-beds, which are regarded as sedimentary deposits from water, subsequently more or less altered by heat pressure, and chemical waters acting during immense periods of time.

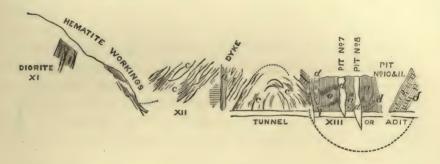
The Lake Superior-Barnum deposit evidently has a bottom, which will be reached within a period, of which it is worth while for the present generation to take some heed. So of many other deposits in the region.

As we go westerly from these mines the basins become, as we have seen, wider and correspondingly deeper. A depth of 300 feet in the Edwards mine reveals no essential change in the dip of the deposit, as will be seen by reference to the plans of the mine. The same is true of the Champion mine.

The time may come when, having worked out the steep upturned edges of the basins, and the flatter or deeper portions of the deposit are reached, ore properties will be valued somewhat according to the number of acres *underlaid by ore*, as coal now is.

Passing to the east portion of the Lake Superior mine, I confess myself unable to give any intelligent hypothesis of its structure. The facts observed are in part recorded on the Map of the mine on section E—E', and on the accompanying sketch, in part ideal, which represents on a small scale a section near E—E'. There seems to

Fig. 9.—Sketch showing Geological Section of the Lake Superior mine (looking west), near Sec. E—E', Map IX.



a. Chloritic schist. b. "Mixed ore." c. Limonitic schist (hematite rock). d. Pure Ore.

have been such a gathering together, crumpling, squeezing and breaking of the strata, as to nearly obliterate the stratification. An attempt has been made to represent the present condition of things, so far as revealed, by the workings. The remarkable features are the great masses of light grayish-green chloritic schist, having a vertical east and west cleavage, no discernible bedding planes, and holding small lenticular masses of specular ore, which conform in their strike and dip with this cleavage, and which seem to have no structural connection with the main deposits. They appear like dykes of ore, squeezed out of the parent mass, which we may suppose to

have been in a comparatively plastic state, when the folding took place; or they may have been small beds, contained originally in the chloritic schist, and brought to their present form and position by the same causes, which produce the cleavage in the schist. A comparison of these sections, showing effect of the folding on a large scale, with the figures (19 to 29, Vol. II.) representing the contorted lamina of the mixed ore of Republic mountain, will be found instructive. Indeed the same phenomena may be observed abundantly at the Lake Superior mine, and still better at the Cleveland knob.

Lake Superior mine sections E—E', and Fig. 9, may almost be said to represent a huge breccia.

The peculiar nature of the hanging wall of the Lake Superior mine deserves further notice. Instead of the quartzite, which we have hitherto found overlying all the deposits of rich ore, we have here a magnesian schist very similar to, if not identical with, that already mentioned as being associated with the ore, as will be seen by reference to the geological sections, and to Spec. 55, State Collection, App. B, Vol. II. These rocks are given, however, different colors on the maps. The hanging wall of pit No. 25, Section A—A', it will be observed, is made up of this schist and of layers of quartzite. Whether the Upper Quartzite is replaced by this schist, making it belong to XIV., or whether it is a member of the ore formation XIII., in which case XIV. would be wanting at this locality, I am not able to determine, but incline to the first opinion.

The hematite formation XII. is fully developed at this locality, producing an excellent ore which is extensively worked. The relation of this formation to the overlying and underlying rock is obscure, as has already been pointed out. This relation was very plain, it will be remembered, on the Saginaw-New England range.

The structural hypothesis by which I have attempted to connect the Lake Superior deposit with the Lake Angeline on the south, and Marquette, Cleveland and New York mines on the east, need not be further described here, but will be understood I think, by those interested in the question, from an examination of the following figure in connection with the maps.

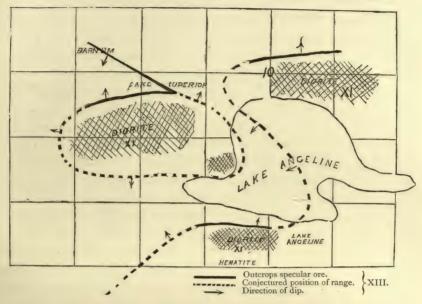


Fig. 10.—Sketch (part ideal) showing position of ore basins at Ishpeming.

New York, Cleveland and Marquette Mines.

The geological facts to be observed, the general structure, nature and extent of the workings of the New York mine, which is one of the most regular deposits in the district, are so plainly set forth on the accompanying Map, No. X., that but few words of description are necessary. It will be seen to be a monoclinal deposit, in every essential particular, like the Barnum, Champion and Spurr. interesting facts will be observed: 1st. The absence of formation XII.; the pure ore, with its associated chloritic schists, seems to occupy the whole space between the Upper Quartzite, XIV., and the diorite, XI. It may be here observed that, as a rule, the purest ores are found in the upper part of the ore formation, that is, nearest the Upper Quartzite; the New York mine presents an exception. 2d. The deposits on the north side of the railroad, worked by Pits No. 3 and 4, have a striking resemblance to the small deposits, Pits 16 to 21, of the Lake Superior mine, just described. The facts to be noted at the Collins location, just east, taken in connection with Pits 3 and 4 of the New York, point plainly towards the existence of a small independent trough, north of the Cleveland-New York

deposit. Explorations and mining operations so far, do not indicate the presence of a large amount of first-class ore here.

I made no special survey of the Cleveland mine, the fund at my disposal not permitting it; the main object of the survey in this direction being, to represent in detail a sufficient number of typical mines, to cover the various structural phenomena to be found in the district. The sketch of the Cleveland and Marquette mines, Plate II., from A. Heberlein's map, in connection with the New York mine (Map No. X.), will give a good general idea of this group. It will be seen, that the most northerly pit (Gents, No. 3) of the Cleveland mine, is a continuation of the New York deposit, having the same strike and dip. Gents pit is in one of the largest deposits of pure specular ore in the whole Lake Superior region. It dips south, forming the northerly edge of a narrow synclinal basin, which immediately comes to the surface again in the Swedes pit, where the ore has a northerly dip. These two pits produced in 1872 over 100,000 tons of ore. The ore basin widens and deepens to the west in a similar manner to the Lake Superior, and undoubtedly underlays the swamp, on which the village of Ishpeming is built. The connection of these deposits with those worked in the more southerly Cleveland and Marquette openings, has not received that attention which would enable me to express an opinion on the subject.

There can be little doubt, but that the Cleveland mine promises as well, if not better, for the future production of first-class specular ore, than any one of the older mines.

Jackson Mine and Negaunee Hematite Deposits.

No special survey was made of the Jackson mine; but the accompanying Plate (iii.), from O. Dresler's map and Atlas map of the iron Mines at Negaunee (No. V.) will make known the general structure of the mine, which is essentially similar to that of the Cleveland and Lake Superior. This mine, although it produces first-class specular ore, will be here considered in connection with the hematite deposits, because they are adjacent, and their geological structure can be most conveniently described together. The Jackson mine, so far as is known, is the extreme east end of the



North



rich ore basin formed by bed XIII. No workable deposit of ore of any kind has been found north and east from this locality, and the ores to the south are believed to belong to a lower horizon, and to be, on the whole, inferior in quality.

Looking back over the field we have now hastily surveyed, and assisted by the map of the Marquette iron-region, it will be seen that, while there are many minor irregularities, on the whole the ore basin gradually widens towards the west, from a mere point at the Jackson mine to a width of fully five miles at the west end of Michigamme lake, beyond which too little is known, to enable us to accurately define its limits. It follows, therefore, that all the Huronian rocks north, east and south from the Jackson mine, are below, or older than the ore formation (XIII.) and all the rocks to the westward and inside of the ore-basin are younger, hence above it.

The large amount of exploration work, done in the vicinity of Negaunee, in searching for hematite within the last few years, has aided greatly to develop the geological structure of that locality. But unfortunately, the money I had to expend here was more than exhausted, before this work began, so I have been enabled only in part to avail myself of it.

The facts observed are mostly recorded on the local map, mentioned above, and on the general map of the region. By reference to the former it will be seen, that a belt of country, about one mile wide, extending southeast from the Jackson mine, is dotted over quite irregularly with hematite workings, which are mostly on lands leased from Edward Breitung, as is explained in a note on the map. These mines produce dark-colored earthy hematite, containing metallic manganese, often up to an average of 5 per cent., varying considerably in the amount of metallic iron, but on the whole averaging lower, than the hematite ores heretofore mentioned, as will be seen from the chapter on analyses. I believe these ores all belong to one formation, No. X., in which, up to this time, no merchantable ores, except the Lake Angeline hematite, have been mentioned as occurring; it is at least certain, that they are older than formation XII., which embraces the Lake Superior and Winthrop deposits.

The geological sections A—A' through the Himrod and Green Bay mines, and B—B' through the Jackson Company's new hematite and old specular ore workings, fully illustrate the hypothesis of

structure adopted. It will be seen, that the ore is contained between two beds of diorite, IX. below and XI. above, and that there is associated with the ore, chloritic schists and various ferruginous schists and flag ore. These last-named rocks, it will be remembered, made up this entire formation in the Michigamme district, where hematites are wanting, as are magnetic ores in the district we are describing. Underlying the lower diorite mentioned, is a clay slate, which is in turn underlaid by a gray quartzite, to be seen outcropping near the centre of the north half of Sec. 8, and represented in Sec. A-A' under the number VIII. This is undoubtedly the same quartzite to be seen in the railway cut near the northwest end of Goose lake, where it is overlayed by a soft schist. See formations VIII. and XI., Geological Sec. No. 1, Map III. The clay slate on south shore and near west end of Teal lake, and exposed in railroad cut one mile east of Negaunee, is also believed to be of the same age.

The lithological character of the several formations, mentioned above, will be better understood by an examination of the following specimens of the State Collection: No. 21 quartzite from VIII.; No. 20 is a clay slate also from VIII.; No. 31 is from diorite IX.; Nos. 24 and 25 are hematite ores from formation X.; No. 26 is a specimen of ferruginous silicious schist from the Foster mine, which is also regarded as belonging to the same formation (X.); Specimen 28, from the same formation, is a magnetic, chloritic, silicious schist.

Referring again to Map No. V., it will be observed, that the Jackson Company's hematite workings, the McComber, Maas and Lonstorf's most northwesterly opening, the Rolling Mill, Himrod, Spurr and Calhoun, and Iron Cliff Co.'s Sec. 18 mines, are all in a rude curve, skirting the great development of diorite, which seems to limit these deposits on the southwest, and under which they all dip. The remaining openings are mostly contained in a narrow belt, which extends east-southeast from the Grand Central, diverging from the other range, which curves to the south. The diorite ridge which runs through the centre of the latter range is apparently a synclinal ridge underlaid by ore, which should therefore dip towards it from all directions, as is the fact so far as known. Undulations in the bed now unknown, may very likely bring the ore to the surface at several other points.

There can be no doubt of the great extent of this ore; it cer-

tainly can be on the average more cheaply mined and shipped than any other ore in the region, except perhaps the hematites of the Taylor and S. C. Smith mines. Location at the junction of two railroads, and contiguity to a prosperous village, are additional advantages, which will go a long way towards offsetting the disadvantages of lower percentage. The presence of several per cent. of manganese in this ore helps its working in the furnace, rendering it a desirable mixture. The McComber mine was first opened, and its ore is well and favorably known to many furnacemen. My analyses indicate, that this is a richer ore than the other mines of this group, but this cannot be established without further developments, as work has but just begun at most of them.

The **Teal Lake** ore deposit belongs to the same formation, as may be seen by an inspection of the map and sections. I have not been able, however, to find any good hematite in the old exploration pits, now nearly filled; a lean flag ore is very abundant.

The **Foster mine**, near southwest corner of Sec. 23, T. 47, R. 27, is another hematite deposit belonging to formation X. It has produced a considerable amount of hematite ore of medium grade, which contains no manganese; the deposits, or rather pockets, are pre-eminently irregular in form and uncertain in extent. The geological position of the Foster range is shown on Map No. III. and accompanying sections.

The Cascade Range.—The deposits on this range are the only ones now wrought, which remain to be described in the Marquette region. Like nearly every other described in this report, this ore was known to the United States linear surveyors, and afterwards examined and commented upon in considerable detail by Foster and Whitney. The range extends east and west through the south part of T. 47, R. 26. See Map III. The locality known as the Gilmore mine, at 1/4 post between sections 23 and 26, is the most easterly point at which ore has been seen in quantity. This, it will be observed, is about three and one-half miles east, and two miles south, of the Negaunee hematite mines. The range has been traced west by south from this place for five miles, or to a point just four and one-half miles south of the Jackson mines. This country has recently been opened up by a branch of the C. and N. W. road, which closely follows the ore range. The principal open-

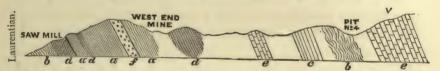
ings have been made by the Cascade, Pittsburg and Lake Superior, Carr and Gribben Iron companies, who shipped an aggregate, in 1872, of over 40,000 tons, nearly all of which was by the firstnamed company and its lessees. The last two named companies -Carr and Gribben—have done too little work, to enable us to speak with much certainty about their deposits. (See tables, on Sheets XII. and XIII., Atlas.) By reference to the chapter on analyses, which is quite full regarding these ores, it will be seen that they have, on the average, less metallic iron and more silica, than the standard hard ores of the district. The West-End mine, however, worked by the Cascade company, and which produced last year about one-third of their product, appears to be an exception to the above rule, and to rank nearly with the first-class specular ores; certainly considerable amount of high grade ore was taken from this pit last year, but whether it was kept separate from the leaner varieties in the shipments I do not know. The ore which largely prevails is a silicious or quartzose, or jaspery (practically these words have the same import) red oxide, having a characteristic coarse, slaty, or flaggy structure; hence the name by which they are known throughout this report. They correspond nearly in composition, although not in their appearance and geological position, with the second-class ores of the old mines, as the analyses referred to prove. See Iron Ores, Chap. III. Some varieties closely resemble, if they are not identical with, certain varieties of the high grade ores; but as a rule they are lighter in weight, duller in color and lustre, are harder under the knife, and pre-eminently flaggy or slaty in structure. I have not been able to obtain a statement of the working of these ores in the furnace. Further information regarding their lithological character may be obtained from descriptions of Specimens 5 and 6 of the State Collection, App. B, Vol. II.; the latter is the beautiful "Bird's-eye" slate ore from the Bagaley and Wilcox pit. Specimen 7 is from the diorite bed, which overlays the West-end mine, and is interesting from its resemblance to granite in outcrop. *

The structural position which these ores seem to me to occupy is shown on geological section No. 2 of Map No. III. They are near the Laurentian, and the whole series is overlaid by a talcy quartzite, which I believe to be the equivalent of No. V. of the Re-

^{*} Mr. Julien has determined the feldspar in this rare variety to be orthoclase.

public mountain series, and to be the same bed, which outcrops so conspicuously on the north side of Teal lake, and is calcareous at the Morgan furnace and at the Chocolate flux quarry, where it strikes the shore of Lake Superior. This rock varies more widely in its lithological character, than any other in the region, as will be pointed out elsewhere. If this hypothesis is correct, it will follow, that these ores are the equivalents of the Michigan and Magnetic ores of the Michigamme district, and are older than any iron bed made out in the Republic mountain series. The fact, that no iron in quantity has been found north of Teal and Deer lakes under quartzite V., where we should expect to find the opposite cropping of the Cascade series, is to be regarded in considering this question. The shortness of this range, which appears to terminate abruptly to the west, has not been found far east, and has altogether a local and isolated character, is significant. A hasty examination will satisfy any one that the quantity of ore in these deposits is very great, and that it is very favorably situated for mining and transporting. The accompanying north and south sec-

Fig. 11.—Geological section across Cascade range, looking west.—Part Ideal.



a. Flag ore or silicious hematite schist, in places quite rich. b. Banded jasper and specular ore with flag ore. c. Hematite rock or hematitic silicious schist. d. Diorite and dioritic schist. e. Quartzite. f. Conglomeritic and brecciated quartzite.

tion represents the different rocks to be seen outcropping on this range, projected on one plane. No attempt has been made to group them under formations I. to IV., to which they are supposed to belong. The general section No. 2, Map III., which has been mentioned, should be examined in connection with this sketch.

The Iron Mountain; Ogden and Tilden mines, not now worked, produced flag ores similar to those of the Cascade range, but not so rich on the average. These deposits belong, as will be seen by Map No. III., to formation X.; the Iron Mountain and Tilden mines being in opposite croppings of the same basin. The Foster mine, as has been observed, is also in the same formation, being overlaid and underlaid by flag ores. The Negau-

nee hematite and Teal lake ores being also in X., make that formation remarkably fruitful in the quantity and variety of ore, which it contains; but it does not, so far as known, hold the high grade specular ores in quantity.

Lower Quartzite, embracing Marble and Novaculite.

A brief consideration of the question of materials for furnace flux may come within the limits determined for this report. The subject, so far as the Silurian limestones are concerned, has been fully considered by Dr. Rominger, in Part III., who gives many analyses. The Menominee marbles will be mentioned in Chapter V. on that region. No calcareous, or other rock suitable for flux, has yet been found in the Laurentian system of the Upper Peninsula, although in Canada large beds of marble occur in this oldest series. It remains only for us to consider the silicious variegated marbles, found in the eastern part of the Marquette region, none having been worked west of Goose lake, which happens to mark the most easterly show of iron. The purest stone is found at the Morgan furnace, seven miles west of Marquette, where a heavy east and west bed of silicious marble, with vertical dip, and having associated with it clay slates, is prominently exposed. The prevailing colors are light-gray and pink. Specimens 11 and 12, State Collection, are from this locality; and Specimen 70, from the Gorge, represents the chloritic schist, which underlies the marble on the north.

The Chocolate Flux quarry on the shore of Lake Superior, three miles south of Marquette, is another locality, from which a small amount of furnace flux has been obtained. But the admixture of quartzose matter is here so great, that its use has been abandoned. Specimens 9 and 10, State Collection, represent the so-called "marble" and slate from this locality. It and the associated rocks are fully described in the extract from Dr. Houghton's unpublished notes, given in Appendix E, to which a sketch is appended. Mr. Julien examined a full suite of specimens from this locality, which are described in App. A, Vol. II., Nos. 106 to 113. No other marble locality possesses sufficient interest, to warrant mention, although flux has been quarried at several points near

Goose lake. It has been mentioned that the *novaculite* quarry, just east of Teal lake, from which whetstones were taken more than twenty years ago, is in the same formation. These stones are not now worked. See Specimen 13, State Collection, App. B, Vol. II.

During the past season several car-loads of quartzite were quarried in the same vicinity, and used as lining for Bessemer steel converters, at Capt. E. B. Ward's works, for which purpose it answered well.

The various marbles, slates, and quartzose rocks described above, are all believed to belong to one and the same formation, the Lower Quartzite (No. V.), which, it will be remembered, underlies the Republic mountain series, and overlies the Cascade series. This formation is one of the most interesting, geologically, in the Marquette region, and is worthy of a far more careful study than I have been able to give it. Specimens 8 to 13, inclusive, State Collection, App. B, Vol. II., represent several varieties of rock from this formation; as many more varieties could easily be procured, including some very fair specimens of iron ore from south and east of Goose lake.

A brief description, in addition to what has already been given, of the great geological basin formed by this quartzite, which embraces within its folds the great mass of the Huronian rocks, and nineteen-twentieths of all the ore, will possess interest. Like the ore horizon XIII., which we saw came to a point at the Jackson mine, and widened to the west, so the opposite croppings of this quartzite converge to the east and come together at the Chocolate Flux quarry, already described. From this starting-point the south rim of the basin bears away towards Goose lake, where some minor folds and low dips make it the surface rock for a large area northeast of the lake. From the south end of the Lake west, the formation has a prevailing talcky character, often argillaceous and sometimes conglomeritic; it has a great thickness and strikes west by south. West of the Cascade it seems to assume more the character of a chloritic gneiss and protogine, or at least a well-defined bed of protogine rock occupies the position in which we would expect to find the quartzite. See Map No. III. and sections.

The northerly rim, starting also from the Chocolate quarry, maintains a nearly due west course, crossing the railroad at the

Morgan furnace (where it holds the maximum amount of lime), forms the barrier rock in the Carp at the Old Jackson forge, passes north of Teal lake and south of Deer lake, occasionally at various points further west, and last, so far as I know, north of the Spurr mountain, nearly 40 miles west of Lake Superior.

3. ESCANABA DISTRICT.

The most southeasterly deposit in the Marquette region, and one which is entirely isolated from the localites already described, is the S. C. Smith Mine, producing soft hematite ore; it is located on Sects. 17, 18, and 20, T. 45, R. 25, and connected by a branch with the C. and N. W. railroad. It is but 42 miles from Escanaba, giving it a great advantage in distance over any mine, now shipping ore through that port. The geographical position is less remarkable than what might be called its geological isolation, for it appears to be in a small patch of Huronian rocks, in the midst of a great area of barren territory, underlaid by the Laurentian and Silurian systems. See Map III. The discovery of this deposit, a few years since, by Silas C. Smith, Esq., reflects great credit on his knowledge of the nature and distribution of ore deposits, and his perseverance in searching for them. Mr. Smith also first directed attention to the Republic mountain, which was, until within a few years, called by his name; he also made the first explorations in the Menominee region.

The few outcrops about the S. C. Smith mine, and the small amount of work done, when my examinations were made, enable me to say very little about its geological structure. The ore range runs northwest and southeast, approximately parallel with the Escanaba river, and cuts the southwest corner of Sect. 17. Contiguous on the northeast (whether underlying or overlying I am unable to say) is a bed of black clay-slate, in places identical with the so-called "plumbago" of the L'Anse range, which has been heretofore considered. Numerous fragments of a similar slate, probably belonging to the same formation, are found on the east side of Sec. 29. Laurentian granite is seen on both sides of the river, just east of this locality, away from which we have a right to assume the slate dips, rendering it probable, that the whole series dips

southwesterly, in which case the slate would form the foot-wall of the ore deposit, as on the L'Anse range. On Section 20, west of the river, a talcky schist, holding grains of quartz, was observed, but its relations with the other rocks were not determined.

Near the west ½ post of Section 20, and at other points in the vicinity, a flag-ore of good quality has been found; a specimen from one of the test-pits gave Mr. Britton 56 per cent. of metallic iron; whether there is any considerable amount of ore of this degree of richness has not, I think, been determined. Hand specimens of very fair specular ore could be found, but, as a whole, it seemed to me to be much more closely allied to the flag ores. Small boulders of this kind of ore had been found in this vicinity by C. E. Brotherton, some years ago.

Lapping over the upturned edges of the black slate on Sec. 17, and extending towards the east, is a horizontal Silurian limestone, which is, however, cut off by the river, beyond which numerous outcrops of granite and gneiss rear their heads above the flat sand plain. Silurian rocks are also seen on parts of Sec. 19, but west and northwest the country is all Laurentian, so far as I have been able to learn. South and east is a great plain, undoubtedly underlaid by Silurian rocks, but affording no outcrops, except near Little lake, where an isolated hill, apparently Huronian, rises out of the plain; I have not learned that any indications of iron have been found there.

I regret not having had the time and means to make a re-examination of this interesting and important district, after last season's extensive developments, and reluctantly present this imperfect sketch for want of fuller and more complete data.

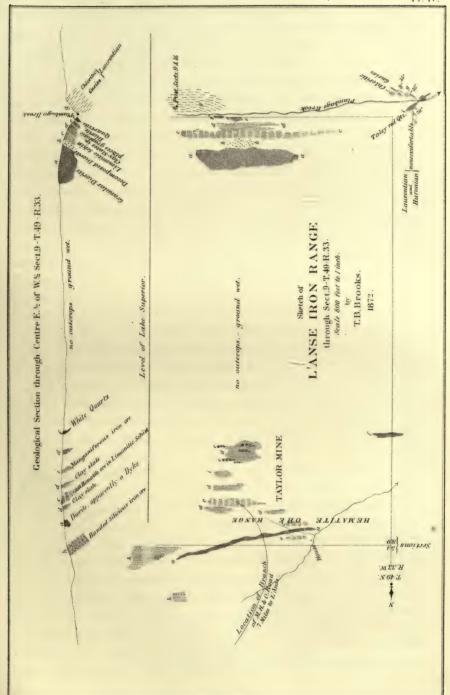
4. L'ANSE DISTRICT. (See Plate IV.)

The United States surveyors marked "iron ore" in two places on the line between Sects. 4 and 9, T. 49, R. 33. A quartzose or silicious brown and red ore can be seen outcropping, at several points in this vicinity. These facts early drew the attention of explorers to this district, and a considerable amount of land was bought from the government, for iron, as early as 1864. The fine harbor at the head of Keweenaw bay, only seven miles distant, and the abundance of excellent hard wood, tributary to this bay, have long

caused it to be regarded as one of the best points in the northwest, at which to make charcoal pig-iron, and establish other manufactories related thereto. The soil along the protected shores of Keweenaw bay is good, which led to the establishment of Indian missions there many years ago. A circle having the village of L'Anse as a centre, and a radius of 35 miles, would embrace the Washington, Edwards, Champion, Republic, Michigamme, Spurr Mountain, Magnetic and Taylor mines, with others less promising, together with all the copper mines in the Portage Lake district, the Hecla-Calumet mine, as also the principal mines in the Ontonagon district. It would also embrace all the roofing slate territory to which attention has already been directed, and an immense sandstone area, about which little is known. The amount of hard wood within the circle would be surpassed by very few equal areas on the Upper Peninsula, and the quantity of pine is large. A railroad running west, tapping the Ontonagon copper region, and continuing through the Gogebic and Montreal river mineral region, so as to connect with the Northern Pacific road, would, with existing roads and the excellent water communication, make the greater part of the area described easily accessible from L'Anse. If the advantages of the geographical position of L'Anse have not been here overstated, it is somewhat remarkable that the locality should have remained so long undeveloped. The want of railroad communication with the outside world was, undoubtedly, the main reason. What effect the very heavy grades, encountered within ten miles of the town, will have on the amount of ore which will be carried there from the Michigamme district, remains to be seen. The ore from the Taylor mine, and others that may be opened on the L'Anse range, can be put on board vessel at L'Anse at less cost for transportation, than any equally good ores with which I am acquainted, on the entire chain of the Great Lakes.

As has been before remarked, the L'Anse iron range, so far as made out, lies in the north part of T. 49, R. 33, the best ore being in Secs. 9, 8, 4, and 5; it has a general easterly and westerly trend, like nearly all of the iron ranges of the Upper Peninsula.

The **Taylor Mine**, the only point where the existence of a workable deposit has been demonstrated by actual exploration, is near the centre of the northeast ½ of northwest ½ Sec. 9, T. 49, R. 33.





This ore deposit is 950 feet above the surface of Lake Superior, and seven miles from L'Anse by railroad, built or building. The ground slopes gently to the west, affording an excellent opportunity for attacking the ore, which is covered by but a few feet of earth. The timber in the vicinity is first-rate hard wood.

The prevailing variety of ore at the Taylor mine is a soft hematite, similar in character to that of the Lake Superior and Winthrop mines. A number of analyses of average specimens, the results of which are given in full in Chapter X., varied from 44 to 57 per cent. metallic iron, with a remarkably small percentage of silica for an ore of this class. I see no reason to doubt but that a hematite can be mined here, which will yield an average of 55 per cent. of pig-metal in the furnace. Cross trenches and drifts show the deposit to have a maximum thickness 20 to 25 feet free from rock, and three or four times this thickness of such mixtures of ore and rock, as usually occur at hematite mines. The distance between the most easterly and westerly points at which ore has been found, is about 1,000 feet, but up to this time the explorations made have not demonstrated the deposit workable, as to quantity and quality, for more than about one-fourth of this distance. The oft-mentioned irregular pocket-like character of these deposits makes it difficult to predict, with any degree of certainty. regarding them, beyond what can be actually seen. But the heavy bed of hematitic rocks, which show a constant tendency by their decomposition to pass into ore, together with what has been actually developed by the workings, leaves no reasonable doubt but what there is here a large workable deposit of ore.

About 200 feet south of this ore deposit, and overlying it (the whole series dip south), is a bed of highly manganiferous iron ore, average specimens of which have yielded as much as 44 per cent. of the oxide of manganese; such ore must, of course, be comparatively poor in iron; this subject was considered under iron ores in Chapter III. The deposit is of uniform quality for a thickness of ten feet, and was penetrated by a shaft for the same distance. One per cent. of oxide of manganese was reported in some of the analyses of soft hematite mentioned above, showing the general dissemination of this substance, which seems to have its greatest concentration at the point we are describing. Whether this ore would possess value in the manufacture of metallic manganese, I am not

able to say, but its presence, undoubtedly, gives additional value to iron ores, in improving the quality of the metal produced, and causing the ore to work more easily in the furnace, besides especially adapting the metal for steel manufacture.

Several other "shows" of iron in this vicinity are worth mentioning. Near the south ½ post of Sec. 4, being on the north face of a high hill, is an extensive outcrop of several varieties of flag ore, more or less mixed with rock, in the vicinity of which considerable exploration work has been done. Some rich hand specimens of specular ore have been procured at this locality, but the great mass of the material to be seen is made up of layers of silicious ore, banded with quartzose material, the latter greatly predominating. The indications of hematite to be seen here are not promising. I see no reason why a flag ore yielding from 40 to 50 per cent., may not be sought for with reasonable chances of success. A similar ore was found several hundred feet farther north. The quantity of this mixed material existing in the S. ½ of S. ½ Sect. 4 is undoubtedly very great.

In the S. ½ of the N. E. ¼ of Sect. 8 are outcrops of hematitic rocks, which point towards the continuation of the Taylor mine series, making this a promising ground for exploration. Further west and southwest the ground falls off, the drift deepens, and no outcrops of any rock, so far as I know, are to be found, except in the immediate valley of Plumbago brook, where in Sect. 13, Town 49, R. 34, is an outcrop of argellite, which suggests a possibility of there being roofing-slate in the vicinity. Three miles west of the Taylor mine is the east edge of a treeless, sandy plain, which occupies nearly the whole of T. 49, R. 34, and extends into the townships south and west.

A similar desert country is passed through by the Peninsula Railway, commencing 7 miles from Negaunee. This latter, however, is underlaid chiefly by Silurian rocks, while the other is believed to be Huronian.

On the south side of Sect. 9, between Plumbago brook and the diorite ridge, which extends easterly and westerly more than one-half way across T. 49, R. 33, is a range of hematitic rock, similar to that at the Taylor mine, but which is not so promising for ore, so far as explorations have revealed. It has been traced for a distance of more than half a mile, and is the rock which immediately under-

lies the diorite, being itself in turn underlaid by clay-slate, the whole series dipping to the north, as will be seen on Plate IV.

Before dismissing the economic consideration of this district, it would be proper to notice the so-called "plumbago," found so abundantly in the north bank of Plumbago brook; but as this subject has been fully treated under the head of Carbonaceous Shale, Chap. III., it need not be further referred to here.

The **Huron bay slates** with associated rocks, may be regarded as belonging to the L'Anse series, although more than ten miles away in a northeasterly direction.

This district, which is now being explored for roofing-slate, affords indications of iron at several points, which I have not had such opportunity to examine, as would enable me to make any definite statement about them. So far as I can learn, those best acquainted in the district are not sanguine as to the existence of workable deposits of merchantable ore. At the end of Chap. I. will be found brief statements, regarding the slate companies now at work in this little-known district.

An inspection of Plate IV., in connection with what has been said, makes it necessary to add very little, regarding the structure of this range. The absence of outcrops through the central portion of Sec. 9, leaves the geological section quite incomplete. There can be little doubt, however, but that the quartzites, diorites, clay-slates and hematitic schists, so well exposed on the north side of Plumbago brook, where they dip north, are the equivalents of the Taylor mine series, which dip south, although the sequence is not exactly the same; and the diorite, so conspicuous on the south rim, is not exposed on the north side of the basin, unless the dyke-like mass of greenstone north of the Taylor mine represents it, which I do not think probable. The absence of outcrops also makes it impossible to determine whether there are any minor folds between the two croppings of the basin. If there are no such folds, then there is room for a considerable series of rocks above or younger, than those enumerated; and among them should occur, if it exists here at all, the rich hard ore of the Marquette district. It is assumed in this hypothesis, that the rocks to be seen are the equivalents of formations I. to X. of the Marquette series; this assumption is based chiefly on lithological grounds. Any rich hard ores found must be specular or red oxides, as there

is an entire absence of magnetic attraction in the L'Anse district. Magnetic ores have not as yet been found associated with soft hematites, so far as I am aware, in the Upper Peninsula.

The diorite immediately north of the Taylor mine has been mentioned as dyke-like. Whether it actually cuts the series of clay and ferruginous slates and schists at an acute angle, was not determined, but in places it certainly has that appearance. If it does so, it is the only case that has come under my observation, in which the Huronian diorites (often termed greenstones and traps) do not conform with the schistose and slaty strata, with which they are associated. This locality, in connection with others which show unmistakable dykes of magnesian schist cutting various rocks, is worth the study of the geologist, but is comparatively not of much importance to the explorer and miner. Mr. Julien, as will be seen by reference to App. A, Vol. II., Specs. 342 to 353, regards the L'Anse greenstones as a peculiar variety of diorite.

Another point of considerable interest, in connection with the diorites of this locality, is the *dioritic sand*, which forms the base of the great south bed, and separates it from the underlying hematitic schist on the south. This material is an angular, coarse, dark, greenish sand, and has evidently been produced by the disintegration of the rock, which is in places quite friable.

But by far the most interesting geological fact to be observed at this locality, and one, the importance of which can scarcely be overestimated in considering the grand subdivisions of the Azoic rocks, is the nonconformability of the Huronian, or iron-bearing series, with the older Laurentian, which can be observed in the gorge formed by Plumbago brook, about 400 feet southwest of the southwest corner of Sec. 9, T. 49, R. 33 (See Plate IV.). Here a talcky, red, quartzose rock, dipping at a low angle northwest, and which is unmistakably Huronian, is seen nearly in contact with a Laurentian chloritic gneiss, which dips at an angle of about 35° south-southwest. The same phenomena can be noted at a point near the Republic mountain (see page 126); and the nonconformability is further proven by the fact that the Laurentian generally abounds in dykes of granite and diorite, which are almost entirely absent from the Huronian.

CHAPTER V.

MENOMINEE IRON REGION.*

THE centre of this region is about 40 miles west by north from Escanaba, 50 miles south-west from Marquette, and 50 miles north from Menominee, as the bird flies. (See Map, No. II.) The area known to bear iron is embraced within a square of 16 miles, being portions of Towns 39, 40, 41 and 42, Ranges 28, 29, 30 and 31. This does not include the iron deposits west of the Paint river, nor the Michigamme mountain, owned by the Republic Iron Co., in Sect. 4, T. 43, R. 31.† The iron ores in the Menominee region occur in two approximately parallel E. and W. belts, each probably composed of two distinct ranges or horizons of ore; these belts are separated by a broad granite area, in which a little unpromising iron has been found on Sects. 10 and 15, T. 41, R. 29.

This granite area narrows towards the west, caused by the convergence of the iron belts, and has nearly the shape of a flat-iron. The region is drained by the Menominee river, which skirts its W. and S. sides, and by the Sturgeon, a branch of the Menominee, which winds through the eastern part of the iron-fields.

^{*} The facts contained in this chapter, as well as on Map No. IV. of Atlas, are largely from the Surveys and Explorations of Prof. R. Pumpelly and his assistant, Dr. H. Credner, made for the Portage Lake and Lake Superior Ship Canal Co. Prof. Pumpelly placed his private notes and sketches at my disposal, and added most valuable explanations. A valuable paper on this region is "The pre-Silurian formation of the Upper Peninsula of Michigan, in North America, by Dr. Herman Credner, Leipsic, illustrated by maps, diagrams and geological sections found in Plates VIII. to XII. (from the Journal of the German Geological Society)." Prof. Pumpelly and Dr. Credner are not in any way responsible for the hypothesis of structure here employed, nor for the views expressed as to the quality of the ores.

[†] A large amount of silicious iron ore occurs at this locality on the S.W. side of a high hill. Marble is found south and west, but in greatest abundance to the north, between Deer and Fence rivers, and on the upper waters of those streams. This district possesses much geological interest, and quite possibly economic importance, but means were not available for its examination.

I. SOUTH IRON BELT.

The South and, geologically, uppermost iron range of this Belt is probably the most regular and one of the most extensive iron deposits on the Upper Peninsula. The most easterly exposure of ore in this range is at the Breen mine on N. ½ of N. W. ¼ of Sec. 22, T. 39, R. 28. This location is 34 miles from Escanaba, and 45 miles from Menominee, in a bee line. The air-line distance from the elbow of the C. & N. W. R., now in operation, is 12½ miles.

Travelling from the Breen mine on a course N. 74° W., which is parallel with the general course of the river, we find on S. ½ of Sects. 11 and 10, N. ½ of Sect. 9, and S. ½ of Sect. 6, T. 39, R. 29, large natural exposures of ore, which have been still farther developed by recent explorations.

In the N. ½ of Sect. 2, T. 39, R. 30, are boulders of iron-ore, and near the S. ¼ post of Sec. 34, T. 40, R. 30, magnetic attractions, which indicate the presence of the iron range. Near the S. ¼ post of Sec. 30, T. 40, R. 30, is a large exposure of ore; thence, following a line of magnetic attraction which leads about W. by N., we find in the centre of the S. E. ¼ of Sec. 25, T. 40, R. 31, another exposure of ore, and a continuation of the local magnetic variations, westerly towards the Menominee river, two miles distant. A range of iron ore, corresponding with this and probably its continuation, has been made out in Wisconsin, between the Brulé and Pine Rivers. Here are no less than nine large exposures of ore, the extreme ones 16 miles apart, which lie in one straight, narrow belt.

Immediately N. of this iron range is a broad belt of impure marble, equally regular, of greater thickness, but which apparently widens towards the W.

North of this, in the vicinity of the Sturgeon River, on Secs. 7 and 8, T. 39, R. 28, and Sec. 12, T. 39, R. 29, are local magnetic attractions and iron boulders, which are believed to mark the position of another geologically lower iron range, although no outcrop has been seen in this vicinity; but near the centre of N. ½ of Sec. 20, T. 40, R. 30, just N. of Lake Antoine, is an outcrop of silicious ore.

Strong magnetic attractions can be observed near the S. W. cor. of Sec. 22, and iron boulders in Sec. 27, and also on north shore N. of Lake Fumée, in T. 40, R. 30.

These indications make certain the presence of a second iron range, although it cannot be demonstrated that these several shows belong to one horizon.

These two ranges, separated by the marble, constitute the South iron belt. North of and underlying both, is an immense bed of quartzite, which is well exposed at the falls of the Sturgeon river, Sec. 8, T. 39, R. 28; also on Sec. 1, T. 39, R. 29, and Sec. 28, T. 40, R. 29, and at the southwest ¼ of Sec. 23, T. 40, R. 30, as will be seen by the map. This quartzite, although believed to be geologically conformable with the ore formations, is not parallel with them, running more northwesterly, and dividing in T. 40, R. 30, into two and perhaps three ranges.

North of this quartzite, and underlying the whole series already described, are the Laurentian, granites, gneisses and schists, which make up the *granite area*, already referred to as probably being barren in workable deposits of ore, and which, therefore, our investigations do not embrace.

South of the south iron range, already described, is a bed of chloritic schist, well exposed on the south shore of Lake Hanbury, Sec. 15, T. 39, R. 29, and on the Sturgeon river in Sec. 13. Immediately south is a second quartzite, which is quite different in its character from the bed already described.

Next south is a broad belt of argillaceous slate, running parallel with the iron range, and exposed at several points in T. 39, Ranges 28 and 29. (See map.) South of this, and embracing portions of the Menominee river, is a broad well-defined belt of chloritic, hornblendic and dioritic rocks, running parallel with the iron range, the harder members of which form the barrier rocks of all the falls in this part of the Menominee, and probably those of Pine river in Wisconsin. This series are perfectly exposed at Sturgeon Falls, Sec. 27, T. 39, R. 29, and at the great and little Bequensenec Falls, and Sand Portage, in T. 39, R. 30.

2. NORTH IRON BELT.

The North iron belt or range has a course nearly due east and west, and is all embraced, so far as known, in the south tier of Secs. of T. 42, Ranges 28, 29, and 30. The most easterly dis-

covered exposure of ore, known as the Felch mountain, is in the N. 1/2 of Secs. 32 and 33, T. 42, R. 28, and is sixteen miles north and three miles west from the Breen mine, the position of which has been defined. Travelling due west, fragments of iron ore are found in N. E. 1/2 of Sec. 31, T. 42, R. 28; after which no absolute proof of the presence of iron is found (although it is probably continuous) until we reach Sec. 31, T. 42, R. 29, where, in the centre of the section, is an immense exposure of iron ore in an E. W. ridge, which can be traced westerly halfway across Sec. 36 of the next Township. The natural exposure of ore on Sec. 31 is larger than at any other point in the Menominee region, and the quality is as good, if not better, so far as can be judged by surface indications. Magnetic attractions and iron boulders, found farther west and southwest on this range, prove its extension in that direction. Whether the westerly course continues, or whether it curves to the southwest, as seems probable from the position of the lower quartzite and local magnetic attractions in the northwest part of T. 41, R. 30, has not been determined. The latter hypothesis is most in accordance with the known facts, although the southeast dip of the quartzite on Secs. 17 and 18, observed by Dr. Credner, is not explained. If this hypothesis is true, the iron range should cross the Menominee somewhere in Secs. 24 or 25, T. 41, R. 31, into Wisconsin. There can be little doubt but that the North and South belts belong to one geological horizon, hence somewhere come together.

The existence of two distinct iron ranges in the North belt, does not admit of so easy proof as in case of the South belt. The facts which point towards this are the following: About one-fourth of a mile north of the iron range, already described as existing on Sec. 36, T. 42, R. 30, is a bed of marble running east and west, parallel with the iron, on both sides of which are slight magnetic attractions. Prof. Pumpelly found, "about 80 paces south of this marble, an outcrop of strata made up of layers of quartz, magnetic iron and chlorite," probably of no economic value.

Again, in the E. ½ of Sec. 35, are two parallel lines of feeble magnetic attractions, several hundred feet apart, and to the north are some large, angular boulders of magnetic ore; similar smaller boulders are found between Secs. 33 and 28, still farther west.

South of the iron deposits on Secs. 31 and 36, is a bed of mar-

ble, somewhat similar to the one already described as underlying the south iron range of the South belt, and possibly the equivalent of it, as the two have the same relative geological position. Farther south, immediately adjacent to, and overlying the granitic rocks, is a heavy bed of quartzite, which is undoubtedly the equivalent of the lower quartzite, already described as forming the base of the South belt. This quartzite at the S. ¼ post of Sec. 31, T. 42, R. 29, is characterized by the presence of mica scales in the bedding planes, and might be denominated a micaceous quartz schist. It has considerable resemblance to the rock, associated with the Cannon ore in the Marquette region. This fact possesses considerable geological interest in connection with the relative age of the Felch mountain ore deposit, which, I think, belongs in this lower quartzite. See Chap. III., Group H. mica schists, and below.

The Huronian rocks in the N. ½ of Sec. 31, are covered with horizontal layers of Silurian sandstone, hence cannot be seen. North of the iron on Sec. 36, is the marble already mentioned, which is peculiar in being filled in places with crystals of kyanite, giving the gray weathered surface of the rock a rough jagged character, like a coarse rasp.

Just N. of the N. 1/2 post of Sec. 31, T. 42, R. 29, is an east and west range of gneiss rock, and still farther north a heavy bed of hornblendic schist. At numerous points east and west, through the centre of T. 42, Ranges 28, 29, and 30, are outcrops of similar hornblendic rocks, together with beds of mica schist and gneiss, traversed in places by dykes, and perhaps by beds of granite. This broad belt of hornblendic rocks is apparently represented in its westerly extension, where it crosses the Michigamme river, by the mica and chloritic schists and gneisses, so well exposed at the Falls of the Michigamme, Cedar Portage, Long Portage, Norway Portage and intermediate points in Towns. 41 and 42, R. 31. Similar rocks cross the Paint river, a few miles farther west. This series would correspond in their geological position, as they do partially in their lithological and topographical characteristics, to the hornblendic and chloritic series, already described as forming the southernmost formations of the South belt, and which there, as here, produce numerous waterfalls.

Near the centre of this hornblendic belt, in the north part of Secs. 21, 22, 23, and 24, T. 42, R. 29, is a line of comparatively feeble magnetic attractions, which seems to have no equivalent in the

South belt, unless it be in Sec. 28, T. 39, N. R. 18, E. Wisconsin; or in one of the beds of hornblendic rock at Little Bequensenec Falls, to be described hereafter, which contains many specks of sulphuret of iron and of magnetic ore.

This line of attractions, noticed in T. 42, R. 29, may represent the north edge of a basin, of which the North iron belt, already described, is the south edge; but I incline to the hypothesis, that it is an independent ferruginous range. No outcrop or boulder of iron has been seen upon it in Michigan, and it is doubtful if it is of any economic importance, although of much geological interest, as helping to elucidate the structure.

Returning to the most easterly exposure of iron on the North belt, the Felch mountain, we find a different and less complete sequence of rocks. Except some boulders about one mile west, no marble can here be seen. The Felch mountain ore rests immediately upon, and is bounded on the south by hornblendic, micaceous and gneissoid rocks, which are undoubtedly Laurentian, thus shutting out the marble and quartzite, already described as existing under the iron to the west. No indications, which would suggest the presence of a second iron range, can be found here. Within half a mile north the hornblendic schists are to be seen. At the N. ½ post of Sec. 31, about 1½ miles westerly, is a large exposure of quartzite, running east and west, and apparently dipping to the north, although the bedding is indistinct. This may be the equivalent of the north marble range, Sec. 36, T. 42, R. 30, for quartzites sometimes pass into marbles in the Marquette region.

The Felch mountain ore, so called, is in reality a dull red jasper-like quartzite, containing numerous thin lamina and minute gash veins of very pure specular ore. It has somewhat the appearance of the "mixed" or second class ore of the Marquette region (see Chap. III. A), differing in containing less iron, and in the fact, that the ore lamina have less continuity. Considerable amount of a similar rock can be seen on the Penokie iron range, Wisconsin. I have a two pound specimen of specular ore from the Felch mountain, which is as rich as any I ever saw. The deposit is somewhat magnetic, the east and west belt of magnetic influence having considerable breadth.

It is not at all improbable, that better ores may be found adjoining this on the north, or possibly still further north, in a geological position corresponding with the ore on Sec. 31, T. 42, R. 29.

In the south half of Sec. 36, T. 42, R. 29, about two miles west of the Felch mountain, Prof. Pumpelly and Dr. Credner observed a variety of the lower quartzite, the character of which is important in connection with the age of the Felch mountain deposit. It has been described as containing mica enough on its planes of stratification, to make it semi-schistose, is porous, and contains thin streaks of magnetic iron in crystals, with here and there cubes of iron pyrites.

The above facts lead me to accept the hypothesis already advanced, that the Felch mountain ore deposit is itself in the Lower Quartzite. If we suppose the mica contained in the quartzite exposed at S. ¼ post of Sec. 31, and in the S. part of Sec. 36, to be replaced entirely by specular ore, a Felch mountain ore would be the result. This hypothesis is supported by the fact, that the Cannon ore, Sec. 28, T. 47, R. 30, is a quartz schist, having specular ore in its bedding planes, and which in a short distance changes into mica. (See Chap. III., Mica schist.) It should be noted, however, that while the Cannon ore is micaceous, the Felch mountain is eminently granular. The Cannon, like the Felch deposit, is at the base of the Huronian series, resting immediately on the Laurentian.

It has already been mentioned that Silurian sandstone capped the iron bearing rocks on N ½ of Sec. 31, T. 42, R. 29; the same is true in places on Sections 34, 35, and 36, in same Township, as also in Sections 31, 32, and 33, in the Township east. Passing to the South belt, we find the sandstone covering the iron series in Section 25, T. 40, R. 31, in Secs. 30, 29, 23, and 36, T. 40, R. 30; also in Sections 9 and 10, T. 39, R. 29, and in Sec. 15, T. 39, R. 28 immediately north of the Breen mine, as well as at numerous other points, which it is not necessary to mention.*

Explorations eastward on the two iron belts of the Menominee region, reveal the presence of this sandstone and its accompanying overlying limestone (calciferous sand rock), in greater quantity, even to the point of entirely covering up the Huronian and Laurentian rocks, which is done, so far as known, from near the east side of the Menominee iron region, all the way to the Canadian line at the Sault Ste. Marie. Local magnetic attractions, discovered by

^{*} These irregular patches of sandstone are not represented on the maps.

United States surveyors at various points in this Silurian area, render it likely that the iron-bearing or Huronian rocks extend far to the eastward, connecting probably with the similar rocks of the north shore of Lake Huron, where they were first studied and named by the Canadian geologists. Pine explorers inform me, that they have observed dark-colored heavy rocks, which were somewhat magnetic, in the eastern portion of the Upper Peninsula. These may have been Huronian islands in the sea, in which the sandstones were laid down. This subject is discussed in Chap. II.

Like their equivalents in the Marquette region, the ore strata and accompanying rocks of the Menominee region usually conform in their strike with the general trend of the belts and ranges, and dip at high angles, thus presenting their upturned edges to the observer, and affording, where exposed, the best possible opportunity to observe the thickness of the beds and their mineral composition. But highly inclined strata, especially if they should be overturned, as is occasionally the case, are not favorable for making out the structure and sequence of the various beds. This question is farther complicated by the difficulty of distinguishing, in the case of the clay and chloritic slates, between the cleavage and bedding planes. The latter are sometimes very obscure, and have been confounded with the other, thus leading to erroneous results.

The geographical distribution of rocks in the Menominee region which has already been given in a general way, in connection with what has been said in Chapter II, concerning the structural relations of the Laurentian, Silurian and Huronian systems, leaves but little more to be said regarding the structure. The Laurentian area is the broad backbone of the great E. and W. anticlinal, on and against the north and south sides of which the iron series repose, dipping away from the axis; that is, the South belt south and the North belt north. This general structure, it will be observed, is similar to that presented by the Michigamme district on the south and the L'Anse-Huron bay districts on the north of the Marquette region, separated as they are by a great Laurentian anticlinal. is probable that the Laurentian area of the Menominee region may wedge out at a point just west of the Menominee river, in the same way as do the Laurentian rocks of the Marquette region in the west part of T. 49, R. 33. (See Map I.)

In order to bring out the structure more fully for the information of the explorer and miner, three geological sections will be given, two on the South and one on the North belt. Like most geological sections, they are to a certain extent ideal, but are intended to correctly present the facts, together with such inferences as seem to be warranted. I should note that Dr. Credner's corresponding sections differ considerably in the hypothetical parts from mine, as will be seen by reference to his paper already mentioned.

Geological Sections, Menominee Iron Region.

Section A.

Projecting the more important rock exposures of the eastern portion of the South belt on one plane, which may be taken at right angles with the strike of the rocks, that is, N. 16° E, through Sturgeon Falls, Sec. 27, T. 39, R. 29, the following series will be found (See Map No. IV.):—

At the falls of the Sturgeon, Sections 8 and 9, T. 39, R. 28, is a group of strata, which divide rocks unmistakably Laurentian on the N., from the lower Huronian quartzite on the S., and which Prof. Pumpelly and Dr. Credner regard as of Laurentian age, but which seems to me to admit of some doubt, as they conform with the bedding of both systems (all being conformable) and have lithological affinities with both.

Prof. Pumpelly describes them as follows, beginning with the uppermost strata:—

- 1. Talcose slates, soft, light-greenish, gray, with distinct ripplemarks.
- 2. Four beds of conglomerates, consisting of more or less rounded fragments of quartz, granite and gneiss, 15 to 30 feet wide. See Spec. 65, State Coll., App. B, Vol. II. This conglomerate has not been observed elsewhere, although a somewhat similar rock outcrops on Sec. 10, T. 42, R. 28.
- 3. Underlying the series are two beds of protogine gneiss, of reddish color, separated by a bed of chloritic schist; the upper one of the beds of protogine encloses a segregated vein, two feet wide, of a mixture of magnetic iron and sulphuret of iron, which does not promise to make a workable deposit.

North of this series, at the head of rapids on Sec. 9, T. 39, R. 28, unmistakable Laurentian rocks occur, but which appear to be conformable with the Huronian. The chief varieties found here as well as elsewhere in the Menominee region are,—a granite (in places porphyritic) syenite, mica-gneiss, with some mica-schist, hornblendic-gneiss and schists, chloritic and talcose gneiss, with some chloritic and talcose slates.

I. The lowest, geologically, and most northerly formation which is unmistakably Huronian in the South iron belt, is a *quartzite*, which outcrops conspicuously at the Falls of the Sturgeon river, Sec. 8, T. 39, R. 28 (not Sturgeon Falls), where it is not far from 1,000 feet thick, and rises to an elevation of over 200 feet above the river. It is usually light-gray, massive, compact, and often semi-vitreous, with indistinct bedding; has more the appearance of vein quartz than the Marquette quartzites. In places it shows ripple-marks with great distinctness; the weather has no appreciable effect on it.

This formation outcrops conspicuously, forming high ledges on Sec. 9, T. 39, R. 28, on Sec. 1, T. 39, R. 29 and Sec. 28, T. 40, R. 29. A quartzite, believed to be the equivalent of this, outcrops near the N. W. cor. Sec. 26, T. 40, R. 30. The Felch mountain iron deposit is also supposed to belong to this formation, as has already been explained.

II. A quartzose sandstone and conglomerate rock, which has a lithological character more allied to the Silurian than the Huronian, seems to overly this quartzite on the S., outcropping near the S. W. cor. of Sec. 2, T. 39, R. 29, and on the E. bank of Sturgeon river, on Sec. 8, T. 39, R. 28. But little is known about it, and its existence as a member of the iron series is not absolutely proven. From its soft, friable character it would more likely be found under swamps than on elevations.

The marble outcropping in Sections 24 and 25, T. 40, R. 30, would appear to occupy the same horizon. The same marble may exist on this geological section, but it has not been seen; the formation we are describing may be its equivalent.

III. The existence here of a range of slightly magnetic ore is indicated by angular boulders of lean ore in the valley of the Pine river, Sec. 12, T. 39, R. 29, and by magnetic attractions, Secs. 7 and 8, T. 39, R. 28. It does not, however, outcrop in this vicinity.

The hypothesis assumed for the structure of the South belt would make this ore the equivalent of the range known to exist north of Lakes Antoine and Fumeê, in T. 40, R. 30. It is possible, as will be seen hereafter, that this conjectured iron range may be the equivalent of the main iron deposit of the North belt.

IV. Crystalline *limestone or marble*.—This formation has an immense development in the South belt, far greater than in the other, its thickness being probably greater than that of the quartzite I. It is generally thinly bedded, and usually of a light-gray color, but is sometimes reddish, yellowish, or bluish.* The upper portion contains thin bands of slate, in which it resembles the marbles of the Marquette region, but differs from them in being freer from silica, less variegated in color, having fewer joints, as well as in being immensely greater in its extent, and more dolomitic. The Marquette marbles are indeed but calcareous beds in the Lower Quartzite (V.) of that series, there being no proper marble formation in the rocks of that region.

A piece of marble from near the Breen mine gave Dr. Rominger carbonate of lime, 61 per cent.; carbonate of magnesia, 34 per cent.; hydrated oxide of iron and manganese, I per cent.; and silicious matter, 0.25; which composition would make the rock rather a dolomite than a limestone. Specimen No. 66, State Collection, App. B, Vol. II., came from Sec. 11, T. 39, R. 29. Five specimens from this locality gave an average specific gravity of 2.81, approximately determined. Dr. Rominger gave attention to the value of this rock for building. (See his Report, Part III.) Large outcrops of marble occur on the south side of the Pine river on Secs. 11 and 12, T. 39, R. 29, and on the Sturgeon river, Secs. 17 and 18, T. 39, R. 28.

V. The principal *iron ore formation* of the South belt overlies, on the south side, the formation just described. It is made up chiefly, so far as is now known, of silicious specular slate ores, corresponding nearly with the so-called flag ores of the Marquette region. There is generally such admixture of magnetite as to produce moderate variations in the needle, but no evidence of the existence of a large body of magnetic ore. Specimen 68, State Col-

^{*} The weathered surface is often rough, from minute ridges, caused by the more silicious layers, which best resisted the weathering.

lection, App. B, Vol. II. is from Sec. 11, T. 39, R. 29. At the Breen mine some very good soft hematite occurs in the same formation, which promises to be in workable quantities. See Specimen 67, State Collection, App. B., Vol. II. This ore would probably be found elsewhere if sought for, but it never outcrops. A blackish, porous ore, hematitic in its character, containing 56 per cent. of iron and nearly 1 per cent. of manganese, was found in a pit at the ½ post between Sections 9 and 10, T. 39, R. 29, but its extent was not determined. Boulders of the same ore were seen in other places on the range.

The best exposures of the hard ores of this formation in the vicinity of the Sturgeon river, besides the Breen mine, are in Secs. 11, 10, 9 and 6, T. 39, R. 29. These ores will be described more fully, and analyses given hereafter.

VI. On the south shore of Lake Hanbury, which lies in Secs. 9, 10, 15 and 16, T. 39, R. 29, is an extensive outcrop of chloritic schist, the most easily splitting planes of which strike west by north, and dip south at a high angle. A similar rock, believed to be the same bed, can be seen on the Sturgeon river, near centre of Sec. 13, T. 39, R. 29. South of Lake Hanbury, 200 steps, is a rock partaking of a dioritic character, but which is probably a harder granular form of the same schist. Such rocks often graduate into each other in the Marquette region (Chap. III.). This schist may probably underlie Lake Hanbury and the swamps easterly and westerly from it.* It is represented on the section as following in its foldings formations VII. and VIII., described below. It is at least possible that this formation may be the same as the Menominee river diorites and chloritic schists, IX. and X., there brought to the surface by another series of more southern folds. But this hypothesis is not assumed in this discussion.

VII. Clay-Slate.—At 350 steps south of Lake Hanbury, on lines between Secs. 15 and 16, T. 39, R. 29, is a bluish and greenish gray slate, showing indistinct contorted bedding, with prevailing dip to *north*; the cleavage planes of which strike about north 70° west, and dip 80° to south. Veins of white quartz occur in

^{*} Since the above was written Professor Pumpelly has informed me that he observed a large outcrop of marble south of the iron formation III., in T. 40, R. 30, which will be described below under Section B. This marble may fill the apparent blank existing at Lake Hanbury.

these planes. At 550 steps south of the lake, a similar slate is found dipping *north* under the quartzite VIII., next to be described. It is believed that these two outcrops of slate, are the opposite sides of a synclinal trough, which holds the quartzite.

In the N. E. $\frac{1}{4}$ of Sec. 20, T. 39, R. 29, is an outcrop of talcose clay-slate. In Secs. 29 and 39, T. 39, R. 28, are several outcrops of dark colored, finely cleavable, but indistinctly bedded clay-slates. It is assumed that all these outcrops are parts of bed VII., which is folded into a synclinal and partial eroded anticlinal, as represented on section A of Map IV.

I am not in possession of sufficient facts to demonstrate the precise relations of these beds to each other, but the general fact is established by the northerly dips observed by me on Secs. 14, 15 and 16, that there are at least two folds between the iron range and the Menominee river, which probably reduces the estimated total thickness given in Dr. Credner's paper (18,000 feet), one-third. See page 175.

VIII. Associated with the clay-slates south of Lake Hanbury, is a bluish gray quartzite, which weathers into a brown, friable sandstone,* and in places reticulated with fine veins of quartz. At 550 steps south of Lake Hanbury, on line between Secs. 15 and 16, T. 39, R. 29, this quartzite is underlaid, as has been mentioned, by the clay-slate, VII., the division plane dipping plainly to the north at an angle of from 45° to 75°; the same rocks with the same northerly dip were observed farther east, on Secs. 15 and 14. This quartzite may be simply a local bed in the clay-slate formation, hence not entitled to a distinct number. The marked contortions both in the clay slate and quartzite are noticeable, and point unmistakably to the presence of a great fold. The cleavage planes maintain their east-west strike and southerly dip.

IX. This number is intended to include the soft magnesian schists (chloritic, talcose, and probably argillaceous) occurring so abundantly along the Menominee river, in the vicinity of the mouth of the Sturgeon, as well as at the several falls above. They will be more particularly described under geological section B.

^{* &}quot;Iron slate" is marked on the United States plats at this locality. The brown color of the quartzite has something the appearance of iron rust. The very feeble magnetic attractions existing along this range, indicate the presence of magnetite.

X. This formation is designed to embrace the granular dioritic rocks which form the barrier of the Sturgeon and other falls above, for 20 miles. It varies considerably in character, but on the whole bears a strong family resemblance to the granular diorites of the Marquette region. A peculiar gray variety, occurring at Sturgeon Falls, Sec. 27, T. 39, R. 29, is illustrated by Specimen No. 65, State Collection, App. B, Vol. II. This is the formation, it will be remembered, which in its supposed westerly prolongation into Wisconsin, produces the falls in the Pine river, and near them becomes iron-bearing. If the hornblendic schists mentioned as occurring in T. 42, are Huronian, they are probably the equivalents of this formation.

XI.—South of X., on or near the Menominee river, in south part of T. 39, R. 29, are several exposures of what appear to be magnesian schists and protogine, the structural relations of which to the rocks already described have not been made out. A rock similar to the protogine was observed in Sec. 13, T. 42, R. 30, and would there seem to have about the same relative position to the North belt that this has to the South belt.

Geological Section B runs northeast by north, across T. 40, R. 30, cutting Lake Antoine, and passes near the head of Great Bequensenec Falls. (See Map IV.)

I. Lower quartzite.—This formation appears far more conspicuously in this section than in A, owing to the double fold hypothetically introduced to cover the facts observed in the N. ½ of T. 40, R. 30. The large exposure of quartzite lying against the Laurentian, on Secs. I and 2, and the numerous angular boulders on Secs. 7 and 8, with the outcrop of quartzite near S. W. cor. of Sec. 23, taken in connection with the granite exposures on Secs. 4 and 9, lead one to the conclusion that one bed of quartzite, forming a synclinal basin under the Pine river and an eroded anticlinal to the south, best reconciles the facts observed. The lithological and topographical characteristics of this quartzite have already been given under A, and need not be repeated.

II. This formation was represented on A by friable sandstone and conglomerate, not observed near this section; the blue and pink *marble* outcropping near centre of Sec. 25, and the marble at the N. W. cor. of Sec. 24, are assumed to belong to one horizon

(as shown by map and section), which is supposed to immediately overlie the quartzite. There is no reason to believe that this formation has any great thickness.

III. The "shows" and "signs" of ore to which this number was attached on section A, have developed into certainty on this section, where, near the centre of the N. 1/2 of Sec. 20, T. 40, R. 30, a considerable outcrop of iron ore is seen in the bottom of a small ravine. It is a silicious, red oxide, resembling in its general character the great ore formation of section A. Its continuation eastward is made certain by the magnetic attractions on the south line of Sec. 22, by the iron boulders of N. E. 1/4 of Sec. 27, and on the north side of Lake Fumeê, on Sec. 26. Except the slight attractions noted by United States surveyors, at N. E. cor. of Sec. 30, T. 40, R. 29, there is no connecting link, so far as known, between this deposit and the indications of this bed on A. It is not proven that they are identical. Dr. Credner, as will be seen by reference to his paper, believes the ores on the north side of the lakes are the equivalents of those on the south, the two being connected by a synclinal fold.

IV. Crystalline *limestone or marble*. There are immense outcrops of this rock in the S. part of Secs. 34 and 35; large exposures on the S. shore of Lake Antoine; boulders on the W. side of Sec. 30, all in T. 40, R. 30, and a continuation of the boulders in Sec. 25, in the Township west. The apparent thickness is greater than was shown on A., which may be owing to a crumpling or short abrupt folding of this part of the formation; or, it may be due to an actual thickening of the formation to the westward.

Two outcrops referred to, deserve especial mention: that in the N. W. fractional ¼ of Sec. 29, contains beds of a sandy and almost conglomeritic rock, which is associated with thin beds of dark-gray argillaceous limestone. The outcrop on Sec. 35 is the largest marble outcrop in the Menominee region, it being over 1,200 feet wide. As the dip is at a high angle to the S., the perpendicular thickness of the bed cannot be less than 1,000 feet. The S. part of the outcrop shows bands of limestone alternating with thin seams of quartz.

V. The main iron formation is marked by an outcrop in the centre of S. E. ¼ of Sec. 25, T. 40, R. 31, and by another which forms the west end of a high ridge on line between Secs. 30 and 31, T. 40, R. 30, the two being connected by a line of magnetic influence.

Attractions also exist near the south ½ post of Sec. 34, T. 40, R. 30, and in the N. W. ½ of Sec. 2, T. 39, R. 30, are iron boulders. There is at present (October, 1872) no reason to believe that the ore in Towns 39 and 40, R. 30, is less in quantity, or differs in quality from that already described under the corresponding formation of geological section A.

VI., VII. and VIII. The hypothetical place of these formations on section B, is covered by deep drift—constituting the sandy terraces of the Menominee river. No outcrops of any kind can be seen on this belt of rocks, either in Ranges 30 or 31, except a large exposure of marble observed by Prof. Pumpelly, just south of the ¼ post, between Secs. 32 and 33, which corresponds in strike and dip and in general lithological character with marble formation IV. Reference to the map will show that this rock has no observed equivalent on A, where, if it exists at all, it should be found under Lake Hanbury.

I must confess that the existence of this marble, but lately made known to me, points to the existence of folds in the neighborhood of Lake Antoine, not suggested by my geological sections.

- IX., X. The chloritic, hornblendic, and dioritic rocks embraced under these two formations are well exposed at the Great and Little Bequensenec Falls, and at Sand Portage, all in T. 39, R. 30. These falls afford an unsurpassed opportunity to study this series, which was carefully done by Dr. Credner, who made out the following section at the upper fall from north to south:—
- a. Crystalline hornblendic rock, consisting of light to darkgreen hornblende in crystalline masses, white feldspar, a little chlorite and some quartz.
- b. Talcose rock, consisting only of fibrous talc, which forms a kind of soapstone in three heavy beds.
- c. Fissile talcose silicious slates, of a reddish color, with small crystals of orthoclase.
 - d. Soft talcose slates of light green color.
- e. Chloritic slates, dark green, with spots and layers of clayish red oxide of iron.
- f. Hornblendic rock, dark green, crystalline, coarse-grained to aphanitic, with specks of sulphuret of iron.

By the Little Bequensenec Falls the following series of strata is laid open, from north to south:—

- a. Talcose chloritic slates, with a great many segregations of quartz.
- b. Hornblendic rocks, with much dark-green chlorite, and many specks of sulphuret of iron and magnetic iron ore, 35 feet.
- c. Soft fibrous soapstone in two heavy beds, with some sulphuret of iron, 8 feet.
- d. Talcose slates, fissile, with many layers and segregations of white quartz and red limonite.
 - e. Chloritic slates, 10 feet.
 - f. Bed of hornblendic crystalline rock, 12 feet.
 - g. Chloritic slates with seams of iron pyrites, 30 feet.
 - h. Fibrous talcose slates, reddish, with bands of green color.
 - i. Chloritic slate.

Geological Section C. (North Belt). On line between Ranges 29 and 30, T. 42.

I. A quartzite, which is micaceous at S. ¼ post of Sec. 31, and in south part of Sec. 36, T. 42, R. 29, and ferruginous at the Felch mountain. The lithological character and stratigraphical position of this formation have been fully considered. Although it differs considerably in its character from the equivalent formation of the South belt, there can be little doubt but that it is the same.

North of this quartzite is a considerable breadth of low damp ground, with no outcrops.

II. Crystalline limestone or marble, of a quite pure snow-white, to reddish granular variety, outcrops immediately south of the iron on Sec. 31. In the southeast ¼ of Sec. 35, T. 42, R. 30, is an outcrop of marble presenting very distinct bedding planes, which dip to the north. These two outcrops define a range parallel with the quartzite, and probably belong to this bed, II. Another outcrop of marble near the centre of Sec. 35 cannot be reconciled as belonging to this formation, and there is some uncertainty as to whether it lies above or below the iron formation. If below, then it would have the same relative position to the iron as the outcrop first mentioned above. More facts are needed to establish the relations of these marbles. As will be seen by comparing sections C and B, it is assumed that the limestones marked II., on each, are equivalents of this bed.

III. The great iron-ore formation, which extends easterly and westerly across Sec. 31, half way across Sec. 36, and probably much farther each way, has already been partially described. This bed is apparently the equivalent of III. of the South belt, but it is certainly more extensive, and, so far as can be seen, contains better ore. If this hypothesis be correct, then the upper and main iron formation of the South belt has no representative in the North belt, unless it be indicated by the slight magnetic attractions already mentioned as having been observed in the north part of Sec. 36. The strongest indication of the continuance of this formation eastward is to be found, so far as known, just six miles due east, in the N. E. 1/2 of Sec. 31, T. 42, R. 28, where Prof. Pumpelly observed numerous large angular fragments of specular iron ore, associated with fragments of marble. This deposit should, on this hypothesis, pass just north of the Felch mountain, in its eastward prolongation.* The quartzite near the north 1/2 post of Sec. 31, T. 42, R. 28, would, on this hypothesis, be the equivalent of the before mentioned marble in Sec. 36, seven miles west.

IV. Crystalline *limestone or marble*, containing crystals of kyanite, outcrops about 300 steps south of the north ½ post of S. 36, T. 42, R. 30. Several outcrops of the same rock occur a short distance to the west, and a little south, indicating the probable existence of a large deposit of this rock. Except in the presence of the kyanite crystals, which gives to a weathered surface the rough character heretofore described, this rock has much the character of the marble, with corresponding number of geological sections A and B. Whether these marbles are equivalents is not proven, but it is assumed as being more in accordance with the facts than any other hypothesis.

V. An interesting fact in connection with the limestone outcrops on Sec. 36, just described, is the presence of a very noticeable magnetic attraction on both sides of the marble, or rather associated with it.

Prof. Pumpelly observed south of one of these outcrops of marble "strata made up of layers of quartz, magnetic iron and chlorite, containing garnets, and resembling some of the strata at Republic Mountain, Marquette region." These attractions

^{*} The blank space north of and above the iron formation III., on section C, is marked by no outcrops except Potsdam sandstone, which covers the Huronian rocks on Sec. 31, as has been already stated.

are probably due to this rock, which is certainly but a poor representative of the great upper iron bed of the South belt.

VI., VII., VIII. No other rock was observed on this section for several hundred paces; this space may or may not be filled by these formations, which, so far, have only been seen on geological section A. The numbers are introduced here, in order to carry along the hypothesis of structure which will best reconcile and present the observed facts.

IX., X. Just north of the north 1/2 post of Sec. 31, T. 42, R. 29, is a large outcrop of gneiss, with thin layers of granite, and adjoining this on the north is the most southerly observed outcrop of the great hornblendic and mica schist series, the geographical extent and general structure of which have been fully considered. Whether this series of schists are the equivalents of beds IX. and X., which occupy the immediate valley of the Menominee, cannot be estab-They have the same relative position to the iron ore, marble and quartzite series, and similarity in their lithological It must be admitted, however, that the lithological affinities of this series of rocks of the north belt are decidedly Laurentian rather than Huronian. The gneiss and granite outcrop, above described, may be almost regarded as a typical Laurentian rock in its appearance. If future investigations prove them to be Laurentian, a very troublesome structural problem would be presented here, as we would have Laurentian rocks conformably overlying beds, unmistakably Huronian. There seem to be fewer difficulties in supposing that the Huronian rocks of the Menominee region embrace lithological families not, so far, found represented in the equivalent series in Marquette region.

An important observation may be made here bearing on the variable thickness of the Huronian series, or else pointing unmistakably to tremendous folds in the rocks of the South iron belt,—it is this: the superficial breadth occupied by formations I. to VIII. inclusive, is nearly four times as great in the South belt as in the North. A portion of this difference may be accounted for by the thinning out of this series to the north; but the folds figured in geological section A, and possibly others not determined, would, I think, account for the greater part of this discrepancy.* There are no evidences of any folds in the corresponding series in the North belt.

A range of marble associated with quartzite, chloritic and talcose rock, and overlaid by a chloritic gneiss, with beds of chloritic schist and gneissoid conglomerate, the whole dipping at a high angle to the south, passes about five miles north of the North belt. These may represent the north side of the trough or basin, of which this iron belt is the south outcrop. No iron has, however, been found, as far as I know, on this range.

Along the Menominee river, where it crosses this broad schistose belt which lies north of the North belt, is a series of north and south dips, observable at the Cedar, Long, and Norway portages, which point unmistakably to intermediate folds in these rocks, whose thickness, therefore, may not be very great.

Nothing remains to be said regarding the Menominee iron region which is of practical importance to the explorer, miner, or capitalist, and which would properly come within the scope of this work, except a statement as to the *quality of the ore*. The quantity has already been described as great, and the chances to mine all that could be desired. The distances by rail from shipping port and grades are most favorable. If the ores are of first quality, this region has a future which will only be surpassed, if it is surpassed, by the Marquette region, now developed to that extent that its ores produce nearly one-fourth of all the iron made in the United States.

Unfortunately at this time the question of quality cannot be fully answered, for the simple reason that up to the date of my last visit, in October, 1872, comparatively little exploring had been done, and iron deposits very seldom expose naturally their best ores; these have to be found by digging. This subject is fully treated in Chap. VII.; but I will repeat here that ninety-nine hundredths, if not nine hundred and ninety-nine thousandths, of all the ore outcropping in the Marquette region (and there is an immense amount of it) is not merchantable, according to the present standard for shipments. Soft hematite ores never outcrop; therefore if pure high grade ores be abundant in the Menominee region, they might not yet have been found from the little work that has been done.

The facts observed by me are as follows, taking the several iron locations in succession:—Ist, The *Breen mine* on N. ½ of N. W. ¼ of Sec. 22, T. 39, R. 28, **South belt**. Three kinds of ore occur at this locality, the predominating variety (constituting perhaps

four-fifths of all exposed) being a lean, silicious, slaty or flaggy ore, resembling the Iron mountain and Teal lake ores of the Marquette region. It varies in quality from a ferruginous quartz schist, containing but a few per cent. of iron, up to masses as good, if not better, than the second-class or flag ores of the Marquette region, with occasional richer streaks. Careful mining and selecting would produce an ore of this kind that should yield say 45 per cent. in the furnace, but it would be apt to "work hard," from the large amount of silica, and produce a hard iron, suitable, perhaps, for rail-heads. (See Iron Ores, Chap. III.) What percentage of the whole mass would be of this degree of richness, practical mining only can determine; from what could be seen in October, 1872, I should say not exceeding one-third.

The next variety in abundance is a soft, earthy, dark-colored hematite, resembling in its general appearance the Negaunee hematite ore of the Marquette region. A sort of irregular pocket of this ore was found lying in the first described variety, appearing as if it may have been produced by a partial decomposition and disintegration of the flag ore,—that is a secondary form of it. This hematite pocket, so far developed by the shafts and trenches, is of sufficient size to work advantageously, but is divided through the centre by a bar of very silicious ore. Several "shows" of this ore were found in other places, but none were proven to be of workable extent. See Spec. 67, State Coll., App. B, Vol. II.

The third variety of ore is best in quality, but, so far as known, least in quantity. It can be seen near the mouth of a drift on the south side of the ridge next the swamp, where a bed two or three feet thick was passed through, flag ore being found to the north of it. This is a hard, more or less porous, bluish, heavy, red ore, of a hematitic character, and has considerable resemblance to the so-called Jackson "hard hematite." It would undoubtedly work well in the furnace, and would yield not less than 60 per cent. of metallic iron. There are reasons to suppose that there may be a workable bed of this ore on the property; but judging from what is to be seen at the drift above mentioned, it may be under wet ground.

On the whole, it may be said of the Breen location, that the great amount of ferruginous schist there developed, and the tendency shown by it to pass into soft hematite, render it very probable that a considerable quantity of workable ore of this kind

may exist. The absence of local magnetic attractions, and of boulders of rich hard ore, leads me to consider it doubtful whether any rich specular and magnetic ores, such as are now produced in the Marquette region, will be found here.

The ore range probably extends east and west, the entire length of the "80," or one-half mile, forming a ridge where the explorations have been made, from 20 to 30 feet high, bounded by a swamp on the south side. The whole iron series dip south, and are underlaid on the north by soft shaly magnesian and argillaceous rocks.

Sections 6, 9, 10 and 11, T. 39, R. 29. The ores on these sections form what appears to be a continuous deposit, and are so much alike in their general character that they can be more commonly and briefly described together. Except a few trenches dug by the Canal Co. on Secs. 9 and 11, and some test-pits sunk this season on Sec. 6, no work had been done on this range at the time of my last visit. Here, as at the Breen, the prevailing variety, in fact the only variety which I saw in quantity, was the silicious flaggy ore already described. The quantity of this ore is enormous, forming as it does the south face, and, perhaps, the great mass of a considerable ridge running west by north. The opportunity for attack by open cuts into the south face of this ridge is unsurpassed. Like the hard ores at the Breen, they vary greatly in richness,—from a quartz schist slightly impregnated with iron up to specimens, and even considerable masses which will yield 50 per cent., and occasionally a specimen that contains 60 per cent. of metallic iron. The prevailing variety, however, is represented by Specimen No. 68, App. B, Vol. II., from Sec. II., which contains from 25 to 45 per cent. of iron.

Dr. Credner reports having found, in "Cut D, on Sec. 11, 28½ feet of good fine-grained, steel-gray iron ore, with here and there a narrow streak of silicious ore, but in such a small proportion as not to spoil the good quality of the mass. The whole series gives a dark-red streak." Specimens designed to represent the average of this deposit gave Dr. C. F. Chandler 52 per cent. of iron. In another place he found a bed "6 feet thick, supposed to be very rich ore." I did not find these trenches (as afterwards appeared), although I designed to see all, and had with me two men, who helped to dig them. Dr. Credner further reports an aggregate of 139 feet

in thickness of "workable ore" on Sec. 11, but my own observations lead me to question this, unless the standard of furnace-yield be put considerably lower than at present. It is unwise, however, to predict at this time what thorough explorations may reveal.

The ore on Sec. 9 is very similar to that on 11, but on the whole (so far as can be seen) not so good: the same may be said of that on Sec. 6. Two smaller boulders of rich specular slate ore were found on the latter section, but no large ones. Occasional narrow seams of tolerably rich ore were found, one of them over one foot thick, but nothing that looked like a workable deposit. At the 1/2 post between Secs. 9 and 10, north of Lake Hanbury, are to be seen several boulders of a black, porous earthy ore resembling somewhat varieties of the Negaunee manganiferous hematites; the same ore was found in place in a pit near by, and a large boulder of it near the center of S. 1/2 of N. W. 1/4 of Sec. 6, and at other points. A hand specimen gave Mr. Jenney 56.44 per cent. of metallic iron, less than 16 per cent. of insoluble silicious matter. and nearly I per cent. of manganese. It is unlike the Breen mine hematite, and, in fact, unlike any Lake Superior ore I have seen. It is not improbable that workable deposits of it may exist, which being soft would not be likely to produce outcrops or boulders. I think it is well worth investigation. I have some reasons for supposing that this ore may be Silurian.

The next exposure of ore west of Sec. 6 on the south range is near the ½ post between Sections 30 and 31, T. 40, R. 30. This ore is softer and more slaty than those already described, although belonging to the flag ore family. It is apparently more argillaceous, and outcrops conspicuously in several places west of the ½ post, dipping at a high angle to the north, which would necessitate an overturned dip in order to harmonize with the hypothetical geological sections given on the map. The exposed bedding-planes are bright and specular, giving the ore the appearance of being richer than it really is. The ore exposed here may yield 45 per cent. in the furnace; see analysis No. 254, Chap. X.

From this locality we are led by a broad belt of very moderate magnetic attractions west by north for half a mile, to the iron ore exposed in the centre of S. E. ¼ of Sec. 25, T. 40, R. 31, where the Canal Company have done some trenching; the exposure

here is not great, the ore being in a small ravine on high ground. It is intermediate in character between the flag ores noticed, but most like the last. I followed the attractions about one-eighth of a mile west, to a point where the hill seemed to be capped with Silurian sandstone.

I have now mentioned in order, beginning at the east, all the main exposures of ore in the south range of the South belt, which has already been referred to as the most regular and one of the most extensive deposits of ore in the Lake Superior region; whether it is absolutely continuous for the 16 miles intervening between the extreme exposures, can only be determined by expensive explorations or actual mining.

Passing from the south to the north range of the South belt, we have but one exposure to consider, that near the centre of N. ½ of Sec. 20, T. 40, R. 30. This is in a small ravine, down which, to the south and toward Lake Antoine, a rivulet has its course in wet weather; the water has uncovered a narrow surface of flag ore similar to that seen on the south side of Sec. 30, but less slaty. Iron boulders are strewn along the ravine for over 100 feet. This ore is a red oxide, but holds enough magnetite to give it a moderate magnetic power.

Ten miles northerly across the granite region, from the last mentioned locality on Sec. 20, bring us to the main deposit of ore in the North belt—that on Secs. 31 and 36, of T. 42, and Ranges 29 and 30. The great extent of this deposit, and its favorable situation for mining, have already been commented on; it only remains to notice the quality of the ore. It is more granular and massive than the flag ore of the south range, and, as a whole, contains less silica and more metallic iron. The natural exposures of ore in the ledge are greater, no digging or uncovering at all being required to reach a great quantity of the ore. The best ore to be seen outcropping, is just southeast of the centre of Sec. 31: the top of the cliff is here about 100 feet above the low ground at its base on south side; and for about one-third of this height is a ledge of ore, from the foot of which the surface slopes rapidly to the low ground, affording the best possible opportunity for mining. This outcrop was carefully examined for a distance of several hundred feet in length, and from

the richest places to be found in it, 29 specimens of ore, of about one pound each, were collected, no two being broken from the same place. The specific gravity of these specimens was approximately determined on the ground, and was found to vary from 3.26 to 4.15, the mean of the 29 specimens being 3.71; this multiplied by 12, according to the empirical rule given under Explorations (Chap. VII.), gives 45 as the average percentage of the whole. An ore which actually analyzes 45 per cent. of metallic iron should yield say 47 ½ per cent. in the furnace, which is about what I consider this ledge of ore would work, if mined and sorted with ordinary care. Several ounces, chipped from five of the best hand specimens I could find, gave Dr. Wuth, of Pittsburg, 54.81 per cent. of metallic iron (See Analysis No. 98, Chap. X.). Separate analyses of ten hand specimens, selected from same locality by Prof. Pumpelly and Dr. Credner, gave Dr. Chandler from 49 to 64 per cent. of metallic iron, the average being 53.74 per cent. If this higher grade can be found in workable quantities (which is probable), then we should have a 55 per cent. ore, which, considering its granular and semiporous nature, and the fact of its being a red oxide, would indicate an ore not difficult to reduce, and one which would sell in the present market.

No boulders were observed in this vicinity which would indicate a richer ore than the above of the red oxide variety, and no magnetic attractions were observed which would suggest a workable deposit of magnetic ore, although all the ores of this region are slightly magnetic. As hematite ores do not outcrop, and as no explorations have been directed to finding such ores, nothing can be said regarding them. My impressions are that they will be found on Secs. 31, 32, or 36 of the North belt.

The Felch mountain ore was fully described when considering the lower quartzite. It is totally unlike either of the preceding varieties, and more closely resembles the "mixed ore" which accompanies the rich specular ores of the Marquette region. The laminæ of ore are very rich, analyzing from 63 to 67 per cent. of metallic iron; but the large admixture of quartzite (at least three quarters of the whole) would render it unmerchantable at present. It is by its constitution particularly well adapted to *stamping* and washing, and on account of its proximity to several rapids and falls in the Sturgeon river, is well situated to be worked in this way,

when the market drives miners to this means of production, as it will sooner or later.

3. PAINT RIVER DISTRICT.

Too little is known about the remote Paint river district, in Towns 42 and 43, Ranges 32 and 33, to enable me to give anything of interest regarding its geological structure. The Huronian rocks are extensively developed there, and contain deposits of hard hematite ore. I had the opportunity to examine only two localities, at the Paint River Falls, Sec. 20, T. 43, R. 32, and on Sec. 13, T. 42, R. 33. The ores are identical, and unlike any in the more easterly part of the Menominee region, in being richer in iron, freer from silica, and in containing more water. (See Analysis 68, Chap. X.)

Explorations now in progress will determine many of the unsettled questions regarding the ores of the Menominee region, especially of the South belt. I regret that I cannot embody their results in this Report, and thus give it a completeness that in the present state of my information is impossible.

CHAPTER VI.

LAKE GOGEBIC AND MONTREAL RIVER IRON RANGE.

An examination of this but little known iron-field was not contemplated in the original plan of the survey. But, having had occasion in the line of my profession to make some explorations there, a few of the general results obtained will be given, with a view of aiding future explorations, and of calling attention to a comparatively unexplored region. The probability of there being early railroad communication through this country, connecting the existing system of roads of the Upper Peninsula with the North Pacific, Minnesota and Wisconsin systems, now radiating from the west end of Lake Superior, attaches additional interest to this most western portion of the Upper Peninsula.

The facts observed and conclusions formed are the joint work of Prof. Raphael Pumpelly and myself, and have, so far as they bear on the stratigraphical relations of the four great systems of rocks, been in substance given to the public, in the American Journal of Science and Arts, Vol. III., June, 1872. Many rock specimens, gathered by us are minutely described by Mr. Julien, in App. A, Vol. II.

The iron range under consideration may be regarded as the eastern prolongation of the Penokie range of Wisconsin, as well as the western extension of the Marquette series, the whole being Huronian. The position of the range is tolerably well defined by magnetic observations and notes on the U. S. land office plats; on these we find mention of iron and magnetic attractions on Secs. 7 and 8, T. 47, N., R. 45, W., as also in Secs. 13 and 14 of the Town west. The belt of Huronian rocks, as made out by us, extends nearly east and west, through the north part of T. 47, Ranges 44, 45, 46 and 47, crossing the Montreal River in Secs. 16 and 21, of the last-named Township. Going east, the range was lost before it reached Lake Gogebic.

The geological boundaries of this range are fortunately of the most unmistakable nature, and render a detailed description of its position unnecessary. (See Map I.)

On the north is the high, broad, irregular ridge, or series of ridges, constituting the South Copper Range, the rocks of which are greenish and brownish, massive and amygdaloidal copper-bearing traps, their bedding being exceedingly obscure, with occasional beds of sandstone and an imperfect conglomerate. The strike of these rocks, so far as it could be made out, was east and west, with a dip to the north at a high angle, thus *conforming* with the Huronian rocks underneath.

Against and over the copper series on the north, abut the horizontally bedded lower Silurian sandstones, which are beautifully exposed on the west branch of the Ontonagon river, in Sec. 23, T. 46, R. 41. These sandstones form the surface rock, and occupy the broad belt between the two copper ranges from the region we are describing to Keweenaw bay, but taper to a point before reaching the Montreal river, in going west.

On the south of the iron-bearing rocks are a series of granites, chloritic gneisses and obscure schists, which, except the latter, are unmistakably Laurentian in their lithological character, and are non-conformably overlaid by the Huronian rocks. The general structural relations of the four great systems here enumerated are shown in the accompanying diagram. As the non-conformability

Fig. 12. Sketch showing Geological Section—looking west, between Lake Gogebic and Montreal River (in part ideal).



L. Laurentian rocks—gneiss, granite and schists, which are non-conformably overlaid by, H. Huronian—Clay slate, ferruginous and jasper schists, flag ores, quartzites and diorites, say 4,000 feet thick, which are conformably overlaid by, C. Copper-bearing rocks, chiefly greenish and brownish, massive and amygdaloidal traps, with occasional sandstones and conglomerate layers, which are non-conformably overlaid by, S. Lower Silurian sandstone, coarse quartz sandrock.

of the copper-bearing rocks and sandstones is doubted by some geologists, it should perhaps be stated that the actual contact was not seen. But the sandstones were observed lying horizontal, and affording not the slightest evidence of disturbance, within a few miles of highly-tilted copper rocks, which gave every evidence of having been elevated before the deposition of the sandstones. So far as my observation has extended, this rule is general; that is, no Lake Superior sandstone, which is unmistakably lower Silurian, has ever been found in any position other than nearly horizontal; and no rock which was unmistakably of the Copper series has been seen which was not considerably tilted. The fact that certain sandstones belonging to the copper series are very similar, if not lithologically identical with some of the lower Silurian sandstones, has helped to complicate this question. An interesting locality for study in this connection is the west fork of the Ontonagon river, just south of the Forest Copper Mine. I am not sure but that it affords an exception to the rule above stated, as at that point sandstones, apparently Silurian, dip south at an angle of 45°.

The best locality in which to study the character of the iron series in the West region, is on Black river and its tributaries, especially on the outlet of Sunday lake, T. 47, Ranges 45 and 46. Here will be found banded ferruginous jaspery schists, chloritic greenstones, brown ferruginous slates, black and gray banded silicious slates, silicious flag ores, several varieties of quartzites and clay slate. The whole series strike east and west, and dip north away from the granites and gneisses and under the copper rocks, at an angle of from 40 to 90°. Several varieties of the Huronian and Laurentian rocks of this vicinity have been examined by Mr. Julien, for descriptions of which see Appendix A, Vol. II. It will be observed from these descriptions that these rocks, although somewhat different from the Huronian series of the Marquette region, are still essentially the same; and I know of no good reason why merchantable ores may not be found amongst them. No ore, however, was found either in place, or in the form of boulders, which would pass for shipping ore in the Marquette region at this time. The absence of strong magnetic attractions renders it improbable that pure magnetic ores will be found here. The most encouraging indications observed pointed towards the existence of soft hematites, which may very likely be found of a quality and in quantity to pay for working. The best "show" observed was in the south ½ of the S. W. ¼, Sec. 18, T. 47, R. 46. It is on the north-easterly side of an east and west ridge, where there is a large exposure of highly ferruginous quartzite in places holding hand-specimens of hematite ore of fair quality. As this kind of ore never outcrops, on account of its soft, earthy character, and as we had no facilities for digging, nothing more definite was determined.

CHAPTER VII.

EXPLORATIONS (Prospecting for Ore).

1. How Failures have occurred, and how to avoid Them.

THE history of the development of a good many of our iron mining enterprises has been somewhat as follows:—The deposit is found, sometimes by accident, but often by systematic explorations made at the expense of corporations, firms, or individuals, by a class of men known as *explorers*; who are acquainted with woodcraft, are often miners, and who always have some knowledge of structural geology, the different varieties of ore, and the use of the miner's compass. A boulder of ore, red soil in the roots of a fallen tree, the variation of the magnetic needle, the proximity of rocks supposed to belong to the iron range, and often the outcrop of the ore itself, determines where digging shall be commenced.

If the indications are promising, before many marks are made the land is secured, if not already owned or controlled by those interested in the explorations. If government land, it is "entered" at the land office at \$1.25 per acre, or \$2.50 if within the limits of some railroad grant. If the land is "second-hand," already entered, it may be bought outright, or if the price be regarded as too high, a refusal is often taken with the privilege of exploring.

If the discovery is on the land of some railroad or mining company, it usually cannot be bought. In this case, all trace of the work done is often concealed, secrecy enjoined on all concerned, and the explorer lives in the vain hope that he may sometime have the opportunity to buy the land, an expectation in which he usually dies, as large corporations do not often sell iron deposits for small prices, if at all. Instead of this unwise course, explorers often sell their information to the companies owning the land, which they can usually do at a fair price. Our supposed exploring party having secured the land, begin to dig test-pits and trenches openly

and systematically. The solid ledge is usually soon found, which may prove to be some variety of iron ore, perhaps pure, but far more likely a "mixed ore" or lean flag ore, hence not merchantable.

Specimens (which I am sorry to say are apt to be the best that can be found) are sent in as *averages* of the deposit. Experts pronounce them shipping ore, and common talk asserts that So and So have a "good show" for a mine.

Soon the test-pits, trenches and drifts develop a workable width and length of what seems to the explorers to be merchantable ore. "Mixed with a little rock perhaps in places," but this occurs in most mines at the start. Experienced mining men visit the new deposit, examine it carefully, and assert honestly that "it looks better than did the Champion or Barnum locations when they first saw them."

The explorers select what they believe to be strictly average specimens of the ore (an impossible thing as will appear), which are sent to some distinguished chemist who reports, perhaps 65 per cent. of metallic iron, and only traces of sulphur and phosphorus, and expresses the opinion that the ore will work well in a blast furnace, and is identical with other well-known Lake Superior ores. This report, with the certificates of good practical mining men, and the opinion of some geologist who may have examined the locality, satisfies the owners that they have a workable deposit of "shipping ore."

Next in order, if it has not proceeded simultaneously with the above, is the organization of a company under the general mining law of Michigan,* which prescribes not to exceed 20,000 shares at \$25 per share, par value. The property above mentioned is put into the new company at a moderate price; some prominent man of character and means is found to take the presidency of the company, his friends, with others, being "let in" on the "ground floor," and the None-such Iron Co. is organized and at work.

Building up a location is the next thing in order. To this end a contract is usually let to some French Canadian to build a dozen log houses for miners' families, a company's store, barn and shop. For this purpose the contractor lays out fifteen different lines on which to put the buildings, being governed in each instance by the ease with which the logs can be got together. In clearing for the foundations it is usual for the Frenchman to find a new deposit of ore

^{*} App. I., Vol. II., contains an abstract of the Mining Laws of Michigan.

better than the one first found, to which a part of the mining force is at once transferred, the location of the buildings being changed so as to avoid the fragments which blasting has already begun to throw. The condition of affairs at the new location is at this period about as follows:—houses are going up rapidly, stripping is being pushed to the utmost, several "pairs" of Cornish men are sinking shafts or blasting off the "cap rock" so as to get at the ore. The contract for a first-class wagon road to connect with the State road has been let at \$2 per rod, and a party of engineers are at work locating a branch railroad to the mine, and it is confidently predicted that a considerable amount of ore will be shipped from the mine that season.

About this time the president of the company—an old iron man, who has made a fortune by smelting 40 per cent. ores with anthracite coal in Eastern Pennsylvania-and a part of the board of directors visit the mine. One of the directors is an eminent lawyer who helped to "place" the property, another is a stockbroker who had made a fortune in Wall Street, a third is a railroad king, and another a successful whisky distiller. None but the president knew anything of iron before they came into the company. He is of course amazed at the richness of the ore, and tells the captain in charge of the mine truthfully, that he is throwing away as good ore as he ever used in his Pennsylvania furnaces. All collect and examine numerous specimens, which are submitted to the president and captain for their judgment as to richness. Nothing less than 50 per cent. is found, and the average is much higher. The lawyer who has fine muscular sense and a consciousness of its possession, soon discovers that he can judge accurately of the percentage of iron by handling the pieces of ore, and speedily becomes an authority with the broker and distiller. Specimens are hefted which contain 59, 61, 621/2, 68, and finally one fine-grained fragment of steely ore, which, after careful manipulation in each hand, it is decided contains 75 per cent. of metallic iron. The captain unhesitatingly admits that to be richer than anything in the Jackson mine. Rock is found in several pits, but the captain explains that it is only greenstone which "caps" the ore, and proves by the magnetic needle which is "dead 90," that the ore is there. Being in a hurry he may not have faced the instrument exactly east and west.

Having spent one half-day in the examination of their property,.

and becoming satisfied that it is first-class and will prove a profitable investment for themselves and friends, the company leave, having first instructed their superintendent to bend all his energies to getting out ore, without reference to quality, cost, or future condition of the mine—though the whole is not, of course, directly expressed. On their way East, the president perhaps sells a thousand tons or more to some furnace man who is a stockholder in the new company, and telegraphs back to the superintendent to ship it at once.

The foregoing sketch contains the elements on which many Lake Superior iron mining enterprises have been organized, and at the start operated. It is needless to remark that many such undertakings result in utter failure. In the copper region the proportion of failures is far greater, and in oil, gold, and silver enterprises overwhelmingly so. The average human imagination becomes temporarily diseased when stimulated by the chances of possessing hidden mineral wealth. Iron, being the least valuable of the metals, has less of this influence than the others, but is not entirely free from it.

It may interest those who are disposed to identify themselves with Lake Superior iron mining enterprises (and I believe no equal investment has paid better in past time or promises better for the long future) to know the cause of failure in such enterprises. Classifying them carefully, I find that about two thirds of the disastrous enterprises were based on deposits of ore the *quality* of which was not merchantable: they were not rich enough in metallic iron. The extraordinary richness of Lake Superior ore is not generally known. I have reports from 40 furnace stacks in which these ores are smelted, which show that the average furnace yield of 250,000 tons of magnetic and specular ore for 1870 was 65 per cent.

The amount of high grade hard ore is so great that consumers can usually get all they require, and will not buy an inferior grade. For this reason experienced iron men from other regions have often been deceived; they had not a sufficient realization of this question of quality. Marquette ores—which were rich compared with what they were used to—could not be sold on account of their leanness. The soft hematite ores are not considered in this connection.

The remaining third of the failures have come from a lack of quantity, the quality of the ore being satisfactory. It follows, therefore, that the question of first importance in a new iron mining

enterprise is to know—First, the average percentage of metallic iron in the deposit. What will the ore, mined in the usual way, yield on the average when smelted in the blast furnace? Second, approximately or relatively, how much is there of it? The failure to answer these questions correctly at the start has caused the loss of over one million dollars in the Marquette region during the last ten years, and the business is still going on. Experience is an expensive school, but is always full; no sooner does one class graduate than a new crop of "freshmen" take their places.

I believe it is not impossible nor even difficult to ascertain, at a moderate cost, the average amount of metallic iron, in any given deposit, sufficiently near for all practical purposes, and whether there is enough ore to pay for working.

It is the business of the explorer to find ore deposits and to determine approximately their extent and richness, thereby avoiding such failures as have been described above. This subject will now be considered under the several following heads:—

2. PROSPECTING AND WOODCRAFT.

As considerable part of the iron exploration work now being carried on in the Lake Superior region involves camping out and a knowledge of woodcraft, some facts regarding this part of the business will not be amiss here, and are the more necessary because very little reliable information on this subject can be found in any book with which I am acquainted. There are no roads through large districts of country, which, in consequence, can only be reached by boats or walking; in either case a considerable part of the labor is packing, which means transporting everything on the backs of men. This mode of transportation costs about \$9 per ton per mile at the present time, which is twenty-seven times as much as it costs to move freight on wagon-roads; it is, therefore, important to carry only such articles as are needed. Many an exploration enterprise has practically failed because the chief energies of the party were expended in carrying supplies and material which were not needed, while necessary things were left behind. It is safe to say that two times out of three, even in the case of experienced explorers, supplies do not come out equal. The party will be out of pork and have an abundance of flour, or the converse; will travel in a leaky

canoe for the want of a little pitch, or be barefooted because they had no awl; or ragged for want of thread; or suffering for food, where there is plenty of fish and game, because the salt had failed; or have their supplies wet for want of a piece of oilcloth. I have been in all these straits.

Organization of the Party.—Take the ordinary case of searching for mineral or timber, when an explorer and two men constitute the party. As packing is the heavy work, it is indispensable that all hands understand it. An average packer will carry 70 to 80 pounds and his blankets, but loads of 50 to 65 pounds are more common; across portages men often carry 100 pounds, and sometimes a barrel of flour weighing 200 pounds; but the packer who carries 70 pounds and his blankets, 10 to 15 miles per day, on a trail, or 5 to 10 miles through ordinary woods, has earned the \$2.25 clear per day, which is the present average wages.

Next to packing, cooking is an indispensable qualification. No man is fit to go in the woods who cannot cook; and many a woodsman, with a frying-pan and two tin pails, will, over his campfire prepare a better cooked meal, and in less time, than can be produced in one-third of the kitchens of the country, with all the appliances that belong to modern housekeeping.

An ability to handle a canoe in rapid water is almost as indispensable as the others. Three men with a month's supplies will require a 16-foot canoe, which will weigh, when dry, about 125 pounds, and can easily be carried across a portage by one man; such a canoe will cost, in the Menominee waters, at this time, \$15 to \$30. The Bad water Indian village is the chief source of supply.

Next to packing, cooking, and canoeing, an ability to travel through the woods, and locate himself, by the United States Land Office plats, or maps made from them, aided by a pocket-compass, is essential. A man who possesses these qualifications is a woodsman, and has a calling which, if he is honest and intelligent, will be profitable in the Lake Superior region for a long time to come. If, in addition to these requirements, he is a judge of timber, and can keep simple accounts, write letters, and locate himself by the "40," then he is fit to lead a party, and become a "pine-looker," or "cruiser." If he add to this, a knowledge of the more common rocks and minerals, and an ability to make rough maps or

plans of ground, then he is an explorer. Such men can command from \$4 to \$6 per day clear, with full time, and often an interest in what they find besides; or if they choose to examine lands (either timber or mineral) on their own account, they can usually sell their "notes" at so much per acre, subject to re-examination; or some one may purchase the land, paying the explorer for his services in an undivided interest in them. Notes of pine lands now sell readily at from 50 to 75 cents per acre.

Supplies.—Pork, flour and tea embrace all that is absolutely essential in the way of supplies, though sugar, beans and dried fruit are usually added; rice, oatmeal or wheat grits are also generally carried, and a little hard bread is convenient, to which a few pounds of cheese may be supplemented. Pickled ham, especially in summer, may take the place of part of the pork, and smoked beef is sometimes used.

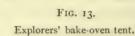
The following table of supplies has been prepared with considerable care from actual experience:—

Rations Required for Three Men. One Month.

Rations,		Amount in percentage of the flour.*
Flour, biscuit or crackers, rice, grits or oat-		
meal, but at least 3/4 self-raising flour (equal }	125	I.
to $1\frac{1}{3}$ lbs. per man, per day)		
Extra heavy clear mess pork about 3/4; pickled	l	
ham, say ¼	82	.650
Beans or peas	20	.160
Sugar (coffee A)	18	.140
Tea (good young Hyson)	3	.024
Dried apples	IO	.080
Cheese	. 4	.032
Salt	2	.016
Pepper	1/4	.002
Baking powder (Durkee's or Royal), if self-raising		
flour is not used	2 1/3	.020
Equal to $2\frac{85}{100}$ lbs. per man per day, or total	2663/4	lbs.

^{*} Supplies purchased in the proportions given in this column should come out even.

Equipment.—A shelter or bake-oven tent is preferable, although a closed A tent is often used in "fly time" (June and July): the former is more cheerful, healthier and warmer, because it lets the fire shine in. The style sketched will hold three men with supplies: it requires 12 yards of cotton drilling, 36 inches wide.





Two light explorer's axes, weighing with handles 23/4 lbs. for summer use and 5 lbs. for winter, each, are needed; if the exploration is for mineral, the backs or poles should be of steel. For three men a nest of two or three oval tin pails with covers, the largest holding 5 quarts, one frying-pan with socket handle, one 2 or 3 quart tin basin, one large spoon, one butcher or sheath knife, and a tin cup, plate, knife, fork and spoon to each man in the party, is all that is required.

If the party be large, a tin bake-oven will pay; it should be hinged so as to fold up. Canoes have already been mentioned: they are best for most kinds of river and lake service on account of their lightness, which makes them easy to portage, the ease with which they can be repaired with canvas and pitch (or resin and pork fat), and their suitableness for running rapids. But sometimes they

NOTE.—The stools shown in Fig. 13 do not belong to a camp outfit. They were introduced inadvertently by the engraver.

cannot be procured, and in low water are more liable to injury from rocks than are boats; a skiff 21/2 fathoms long, pointed at both ends, with flaring sides, and made of 1/2 inch boards, is a good substitute. Each man in the party should have a pocket compass. water-proof match box, and sheath knife, and there should be at least one leather and one tin map case in each party. Should the exploration be for minerals, a dip compass, and at least one exploring pick ought to be added. A small shovel will pay in such a party, but is seldom carried. A dial compass for use in traveling when there is local attraction, or in discovering the same, is often advantageous. I have found a small horse-shoe magnet and a pocket lens useful. Every party going in the woods should be supplied with the best maps that can be procured (Farmer's are the best I have seen), and always with exact tracings of the U. S. Land Office plats or maps of the Townships they propose visiting; these plats can be obtained at any U.S. Land Office and cost, if they show variations of the needle and geological notes, about \$2.25 each at Marquette. The following are the locations of all the U.S. Land Offices in Michigan, with names of officers.

U. S. Land Offices in Michigan.

Detroit ... Detroit ... F. Morley ... J. M. Farland.

East Saginaw .. East Saginaw .. Wm. R. Bates .. A. A. Day.

Ionia ... J. H. Kidd ... J. C. Jennings.

Traverse City ... Traverse City ... Morgan Bates ... Perry Hanna.

Marquette ... Marquette ... A. Campbell ... J. M. Wilkinson.

The explorer cannot too carefully study his maps; next to personal examination in the field they are his great original sources of information. The surveys of the Upper Peninsula, as is explained in Chapter I., were made with great care, and embrace topography, timber, soil and geology.

Under sundries which will be found useful in camp, may be mentioned: Soap and towels, thread and needles, buttons, awl, strong twine, some cotton cloth, a file to sharpen axes, a few wrought nails if a boat is used, some extra pairs of moose-skin moccasins (for summer), fish-lines and hooks, extra compass, resin or pitch, blank U. S. plats, and fly-nets or "fly-medicine," or

both in "fly-time." A large, stout, water-proof, tin match-box, extra note-book and pencils, paper and envelopes, are desirable. A short, light, single-barreled shot-gun, with bore large enough to chamber buck-shot, may be carried to advantage after the middle of August.

Mode of Working.—Mineral explorations, and especially those for iron, will only be considered under this head. The leading idea is, of course, to make a systematic and exhaustive examination of the surface for the mineral sought: to this end all outcrops of rock of whatever kind, and all boulders must be examined for some "sign" or "show" of mineral. As has been elsewhere remarked, the upturned roots of trees afford one of the best sources of information: the beds of rapid streams, which usually contain boulders and often expose the solid ledge, should be carefully examined. Any indication at all favorable should be followed up by digging. Next in importance to this kind of search is the use of the magnetic needle in discovering local attractions due to iron-ore; it is safe to assume that more than one-half the iron in the Lake Superior iron region is sufficiently magnetic to produce appreciable variations in an ordinary compass; and as magnetic ore will attract the needle at the same distance with equal strength when covered by rock, earth, air or water, this instrument is of great service to the explorer. Its use is fully considered elsewhere, as well as the geological principles applicable to this kind of work.

An explorer should make a careful sketch or map of each section examined, on a scale of 4 inches to 1 mile: on such a scale "a 40" is one inch square. On this should be marked in their proper places all streams, lakes, swamps, hills, etc., and all outcrops, with a name or sign indicating the kind of rock; colored pencils are convenient for delineating the different varieties of rocks. Opposite each such sketch should be a full written description of the rocks and minerals found, as well as notes on timber and soil.

The accompanying sketch (Fig. 14) of Sec. 29, T. 50, R. 30, from the note-book of the late A. M. Brotherton, a perfectly honest and thoroughly competent explorer, will serve as an illustration. To it is appended a map of the same section (Fig. 15), from the U. S. Surveys, which shows, valuable as these surveys are, and reliable, so far as the section lines go, they often are considerably in error in their representations of the interior of sections.

Fig. 14. Sec. 29, T. 50, R. 30. Explorer's Sketch.

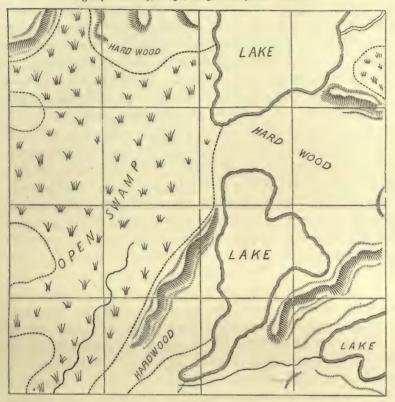
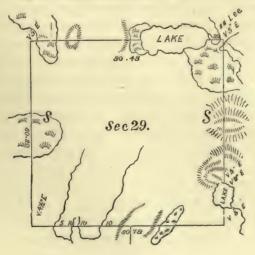


Fig. 15. Same Section, from U. S. Linear Surveys.



How to Recognize Iron Ores.—As a large majority of the explorers now employed are timber-hunters, they need not necessarily have a knowledge of minerals. I have, however, generally found these men more or less interested in rocks, and often very desirous of knowing how to determine the more common ores, so as to be able to note any they might find. To obtain a good knowledge, the study of a complete collection, or a residence at a mine, is indispensable. A few brief characteristics, only, will here be given, by which explorers may generally recognize iron ores in the woods.—First.—They are considerably heavier than any other rocks with which they are associated. Rich, magnetic, and specular ores, like those of the Marquette region, are nearly twice as heavy as the same bulk of the more common rocks, and fivesevenths as heavy as a piece of iron or steel of the same size. The soft hematites are much lighter, but are still appreciably heavier than the heaviest rock. As fine muscular sense and much practice are necessary in this business, the inexperienced explorer is advised, in every instance, to break pieces of rock of the same size as the supposed ore specimen, when, by lifting them together and changing from one hand to the other, the difference in weight will at once be felt if one of the specimens be iron ore. If the explorer is provided with a pair of balances, as is explained hereafter, he may determine, not only as to whether the substance is iron ore or not, but also approximately the percentage of metallic iron.—Second.—As to color, magnetic ores are black, and when pounded with the axe give a black powder, which will adhere to the axe or pick. Red specular ores are often bright and shining on their weathered surface, almost like polished steel; they give a red powder when pulverized, which does not adhere to the axe. Soft hematite ores are reddish and brownish in color, are generally porous, and often soft and earthy, in character; when pulverized they give a brownish and sometimes a yellowish powder, which does not adhere to iron or steel. These characteristics are possessed by none of the rocks of the Marquette region.—Third. Magnetic ores attract the needle of the compass strongly, often causing the north end to point south. Other ores and rocks do not attract it, but a little magnetic ore is often disseminated through rocks, especially greenstone, thereby producing more or less variation of the needle, which may not indicate valuable ores. The rock which is oftenest mistaken for iron ore is Hornblende, and the related Diorites or Greenstones. These rocks are heavy and dark colored, and often contain enough magnetite to give them some influence on the needle. Many an explorer has carried heavy pieces of this rock many miles through the woods, only to throw them away in disgust on meeting some one who had, perhaps, only so much knowledge of ores, as it is expected these few facts will impart. Some have persisted in their folly, and bought lands on which experienced iron explorers could only find hornblendic rock. This rock differs from the ore, which it most resembles, in being lighter, and in giving a light colored powder, which does not adhere to iron or steel, as well as in other less important particulars, as may be seen by comparing the two, which should be done.

The text relating to the magnetism of rocks and use of the needle in finding ore might properly have been inserted here as a division under Exploration, of which subject it forms properly a part. But the amount of material which had been prepared on that subject, and other reasons, determined me to place it in a distinct chapter (VIII.), which follows.

3. DIGGING FOR ORE.

The exploration work above described is superficial, and will not usually determine whether a certain piece of land contains workable deposits of ore or not. Such examinations are usually made to determine whether lands are worth buying at government price, or as preliminary to a more thorough exploration. When we consider that soft hematite ores never outcrop, and that pure hard ores rarely do, it is evident that something more than looking over the surface is necessary. The excavations of earth and rock required in an exhaustive exploration of a piece of land are mining operations, and will be considered in another chapter. Only a few points will be presented here which bear especially on work of this kind.

This work is simply sinking test-pits and shafts, and opening trenches (costeaning) and drifts to expose the solid ledge. It rarely happens that such work need be prosecuted into the solid ledge. As has been before remarked, if there be pure ore at the locality, i

will be almost certain to come to the surface of the ledge somewhere, and will there be found by digging through the earth. This may not always be the case, but it is safe to say that, as a rule, nine-tenths of all the money to be expended in exploring at any given locality, had best be expended in earth excavation.

There is a great deal of vague talk among miners and explorers of the Marquette region about "cap rock;" one would get the impression, from much that is said on this subject, that pure ores were always overlaid by rock. The fact is, however, that there are very few workable deposits of ore but what come to the surface, or, at least, connect with those that do. I should distrust any locality where "cap rocks" prevailed to any great extent; our ironore deposits are comparatively thin beds, which sit on edge, and come to the surface without wearing any "cap." * There are places, however, where the solid ledge has to be penetrated; when this is necessary, I think it had usually best be done by drilling. means of hand drills, holes can be sunk 22 feet, and by means of the appliances used in sinking oil-wells to any required depth; an experienced miner will have little difficulty in judging of the material passed through by the drill mud, and if there is any question as to richness, it can easily be settled by an approximate analysis which will be described hereafter. The diamond drill gives the most valuable results, and has been used to some extent in this region, and still more extensively in the Lake Champlain region.

Exploring excavations should always be done by contract; a large amount of "test-pitting" has been done in the Marquette region at seventy-five cents per foot in depth for a 4×6 shaft, the miner being paid only for such shafts as were "bottomed," *i. e.*, the solid ledge reached and uncovered, whatever the depth or difficulties. For drifts 3×6 which bared the ledge, \$1.50 was paid, and for open trenches a price proportionate to depth and width. Good miners can find themselves and make good wages at these prices in much of the ground in the Marquette region. Pits are sometimes sunk 35 feet, but the average depth does not exceed 12 feet. Mr. Colwell sunk 67 feet through sand on Section 24–47–28. Large

^{*} In the Menominee region true "cap rocks" are found in the horizontal sandstones which overlie some of the ore, see page 68.

boulders and water are the difficulties usually encountered; beyond 10 feet a windlass is necessary. A portable forge and mass of iron for an anvil are desirable, but picks can very well be heated in a hard-wood camp-fire and sharpened on a rock.

With regard to the significance of the material passed through, but one remark will be made; mixed drift, that is, large and small boulders, sand, clay, etc., is usually not very deep, 40 feet being the greatest depth I have observed, the average being less than 10 feet. Sand with no boulders is usually deeper and sometimes very deep.

4. QUALITY AND QUANTITY.—SAMPLING.—APPROXIMATE ANALYSIS.

Up to this point we have considered chiefly the question of finding ore regardless of quality and quantity. These are, after all, the vital questions, and their importance is rendered still more conspicuous by the statement, that there is at least twenty times as much ore in the Lake Superior region that is worthless from a lack of metallic iron, as there is of merchantable ore, according to the present standard for shipment; and further, it is easy to find specimens of pure ore in almost any body of worthless ore.

To determine approximately the average percentage of metallic iron, proceed as follows:—Open two or more trenches or drifts entirely across such portion of the ore formation as is regarded fit to work. In the region we are considering, the ores usually dip at a high angle, so that the edges of the beds or strata are exposed by such cross cuts; free the solid ledge from all earth and loose material; then, with a heavy hammer, break off small fragments every two inches across the entire bed, without reference to whether the pieces are ore or rock. Wash all of those pieces, break them all into fragments of the size of grains of wheat, mix them up thoroughly, send a tea-cupful to a reliable chemist, and his return will be the practical average of metallic iron in the whole bed from which the pieces came.

Of course, in mining, the ore is sorted, so that we should expect to get a somewhat better yield from working the ore, than that found as above, but it is not wise to count much on this. If, after trying, say half a dozen cross cuts in this way, an average yield of fifty per cent. (50%) of metallic iron is not found, the deposit is doubtful; if less than forty per cent. (40%) it is of no value in the present market, should the ore be specular or magnetic. Nineteen times out of twenty, such *mechanical averages*, when honestly taken, would show a yield of less than forty per cent. (40%.)

The plan above described is somewhat expensive and consumes time, which is an important element where one is maintaining an exploring party in the woods. A method which can be used on the ground, and which will give results, according to my experience, within a few per cent. of the above in the case of the silicious or quartzose hard ores (the kind usually found), is the following: Provide an ordinary swing balance which will sustain at least two pounds, and weights, the smallest of which should not exceed five grains, the whole costing less than \$5. Break up numerous hand specimens across the ore deposits as before, wash and dry them. Suspend each in turn by a fine fish-line and weigh it in the air, afterwards weigh it when immersed in water. Divide the weight in air by the difference between the weight in air and the weight in water. The quotient will be the specific gravity of the specimen, and will range from 3.17 for very lean ores to 5.13 for very rich compact ores. The specific gravity so obtained, multiplied by thirteen, if the ore be rich (i.e., above 55%), and by twelve, if the ore be lean (i.e., from 40 to 55%), will give the approximate percentage of metallic iron in the specimen.

The mean of a large number of determinations, made with specimens selected promiscuously from the deposit, will give a close approximation to the average percentage of metallic iron in the bed. According to my experience, the error will fall within five per cent., which is nearer the truth than any man can determine by simple inspection. It must be borne in mind that this purely empirical rule applies only to Lake Superior magnetic and specular ores, and only to such as contain some form of quartz as gangue, which is true of nearly all. The numbers 12 and 13, given above, as multipliers, were derived from numerous analyses and specific gravity determinations made by Dr. C. F. Chandler, of New York, and J. B. Britton, Esq., of Philadelphia. This plan is not offered as a substitute for chemical analysis, but I believe will often prove useful in the woods, and may sometimes help in deciding whether it is worth while to have an analysis made. As has been before

stated, unless the deposit is proven by analysis to contain an average of 50% of metallic iron, if specular or magnetic, and not less than 40%, if soft hematite, it is of doubtful value at the present time.

It would seem as if sufficient experience should enable us to judge of the quality of an ore at sight, or at least enable us to select an average specimen for analysis, without the laborious plan above described; but this is not the case, as is well known to those who have had experience in iron ores. It may be stated as an economic and psychological axiom, that no man, however honest or skilled, can, on his judgment alone, select an average specimen of ore from a deposit; he will always choose a richer specimen than the average. This would, of course, be very difficult from the technical stand-point, on account of the delicacy of muscle and skill of sight required; but the greater and insurmountable difficulty is in the human mind. We cannot help feeling that at a new opening there must be somewhere under our feet, or near by, better ore than we can see and the specimen selected is designed to be rather what we suppose, believe or hope the deposit to be, than an average of what we actually see and feel. I have numerous facts under this head, and am able to give an approximate mathematical expression to this form of human hopefulness. eleven instances the difference between the average by judgment, and the mechanical average obtained as above described, varied from 6 to 24 per cent., averaging II; the mechanical average being least in every instance; in each case I had reason to have confidence in the honesty and skill of the parties. It does not seem possible that such errors in average could exist, but they are constantly made, and will continue to be as long as iron ores and human minds are constituted on the present plan.

One of the fallacies which have caused innumerable disappointments in iron mining is the belief, almost universal, that ores grow richer in depth. This may be true of certain ores in some regions, but it is not true of the iron ores here being considered. They are just as good on top as in any part of their extent, and it may be stated as an invariable rule that if there be any good ore in a given deposit which is available for mining, it will somewhere come to the surface, except the earth covering in the Marquette region and the sandstone in the Menominee, which of course have to be removed when found. Hence a sufficient number of earth test pits,

trenches and drifts will usually find it, if it exists, without penetrating the rock. I do not mean to say that a deposit of ore may not grow thicker in depth; they often present this feature, and on the other hand sometimes grow thinner, and wedge out entirely. As has been before stated, by far the larger part of the money available for the exploration of any given locality should be spent in earth work.

While it is not difficult to determine with sufficient accuracy for all practical purposes the quality of a deposit of iron ore, as has been above shown, it is often impossible within a reasonable cost, to form so reliable a judgment as to the quantity. But a sufficient amount of judicious exploration will usually settle the all-important question as to whether the deposit is large enough to warrant development as a mine, future operations alone determining whether it will prove a great or small one. The method of doing this is obvious; many test-pits and trenches must be dug and drifts made where the earth is deep, the ledge of ore being thus laid bare in as many places as possible. No one engaged in making an exhaustive exploration of an iron-ore property should neglect the advantages of deep drill-holes; these can be sunk 20 feet with the ordinary drills employed at the mines. An inspection of the mud, and especially an analysis of an average of it, will prove of great value.

The annular diamond drill was introduced in 1870, at the Lake Superior mine, and gave very satisfactory results; the core gives almost as good an idea of the nature of the rock passed through as a shaft, and the cost is far less,—about \$5 per foot. But being propelled by a steam engine, it is only adapted to work near communications; it cannot be taken into the woods.

In the case of magnetic ores great assistance in determining the extent and position of the bed can be derived from a proper use of the magnetic needle, which subject is considered in the following chapter. Attention will, in this connection, only be directed to one important fact; worthless ores often attract the needle just as strongly as merchantable ones. Now, as there are many times more lean magnetic ores than rich, it follows that a variation or dip of the needle may not, probably does not, signify a workable deposit.

CHAPTER VIII.

MAGNETISM OF ROCKS, AND USE OF THE MAGNETIC NEEDLE IN EXPLORING FOR ORE.*

1. Elementary Principles.

A FEW of the elementary principles of the science of magnetism, made use of in the following investigations, will first be given.

Magnetite, or magnetic iron ore, contains, when pure, about 72 per cent. of iron and 28 per cent. of oxygen. The unmixed mineral is black, or blackish in mass and streak, has a specific gravity of 4.9 to 5.2, and hardness of 5.5 to 6.5, which is somewhat less than that of quartz; its crystals are usually octahedrous, and in the massive state it is often granular, and sometimes friable. Magnetite is one of the most abundant ores of iron in the United States, and, besides occurring in workable masses, is often disseminated through certain rocks, in grains, or in bunches and thin seams or laminæ, thus constituting what will be called "magnetic rocks" in this paper.

Its home is in the oldest rocks:—the primary (azoic, eozoic or archæan), as they have been successively termed. When it occurs in younger rocks, its origin can generally be traced to local metamorphism. The characteristic property of this mineral is its magnetism, with reference to which it is sometimes called lodestone. When brought near to pieces of iron or steel it often manifests an attraction for them, as it always does for another magnet. It hence causes the magnetic needle to deviate from its normal direction when brought near it. This property does not belong, in any marked extent, to any other mineral, and is the one which we have here chiefly to consider.

A piece of magnetite, broken from its parent bed, and suspended

^{*} A part of this paper was read before the American Philosophical Society, Philadelphia, and published.

by a thread, will take a position, as near as the mode of suspension will permit, corresponding with its original one. If a north and south line be marked on a specimen thus suspended, it would rudely and imperfectly answer the purpose of the magnetic needle; if with this piece of magnetite we rub, in a certain way, a slender bar of hardened steel, it in turn becomes magnetic, and, if properly mounted, will point north and south, and constitute a compass. Mounted in another way, so as to admit of vertical motion, the magnetic needle will, while pointing north, incline downward at an angle of about 76° at Marquette. This "dip," as it is called, increases to the north and decreases to the south.

Two magnetic needles made in this way present these phenomena: their north poles or south poles repel each other, while the north pole of one will attract the south pole of the other, and conversely. The same is, of course, true of two pieces of magnetite, or of a piece of magnetite and a magnetic needle; opposite poles attract, and similar poles repel. This property is termed polarity. From this it appears that the north magnetic pole of the earth must, in the light of the science of magnetism, be regarded as a south pole, because it attracts the north end of the magnetic needle. The poles of any magnet are understood to be those points opposite each other, and near its surface, where the attractive and repulsive power may be supposed to be concentrated. Any magnet, natural or artificial, exerts its influence or sends out its rays in every direction, like a luminous point. The limit of this influence may be designated as the sphere of its attraction. A magnetic needle within this sphere, and uninfluenced by other force, would point directly to the centre of the sphere or focus of attraction. The force which holds it in this direction varies inversely as the square of the distance from the centre; hence practically (on account of this rapid diminution of power) we soon get beyond the influence of even a great natural magnet, like a hill of magnetic ore.

All the properties above designated, and numerous others not necessary to our purpose, appertain in general to a mountain of magnetic ore or rock, as well as to the delicate needle of a miniature compass. It is therefore evident that the magnetic needle should assist in determining the position and magnitude of rock formations containing magnetite. It has been extensively used in numerous places in finding iron ore, and to a far less extent, if practically at all, in this

country, by field geologists, in determining the geographical extent, and, in part, lithological character of formations containing too little magnetite to give them commercial value, and which have already been designated magnetic rocks. The fact that all substances usually encountered in magnetical observations are transparent to the magnetic rays, or permeable by them, enables us to be certain of the existence of magnetic rocks or ores, though they be covered with water, earth, or non-magnetic rocks, to the depth of many feet, or even fathoms. A given magnetic force affects the needle just as much through one hundred feet of granite as through the same distance of the atmosphere. Dr. Scoresby gave a fine illustration of this fact, and an important application of the science of magnetism, by measuring, with great precision, 126 feet through solid rock, by observing the deviations in a needle, caused by an artificial magnet.

The earth itself may be regarded as a great magnet, which has the power of inducing this force (all magnets have a similar power) in masses of magnetite, and in all forms of iron and steel. We may suppose the force we have described above, as existing in the magnetic rocks and artificial magnets, to have been derived from the earth. An unmagnetized mass of steel or iron always manifests polarity induced by the earth, the upper or southerly portion being the south pole, and the lower or northerly end the north pole, in accordance with the law already stated. If the mass of iron or steel be elongated in form and made to stand nearly vertical, or to lie nearly in the plane of the meridian, this force is more manifest. To illustrate:—The upper end of all cast-iron lamp-posts attracts the north end of the needle, and the lower end the south. The magnetism thus induced in the wrought-iron pipes, lining the so-called magnetic wells of Michigan, would probably explain all the phenomena actually observed there. The law is, briefly: the upper part of every mass (of whatever form and size) of iron, steel or magnetite, is a south pole, and the lower part a north pole. This is, of course, true of magnetic rocks; hence almost universally the north end of the needle is attracted by such rocks, because it is the south pole of the rock which is uppermost and nearest. South pole or negative attractions, which are occasionally observed, come usually from faults or other divisional planes in the rocks; opposite poles being produced on opposite sides of such

breaks which sever the mass; a precisely similar phenomenon can often be observed on opposite sides of the joints in railroad tracks.

From this cause several natural magnets are often encountered in a short distance; and a needle, passing in a few feet from the sphere of the attraction of one of them, will turn round and point toward the pole of a neighboring mass which more strongly attracts it. Hence, in magnetic surveys, we have not the simple focal point first considered to deal with, but often several local centres of attraction, positive and negative, in addition to the directive force of the earth, all influencing the needle at the same time. The recent investigations in the use of "magnetism in testing iron for flaws" would undoubtedly aid in the study of the effect of faults on the magnetism of rocks. See Engineering (London), 1867, p. 550, and 1868, pp. 297 and 440. The magnetism of iron ships should also possess interest in the same connection.

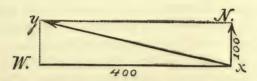
The direction which a magnetic needle takes (allowing it to have universal motion), under the circumstances supposed above, and the power with which it holds to that direction, must be the mechanical resultant of all the forces acting on it. It cannot point in two directions at the same time, hence stands between, inclining to the greater force. The principle of the parallelogram of forces makes it easy to determine the direction of this resultant, and to measure with mathematical precision the power which urges it. To do this we must know the direction and intensity of all the forces.

As an example, suppose a magnetic needle which, uninfluenced by other force than the earth's attraction, points due north and vibrates 10 times in one minute, to be placed due east from a south pole in a magnetic rock; and that, in this position, the earth's directing force be exactly neutralized by an artificial magnet, placed south of the needle,—it is evident that a needle so situated will point due west, urged by the local force alone, and that its vibrations will be solely due to this force. Suppose, for example, these vibrations to number 20 in one minute, or twice as many as were due to the earth's force. Now remove the artificial magnet; what will be the direction of the needle, and what number of vibrations will it give, urged by the local and cosmical forces?

It is a law of magnetism that the force urging a magnetic needle is proportional to the square of the number of vibrations made in a given time; $10^2 = 100$ and $20^2 = 400$, hence the local force is four

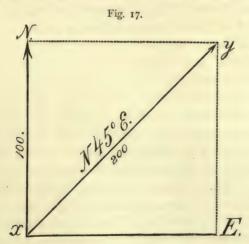
times as great as the earth's. Lay off in Fig. 16 the line x N due north, making it equal 100 on some chosen scale: lay off the line x W due west, making it equal by the same scale 400; complete the parallelogram by drawing the lines N y and W y parallel with the first lines. Draw the diagonal x y, it will be the resultant

Fig. 16.



sought. Applying the protractor and scale we find its course to be N. 75° 53 W., and length to be 412.31, the square root of which is 20½, which would be the number of vibrations.

Suppose that in another locality the same needle pointed N. 45° E. and vibrated 14¼ times in one minute, what would be the direction and intensity of the local force? In Fig. 17 lay off the line x y N. 45° E., its length equal to the square of the number of vibrations = 200; complete the parallelogram as before. It is evident that the line x E represents the direction and intensity of the local force, which in this case is due east, and has a power just equal to that of the earth. Unfortunately the simple cases here presented



seldom occur,—usually two or more local forces act on the needle at the same time.

In a similar manner any number of forces acting in as many different directions can be resolved. It follows that a magnetic needle, influenced by the earth's force, can never point directly toward a local magnetic pole, but will, with two exceptions which need not be named, always incline to point to the north of it.

It is evident that the degree of magnetism possessed by a needle, while it makes no difference with its direction, will affect the number of vibrations. Take the needle in the last case, and suppose it more highly charged; it will still point N. 45° E., but its vibrations will be increased in number just in proportion to the additional power imparted. Hence, in determining absolute terrestrial or local intensity, a standard for comparison is necessary; but this is not required in the work under consideration.

2. Magnetic Instruments—Dip Compass.

As the instruments employed in these observations are quite different from those used in Terrestrial Magnetism, which are described in the works on this science, a brief account of them will be given.

The Dip or Miner's Compass is a circular brass box, a common form being 33/4 inches in diameter, and 3/4 inch thick, having a circular glass on each side, which permits a perfect view of the needle. The needle is 21/8 inches long, weighs 133/4 grains, and is counterpoised so as to stand horizontal where there is no local attraction, the needle being permitted to swing in a north and south vertical plane, which is the position in which it is ordinarily used. The axis of the needle is of hard steel, its points resting loosely in conical cavities in agates, fixed in two arms projecting from the sides. Outside is a ring for supporting the instrument when observations are made, so placed that the weight of the suspended instrument brings the zero line of the graduated circle to a horizontal position. Although designed to be used chiefly for determining dips or inclinations of the needle due to local influences, it answers passably well for taking magnetic bearings when laid on its side, and is frequently used in this way in rough work.

As there is usually no means of throwing this needle off its points of support, the wear is great, and the instrument is often out of order. A person going out of the way of shops where repairs can be made, would do well to take two, and then have the means at

hand for making ordinary repairs. These compasses generally possess each an individuality of its own, and one must know his instrument before placing much confidence in his results: they will seldom reverse, 30° difference in the two readings being not infrequent. A New Jersey iron explorer informed me that his Dip Compass always indicated 90° when faced west, and the true dip due to local attraction when faced east. He is said to have used one position in buying and the other in selling iron lands very successfully.

My compass was made by Messrs. W. & L. E. Gurley, of Troy, N. Y. I have since seen one made by H. W. Hunter, of N. Y., which promises well. A reliable dip compass is a desideratum.

This is exclusively a hand instrument, and has no support; nearly all the magnetic observations recorded in this paper were made on instruments held in the hand. This may seem rude and unscientific to precise observers of physical phenomena; but it was found by trial that the average error by this mode of observation was less than 3°, which was comparatively small in localities where changing the position of the instrument only a few feet often made 50° difference in the direction of the needle, and deviations of 180° from the normal direction were common. It is not necessary to observe the direction of the wind to the degree to construct a useful theory of storms. Had the accurate instruments and precise methods of terrestrial magnetism been employed, not more than 50 stations could have been occupied with the time at my disposal, while with my rude methods over 1,000 stations were observed at.

The miner's compass above described is now in very general use in the magnetic iron-ore regions of the United States. The object here sought is to endeavor to point out new and perhaps better modes of using that instrument in finding iron ore, and incidentally to ascertain if it has any place in general geological field work. I have long believed that the magnetic needle can be so used as to give more definite information regarding magnetic ores and rocks than has yet been done to my knowledge. I did some rude and incompleted work in this field, at the Ringwood Iron Mines and elsewhere in New Jersey and Southern New York, the results of which are in part published in Prof. Cook's Report on the Geology of New Jersey. The observations of Prof. Cook and Dr. Kitchell on the magnetism of the iron ores of New Jersey, and the use of the magnetic needle in finding them, possess interest; see pp. 532–538 of

their report. The map of the Ringwood Iron Mines, accompanying that report, exhibits a part of my own observations above referred to.

The idea of applying Magnetic Science to Geology is not at all new; years ago Bischoff, after citing numerous observations that had been made in various parts of the world by different observers in regard to the influence of mountains on the magnetic needle, concluded as follows: "Assuming that it is magnetic ore alone, either as masses or disseminated through the rocks, to which the magnetic influences are to be ascribed—and in my opinion this is quite unquestionable—it would seem that magnetic observations instituted with the same degree of care as those made by Reich. would be well adapted for the discovery of hidden beds of magnetic iron ore. Such observations might therefore prove eminently serviceable to the iron industry. Certainly it would be requisite first to ascertain whether mountain masses containing only disseminated magnetic iron ore, but extending over a considerable surface, would not produce as great an effect as beds of magnetic iron ore. Sabine's observations do not appear to favor this; but, however this may be, the magnetic needle indicates the presence of magnetic iron ore where it cannot be recognized mineralogically, and demonstrates the very general distribution of this mineral."

My mode of observing was as follows:—To determine "variations" east or west,* the bearings of a standard line were taken as in ordinary surveys. Sometimes a solar compass was used, but oftener a pocket compass. The variations as shown by the miner's compass, termed "dips," were observed on this compass held in the hand generally in the plane of the meridian, hence the instrument would face east and west. Sometimes observations were made with the compass held at right angles with this position; that is, facing north and south. The instrument was always held in the hand and levelled by its own weight.

The *intensity* of the magnetic force for the three positions of the compass above designated, was measured by the number of vibrations† made by the needle in a unit of time, usually taken at ¼ of a

^{*} Declination, or the cosmical deviation of the needle from the true meridian, is not here considered.

[†] Half-vibrations would be the proper term, as the time from one point of rest to the next was counted and not the complete vibration.

minute. The vibrations varied from 0 to 60 in this time, 6 being the normal for my compass, due to the earth's influence. No attempt was made to eliminate the earth's attraction by neutralizing it with a magnet when the observation was made, or by computation. Of course, when the compass faced north or south, this was partially accomplished, because the earth's attraction would then be nearly in the direction of the axis of the needle. It must be borne in mind that the great amount of friction in this form of compass renders the number of vibrations only a rude approximation to the number which would be indicated by a delicately mounted needle.

The short needle of an ordinary pocket or dip compass, if in good order, will vibrate quickly and for some time where there is no local attraction. This motion is sometimes termed "working," and such normal "working," due simply to the earth's attraction, has often been mistaken by inexperienced persons for an indication of ore.

There is no better instrument for observing variations accurately than Burt's Solar Compass; but it is too heavy for explorers' use. I have found a convenient substitute for rough observations in the Pocket Dial Compass, which, used with a watch indicating local time, is rapid and sufficiently precise. This instrument, or an ordinary portable sundial, can also be used for running lines where there is local attraction; for rough work I have used it instead of the Solar Compass.

I hoped to have made some observations with properly constructed instruments, such as are used in determining the elements of terrestrial magnetism, in order to institute a comparison between accurate results and my own rude work; but the nature of such investigations requires more time than I have thus far had at my disposal. Fortunately Dr. John Locke made complete magnetic observations at several points in the Marquette Iron Region, which are recorded in "Smithsonian Contributions to Knowledge," vol. 3, pp. 25-27. One station was over magnetic rocks in Section 18, Town 47 north, Range 26 west, the geology of which he thus describes: "A loadstone in place broken into sharp angular fragments; here were two poles, 17.67 feet apart, one attracting the north, the other the south pole of the needle." Dr. Locke found the dip to be 42 deg. 53 min., when it should have been about 76 deg. The duration of 500 vibrations was 822 sec., when it should have been about 1,500 sec., and the calculated horizontal intensity was more than four times the normal force computed for that station. If Dr. Locke had occupied 500 stations on that section of land, he would have obtained different results at each, often differing more from each other than the foregoing do from the normal forces.

These observations, like all recorded ones that have come under my notice, have had terrestrial magnetism as their chief object; therefore the observers have avoided the very localities which to the geologist and explorer possess the greatest interest—those where local magnetic attractions exist. Dr. Locke calls attention to the importance of magnetic science to the geologist, and gives many interesting isolated facts bearing on the subject, particularly regarding the existence of magnetite in volcanic rocks, where it usually occurs.

Before dismissing the subject of instruments suited to magnetic surveys, I will call attention to a patent mariner's compass made by E. S. Ritchie, Esq., of Boston, in which the needle is entirely supported by a liquid having the same specific gravity, thus giving it universal motion. A needle so mounted and having the earth's attraction neutralized by a magnet, should point directly towards a local magnetic pole when brought within its influence, thus accomplishing with one observation and no calculations what requires at least two with the ordinary compass. For intensity Mr. Ritchie, suggested the following mode:—Time the needle from the instant of its being let off at 90 deg. to its passing the resting point. I am of the opinion that a valuable instrument for miners and explorers could be made on Mr. Ritchie's plan.

A modification of the ordinary compass has been made which accomplishes the same thing in part. The agate support is fitted to the needle by a sort of universal joint, which gives the needle a vertical range through half a quadrant in addition to its horizontal motion. The only one I ever saw was made from the design of the late Wm. J. Amsden, Esq., of Scranton, Pa., who made some valuable magnetic surveys.* A pocket compass on a similar idea has lately been patented. A somewhat similar instrument has, I understand, been used for a long time in Sweden and Norway. On the same principle the ordinary surveyor's compass indicates dips rudely. At the west quarter post of Section 7, Town 46

^{*} Messrs. Gurley now make a dip compass which gives the needle limited lateral range.

north, Range 29 west, being on the east side of Republic Mountain, I find marked on the U. S. Survey plat: "End of needle dips ¼ inch, variation 62 deg. west."

C. F. Varley, Esq., the English Electrician, suggested to me that a portable electro-magnetic apparatus could be constructed, with which might be determined the direction and distance to the pole of a magnetic rock by some simple observations and computations. An instrument of this kind would have considerable value in connection with magnetic needles, especially where the magnetic ore or rock was covered with considerable thickness of other material. In 1867 Mr. Varley, with a view to detecting electric currents, if any existed, made some observations both in the copper and iron-bearing rocks of Lake Superior; he found such currents in the mines of native copper, but none in the iron mines. The instruments employed were rude, having been extemporized on the spot. I do not know whether he has published anything on this subject.

Professor Joseph Henry has suggested in a letter that it is "highly probable that the abnormal variations of the magnetic elements in our iron ores are due to *electro-magnetic* action rather than to magnetic."

3. Geological Sketch of the Magnetic Rocks.

In order to make the perusal of this subject to a certain extent independent of the remainder of this report,* a few facts regarding the geological position and lithological character of the magnetic rocks of the Marquette region will here be repeated, the subject having been more fully considered elsewhere.

Rocks of the four oldest geological epochs yet made out on this continent are represented on the Upper Peninsula of Michigan; two belonging to the Azoic, one to the Lower Silurian, and one between these, of questioned age. The equivalency of these with the Canadian series has not been fully established, but the nomenclature of the Canadian geologists will be employed provisionally.

The Laurentian of the Upper Peninsula is like that of Canada in being largely made up of granitic-gneisses, but differs in containing no limestone so far as I have seen, and little, I may say practically

^{*} Many persons have asked for copies of this chapter who do not expect to get the whole Report.

no iron ore, and very little disseminated magnetite. Next above the Laurentian, and resting on it non-conformably, are the Huronian or iron-bearing rocks; these are also called by the Canadian geologists "the lower copper-bearing series." This series comprise several plainly stratified beds of iron ore and ferruginous rock, varying in the percentage of metallic iron from 15 to 67 per cent., interstratified with greenish tough rocks, in which the bedding is obscure, which appear to be more or less altered diorites, together with quartzites (which pass into marble), clay slates, mica schists, and various obscure magnesian schists. The maximum thickness of the whole in the Marquette region is not far from 5,000 feet.

While the great Huronian area of Canada north of Georgian bay bears, so far as I am aware, little or no workable iron, and derives its economic importance from its ores of copper, the Marquette series, supposed to be of the same age, are eminently iron bearing, and have as yet produced no copper. It is doubtful if in the same extent and thickness of rocks, anywhere in the world, there is a larger percentage of iron oxide than in the Marquette series. In the order of relative abundance, so far as made out, the ores are the flag, the red specular hematites, soft or brown hematites, and magnetites. These all exist in workable beds, and all as disseminated minerals in rocks usually silicious. The geological distribution of these ores of iron in the Huronian series will be considered in another place. The geographical distribution is less understood; so far there seems to be the greatest concentration of magnetic ores in the Michigamme district of the Marquette region. From this, the relative proportion of magnetite seems to decrease as we go east, north, west and south, although there is a considerable magnetic attraction in the Menominee or southern iron region.*

Next younger than the Huronian are the copper-bearing rocks of Keweenaw peninsula, which extend westward into Wisconsin, the age of which has led to much controversy; good authorities having placed them in different epochs, from the Azoic to the Triassic. Recent observations made by Prof. R. Pumpelly and myself go strongly to confirm the view, if we have not 1 ositively demonstrated it, that they are non-conformably overlaid by the Silurian, and are therefore related to the Azoic. The relations of the copper-bearing

^{*} See Appendix H., Vol. II.

rocks to the Huronian are not fully made out. In tracing the dividing line from Bad river in Wisconsin to Lake Gogebic, Michigan, last fall, a distance of sixty miles, we found them nearly, if not precisely conformable, but widely different in lithological character.

With regard to the magnetism of the copper-bearing series, the United States surveyors mark considerable variations at several points on the Land Office plats, due in all probability to disseminated magnetite in the trappean members of the series, although good authorities have ascribed these variations to electric currents. My own observations on the magnetism of these rocks have been limited, but lead me to believe that it is far less in amount and less persistent in character than is usually the case in the Huronian, indicating that the magnetite (to which I ascribe the attractions) is perhaps an accidental rather than essential constituent, and small in amount. Macfarlane found less than one per cent. in one of the Portage lake traps.

The next series of rocks in ascending order are the horizontally-bedded Lower Silurian sandstones, which skirt the south shore of Lake Superior nearly its whole length, called by Foster, Whitney, and Dr. Rominger, Potsdam, and assigned by the Canadian geologists, under the name St. Mary's, to a later period. They have not been proven to be magnetic, although strong magnetic attractions have been observed over this Silurian area, as will be explained hereafter.

To recapitulate, we have: I. The Laurentian granite and gneiss, practically non-magnetic; 2. The Huronian iron-bearing rocks, often highly magnetic; 3. The copper series, slightly magnetic; and 4th. The Silurian rocks, without magnetism. This classification is intended to apply more particularly to the rocks of the Marquette and Menominee regions proper, embracing the central and southern portions of the Upper Peninsula; and even here, as has been noted above, there are exceptions. This sketch of the Marquette rocks, in the light of the distribution of magnetite, would be incomplete, did I not mention the fact that this mineral is very generally present in the form of fine sand in the drift in the region I am describing. If one moves a magnet about in the sand of a creek it is rarely that magnetic sand will not be found adhering. I have never seen it accumulated in quantities

that would point towards its being utilized; nor have I ever observed a local variation which I ascribed to the mineral in this form.

We will now return to the Huronian or highly magnetic series, taking up its structure in some detail. About nineteen lithologically distinct beds or strata make up the series; of these, six and probably seven are generally so magnetic as to cause considerable variations in the needle. These beds vary from forty to several hundred feet in thickness, and strike and dip in all directions, and at all angles. The prevailing strike, however, is easterly and westerly, and the dip at high angles, often vertical. These rocks frequently outcrop, when we have no use for the magnetic needle in their study. Again, they are covered by deep drift, where magnetic observations, or workings, can only reveal them.

In order to study the magnetic characteristics of these rocks more minutely than could be done in the field, two hundred and twenty-two specimens, covering all the more common varieties, were collected and are deposited in the cabinet of the University of Michigan; they are fully described under lithology in this Report. Fifty-four, or twenty-four per cent., were found to possess some degree of magnetic power as manifested by their influence on a magnetic needle; each specimen being in turn made to touch each end of a mounted needle. If it had the power to lead it 20 deg. from its normal direction, the specimen was said to be feebly magnetic, and strongly magnetic when the needle followed the specimen round the circle if held about half an inch from it. Of these fifty-four specimens, thirteen were feebly magnetic, twenty-nine magnetic, five decidedly magnetic, and seven strongly magnetic.* None would, however, lift ordinary carpet tacks. Twenty-four, or nearly one-half, possessed polarity in some degree. Thirty were simply magnetic, with no polarity that could be detected by the rude means employed: in some instances the specimen would repel the needle at half an inch distance, but would attract it if placed in contact. Such specimens were rated as possessing polarity. All of the strongly magnetic specimens were rich in magnetite and possessed polarity, and it is not improbable that

^{*} Appendix H gives the percentage of material lifted by the magnet in twenty-one specimens of Lake Superior ore, together with the color of the powder.

all would have been found to possess it if tested by more delicate means. Von Cotta, however, speaks of magnetic iron ore which possessed no polarity. The specimens generally attracted the south pole more strongly than the north. When examined, they had been collected about three months. Whether they would have shown more or less magnetic power if tested when freshly broken, I do not know. Dr. Kitchell says that under certain circumstances fragments gain magnetism.

In 1860 I saw a powerful loadstone for its size, in the possession of Professor Trego, of Philadelphia, which he had picked up in New Jersey twenty-two years before. I once collected a number of pieces of loadstone in the Bull Mine, New York, which in the mine would lift small nails; in a few days two-thirds of them had lost this power. This may have been due to the fact that in the mine the nails themselves were made magnetic by induction.

Regarding the location of the poles in magnetic rocks, the laws of magnetism would place them near the surface, or next divisional planes or terminations of masses. Observers are generally agreed that iron ore is most magnetic near dykes or volcanic rock. Quoting again from Dr. Kitchell, "Geology of New Jersey," p. 535: "The extent of the magnetic qualities of iron ores depends on their position with respect to the surface; the nearer to the surface the greater will be their magnetic properties. This appears to depend on the action of surface water and atmospheric agents, for it has been frequently observed that ore, when first taken out of a mine at a considerable depth, possessed but slight magnetic properties, but on being exposed to the atmosphere for a few months or years it would increase so much that excellent specimens of loadstone for experimental purposes could be selected therefrom. Seams of ore that contain numerous joints and fissures, through which water and atmospheric agents pass, possess more decided magnetic properties than those which are more compact and free from crevices and fissures." *

These remarks of Dr. Kitchell possess much interest. I have but one fact that bears on this question;—an average sample made up of numerous fragments collected by myself of the Iron Moun-

^{*}If a fact, is this due to the contact of air and water, or is it because the seams necessarily produce small independent magnets.

tain Missouri "surface," or boulder ore, contained only about one-fourth as much magnetite (as measured by the amount lifted with a horse-shoe magnet) as did a specimen of "quarry" (ledge) ore selected at the same time and in the same way.

Classifying the magnetic ores and rocks of the Marquette region economically, the merchantable ores, according to the present standard of richness, would not constitute two per cent. of the whole; the balance being ferruginous quartzites and schists possessing no present value as ores. The merchantable magnetic ores have so far all been found in one formation near the middle of the series, and that is not all pure ore by any means; therefore, when an ore-hunter finds an "attraction" in the Lake Superior region, the chances of his having found a mine are not more than one in fifty. Neither the strike nor dip of the formation seems to affect its magnetic power. This depends, so far as my observations throw any light on the question, chiefly on the percentage of magnetite entering into the composition of the rock. Prof. Cook-"Geology of New Jersey," pp. 537-8—says that the magnetism of iron ores was influenced by the "pinch and shoot" structure so prevalent in the iron mines of New York and New Jersey. He points out the analogy between these regular pod-shaped masses-"shoots" of ore,—pitching downward in a northerly direction and an iron bar in the same position; both become magnetic and have polarity.

The "pinch and shoot" structure exists in the magnetic ores of the Marquette region, but is obscure, and in strike and dip there is no parallelism between our rocks and those of New Jersey, as is shown elsewhere. Yet our ores must usually be more strongly magnetic than those of New Jersey; for Prof. Cook says: "It is generally conceded that ore, covered by thirty feet of earth, will attract the needle, and 'large veins' have disturbed it when covered by fifty feet of earth." Now at five and even fifty times these distances horizontally, the needle is often deflected in the Marquette region, and at the Spurr Mountain the needle indicates a dip of 70 degrees at an elevation of 94 feet above the ore.

With regard to the associations of the various ores it may be said, that magnetic and specular ores are often found together, as are also the specular and soft hematite ores; but so far the magnetites and hematites have not been found in juxtaposition. If we suppose all our ores to have once been magnetic, and that the red

specular was first derived from the magnetite and the hydrated oxide (soft hematites) in turn from it, we have an hypothesis which best explains many facts, and which will be of use to the explorer. As a rule it may be assumed that the hard ores of the Lake Superior region, even although they be rated as red specular, contain a sufficient amount of magnetite to cause some local disturbance in the needle; there are exceptions to this rule, but they are rare. In some instances, especially in the Menominee region, the disturbance is slight, but enough to be noticed by careful observation. It should be noted that the L'Anse Iron Range, so far as known, contains no magnetic ore whatever.

4. Explanation of Magneto-geologic Charts, Plans and Sections.

Having now briefly stated those elementary principles of Magnetism which are involved in our subject, described the instruments employed and their use, and sketched the geology of the rocks whose magnetic forces we are to study, we are fully prepared to examine the results of the observations made, and to draw such conclusions and make such applications as the facts seem to warrant.

It has been found necessary to introduce a few terms which may be new in describing the graphical representations of the phenomena observed. No work to which I could gain access contained expressions such as portions of our work seemed to require. Figures 1 and 2, Republic mountain chart (No. XI. of Atlas), are copied in part from the geological and topographical map of Republic mountain, which see for explanation of geology, relief of ground, and geographical position.

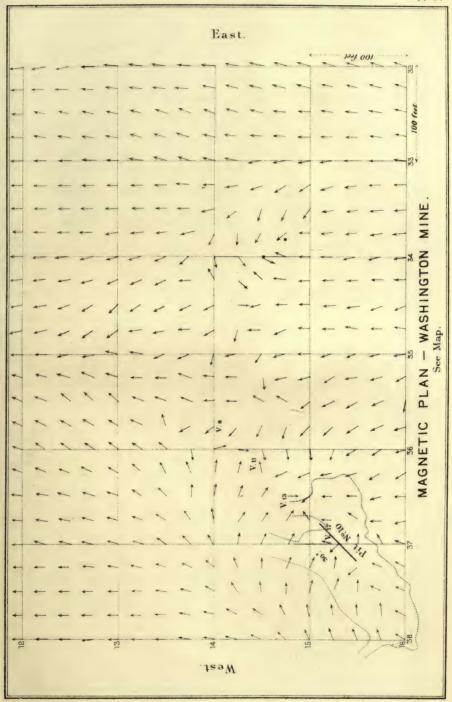
Magnetic observations were made across the entire Huronian series lapping on the Laurentian on each side, along survey lines 26 and 30, which run N. 53° E.; the observations being taken for a considerable part of the distance every 25 feet. The arrows in Fig. 1 indicate the directions which the needle actually pointed under the combined influence of terrestrial and local attraction. The angle between these arrows and the meridian is the *variation* in Azimuth (called simply variation) and ranges, as will be seen, from 0 to 180°. The direction of the arrows, although sometimes irregular, leaves no doubt as to which are the magnetic rocks.

The full significance and value of the common compass in locat-

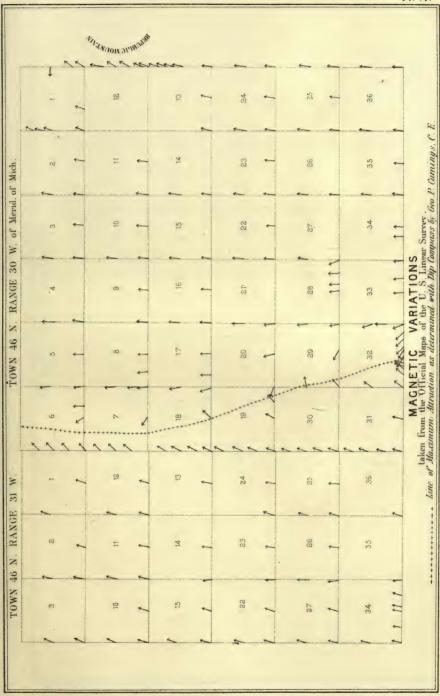
ing magnetic rocks and ores is better shown in Plate V., which represents variations observed at the west end of the Washington mine, embracing the West Cut or Pit No. 10. The stations indicated on the Plate refer to survey lines shown on the map of the mine. No. VIII., to which reference is made for information regarding the geology and topographical features of the locality. A glance at this figure will bring to the mind of all familiar with magnetic experiments, the plumose forms assumed by magnetic sands or iron filings resting on paper and influenced by the magnet. Our figure may be regarded as representing the laboratory experiment greatly magnified. As to the irregularities shown by some of the arrows, it is probable that if the magnetism of ordinary magnets could be studied minutely, as with microscopic needles, that corresponding irregularities would be observed in the directions and polarity of the forces, not unlike those seen on this magnetic plan of the Washington mine. If we admit, as we are forced to do from these facts, that magnetic rocks present phenomena entirely analogous to artificial magnets, then it is not difficult to decide as to the cause of the phenomena exhibited on the sketch before us.

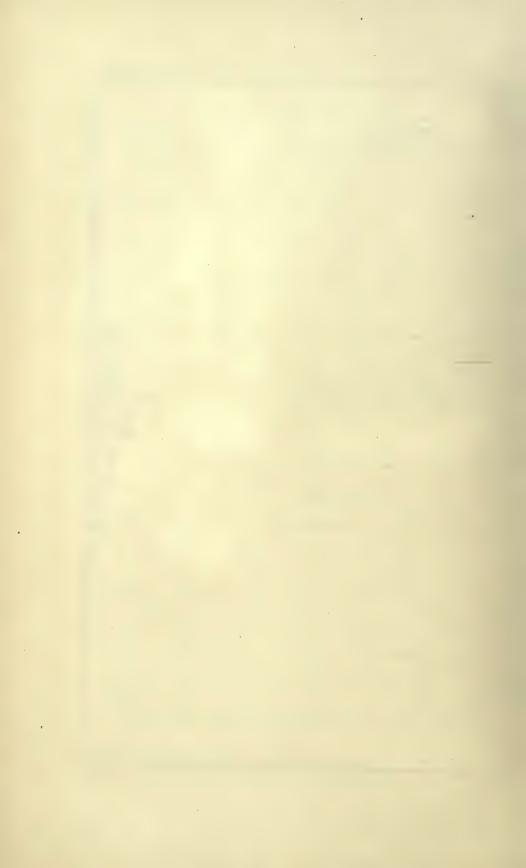
The dotted line is designed to indicate the position of maximum variation, or rather the position of the force which causes the variation. The observations made for intensity along this line, indicated by vibrations (six being the normal number), confirm the indications of the horizontal compass. There can be no doubt but that nearly under this line, at no great depth, is a large amount of magnetite; whether free enough from rock to constitute a merchantable ore, explorations only can establish. Since this plan was made, work has been resumed at Pit No. 10, and a tolerably regular bed of ore revealed, having the strike and dip marked on the plan, which coincides closely with what might have been predicted. The relationship of this deposit with the others constituting the mine will be considered elsewhere. This magnetic plan, as well as Fig. I, Republic mountain chart, shows, that while the variations are governed by a uniform law away from the lines of maxima, within these lines great irregularities of direction exist.*

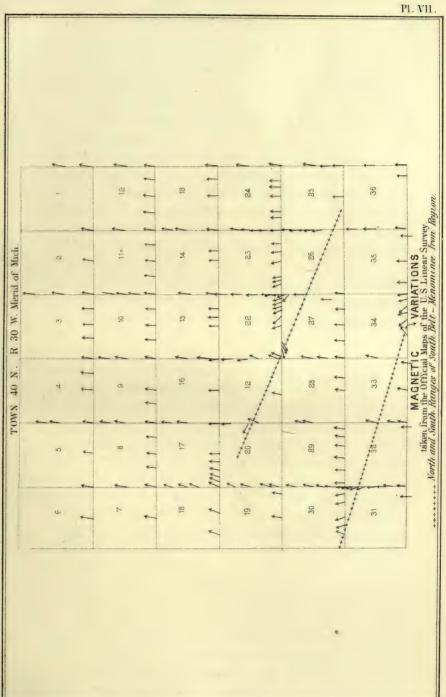
^{*} Since the above was written, I have, by the kindness of Mr. F. Firmstone, of Easton, Pa., been able to inspect some magnetic charts of New Jersey localities, made by the late Mr. Amsden, of Scranton, which are excellent.













Passing from Plate V., which represents but a small area, over which the magnetic observations have been very numerous, to a magnetic plan of a large surface, with widely separated observations, we have in Plates VI. and VII., copied from the United States Land Office books, a fine exemplification of the significance of local magnetic variations.

In Plate VI. the magnetic rocks run nearly north and south,—which direction, as has been heretofore stated, produces the maximum variation. It will be seen that the needle is influenced at a distance of nearly, if not quite two miles, and that the variation diminishes rapidly as we depart from the line of maximum attraction. The disturbances recorded on the north-east part of this plan are due to Republic mountain.

Plate VII. represents one of the iron townships in the Menominee region. The variations are scarcely so great, nor do they extend so far as in the other. As the two iron ranges represented run much more nearly east and west in this case, it is interesting to observe the difference in the behavior of the needle. These plans are additional proof of the value of the Linear Surveys to the explorer, a point to which I have often referred.

Figures 3 and 4, Chart XI., Atlas, are magnetic sections along lines 26 and 30 of plan. The arrows indicate the direction of the dip-needle vibrating in the plane of the meridian. The normal direction is the horizontal line; the arrow head indicating north end of needle should therefore normally point to the right hand side of the chart. It will be seen that the dip, like the variation, often attains the maximum of 180°, that is, the north end points south.

The colored curved lines express approximately the *intensity* of the local magnetic force; their ordinates being the number of vibrations made by the needle in one quarter of a minute, on a vertical scale of eight vibrations to the inch. The *blue* line records the observed vibrations of the *horizontal* needle, the others of the dip-needle. The *black* line refers to the needle vibrating in the plane of the meridian (compass *facing west*). The *red* line refers to the needle vibrating in an east and west plane (compass *facing* south.)

Fig. 2 is a magneto-geological section on the line A—A' of Fig. 1. The upper curve represents a projection on one plane of the maximum intensities of all the curves of Figs. 3 and 4. The lower curve, Fig. 2, has reference to variations and dips, its ordinates being

proportional to the maximum variation in direction of the needle, caused by the magnetic rocks. It is intended as a sort of summary of the facts expressed by all the arrows denoting directions, as the upper curve is a general expression of the intensities. It will be observed that the summits of the lower curve, Fig. 2, which indicates maximum variation, are always northerly from the centre of the magnetic bed. This is as it should be, because the greatest variation takes place before we reach the local magnetic pole, when approaching it from the north. The intensities, on the other hand, are greatest directly over the magnetic rocks. It should be borne in mind that the intensity of a magnetic force is really proportional to the square of the number of vibrations in a given time; but in these investigations the actual number of vibrations has been used in constructing the sections, as being more convenient.

In addition to the facts observed during this survey, which are recorded on the Republic Mountain Chart, and various figures in this volume, certain others, obtained from the United States Land Office, plats of Towns 46 and 47 north, Ranges 29 and 30 west, will be employed, besides those already given from the same source.

The discussion of the facts in our possession falls conveniently under two heads:—First, Regarding the entire Huronian series as a unit, and the comparison of its magnetism with the Laurentian system. Second, A study of the magnetism of the individual beds of the Huronian or iron-bearing rocks, in detail. Republic mountain and vicinity afford an excellent opportunity for both these investigations.

The Magnetism of the Laurentian System or Granitic Rocks.

The Federal township plats above referred to, cover an area of, say twelve miles in diameter, of which Republic mountain is the centre; at least nine-tenths of this territory is Laurentian. The variations of the needle noted are from two to six degrees east, averaging four and a half degrees, which may be regarded as the declination of the needle at the date of the surveys of this locality, due to cosmical causes. From this and similar facts covering the whole Marquette region, we may conclude that this oldest system of all known rocks has here no beds of magnetite, nor does it now contain magnetite as an essential constituent mineral,

nor indeed oxide of iron in any form. Prof. Pumpelly and myself found slightly magnetic rocks in the Laurentian south of Lake Gogebic, and the professor mentions in his report to the Portage Lake and Lake Superior Ship Canal Company "a deposit of iron ore in the Laurentian gneiss and hornblendic schist series on Sections 10 and 15, T. 41 N., R. 29 W.," in the Menominee Iron Region, from which I have seen specimens which do not look very promising. One or two other places are mentioned where magnetic beds occur in the Laurentian, but they are exceptional, the rule being as has been stated. But everywhere in the region we are considering, over or near the Huronian Series, the Government surveyors note variations. The approximate boundary between these two systems of the Azoic in some parts of the Upper Peninsula could indeed almost be delineated from their surveys by magnetic variations alone.

Magnetism of the Huronian Series as a Unit—Republic Mountain.

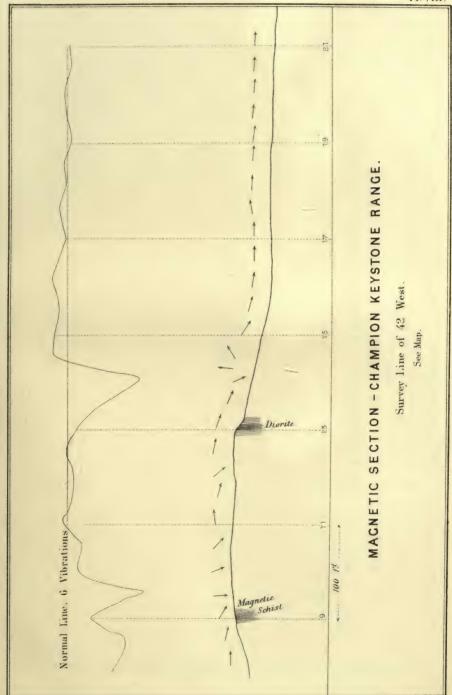
No special observations were made to determine the extreme limit to which its magnetic influence extends. The Federal surveys would make the distance over one mile, and Durocher mentions that he was told in Sweden that "important beds of iron ore produced deviations in the needle up to the distance of nearly two kilometres," or over one mile-Annales des Mines, 5 Series, Vol. 8, p. 220. The Federal surveyors note a variation at the northeast corner of Section 7 (See Fig. 1, Republic Mountain Chart) of 25° west, agreeing very nearly with my observations corrected for the change in declination since the survey was made. This corner is at least 600 feet from the nearest Huronian bed, and probably 900 feet from any member of the series containing magnetite. Judging by the direction and intensity of the magnetic force as exhibited by the needle, as we approach the mountain from the northeast (see Figs. 1, 3, and 4), it seems probable that the bed which chiefly produced the effect was No. VI., and still more distant ores. If this be a correct inference, we have the phenomenon of a magnetic needle deflected 25° from its normal direction by a bed of rocks containing not to exceed 33 per cent. of magnetite, distant 1,500 feet horizontally. The facts from the U.S. surveys given above show that the needle is sometimes influenced to a much greater distance.

Passing to the south-west side of the Huronian basin we find the influence exerted by the magnetic rocks to gradually diminish as we recede from their edge, which is believed to be under the Michigamme river. See Fig. 1. Here we find the needle varying 15° at a distance of at least 800 feet from the nearest magnetic rocks.

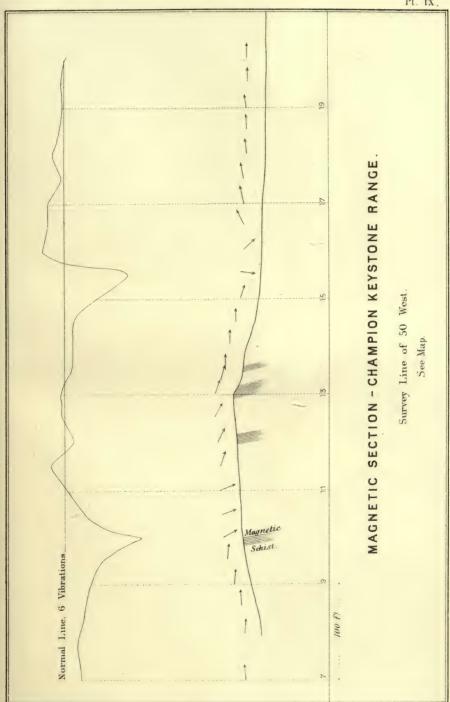
An inspection of Fig. I shows that the variations of the needle are much greater on the north-east than on the south-west side of the mountain, which should evidently be the case from the fact that to the south-west the terrestrial and local forces are more nearly in the same line than on the north-east side; hence in the latter case the mechanical resultant (direction of the needle) would form a greater angle with the direction of the earth's force (magnetic meridian) than in the former.

The question of the distance to which magnetic ore and rocks will attract the needle receives some additional light from the Champion Mine, Plates VIII. to XV. It is evident in this case that the magnetic force of the ore is felt to a distance exceeding 700 feet to the north of the mine. To the south there is less certainty, because of the other magnetic rocks (see sections) which underlie the ore in that direction. It is probable that careful observations would detect the influence of this remarkable deposit of ore through an east and west zone, which in places would attain a breadth of 2,000 feet or more, one-fifth of this area showing a magnetic dip of 90°; but this does not prove the existence of 400 feet of magnetic rocks or ore, by any means, as will be seen below.

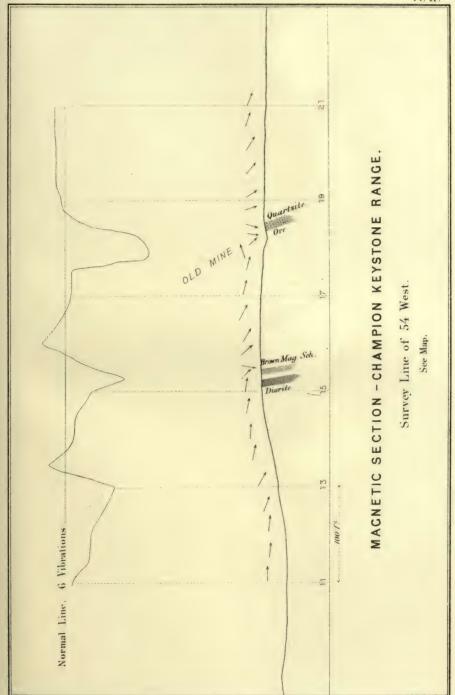
At the Spurr Mountain, which is an east-west deposit of highly magnetic ore like the Champion, Mr. Lawton observed just south of the range 23 vibrations in a quarter of a minute; going south the vibrations diminished somewhat regularly, until at 600 feet the needle vibrated but ten times in a quarter of a minute. At 300 feet north of the mountain the needle settled indifferently in any direction, owing to the fact that the terrestrial and local forces just balanced each other at that point; further north the vibrations increased somewhat irregularly, owing to the presence of slightly magnetic rocks, until at 1,400 feet six vibrations were observed in a quarter of a minute. There must of course be points north and south of all magnetic belts where the vibrations would be equal and normal, but these limits were not reached, the observations proving only that the magnetic belt at Spurr Mountain is over 2,300 feet wide.



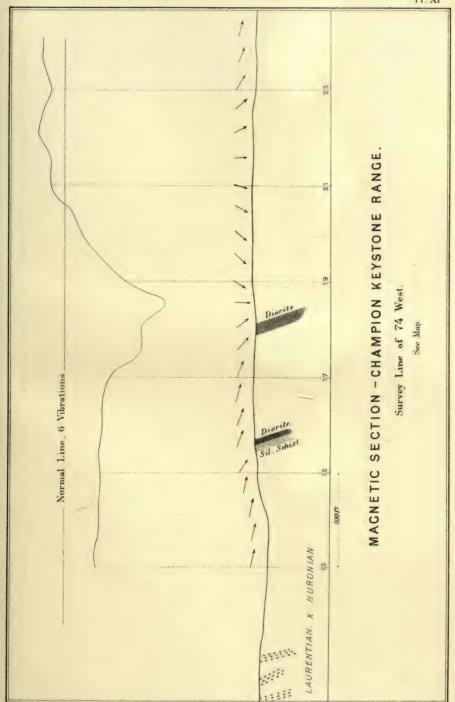










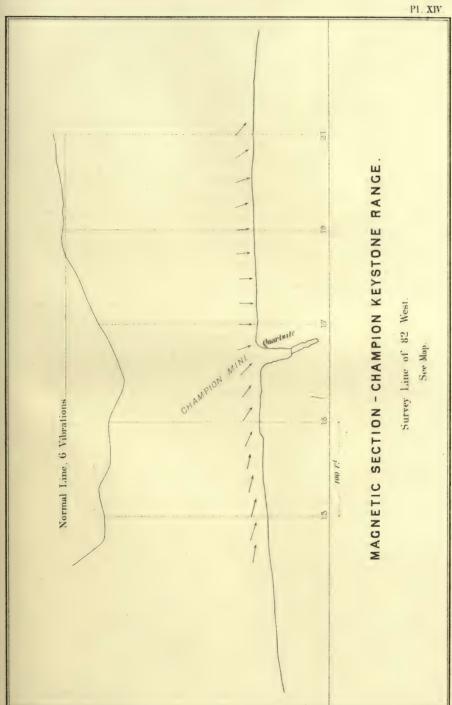




MAGNETIC SECTION - CHAMPION KEYSTONE RANGE. Survey Line of 76 West. See Map. Quartzite 100 Ft CHAMPION MINE Normal Line, 6 Vibrations_ HURONIAN LAURENTIAN









It may be asked why the very silicious magnetic rocks of the Republic Mountain influence the needle at a greater distance than the pure ores of the Champion Mine. It is not at all certain that this is the fact; the limit of the influence has been determined in neither case. The stratigraphical conditions, however, are quite different. The strike of the Republic mountain rocks being north-westerly, is far more favorable for producing variations than is that of the Champion deposit, which is east and west. It is quite evident that a north-south deposit of ore would cause greatest variations (see Plate VI.) and an east-west deposit least. If, in the latter case, we conceive the power to be equally distributed along an east and west mathematical line, there would be produced no variations at all in a horizontal compass. Again, there are four highly magnetic beds at Republic Mountain, while at the Champion there is only one.

Regarding the polarity of the magnetic force: (I) In every instance the north end of the horizontal needle was drawn towards the magnetic rocks; hence, north-easterly of Republic Mountain, the variation was west; and south-westerly, the variation was east. (2) With the dip-needle vibrating in an east and west plane, the north end pointed westerly, or towards the mountain on its northeast side. (3) With the dip-needle vibrating in the plane of the meridian, on the north-east side of the mountain, the south end inclined downward, producing a "negative dip," as shown in Figs. 3 and 4, and this increased as the magnetic rocks were approached until the needle turned entirely over. This apparent negative attraction was probably in reality only the effect of an attraction for the north end of the needle, which inclined to the magnetic rocks by the shortest road. Why the north end of the needle moved upward instead of downward (which was apparently just as short a road) as it approached the magnetic rocks over the non-magnetic Laurentian, I can only explain as follows,—which hypothesis may also explain instances other than this where slight negative attractions have been observed over granitic rocks, for example, south and south-east of the Champion mine. My needles were always counterpoised near Negaunee or Marquette, which towns are built on the Huronian. Of course an effort was made to get away from the magnetic members of the series; but this evidently would be impossible if their influence extends to the distance of one half mile.

Magnetic rocks would probably be found throughout the Huronian belt by boring less than 1,000 feet into the earth, owing to the basin-like structure of the series. It is probable, therefore, that my needles were counterpoised under the influence of some positive magnetic force; hence, when taken over Laurentian rocks containing no magnetite, they would show "negative" attraction. If this hypothesis is correct, then the negative attraction referred to above is explained.

Regarding the intensity of the magnetic force exerted by Republic mountain as a whole, but one observation need in this place be made. The vibrations are greater on the south-west than on the north-east side, or exactly the converse of the variations. The Magneto-Geologic Sections of the Champion and Keystone Range (see Plates VIII. to XV.) present the same phenomenon.* As the needle is carried north from the Champion bed, its vibrations rapidly diminish in number until they become less than the normal number due to the earth's magnetism; after which, on going still farther, the vibrations will increase until the normal number is reached: but in going south, the diminution is far less rapid, and the number of vibrations never falls below the normal number. The same was observed at the Spurr as is noted on page 226.

The obvious reason is this: when the needle is south of the local force, both it and the terrestrial force act in the same direction, producing a maximum effect; but when the needle is north of the local force, it can evidently be influenced only by the greater force less the smaller. In the first case the mechanical resultant is the sum, in the other it is the difference between the two magnetic forces. This readily explains the difference in the slope of the curve of intensity north and south of the magnetic poles, so noticeable in the magnetic sections.

Republic Mountain.

A glance at the directions of the needle as indicated by the arrows in figures 1, 3, and 4 of Chart XI., will impress one with the conviction that there is no direction in azimuth, or inclination which

^{*} The survey lines on the Magnetic Sections, Plates VIII. to xv., refer to Map of the Champion Mine, No. VII., which should be examined in connection with them.

the needle does not assume in crossing the series of rocks. The north end of the needle never points north, often east and west, and sometimes south; while in the dip-compass it turns a series of somewhat irregular somersaults, pointing habitually downward, but often towards the zenith. The needle may be said to "box the compass right and left," as we may suppose that feat accomplished by a drunken sailor. A second glance at the arrows will show us that there is much method in the madness of our ge-go-sence; the needle very generally tends to point toward the blue or red-colored rocks, which contain magnetite, while it is comparatively indifferent to the green, gray, and salmon colored, which contain little or none of this mineral. The particular significance of the variations and dips will be more fully discussed below.

We will leave for the present the consideration of the direction of the magnetic force expressed by the arrows, and return to the subject of the intensity of the force as expressed by the colored curves (see page 223). Nothing is more evident on the chart than that these curves indicate with great certainty the position of the magnetic beds over which they are more or less convex, producing summits; and more or less concave or flat over the non-magnetic rocks, pointing literally as a finger in some instances to the location of the magnetic force. Comparing the three curves in figs. 3 and 4, it appears that:—(1) The red line (compass facing south) oftenest rises higher than any other over the magnetic rocks; and sinks lower away from them. It has also fewer changes in direction than the others. (2) The black line (compass facing east-west) falls lower than either of the others over the magnetic rocks. (3) The blue line (compass horizontal) often has an extreme depression, where the others have an extreme elevation.

These, the most obvious generalizations from the curves, are explained by the principles of the mechanics of forces already mentioned.

Fearing there may be some confusion from representing the same element-intensity by three curves, I suggest the following conception: Suppose an observer to be provided with a horizontal com-

^{*} A Chippewa word for magnetic needle, signifying "little fish," in allusion to its wiggling motion.

pass having a blue needle and two dip-compasses, one provided with a black and the other with a red needle. Suppose, further, these to be mounted for observing at the same station, but so far apart as not to influence one another; the blue needle moves in a horizontal plane, the red needle in a vertical east and west plane, and the black needle in a vertical north and south plane. Suppose, further, a powerful magnet to be placed (1) directly under or directly over the station, it is evident that only the black and red needles will be influenced. (2) If placed north, the blue and black needles only will be influenced. The directive force in this case would be a maximum; because the magnet's power is added to the earth's, both acting in the same line. (3) If the magnet be placed directly south, the red needle will again be uninfluenced, but the black and blue needles will indicate a minimum of intensity instead of a maximum, for their directive power will be the difference between the force of the magnet and that of the earth. (Places have been observed where the needle gave us no vibrations in any position from this cause. A fine illustration occurs in Fig. 3, Chart XI., Station 24, where there must have been a very strong pole to the south of the station; but this pole is evidently north of Station 24, Fig. 4, where the greatest intensity was observed.) (4) If the magnet be placed east or west of our supposed station, the effect will be the same; the red needle will be most influenced, blue next, and black not at all.

We are now fully prepared to explain the phenomena presented by the colored curves.

(I) Why does the red line usually rise higher over the magnetic rocks, and sink lower away from them, and why does it fluctuate least? When the needle vibrates in an east and west plane, its axis points north,—that is nearly in the line of the directive force of the earth, which it thus partially neutralizes; giving the local forces full power. As these are much stronger than that of the earth at short distances, we should expect the result observed over the magnetic rocks. Away from them, the earth's force being nearly neutralized, we should have the minimum of intensity as is shown by the red line. That the changes in direction in this line are less frequent and less abrupt than the others, indicates, I think, that if the earth's attraction was entirely neutralized and the error of observation reduced to a minimum, the curve derived from the magnetic force resident in the rocks on any particular cross-section might be more

regular than any shown in the chart. It is reasonable to suppose that the red curve has most significance in our investigations. (2) Why do the black and blue lines fall as a rule lowest over the magnetic rocks? Suppose a local force, about equal to the earth's, to exist directly south of a dip-compass placed in the plane of the meridian, or of a horizontal compass; we should evidently have a minimum of intensity, because the terrestrial and local forces would balance each other. The marked exception to this rule over formation XI., Fig. 4, is evidently due to the fact that the magnetic power resident in beds X. and XII. just balance each other, and as the directive power of the earth is neutralized in the case of the red line by the direction in which the needle is held, we have a point of comparative equilibrium. (3) Why does the blue curve sometimes present depressions opposite the summits of the others? This is readily explained by supposing the local force to exist directly under the station; its force would then be entirely neutralized by the centre-pin of the horizontal compass, while having its full effect on the dip-needle in both positions.

5. Diminution of Intensity due to Elevation.

All the observations for intensity above considered were taken at an elevation of about 4 feet from the surface. Sometimes the rocks came to the surface, sometimes there were several feet and perhaps yards of drift between; it is therefore an important practical question to ascertain what effect the elevation of the needle has on the number of its vibrations.

The difficulty of attaining any considerable elevation at which to observe intensity, renders our observations on its rate of diminution due to elevation or vertical distance of little value. The theory of the sphere of attraction and law of decrease of force, as the square of the distance from the centre, has been mentioned; but with several local forces acting on the same point (the case usually presented in nature), the law is greatly modified, the decrease being in a less ratio. This subject possesses especial interest in connection with the determination of the depth at which magnetic rocks, producing a given disturbance, will be found; therefore, the few observations made, unsatisfactory though they are, will be

given. At Republic Mountain a staging was erected in the windfall, by means of which eight equi-distant observations were made; the lower one on the magnetic schist, the upper one 14 feet above it. The results were as follows:

Elevation in feet.	Vibra	TIONS.	Remarks.
	Facing west.	Facing north.	
0 2 4 6 8 10 12 14	56 41 33 27 ½ 19 15 ½ 18	53 41 30 30 23 24 24 20	On surface of schist.

At another point near the above, and over the same magnetic rock, the following vibrations were observed:

Elevation in feet.	VIBRA	TIONS.	Remarks.		
	Facing west.	Facing north.	KEMARKS.		
0 3 6 9 12	60 50 36 25 181/ ₂	60 49 37 26 18	On surface of schist.		

The observations have all been represented graphically, but as no law was apparent, and as the figures can be easily reproduced, they are not given. The first table gave the most regular curve, but still too angular to attempt the application of a mathematical formula. They do not seem to me to afford a basis for calculation,

as to how high the appreciably magnetic influence of these rocks would extend. I have an impression, however, without being able to give any reason, that it would be considerably less than one half mile, which was shown to be the distance to which the influence of the same rocks extended horizontally. I cannot consider it probable that a needle would dip where an earth covering of over 2,000 feet exists, if such a case were possible. At the Champion mine, by the aid of shaft house No. 2, an elevation of 44 feet above the ore was attained, and the following observations made:

Elevation in feet.	VIBRATIONS.			REMARKS.			
	I.	п.	111.				
0 18 32 44	18½ 19 16½ 15½	17 ½ 17 16 ½	23 17 16½ 15	Level of surface of ore in shaft. Surface of ground. Girder of shaft house. Girder of shaft house.			

At other points at the Champion mine, 25, 32, 33 and 40 vibrations were observed, the compass being within 5 feet of the ore. The diminution here is quite regular and nearly as the distance. If the rate continue, the vibrations should reach the normal number (six for the instrument used) at about 150 feet; but it is highly improbable that this law would hold for the whole height.

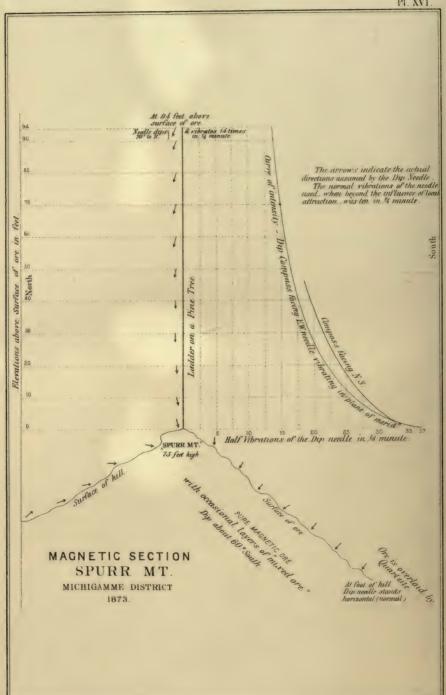
The difference between the rate of diminution at the two localities is very marked; at Republic Mountain an elevation of 12 feet in one instance reduced the vibrations from 60 to 18½, in another 14 feet elevation reduced the number from 56 to 12. At the Champion 44 feet elevation made an average of less than 4 difference in the vibrations. In this comparison the following geological differences must be borne in mind.

The Champion deposit at shaft No. 2 is a heavy bed of nearly pure black oxide running east and west and dipping north at an angle of 68 degrees, and it is the only magnetic rock in the vicinity. The Champion deposit loses its magnetism in going west, specular slate taking the place of the magnetite in that direction. The Republic Mountain bed over which the observations were made (No. X.*) is, on the contrary, a silicious schist, containing not to exceed 33 per cent. of magnetite, (the merchantable ores of Republic Mountain, of which there are large deposits, are in bed No. XIII., and are mostly specular hematites.) This magnetic bed X. is associated with others of a similar character, all striking north-west and south-east and dipping nearly vertical. The specimens of these magnetic schists which were examined possessed marked polarity. The Champion deposit evidently contains far more magnetite within the same sphere of influence than the Republic Mountain.

There is no doubt that variations and dips are a much more delicate and ready means of observing slight magnetic attractions, than vibrations when observed with the hand instruments employed. In one instance at Republic Mountain the dip at 12 feet elevation was 30 degrees, at 9 feet 50 degrees, at 6 feet 70 degrees, at 3 feet 77 degrees, at o or on surface of rock 105 degrees. It appears that the magnetic poles of the Champion bed are more deeply seated than those at Republic Mountain, which seem to be at the surface. This may be due to the fact that the upper part of the Champion deposit is mined out. Sets of careful observations made for considerable heights, both for dip and vibrations, would possess great interest, especially if made over beds of ore or rock, the position and character of which were known. In a record of over three thousand magnetic observations made by me in Michigan, Missouri, New York and New Jersey, I have not in more than six instances found the needle in the dip-compass above described to vibrate over 40 times in a quarter of a minute, and in no instance in which this rate was observed was the needle removed more than 5 feet from the magnetic mineral. Of course in the same needle the vibrations will vary with the degree of magnetism that has been imparted to it, and the condition of the instrument in other respects. I have had a rude standard, and when my needle fell below that it was overhauled, so that the numbers are relatively correct. I do not remember to have observed over 15 vibrations in a quarter of a

^{*} The Roman numerals refer to the order of the beds of the Huronian series, counting upwards from I. to XIX.





minute, or one per second, at a greater distance than 50 feet from a magnetic bed, and usually this number of vibrations would indicate a distance not exceeding 25 feet in the Marquette Iron Region.

Since the above observations were made and recorded, the development of the Michigamme district has permitted observations to be made at the *Spurr Mountain* which throw much light on the subject of the diminution in dip, and intensity due to elevation. The following table records the observations made,* and Plate XVI. represents the general law of diminution graphically. The observations were carried to an elevation of 94 feet by means of a fortunately situated pine-tree, up which a ladder was constructed. While there are minor irregularities, due wholly or in part to errors of instrument, the presence of nails in the ladder, and personal error, the average curve is remarkably regular, and points, as most of the other facts do, to a far more rapid rate of diminution near the surface than at a considerable elevation.

It is not to be expected that the law of decrease of magnetic force would hold at this locality. Had the local force been concentrated in a focal point directly under the tree, and the force of the earth been neutralized, then we might expect the law to be discernible. Some useful practical rules may be readily drawn from the table and plate under consideration.

I. If, in a locality where magnetic attractions prevail, we find considerable difference in the number of vibrations between the compass, when in contact with the ground, or held six feet above it, we may conclude the ore is very near the surface; if there is but little difference, then the ore is probably deep. 2. The amount of the dip gives but little clue to the depth of the ore. If the Spurr Mountain had been covered by a hundred feet of earth, water, or non-magnetic rock, we would have found at the surface a dip of about 70°, and it is probable, if not certain, that if it were possible to make observations to the north of the mountain, at the same elevation, a greater dip than 70° would be found, due to the changed direction of the local force.

^{*} Mr. C. M. Boss rendered great assistance in these observations.

Observations for diminution of magnetic force in vertical direction—(Needle vibrating in north-south plane)—Spurr Mountain.

Height.	Dip.	Vib. in ¼ min.	Remarks.	Height.	Dip.	Vib. in 1/4 min.	Remarks.
0 2 4 ¹ / ₂ 6 7 8 9 10 11 12 13 14 15 16 17 18 20 21 22 23 24	100° 100° 100° 100° 100° 100° 100° 100°	37 33 30 28½ 27 27½ 26 25½ 24½ 24½ 23½ 23½ 23½ 21½ 22½ 21½ 22 21½ 22	On surface of ore. 26 facing south. 24½ facing south. 23 facing south.	25 26 27 28 29 30 31 4 32 1 2 33 34 1 4 2 4 5 6 3 6 8 7 9 9 9 9 4	93° 93° 92° 92° 92° 91° 90° 88° 88° 88° 86° 78° 78° 78° 70°	20½ 20 20 19½ 19½ 19½ 19½ 18 18½ 18 18½ 17 17 17 17 16½ 16 15 14½ 14½ 14	15 vib., faced N., dip 86 E. 15 Vib., faced N. 13 " " " S. 13 " " S.

6. What is the significance of a dip of 90° or "dead 90°."

As there is a general impression among those who have made but little use of the dip-needle in exploring for iron ores, that a variation of 90° signifies merchantable ore directly under the feet, it is important to ascertain the exact purport of such great fluctuations in the direction of the needle. For the present we will leave out of the question the unpleasant fact that in 19 cases out of 20, if not 99 out of a 100, the mineral producing the dip would, if found, prove to be only a ferruginous schist or magnetic rock instead of merchantable ore, and consider the case often presented where there is a dip of 90° at a place which is not underlaid by magnetic mineral, or where there is none within several hundred feet. In such cases there are generally two, approximately parallel lines, of 90° dip, one over the ore, where the vibrations are very

quick (always more than the normal number). The second line (the one we are now considering) will always be found north of the first, and along it the vibrations will be slow, always less than the normal number.

A moment's inspection of almost any magnetic section (Plates VIII. to XIV.) will illustrate the fact and suggest the cause. If we hold a dip-compass over a highly magnetic bed the needle will indicate 90°, pointing directly towards it. Moving north, the needle will continue to point towards the ore, that is, be turned backward, thus varying or dipping more than 90° from its normal direction. Continuing north, we soon get so far from the local influence, that its power ceases to entirely overbalance that of the earth, and the needle commences to return to its normal direction. In doing so it must evidently somewhere stand again at 90°, which means simply, that the local force to the south and the earth's force to the north, are so related in intensity, that the resultant is a vertical Still going north, the dip grows less and less until the boundary of the local attraction is past, and the needle returns to its terrestrial allegiance. It is evident that no such phenomena can occur to the south of the magnetic bed, for the terrestrial and local influences acting in the same direction, no "dead" points could occur.

This "dead 90°" line, then, instead of proving the immediate presence of ore, proves just the reverse if the phenomena are presented as above, which is the case at the Magnetic, Champion, Spurr, and Michigamme mines, and at one place at the Washington mine. There may be ore under this line, but it will always be deep and have little or no influence in producing the phenomena observed. Rule:—When there is a dip of 90°, and the vibrations exceed the normal number, we may conclude that the magnetic mineral is under our feet, or very near us to the south. If, with the same dip, the vibrations are less than the normal number, we may conclude that the magnetic bed producing the effect is south of us, and may be at considerable distance. This rule will evidently apply only where there is one strongly magnetic bed not very deep, which is the most common case. If there be several beds, as at Republic Mountain, the application of the principle is more difficult; but in the nature of force, some modification of the phenomena must be presented by all magnetic rocks.

It is worth remarking that the south belt of 90° dip, is more sharply bounded, especially on its south side, and usually narrower, than the north belt.

The lower curve of Fig. 2, Republic Mountain Chart, illustrates what has been said above; the summits of the curves showing the maximum dips are north of the magnetic beds, while the summits of the curves showing the maximum intensity (see upper curve, same fig.) are over the centres of the magnetic beds.

This subject would be incomplete without considering the case, quite common, where a zone of local attractions has but one line of 90° dip, or to make the case general, but one line of maximum dip whether it be 90° or less. It may be said that this last expression, covers the whole question, but with ore-hunters "dead 90" has a peculiar significance, and it is for them that I am writing. This case (one line of 90° dip) is illustrated in some of the Champion mine sections. A few words will explain how it flows out of the first case.

If we follow the two lines of 90° dip to where the earth covering becomes very deep, so that our distance from the magnetic mineral considerably reduces its influence, our two lines would evidently be merged into one, and continuing on to where the earth was still deeper, which has the effect of raising us above the ore deposit, this maximum dip would become less than 90°.

This maxima line would evidently correspond with the south line of 90° dip in the case first supposed, that is to say would lie nearly over the mineral producing it. With great depth of earth covering, it can be proven that it would lie to the north of the magnetic bed.

An inspection and consideration of the facts presented in the Spurr Mountain magnetic section given above, will, I think, convince any one of this without the aid of the rigid mechanical demonstration which the problem admits of.

I have seen large amounts of money unsuccessfully expended in digging for iron ore for want of a knowledge of the simple principles set forth above, hence I have dwelt longer on this point than its importance to the general subject would seem to warrant.

7. Additional Practical Suggestions and Rules.

The facts given above, with others in my possession, enable us to answer provisionally the following practical questions:

I. Can we by means of the magnetic needle determine the order of superposition or succession of beds of the iron-bearing rocks?

Comparing the magnetic sections obtained at the Republic mountain and Champion mines, it is evident that, while there is considerable variation in the details, the salient features agree remarkably, pointing towards the same order and same lithological character in the rocks. A number of other sections made within 10 miles of the above-named localities, across the same belt of rocks, gave the same general result.

It is therefore asserted with much confidence that where a magnetic section similar to these is found in the Michigamme district, a corresponding geological section will be found beneath the surface; and that, as a rule, there will be less difference in the magnetic sections than in the topographical, which we know depends greatly on the underlying rocks. But whoever expects to find many places where so complete sections can be obtained as these localities afford, will be disappointed, for they present rare opportunities for studying the structure and magnetism of the Huronian series.

In places the covering of drift will be so deep as greatly to reduce the intensity, making it exceedingly difficult to observe with ordinary instruments, as was the case at the Cannon location. Again, the lower magnetic rock, beds VI., VIII., and X., are in places far less magnetic, containing sometimes very little magnetite, as is the case south of the Washington mine. In other places the lower magnetic rocks may be entirely wanting, owing probably to a fault, as at the west end of the Champion. On the north shore of Lake Michigamme there is a magnetic bed above XIII. (the ore formation), being therefore younger than any member of the Republic mountain series. In other places XIII. is wanting, and when present it is sometimes highly magnetic, as at the Champion, and again it holds very little magnetite, as at Republic mountain, the pure ore there being mostly specular hematite, as has been elsewhere observed.

With all these uncertainties, however, the results of magnetic

surveys cannot but be valuable in the exploration and development of iron properties, and in the solution of all questions of structural geology in regions of magnetic rocks. In such rocks, I believe, their value to the geologist is only second to topographical work, and, considering the cheapness of magnetic surveys, they may often pay best if means be limited.

Detailed magnetic observations, if made with precision, ought to throw light on the lithological character and intricacies of structure of these rocks, and on the nature of the magnetic force resident in them. This could not, however, be undertaken; the work done is more than was contemplated in my instructions and more than was justified by the means at my disposal.

II. Is it possible to determine quality—i.e., the percentage of iron—in a magnetic rock by means of the magnetic needle? In other words, can the needle alone make us sure we have a workable deposit of ore under our feet?

This is the most important practical question connected with this subject, and is the one constantly presented to the miner and explorer. Magnetic observations should always be made in connection with topographical and geological surveys; whether these take such names, and are based on instrumentation, or whether they be such rude work as the explorer is constantly doing, but which are as much topographical and geological as the other, and often quite as valuable. A judgment of the commercial value of a bed of magnetic ore should, of course, be based on all the facts available. If nothing more was known than what the magnetic needle revealed, I would not venture an opinion as to whether it was merchantable ore or magnetic rock which produced the phenomenon. In the Marquette Region, as has been before observed, the chances are at least fifty to one that a worthless ferruginous rock is the cause of any observed attraction. But this case never occurs; we always know something more than the needle reveals. One of the most important uses of the needle, and one for which it can within certain limits be depended on, is in tracing magnetic beds in the direction of their strike until some outcrop, which may give us the information sought, is found. I have in this way traced magnetic beds for many miles both in the Marquette Region and in New York and New Jersey.

Preparatory to the examination of any particular range of ore,

the explorer should thoroughly study up, with his own instrument, the phenomena presented at some exposed or developed part of the range he is exploring. This will give him data relating to variations, dips, and vibrations, which can be used where the rocks are covered and unknown. By means of the quickness of the vibrations, or of the rapidity with which they decrease as the compass is elevated, he may judge approximately of the depth of the drift, and so of other phenomena.

III. Does the magnetic needle afford the means of determining the absolute thickness of a bed of magnetic ore or rock?

My observations do not permit an affirmative answer to this question, especially if there be much earth covering. A study of all the magnetic sections which have come under my observation, indicates that, while in some instances the *comparative width* is plainly shown, the boundaries between the magnetic and non-magnetic rocks are not generally brought out sufficiently to warrant a definite expression as to thickness. We should expect this, because the magnetic influence is centred in the poles of the masses, and towards such foci the needle tends to point.

IV. Can we by means of the magnetic needle ascertain the direction and depth of a local magnetic pole? In other words, can we determine the thickness of rock or earth covering which overlies a given magnetic rock?

Often I think we can, with much precision, locate a point in the surface over the pole and determine its depth, by making what may be called a magnetic triangulation. Proceed thus: Remote from any magnetic rocks, neutralize, by means of a bar magnet, the earth's influence on the needle of a solar compass. The needle will then stand indifferently in all directions, and will not vibrate. Record carefully the distance and position of the neutralizing magnet: the compass is then ready for use. Set it up near the magnetic pole to be determined, and fix the magnet in exactly the same relative position it had before. The earth's directive power on the needle will again be neutralized, and the needle will point as nearly towards the local pole as its mode of mounting will permit; mark the line indicated by the needle on the ground; remove the compass to one, or, better, two other positions, and repeat the operation. If there is no other local force to interfere, the three lines must intersect in one point, which will be directly over the pole whose position is sought. By using a dip-compass in a similar manner, it is evident that the data to determine the depth, by the simple solution of a triangle, would be obtained. The fact that several local poles often influence the needle at each station renders this operation difficult in practice;—we should endeavor to find a place where but one strong pole exists.

A magnetic needle having universal motion, like Mr. Ritchie's, would evidently determine both position and depth at the same time; but a solar compass would have to be used to fix the position of the artificial magnet used in neutralizing the earth's force, unless it be fixed by an observation on the North Star, or by a meridian line brought in from a non-magnetic area.

V. When considering the magnetism of the rocks of the four great geological epochs represented on the Upper Peninsula of Michigan, I observed that considerable magnetic variations were noted by the Federal surveyors, over rocks of Silurian age, which had never been observed to be in themselves magnetic. In some instances these variations had been observed over a limestone, supposed to be Trenton, and at a distance of 75 miles from the nearest Huronian, or other (known to be) magnetic rocks.

This phenomenon may be due either: I. To the presence of magnetite in such rocks, due to local metamorphism or other cause.

2. To accumulations of magnetic sand in the drift; or, 3. To the underlying Huronian rocks, which may be supposed to exert their influence up through the overlying Silurian.

Without having made a study of any of these localities, I incline decidedly to the latter hypothesis, as accounting for the known facts better than either of the others.

Should this prove true (and I hope to settle it at some future time) it may lead to a novel and interesting application of the science of magnetism to some important questions in geology—the determination of the thickness of sedimentary rocks by magnetic triangulation in places where it would otherwise be difficult to arrive at such thickness. It might also enable us to work out the structure and distribution, in a rough way, of these oldest rocks which underlie great Silurian areas, which would in no other practicable way be possible, thus throwing light on the nature of the rocky bottom of the ancient seas.

On the same principle we can, of course, trace magnetic iron

belts under water. I have in many instances made very satisfactory magnetic observations from a canoe in the inland lakes of the Upper Peninsula. The bottom of Lake Superior may be thus partially mapped. Silt and sand will make no difference with the needle; it looks through everything but iron.

I have endeavored in the above to set forth plainly just what has been done in this comparatively new field, to give the results obtained, and to call attention to those principles which underlie the use of the magnetic needle in exploration for iron ores. The time and means at my disposal were meagre, my instruments imperfect, and I had no precedent to follow. I am persuaded that the subject is worth the attention of the explorer, miner, geologist, and physicist.

There has been a good deal written bearing on the subject of the Magnetism of Rocks, my references having very much increased of late. I had proposed to examine these authorities before writing this paper, but unfortunately the best libraries of Michigan do not contain any of the works referred to, and not being able to have abstracts made in Eastern libraries. I have derived no benefit from these authorities.* Could I have examined the results of the magnetic observations which must have been made in the great iron regions of Sweden, Norway, and Russia, I should probably have found my meagre results anticipated, and this article might not have been written. I am confident, however, that the Huronian rocks of Michigan have never been magnetically studied, and it may be that the methods that have been used in Europe are not such as would commend themselves to Lake Superior explorers, miners, and surveyors, who require cheap, light, and simple instruments that admit of rapid use.

The State of Michigan, or those interested in her Iron Regions, may at some future time see fit to have this subject thoroughly investigated. To that future investigator I commend my notes, trusting that he may find in them a reconnoissance of his rich field of labor.

^{*} Gilbert's Annalen (German) contains several papers. See volumes 3, 4, 5, 16, 26, 28, 32, 35, 44, 52, 53, and 75.

CHAPTER IX.

METHOD AND COST OF MINING SPECULAR AND MAGNETIC ORES.*

THE iron ores of the Marquette region are mostly extracted in open excavations; hence the process is more nearly allied to quarrying. Several attempts at underground work have been made, which have not, on the whole, been successful. The Edwards mine has been almost entirely wrought by candle-light. The slate ore pit No. I of the New England mine was worked in the same way, as is also the Pioneer furnace pit of the Jackson mine.

The Champion mine was opened systematically for underground work, with two levels, sixty feet apart, and three shafts at distances apart along the bed of about 200 feet; but this idea has been so far modified that one-third of the ore of this mine is now extracted by daylight. The Cleveland mine has recently commenced to mine considerable ore underground.

Several other mines have, from time to time, worked underground stopes, but so far only temporarily; if such stopes could not be opened out to daylight, they have usually been abandoned. In brief, it may be said that no considerable amount of ore has as yet (1870) been mined underground in this region, and of that so mined very little has been taken out at a profit, and I may add that it seems to be the belief of the most experienced mining men that this state of things will hold for some time to come, for reasons which will appear.

Nearly the same remarks may be applied to the mines of the Iron Mountain region, Missouri, the ores of which are very similar in character to those of Marquette. Some of the New York and New Jersey magnetic deposits are also wrought open, but this is the exception, underground mining being there the rule.

^{*} Two papers on this subject read before the American Institute of Mining Engineers and the American Society of Civil Engineers, and published, are embodied in this chapter.



Photo.by Emery

Looking West-towards "Big W." LAKE SUPERIOR MINE

The following brief sketch of the geological structure of the Marquette iron deposits will indicate some advantages of the method of mining employed; the subject being more fully considered in the chapters on the geology of the Marquette and Menominee regions, and illustrated in maps Nos. III. to X. of Atlas. See also Plate VIII., representing Edwards mine.* The iron-bearing or Huronian series of rocks are stratified beds, the principal ore formation being overlaid by a quartzite, XIV., and underlaid by a diorite, or greenstone, XI. This ore formation is made up, first, of pure ore; second, of "mixed ore" (i.e., banded jasper and ore); and third, a soft, greenish schistose, or slaty rock (magnesian), which occurs in lensshaped beds which alternate with ore, thus often dividing the formation into two or more beds of ore, separated by rock. Usually the beds of both ore and rock thin out as they are followed in the direction of the strike from a centre of maximum thickness, producing irregular lentiform masses. Since their original deposition, if we may assume they were laid down under water, the whole series, including the iron beds, have been bent, folded and corrugated into irregular troughs, basins and domes, which often present at the surface their upturned edges of pure ore, standing nearly vertical. A cross-section, finely illustrating this structure, can be seen on the west of the great south-west opening of the Lake Superior mine. It is locally known as the "Big W," which letter is plainly suggested by the sharp folding of rock and ore. See Fig. 7 and View IV.

The fact that, as a rule, the richest ore is found near the upper part of the formation, and the most jaspery part near the base, has led to the separation of this formation into two beds, Nos. XII. and XIII.

This structure, involving sudden changes in the amount and direction of the dip, from horizontal to vertical, would evidently necessitate, in the case of underground work, constant changes in the plan of attacking the ore, as well as in the mode of supporting the roof.

The magnetic iron deposits in the Eastern States may also be regarded as true beds, but are far more regular in strike and dip, extending downward at a high angle to an undetermined depth, and appearing more like veins. If folds exist, they are much deeper and more regular than in the deposits under consideration. The

^{*} Many copies of this chapter will be distributed separately, rendering this geological resumé necessary.

Marquette ore deposits are often very thick, 50 feet being not infrequent, which makes ordinary timbering difficult, if not practically impossible; while the eastern deposits, so far as my observations have extended, are seldom over 20 feet, and average considerably less than that thickness.

The "pinch and shoot" structure, suggesting what are termed "chimneys" and "courses of ores" in some metalliferous mines and which is very apparent in the New York and New Jersey mines (practically dividing the ore into pod-shaped masses, the axes of which "pitch" in the planes of stratification in a direction quite different from the dip), can at this time be best observed in the Marquette region at the Edwards mine, Plate XIX. and Map No. VIII. Atlas. The intervening barren streaks where the hanging and footwalls come near together, and which therefore divide the "shoots," form excellent supports to the overlying rocks and give the mine great security, as all who have worked deposits having this structure will testify.

The soft schist mentioned as occasionally bedded with the Marquette ores, often constitutes the hanging wall in parts of the mine, but does not possess the requisite strength to make a good roof. It is impossible to support such rock with occasional timbers or pillars, for it will scale off between the supports, demoralizing the men, if not actually endangering their lives. Even when the works reach the solid quartzite XIV., which, as has been stated, is the true hanging wall-rock of the ore formation, it is sometimes not safe, particularly near the surface. These facts make open workings a practical necessity at the start, and the great economy of breaking ore from high stopes with heavy charges of powder induces a continuation of the method, even when the rock covering has attained a thickness of many yards, and underground work would seem to be advisable. It is, indeed, hard to say what thickness of solid rock a Marquette mine-superintendent would hesitate to remove if it covered a large deposit of ore. Forty feet of earth and nearly as many of quartzite (as hard as granite) have been "stripped," and the thickness of rock is daily growing greater as the beds of ore are followed in depth.

It may be said, and I do not know but that it is a canon of mining, that all mines, which sooner or later have to be wrought underground, should be systematically opened as mines at the start, but this is not Marquette practice; and I have undertaken to describe,

and, so far as I am justified, defend the methods there employed. It would be difficult to convince our people that, having a large deposit of pure ore before them of unknown form and size, covered often by but little earth, and backed by perhaps a small amount of money in the company's treasury, it is best to incur the delay and cost incident to sinking and drifting to open ground already opened by nature and ready to win. Wrought as open quarries, several of our mines have paid their way from the start, while, had they been opened on a regular system of mining, they would have required an investment of \$50,000 in plant and improvements before shipments could have begun, and at least one year's time. facts settle such questions with American capitalists; and with the uncertainties which attend the opening of new mines in new districts, the high rate of interest in this country, and uncertainty of tariff legislation regarding iron, it may be a question whether this hand to mouth—quick return—let the future take care of itself view of the question, is not in a certain degree defensible.

The appearance of our mines is anything but pleasing. They consist of several (sometimes of ten or more) irregular elongated pits, often very large and generally more or less connected, having usually an easterly and westerly trend imposed by the strike of the rocks. Everywhere are great piles of waste earth and rock, which are often in the way of the miner, and which in some instances have been handled over three times.

There are two principal advantages in open works. First, the preparatory work is all reduced to the simplest and safest kind of pick and shovel, hammer and drill, horse and cart business; such as can be let to the common run of mine contractors. On the other hand, underground mining involves sinking, drifting, timbering and elaborate machinery, all of which require skilled labor and large investments. In an isolated cold country like Marquette, the quality of the labor demanded is an important consideration. The second advantage, already mentioned, is the great economy in cost of drilling and explosives which high stopes in open works permit. These elements of cost are important items in all mining where hard ores are encountered. It is believed that they have been reduced to a minimum in the Marquette iron mines, where holes two inches in diameter are sometimes sunk 22 feet, and 15 feet is common. Such holes are not fired directly with the blasting charge, but are

"shook" several times first, that is, fired with small charges which produce cracks and cavities about the bottom of the hole; when these are large enough to contain a sufficient amount of powder, the lifting charge is put in and the great mass thrown down. Twenty kegs of powder, of 25 pounds each, are sometimes fired at once, and from five to ten kegs is not an uncommon charge for a stope hole. By this method 5,000 tons of material have, in some instances, been removed at one blast, and one-third of that amount is quite common at some of the mines. In this way the entire cost for labor of drilling and explosives has been reduced, for a single blast, to less than three cents per ton. But the average cost is of course much greater, being at some of the mines 50 cents for all the drilling and powder consumed in the mine; about one-third of this is for block-holing the large masses thrown down by the stope holes, which are often so large that they have in turn to be broken by powder. The cost of powder and fuse for the hard ore mines, it is believed, does not exceed ten cents per ton. In some of the New York and New Jersey mines, which are worked underground, I am informed that these items cost much more. In the Persberg mines, Sweden, the drilling and explosives cost 65 cents per ton of ore in 1870.

It may be inferred, from the above description, that Marquette iron mining does not differ essentially from ordinary rock excavation on public works, being work that may be let by the cubic yard or ton. Until quite recently this has been very near the truth, the difference being in the skill and care required in separating the ore and rock which are often mixed together in the deposit. But these palmy days are rapidly passing for most of the mines now worked. An increase of water and greater cost of handling incident to increased depth, and, what is still more costly, the increase in thickness of the rock covering, will soon require, in fact does now (1870) really require, more expensive plants, different methods, and more skill.

The transition from the present system of quarrying to the future method of underground mining, which will have to be made in the Marquette region, will be a critical period, and will possess great interest, as affording a solution of a mining problem such as may not yet have been presented anywhere. Attempts at its solution have already been made, but, as has been remarked, very little ore has as yet been extracted at a profit by candle-light. To recapitu-

late, the system adopted will have to meet the case, 1st, of beds of ore varying, often abruptly, in thickness from 0 to 50 feet; 2d, of beds varying in dip from nearly vertical to horizontal, and passing by a curve of small radius from one inclination to another; 3d, of beds varying in character of hanging wall from a solid quartzite, which will stand with ordinary supports, to a soft schist, which can only be kept in place by a continuous support, or by actual filling in—"remblais." Again, the axes of the folds are not horizontal, but sometimes "pitch" at angles of 30 degs. or more in the direction of the strike, producing a fourth troublesome feature. See Map IX. Now, when we consider that the dressed ore is expected to yield 65 per cent. in the furnace, and is seldom worth on the average over \$4 or \$5 per gross ton on the cars at the mine, including royalty, the general character of the problem will be understood.

In New Jersey, with perfect regularity in the dip, better hanging walls, thickness within the limits of easy timbering, cheaper fuel and labor, and material which breaks easier than that of Lake Superior, the ores of several well-known mines, I am told, cost fully this amount.

Steam machinery for hoisting and pumping, which has cost from six to not less than fifty thousand dollars, has been erected at most of the Marquette mines, as shown by the table at the end of this chapter. In 1870, however, not much more than one-half of the entire ore product of the region was handled by steam, and much less than this proportion of all the material, the balance being done by horses, the use of which, however, is decreasing.

From these facts it may be inferred, that while the cost of breaking ore may have been reduced to a minimum by the system of mining employed, not so much can be said in favor of the methods of handling the ore from the miner's hands to the cars. The expensive horse and cart, swing derrick and whim, are in too general use, and the roads over which the loads are hauled are often not above criticism as to grades and surface. The causes which have led to this extensive use of horses are considered in another place.

The local staff of a Lake Superior Iron Mining Company usually consists of the *agent*, who is often secretary or treasurer of the company, and whose duty it is to take general charge of the company's business, except selling the ore, which is commonly done by a special agent in Cleveland, who may or may not be an officer of the company. This agent supervises the accounts, makes the pay-

ments, attends to shipping the ore and to ordering supplies, and often assists in selling ore. One man sometimes represents more than one company in this capacity. A majority of the agents reside in Marquette. The superintendent, who by custom has the title of captain, always resides at the mine, directs the work, and is in the main responsible for it. On him as much or more than on any other officer of the company does the success of its operations depend.

The offices of agent and superintendent are sometimes united in the same man. Large mines have a chief clerk, who is practically assistant superintendent. Next in order of rank are the foremen. master mechanics, and time-keepers. For names and addresses of agents, superintendents and managing officers, see Statistical Table, Plate XII. of Atlas.

The organization of the force of two large mines in the summer of 1870 is shown below. The first mine (I.) shipped the greatest amount of ore, and the second (II.) did most of its dead work in the winter, the aggregate shipments for the two, for that year, being 300,000 tons. Τ.

TT.

	1.	11.
Contractors engaged in stripping, sinking, etc	77	7
Company account, men, laborers and mechanics		
on miscellaneous work	65	32
Total employed in dead work	142	39
Contractors breaking ore	117	114
Company account, men breaking ore		25
Total at mining proper	117	139
Carpenters and wagon-makers	6	6
Blacksmiths and helpers	17	10
Total mechanics employed in repairing	23	
Drivers and stable-men	20	12
Engineers and firemen	II	8
Loading ore from stock-pile	18	
Total handling ore	49	20
Superintendent and clerks	3	3
Foreman, blaster and watchman	6	. 7
Total staff at mine	9	10
Total force employed	340	224
Total force employed	340	224

This force was employed during the period of shipments, hence of greatest activity; after the close of navigation, in November, it would probably be reduced 25 per cent. Less than one-half of the men employed have families, many single men going "outside" in the fall and returning in the spring.

One large mine, the best managed in the region, expended in 1872, 51,000 days' work all told, of which 48 per cent. was by contractors, and 52 per cent. by the day or on company's account: it produced about 2½ tons of ore for each day's work.

The wages of the men employed in and about the mines, in 1869 and 1870, were about as follows: Common labor was nominally \$1.80 per day for most of the time, but by far the largest part of the mining work was done under contracts. Contractors made, clear of costs, from \$60 to \$77 per month as high and low averages; \$70 is probably near the mean of the whole. It was not uncommon for a "pair" (two or more men working jointly) to make \$100 per month each, and again the earnings will fall so low as barely to pay board; but such are extreme cases. Leaving out the staff of the mine and the contractors, the wages of all others, mechanics, engineers, firemen, drivers, but mostly common laborers, averaged in 1869 and 1870 about \$2.12 per man per day. Mechanics received from \$2.50 to \$4.00. In 1872 the wages of men and contract prices were from 25 to 50 per cent. above the figures here given.

The nationality at three mines, which employed an aggregate of over 600 men, was in 1870 as follows, expressed in percentages:

Irish	. 31
English (Cornishmen)	. 27
Swedes	. 18
Canadians (French)	. 5
Americans	. 5
Germans	. 4
Norwegians, Danes, and Scotch	. 10
	100

The relative proportion of the Irish element is decreasing; a few years since nearly all the men employed at some mines being of this nationality. The percentage of Cornishmen is increasing,

owing largely to a want of work in the copper region. These men are skilled miners, and do a large part of the sinking and drifting. Swedes are rapidly gaining in numbers, many of them having been miners in their own country.

The exodus of Swedes to the United States apparently threatens to depopulate that country. There can be little doubt but that a more genial climate and better food will improve the lower class, from whom the emigrants come. Statistics of the population of the Upper Peninsula are given in App. G, Vol. II.

The unit of measure and comparison in the following table is the gross ton of merchantable ore. The ore is the object of the miner's efforts, and the tons sold measure his business. The items of cost in all that follows express the expenditure per ton of ore mined, prepared for market, and loaded on the cars. In instituting a comparison between these figures and those obtained by the civil engineer on public works excavations, where the cubic yard of vacant space is the ordinary unit of work accomplished, it must be borne in mind that the labor incident to sledging up and sorting out the ore from the rock considerably enhances the cost of mining.

In order to more intelligently follow the methods of working the Marquette mines, we must classify the various items of cost under appropriate heads, and assume some absolute cost per gross ton, as near the actual fact as possible, as a basis of comparison of these items with each other, and with other mining regions.

No discussion of the question which leaves out the cost, would possess much practical interest; but all who have undertaken to obtain such facts for publication, know the difficulty, and will not place implicit reliance on the accuracy of what follows. \$2.64 per gross ton will be assumed as the entire cost of mining the hard ore, and delivering it in the cars ready for shipment (in 1870); but this sum does not include interest on capital, expense of selling, royalty or mine rent, nor depreciation of the mining property. The cost of mining the soft hematite ores is considerably less, and the methods much simpler.

Royalties or mine rents have not become settled; there are not many leased mines; one of the best of its kind (the New York) pays but 20 cents per ton for first-class specular ore. In other in-

stances 75 cents is paid for a lean hematite. Time and experience will settle these prices on an equitable basis. See Atlas, Tables XII., XIII.

Before dismissing the subject of royalty or mine rent, which is not again noticed in the following discussion, I will make a few remarks. Marquette mines, as has been stated, are generally owned and worked by the same parties, hence royalty does not enter directly as an item of cost, but it exists in substance, and may be called depreciation of the mine, an item in the cost of ore often not sufficiently considered. One of the best organized and successfully operated iron companies in eastern Pennsylvania place this item at fifty cents per ton of ore. That is to say, every ton of ore sent from a New Jersey mine (which they own) is charged with fifty cents over and above its cost, as shown by the mine accounts, and a like sum is credited to the capital stock account, or to a sinking fund. This fifty cents stands for the original cost of the ore in the ground, and is all the more real, that it was paid in advance in the price of property and improvements. Any mining company which fails to recognize this principle is doomed some day to serious disappointment. Whoever has had experience with charcoal blast furnaces, which so rapidly sink their capital by the consumption of timber, will be fully alive to the importance of this matter. It is a delusion to suppose that our mines will not eventually be exhausted; iron ores do not grow; a ton shipped from a mine is gone forever, and the property has one ton less remaining, and is therefore worth less money. Continued shipments will eventually exhaust any and all deposits. Abandoned pits, in which no ore can be found, now exist at all of our mines, and in this class are some that two years ago were the best. The Andover mine, New Jersey, once presented as good opportunity to break ore as any pit now worked in the Marquette region; but about 150,000 tons aggregate product exhausted the mine, and to-day the owners do not know where to find a ton of merchantable ore on the property. I do not wish to be understood as predicting the exhaustion of the whole region; I think Marquette will produce iron as long as that article is wanted. New deposits of rich ore will be found, and leaner ones, which now have no value, will be worked, and the old deposits will be followed deeper; but this implies new mines, the building up of new locations, new railroads, new men and more

capital. What I wish to say is, that unless present holders of average Lake Superior iron mining stocks are receiving fair interest on their investments, and in addition are being paid back the capital they have invested at the rate of, say, 50 to 75 cents per ton of ore sold, they are not doing a good business.

Therefore the \$2.64 assumed in the following table should be increased by this royalty, making it \$3.14. Commission for selling, interest and exchange, insurance and expenses of the general office of the company (including salaries), will increase this sum to at least \$3.50, which will more truly represent the actual cost per gross ton of ore on cars and sold. This, from the amount assumed before as selling price, leaves from 50 cents to \$1.50 per ton for interest on all fixed capital invested; in an exceptional condition of the market, like 1872 and 1873, the margins are of course larger.

There may be no better place than in this connection, to speak of another fruitful source of the disappointments which are sometimes experienced by stockholders. I refer to those delusive "permanent improvement accounts," better named permanent disappointment accounts, which are too often kept open, and in which are too frequently placed awkward sums which should properly go to running expenses, and be paid for by the pig-iron, ore, lumber, or whatever is produced. After the necessary real estate is bought, the mining or manufacturing plant built, and the business of production actually commenced, the improvement account should be closed forever. Some kinds of business, in some places, under some managements, may permit an opposite course, but the above is the only safe rule. If in any particular year an extraordinary expenditure is made which is not likely to be repeated, a part of it may properly be held in some open account, in order that it may be distributed over more than one year's product. But this is a different thing from piling up a permanent account under the delusion that the property is enhancing in value.

There are few kinds of business in which there is more danger from this cause than in iron mining, for not only is an iron-ore property depreciating from the exhaustion of the ore, but at any time it may be still more depreciated by unfriendly tariff legislation, for which the iron-master must be prepared.

Table showing the Approximate Cost of Mining the Specular and Magnetic Ores of Lake Superior, made in 1870.

		APPROXIMATE COST OF EACH ITEM.					
General heads under which cost of mining	Elements of cost, not in- cluding royalty or depre- ciation.	In per cent. of the whole.			on a total cost of \$2.64 per ton.		
is classified.		Items.		als.	als.	Amounts.	
				Items.	Totals.	Labor.	Supplies.
I. Dead work (pre- paration).	1. Explorations 2. Sinking shafts 3. Drifts and tunnels. 4. Roads 5. Stripping earth and rock 6. Miscellaneous work and minor improvements *	00.6 01.5 06.1 00.6 13.2	28.1	.015 .040 .160 .017	.742	Eighty per cent.	Twenty per cent.
II. Mining proper (labor). Drilling. Other work.	f. Ledge holes (in stope) 2. Block holes (in fragments) 3. Sledging, sorting, and loading 4. Handling rock 5. Miscellaneous work	04.2	39.8	.110	1.050	1.050	
III. Mining materials and implements ("mine costs").	1. Powder and fuse. 2. Nitro-glycerine 3. Steel (drills) 4. Tools other than drills 5. Blacksmiths' supplies 6. Blacksmiths' labor.	03.6 † 00.7 01.6 01.8 04.2	11.9	.095 † .018 .043 .047 .110	.313	.103	.210
IV. Handon By horses, ore from miners' hands to cars, and pumping. By steam.	I. Teaming, labor of drivers and stablemen. 2. Forage	05.7 04.2 00.2 01.3 04.2	15.6	.150 .110 .006 .035 .112	.413	.272	.141
V. Management and general ex- penses.		04.6	04.6	.122	.122	.062	.060
		100.0	100.0	2.64	2.64	2.107	0.533

^{*} Does not include exceptional permanent improvements.

f No reliable figures obtained.

In order to institute a comparison between American open-excavation mining and the systematic underground work of Sweden, I append the following table, for which I am indebted to Prof. Richard Akerman, of Stockholm:—

COST OF MINING ORE IN PERSBERG MINES, SWEDEN, 1870.

In currency.

General heads under which cost	Elements or items of cost, not including royalty or depreci-	In percei		Based on a total cost of \$2.20 p. ton.	
of mining is classified.	ation.	Items.	Totals.	Items.	Totals.
I. In the mine	1. Boring 2. Powder. 3. Priming reed. 4. Clay 5. Candles, augers, and sledges. 6. Charcoal 7. Auger whetting 8. Shooters' fees.	22.73 5.82 0.84 0.50 2.33 0.44 3.20 2.80	38.66	.50 .13 .02 .01 .05 .01 .07	.85
II. Water drawing (or pumping)	{ r. Water drawing	3.50	3.50	.07	.07
III. Bringing up the mountain (hoisting rock and ore)	1. Putting into the ton 2. Receiving 3. Down freight 4. Hoisting 5. Oil and lines 6. Mine tubs and ladders	4.68 1.11 1.10 2.74 0.28 1.26	11.17	.10 .03 .03 .06 .01	.26
IV.	{ I. Dressing	8.12}	8.12	.18}	.18
V. Picking and washing	{ I. Picking and washing	5.65}	5.65	.12 }	.12
VI. Buildings	{ I. Buildings	16.45}	16.45	.36 }	.36
VII. General expenses	{ I. General expenses	16.45 }	16.45	.36 }	.36
		100,00	100.00	2.20	2.20

Professor Akerman furnished also these explanations:—
a. Our drill holes are about one inch in diameter and cost 7½

to 12 cents currency, per foot, when boring downwards, and twice as much when boring upwards.

- b. Powder costs 11½ cents, dynamite 43 cents, and ammonium powder 40½ cents per Swedish pound (the Swedish lb. equals .93 of the English).
- c. The reason why blasting with us is more expensive than with you, must partly depend upon stronger mountain ground and partly upon the small diameter of our augers.
- d. "Dressing" on the Persberg table is to be understood as sledging and sorting.
- e. "Picking and washing" is a kind of after-sorting by hand of the smaller pieces (of which about a third of the ore consists), got partly by blasting and partly by the first sorting.
- f. "Buildings" include timbering in the mines and all buildings made for pumping and hoisting.
- g. "General expenses" include some benefits for the laborers, such as domiciles, potatoes, gardens, expenses for schools, medicine, administration, etc., etc.
- h. "Down freight" is the cost for bringing down the ore a short distance from the mines to the lake-shore, where it is sold.
- i. Water power is used at Persberg both for pumping and hoisting.
- j. Our miners receive from 48 to 75 cents per day, besides what I above called benefits.
- k. The mining costs at Persberg are among the highest in Sweden.

The titles of the several heads under which mining costs may be divided, and the number of the items, depend on the object sought: the classification employed in the Marquette table, seemed best adapted to the presentation of the facts in hand. It will be observed that the form of the Swedish table differs materially and is of course better adapted to underground work, and to a more careful and laborious selection of ore.

I believe that considerable advantage would accrue to many of the Marquette mines, if the accounts were so kept that cost sheets similar to the foregoing could be prepared from time to time. It is well known that the cost of mining varies greatly in the different mines, some costing twice as much as others. This difference is often largely owing to natural causes, but sometimes it is, in part at least, in the management. There is no better way, in fact there is no other way, of stopping "leaks" of this sort, than by first finding where they are.

A comparison of such cost sheets from different mines, for the same time, or from the same mine for different periods, would indicate at once to which items the excessive cost belongs, and thereby direct the attention of the management to the leak. I therefore venture the opinion, that a carefully prepared cost sheet is one of the first steps in attempting to reduce the cost of ore.

In the detailed description of methods which follows, the items will be taken up in the order of the table.*

I. DEAD WORK.

This general head embraces all the work and costs incident to getting ready to mine the ore, and is subdivided into—I. Explorations (embracing only such searches for ore as are in progress from year to year about the mine). 2. Sinking shafts. 3. Drifts and tunnels. 4. Roads for wagons. 5. Stripping earth and rock, or uncovering the ore. 6. Miscellaneous work and minor improvements. The entire expenditure for dead work is 74 cents per ton of ore produced, which equals 28 per cent. of the whole cost.

I. Explorations.—More or less digging of test-pits, sinking shafts, drifting, trenching, and sinking drill holes is constantly in progress at most of the mines. My facts indicate that this work varies in amount from one-half to three cents per ton at the producing mines, being of course greatest at the new locations. It is not carried on systematically, being pushed when there is an increased demand for ore, or some old pit shows signs of failing, and again entirely discontinued. The price paid for pits 4 feet by 6 feet, and not over 10 feet deep, is from 30 to 60 cents per foot, depending on the ground; when so deep as to require a windlass, 50 to 75 cents and up to \$1.25, if the shaft reach the depth of 30 feet and is wet. Drifting in firm earth will cost about the same per foot, depending

^{*} For detailed descriptions of all the mine workings as they were at the close of the season of 1872, see "Appendix to A. P. Swineford's History of the Lake Superior Iron Region," being a review of its mines and furnaces for 1872, published by the Marquette Mining Journal.

on the depth below the surface and nature of the earth. Drill holes sunk by hand, material 15 feet deep, will cost from 75 cents to \$1.00, and if deeper, considerably more per foot. There seems to be no reason why more use should not be made of the drill in this work. By means of a simple spring pole, such as was used in early days in the oil region, holes could be easily sunk 100 feet, which is as deep as it is usually necessary to go at this time. An experienced miner will judge very accurately of the ground passed through by the mud, and if there was any doubt, chemical analysis would determine the nature of the material; the mud furnishing a strictly average specimen, so desirable in an analysis for practical purposes. As has been mentioned, the annular diamond drill was introduced last season (in 1869) at the Lake Superior mine with success. hole 130 feet deep was sunk at a cost of about \$5-per foot; the core produced furnished very satisfactory knowledge of the substance passed through. The drill did not perform as well at the Washington mine, where several holes were sunk, the deepest 96 feet. In two instances the annular diamond bit got fast in an oblique seam and two were lost; not counting loss of diamonds, the work cost about \$1.50 per foot: whether larger bits, a different setting of the diamonds, or more experience would overcome this difficulty, I do not know. It is a matter of great importance, and is worth thoroughly working out. As the subject of exploration for ore has been fully considered in another chapter, it is not necessary to treat it farther here.

2. Sinking Shafts.—This work, which forms so large an item of cost in some underground mines, varies in the Marquette Region, so far as I have ascertained, from 1½ to 5½ cents per ton of ore. Our open and comparatively shallow workings do not call for many shafts or winzes; the deepest shaft in the region is now (1870) not over 200 feet. The prices for this work range from a mean of \$22.50 to \$31.50 per foot in depth, depending on the hardness of the ground. In some mines, extreme prices range from \$15.00 to \$40.00, and even more if the shaft be very wet. Miners are often permitted to select the size most advantageous to themselves, which may be four feet by six; but eight by twelve feet is more common. The material is generally hoisted with the ordinary hand windlass, but sometimes with a horse-whip or whim, the miner having to deliver the stuff at the mouth of the shaft. From 10 to 15 per cent. of the

price received by the miner for sinking has to be expended in *mine costs*; *i.e.*, powder, fuse, candles, steel, tools, etc. No charge is made against him for smith's work. Sometimes the contract is let at so much per foot of shaft and so much per ton of ore, which gives the miner an interest in separating ore from rock.

3. Drifting and Tunnelling.—This element of cost varied more widely than any other, and might have been divided into two: (1) Drifts designed to open ground for stoping; and (2) Tunnels or adits for drainage and transportation of ore, the latter being of the nature of a permanent improvement. But on the principle that permanent improvement accounts are often permanent disappointment accounts, and to be avoided, and considering the fact that this kind of work is actually going on year by year, and must do so as long as the mine is worked, it does not seem wise to separate it from the current cost of getting ore. Ordinary 4 × 7 drifts cost, in hard ore, from an average of \$22.50 to \$24.50 per foot, the miners delivering the material behind them, and paying their own costs, as in the case of shafts.

Tunnels large enough to admit railroad cars and small locomotives cost from \$30.00 to \$50.00 per foot. The Washington tunnel, now over 1,100 feet long, and timbered a considerable part of the way, cost an average of about \$40.00, not including rails. The timbered portion is twelve feet wide at the bottom, ten feet at the top, and ten feet high in the clear. No machinery has yet been brought to bear on either sinking shafts or drifting; the labor required is more than one-half expended in drilling holes for blasting. The subject of drilling is fully considered under its proper head.

- 4. **Making Wagon-Roads.**—The great amount of team-work employed about the mines requires a complete system of roads for summer and winter use. These are sometimes expensive on account of rock-cuts, costing, in some instances, as high as four cents per ton of ore in the early stages of work.
- 5. Stripping Earth and Rock, or uncovering the ore. This constitutes on the average nearly one-half of the dead-work, and is one of the largest single items in the whole cost of mining. So far as my inquiries extended I found it to vary from 20 to 52 cents per ton of ore. This cost is necessarily increasing at all of the mines worked as open cuts. It is simple rock and earth-work, the material being removed on wagons, carts, or sleds, drawn by horses.

The advantages of light railroads and small locomotives do not seem to have commended themselves for this work. There would, of course, be considerable danger of destroying tracks from blasting, and it often happens that not much work has to be done in one place; still there is no doubt but that a large saving would be effected by substituting steam for horses in portions of this work, as will be more fully considered hereafter.

The aggregate amount of material which has been handled in stripping is very great. Thirty and even forty feet of earth have been removed, and nearly as great a depth of rock; but this is the experience in open workings everywhere. I have seen twenty-one feet of earth and soft, shaly rock stripped from a nearly horizontal bed of 44 per cent. Clinton ore in Western New York, which did not average over thirty inches thick. In South-eastern Kentucky I found the rule among the miners of sub-carboniferous ores to be, that it would pay to remove a foot of earth for the sake of an inch of ore, which does not differ widely from the Western New York practice. In both of these instances the stripping was nearly the entire cost of mining, and labor was much lower than in the Marquette region. The usual contract price for removing ordinary earth (sand, clay, and boulders mixed together) is fifty cents per cubic yard, the digging costing about one-half, and the hauling onehalf. Hauls vary from 100 to 800 feet. The highest price paid for excavating any considerable quantity of rock in open cuts, which has come to my notice, was \$3.00 per cubic yard, equal to \$24.00 per fathom, or about \$1.00 per ton. This was a very hard jasper rock, containing but little ore. Large quantities of rock have been excavated and hauled over 500 feet at the Lake Superior mine for \$2.50 per yard. The soft greenish schist, so common at all the mines, can be moved for from \$1.00 to \$1.40 per yard, including hauling. When a good face can be obtained on the overlying quartzite, which is likely to constitute the greater part of the rock to be moved in future, it should be broken down and loaded on wagons for from \$1.50 to \$2.00 per cubic yard.

The amount of money which it will pay to expend in stripping of course depends chiefly on the quantity of ore uncovered. If we assume fifty cents to be the maximum expenditure per ton of ore for this work (this amount has been greatly exceeded), the problem of what thickness of rock may be stripped admits of an easy theo-

retical solution. One cubic yard of solid ore (allowing for wastage on account of associated rock) may be considered to yield three tons of merchantable ore, which, at the allowance above assumed, would give us \$1.50 to be expended per square yard in stripping a bed of ore only one yard thick. Hence in this case it would pay to remove nine feet in thickness of earth, or about three feet in thickness of rock. But suppose we have a bed of ore twenty-four feet in vertical thickness, which is a more common case, what amount of earth or rock would it pay to remove under the assumed limit of expenditure? Twenty-four feet of ore will yield twentyfour tons per square yard of surface, which, at fifty cents per ton, gives \$12.00 available for stripping per square vard. This sum would remove twenty-four feet thickness of solid rock; or a foot in thickness of rock may be stripped for every foot in thickness of ore uncovered, at a cost of fifty cents per ton of ore. The same expenditure will remove three times this thickness of earth.

An important and often neglected question connected with this subject is, where to deposit waste, that it may be out of the way of future mining operations. Some material has been already handled twice in the Marquette region, and I know of a mine in Southern New York where the same earth was three times handled before it was finally permitted to rest. In a new region, like Marquette, where comparatively little thorough exploring has been done, it is often difficult to decide where waste piles will be out of the way for all future time. If a drill hole were put down for fifty feet in rock, and no ore found, it would be safe to say, that if ore existed under that spot, it would have to be mined under ground; hence, that so far as future stripping was concerned, a waste pile placed there would be out of the way. A very common practice in under-ground work, in some mining regions, is to fill up the worked-out places with the waste, and this can undoubtedly be done to advantage in some instances in open works, although it has not as yet been practised in the Marquette region. The trouble is to find out when a pit is exhausted—it is so common to break through a thin layer of rock and find a bed of workable ore behind it. But there are parts of most mines where the foot-wall has unquestionably been reached, and if any doubt exists, a few deep drill-holes will settle the point. When this is the case, and the foot-wall has a sufficiently gentle slope to permit of its holding materials deposited

on it, it will, I think, be often found advantageous to use it to support a waste pile.

For the sake of illustration, take the New York and Cleveland Mine workings, which are adjacent. In this instance the slope of the foot-wall is so steep that it would probably be necessary to cut in it a rude step on which to rest a rough retaining wall, which could be built of blocks of quartzite swung across from the hanging-wall by means of a derrick. The triangular space thus formed would hold all the waste rock for a long time to come, and would afford a minimum haul. It might not answer to deposit earth in such positions, as heavy rains would be likely to wash it into the pits. The dip of the foot-wall in this, as well as in most cases, will, I think, become flatter in depth, so that a better opportunity will be afforded for a second similar waste receptacle at greater depth, if one should be required.

6. Miscellaneous Dead Work.—Under this head are included several items which were not of sufficient importance to require separate treatment. Improvements such as dwellings, shops, fences, tracks, trestle-works, pockets, docks, whims, skip-ways, pumpingfixtures, etc., etc., occurring from year to year, are embraced here. These items are in part embraced under "Building" in the Swedish table. This head was originally also designed to cover those exceptional expensive improvements which are of occasional occurrence only, and the cost of which might properly be distributed over several years' product. Additional facts, however, lead me to believe that the amount given (16 cents per ton) is too small. expensive pumping and winding plants now being erected, and which will continue to be built for a long time to come, increase the cost of the ore materially unless we charge them to permanent improvement accounts, which is not altogether a safe course, as has been already pointed out.

II. MINING PROPER, OR BREAKING ORE.

This general head embraces all the labor incident to blasting the materials down from the solid ledge, breaking it up into fragments that may be easily handled, the separation of the ore from the rock by hand and loading. The average cost of this is \$1.05 per ton of ore produced, which equals forty per cent.

of the whole. The character of this work will be sufficiently well understood from the table and the following explanation:—

I. Ledge or Stope Holes.—The drilling or rock-boring is now (1870) entirely done by hand. The steel used for drills is 11/4 inch octagon, with a bit 2 inches, making a hole nearly 21/4 inches in diameter. Drills vary in length up to 24 feet. English steel is used at some mines, but a majority use American steel, and the most experienced men who have employed both, inform me that the drill steel made by Hussey & Wells and Parke Bros., Pittsburgh, answers as well as the best imported steel, and much better than the average. The drill is turned by one man sitting and struck by two standing, with eight-pound hammers, at the rate of about thirtysix blows per minute each. In this way from nine to eleven feet of hole are sunk per day, the men working usually on contract. The price of stope holes ranges from 60 to 80 cents per foot in depth, the mean being not far from 75 cents; no mine costs have to be paid out of this price. When there is a large proportion of block holes, which admit of the use of smaller steel, the whole drilling of a pit is often let at from 60 to 65 cents. Very deep holes, say from fifteen to twenty-two feet, are sometimes sunk with still larger bits, which about doubles the cost. In these cases two men are required to turn the drill and three to strike.

The cost of drilling ledge-holes per ton of ore, varies from a mere trifle in the case where one twenty-two foot hole throws down 4,000 tons, as has been done, to a very large item on low stopes with perhaps tight, hard ground. From 3 cents to 25 cents per ton may be regarded as extreme averages, although 35 and even 48 cents have been reached, for short periods, under very unfavorable circumstances. The price given in the table (11 cents) approximates to the average for hard ores; this number divided into 75 cents, the average cost of drilling per foot, gives, say 7, which should represent the number of tons of ore broken per foot of stopehole drilled. The data obtained directly under this head confirm this amount, which is also equivalent to about two cubic yards per foot of hole.

The depth of stope-holes varies from two to twenty-two feet, the short ones being employed in "taking up bottom," that is, in squaring the stope so as to give the best chance for the deep holes. The average of 1,500 holes of all kinds in one part of the

Washington mine was four feet nine inches, but the stopes which furnished this result were below average height. It is believed that nine or ten feet would be nearer the average for deep holes, and say three and a half feet for the short ones.

2. Block-Holes.—The masses of rock and ore loosened by the heavy blasts already described, are often so large that they have in turn to be broken with explosives, which operation is termed block-holing. The amount of this work varies from almost nothing in some pits and in certain mines, to four-fifths of all the drilling required in others, the maximum being reached on high stopes of hard, tough ore. Over two hundred block-holes have been employed to one stope-hole in the Cleveland Mine, one hole being required to every two to four tons of ore. Block-holes sometimes produce fragments so large as to require block-holing in turn, before they are made small enough to be mastered by the sledge. These holes vary in depth from eight to twenty-four inches, the mean ranging near one foot. With nitro-glycerine the holes need not be so deep as for powder. One inch octagon steel is often used in this work, making a hole nearly 11/2 inches in diameter. The drilling is performed as in the case of stope-holes, but usually only one man strikes.

In the same ground, the same drill-gang will sink more than twice the number of feet of block-hole in a day with small steel, than of stope-hole with large steel,—ranging from twenty-four to twenty-seven feet. In open mines of strictly hard ore, this work costs more than stope-holes, and is set down in the table at 13 cents per ton. This amount added to the II cents given as the cost of stope-holes per ton, equals 24 cents for the total cost of the labor of drilling required under breaking ore:—this would also equal about 70 cents per cubic yard, which would pay for one foot of two-inch drill-hole. But this is by no means the whole; the work of sinking and drifting, which is set down as aggregating 20 cents, is more than half drilling; and a part of the cost of rock-stripping is also for this work. I estimate that 40 cents per ton of ore is not far from the actual price paid for this kind of labor in the hard-ore mines, equal to fifteen per cent. of the whole cost. On this estimate, not less than \$300,000 were paid out for drilling in 1870. This work, from the favorable circumstances under which much of it is done in open excavations, no scaffolding being required, is by

far the most purely mechanical labor performed about the mines. While the absolute cost of this item of drilling is very large, and can undoubtedly be reduced by the use of the *power-drill*, it is, as compared with some other mines and regions, small. Our open cuts or quarries afford far better facilities for blasting than under-ground mines. In one Southern New York mine the drilling cost, in 1870, \$1.25 per ton of ore, or forty per cent. of the whole cost of mining; in a large magnetic mine in New Jersey, it cost from 60 to 80 cents per ton of ore. In the Persberg mines, Sweden, when the ore cost, in 1870, \$2.20 currency per ton, the drilling was 40 cents per ton, equal to twenty-three per cent. of the whole cost, being considerably more than ours, absolutely and relatively. When we consider that the average of wages in Sweden is not far from 65 cents per day, or say one-fourth of what is paid Lake Superior miners, it would seem as if Sweden would be a good field for a power-drill.

The facts relating to drilling have been given in much detail in the hope that inventors and owners of rock-drilling machines may become acquainted with the wants of the Marquette region in this regard. I have had my attention called to several of these machines, but have not had opportunity to make such investigation of their respective merits as would justify an opinion. I have no hesitation in saying that a machine which would do the work required at a less cost than it is now done (75 cents per foot) would find ready sale, and every facility would be afforded for experiments.

I need not here remark that a power-drill, adapted to Marquette iron mines, must be portable, as it would have to be shifted every few hours; and I should say that two men, or at most three, should be able to handle it on a ragged rock surface. Again, it must be capable of being set up anywhere, to accomplish which, I think that movable tripod, telescopic legs, like those with which engineers' instruments are often supplied, would be convenient.*

3. Sledging, Sorting, and Loading.—In considering this item, it must be borne in mind that the ore and rock have not only to be broken so that they can be removed, but must be made so fine as to

^{*} Since the above was written the Burleigh Drill has been tried at several mines with varied success. My facts are quite insufficient to enable me to form a judgment as to its fitness to do the required work, or to know whether it has had a fair trial.

be easily separated, and so that the pieces can be fed into a Blake crusher. This work requires more muscle and as much skill and care as any other done at the mine. Eighteen to twenty-three pound sledges are employed, and the difference in results, between the experienced miner who strikes the lump of ore the right blow in the right place, with this immense hand hammer, and the tyro. is very great. Contracts for sledging and loading, which sometimes include a little block-holing and short tramming, have been let at prices varying from 20 to 50 cents per ton. The loading usually costs not to exceed 10 or 12 cents, the balance being chiefly sledging. There is a wide difference in the texture of ore, some kinds requiring five times as much sledging as others. On the whole. Marquette ores break with much greater difficulty than those of the Eastern magnetic mines. With poorer ground worked and the market more in favor of buyers (which makes them more exacting on quality), the cost of this element will be increased.

Drops, similar to those used at foundries to break old castings, have been employed to break very hard lumps of ore, but the expense of getting the lumps of ore to them has caused this plan to be abandoned. In the copper region powerful steam hammers have been used for a similar purpose, but the same objection as that given above would apply to their introduction at the iron mines. It must be borne in mind that a lump of iron-ore is not worth more than about one-hundredth part as much as a lump of copper of the same weight, and therefore will not bear as much handling.

A steam miner who can walk up to the lump of ore and sledge it to pieces where it lies is what is wanted. Nitro-glycerine or duallin breaks the material finer, producing by its explosion more of a smashing effect than powder, and thereby requiring less sledging. There is no doubt, as is elsewhere stated, about the advantage of employing these new explosives in block-holing.

4. Handling Rock.—In addition to the rock which overlies the ore, considered under stripping, at most of the mines more or less rock is found mixed with the ore through the mines, which has to be removed during the process of mining. The proportion varies from none up to one-half of the whole, and often for short periods more than this; the average at this time is believed to be twenty per cent. The 25 cents placed against this item in the table is intended

to cover the cost of sorting out and handling this rock under average circumstances. This cost will be increased as poorer grades of stuff are worked.

5. **Miscellaneous Work.**—The 21 cents opposite this item in the table is no more than sufficient to pay for foremen, repairs of tracks and roads, wheeling, tramming, blaster, sometimes handpumping, and such securing of the workings as may be necessary, etc.

III. MINING MATERIALS AND IMPLEMENTS, EMBRACING "MINE COSTS."

This general head is subdivided in the table into Explosives, Tools, and Repairs, which are in turn itemized, as will appear below. The expense incurred here is 31½ cents per ton of ore produced, equal to about twelve per cent. of the whole cost.

I, 2. **Explosives.**—Powder and fuse and nitro-glycerine. The present (1870) is an unfortunate time to collect statistics regarding the cost of explosives, for the reason that nitro-glycerine is to a certain extent on trial, and most of the mines employ both it and powder in the same pits, making it difficult to separate the results. The place of the new explosive cannot be said to be wholly fixed in our mines. It is more powerful than powder, bulk for bulk, or weight for weight; can be used in wet as well or better than in dry ground, which is very important in some places; it has so far proved no more dangerous than powder, and its fumes have not been found objectionable. As has been stated, the fragments resulting from its use are usually smaller, hence require less sledging, and, it being more powerful than powder, less drilling is needed.

In the case of wet holes intended for sand-blasting, nitro-glycerine can often be used in small charges to produce cracks which carry off the water and thus prepare the way for the powder. Overhanging loose rock can often be advantageously brought down by a flat cartridge of glycerine.

In short holes, 3 to 6 feet, glycerine will sometimes break two or three times as much ground as powder, thus making the saving on the drilling more than balance the extra cost of the explosive.

The quantity of glycerine used per hole, of course, varies with its

depth and other circumstances, and is at the Washington and Republic Mines, according to Captain Peter Pascoe, as follows:—

Deptl	h of hole.	Glycer	ine.
3	feet	3/4	lbs.
4	"	I 1/2	"
5	"	21/4	"
6	"	3 1/2	"
8	•	5	66
10		7	66
12	"	10	66
14	"	14	6.6
16	"	18	"
18	"	21 -	_ "
20	"	24	"

There can be no doubt but that the use of this explosive hastens work. Sinking and drifting can be more speedily done with it than without.

Whether it is suited to breaking the great masses from the solid ledge remains to be seen. Certainly it cannot be used to fill the cracks produced by shaking, where heavy sand blasts are required; and it is doubtful whether drill-holes large enough to contain the requisite amount of the blasting oil can be profitably employed; two or more holes could be used, but this would greatly increase the cost of drilling. It certainly costs *more* per ton of ore mined than powder, but how far this greater cost is balanced by other advantages experience must determine. It is significant that in 1870, being the next year after its introduction, over \$40,000 worth was sold in the Marquette region at \$1.50 per pound. In 1872 about 40,000 pounds were used, the price being \$1.25 per pound. The Painsville Ohio Co. erected (1871) a factory near Negaunee. Duallin and giant powder have recently been introduced.

The figures given in the table, and in what follows, refer exclusively to powder, the nitro-glycerine element having been eliminated as far as was possible. Fuse costs about ½ cent per ton, leaving 9 cents per ton for powder, which, according to the data obtained, varied from 7 to 10 cents. The price of powder ranged from \$3.75 to \$4.50 per keg of 25 pounds. Therefore an average of 45 tons

of ore should have been broken with one keg of powder, or about 1/2 pound of powder to one ton of ore. This, it must be remembered, does not express the actual work of the powder, on account of the amount of rock moved in addition to the ore—in one instance 23,000 weighed tons of material required 320 kegs of powder, or 72 tons per keg. In another instance 31 kegs threw down 3,500 tons (approximate) of quartzite, or 113 tons per keg. One mine, which produced over 100,000 tons of ore in 1869, consumed for all purposes one keg of powder to every 43 tons of ore produced. The waste material in this case did not amount to over 20 per cent., hence about 52 tons, or, say, 18 cubic yards of material, were moved per keg of powder. The consumption of explosives per ton of ore must increase as the mines grow deeper, either by the greater amount required to remove the rock covering, or by the less favorable opportunity afforded for blasting, if the ore be won underground.

In one group of New Jersey mines, the powder and fuse in 1870 cost 18 cents per ton; in another mine in Southern New York, 14½ cents; in Sweden, at the Persberg mines, 15 cents. All of which figures considerably exceed those reached in Marquette, which is proof of the economy in explosives from working iron mines as open quarries as long as possible.

3. Steel.—The use of steel drills has already been described, and reference made to the brands in use. My data, which are far from complete, under this head, indicate that the cost of steel per ton of ore ranges from $\frac{3}{4}$ to $3\frac{3}{10}$ cents, averaging perhaps $1\frac{8}{10}$ cents; the price of steel being 20 cents per pound. This would give about 11 tons of ore, or about 3 cubic yards per pound of steel consumed, which is less than the data obtained direct on this point seemed to indicate.

It is the practice of some mines to charge the ore contractors 2 per cent. on their contracts for wear of steel, which agrees nearly with the above. At other mines the steel is weighed at the end of each month, and the contractor charged with the shortage, whatever it be.

4. **Tools**, other than Drills.—Cost about $4\frac{3}{10}$ cents per ton of ore. The Ames No. 2 D-handled, square, and round-pointed, strapbacked, solid steel shovel is the favorite.

Washoe picks, Nos. 5 and 6, and Powell, same numbers, both

railroad (25 inches long), and pole (19 inches long) are extensively used. Certain mines make their own picks after a fashion of their own.

Solid steel crow-bars, both single and double-pointed, are used. Solid cast-steel sledges, both American and chrome, weighing from 16 to 18 pounds, and often 25 lbs., are extensively used.

Solid cast-steel striking-hammers, 8 to 9 pounds, and in some instances 11 pounds, are employed.

- 5. **Blacksmiths' Supplies.**—This item is largely made up of coal and iron, steel being embraced under another head. Charcoal was formerly used exclusively for working steel; but mineral coal is now employed with good results at most mines. The table shows this item to be a trifle less than five cents per ton of ore.
- 6. Blacksmiths' Labor.—This is largely sharpening drills. The number dulled per day by a gang of three drillers will average about 75, in hard ore. One blacksmith and helper will sharpen about 275 drills per day of ten hours. The 11 cents marked opposite this item embraces all the blacksmiths' work done in and about the mine, for whatever purpose. Therefore strictly, it should have been divided, part going to dead work.

IV. HANDLING ORE FROM MINERS' HANDS TO CARS, AND PUMPING.

Pumping, which has heretofore been a small item in the Marquette region, cannot well be separated from hoisting ore, as the same machinery does both. This item, in the case of some New Jersey magnetic mines, costs 75 cents per ton of ore: at the Persberg mines, Sweden, it costs but 7 cents. The entire cost under this head, in the Marquette region, including hoisting and pumping, is 41 cents per ton of ore produced, which equals 15½ per cent. of the whole. This work is done in part by horses, part by men, and part by steam.

1, 2, 3. The Work of Horses in Handling Ore.—The team work employed at the Marquette mines, apart from the stripping, amounts, according to my inquiries, which have been quite full on this point, to 10 per cent. of the whole cost of mining, or say 27 cents per ton of ore, the drivers' wages being the largest item. This cost is obtained by dividing the total expenditure for teaming, by the

total number of tons of ore produced. If it was figured only on the ore actually handled by the horses, it would be much greater. If to this were added the cost of the team-work employed in stripping, the total would not be less than 30 cents per ton of ore, or, say \$250,000 on the product of 1870, a sum sufficient in itself to supply all the mines in the region with all the additional steam-hoisting and pumping machinery and small locomotives required to do the work now done by horses, and at a very much less yearly cost. We may verify this almost incredible estimate in another way. The total number of horses employed at all the mines in 1870, including hired teams, was about 364, or an average of 30 to each mine, varying from 9 to 74. The best data I can get indicate that to work a lot of horses for one year, including wages of drivers, stable-men, smiths' work, forage, repairs of vehicles, and depreciation, in the years 1869 and 1870, cost an average of \$650 per horse. The wages of hired teams, including drivers, for the same period, was \$6 per day. At this rate, 364 horses would have cost nearly \$240,000, a sum sufficiently near the other to confirm the general truth of the estimate.

These figures surely justify the prediction, that if there ever comes a period when our mines do not pay, it may be due largely to horses. In this age of steam, has a business any just right to prosper which employs horses to do work that can be more cheaply done by machinery? The average number of tons of ore handled per horse employed in and about the mines for all work in 1870 was 2,350, ranging from 1,150 to 5,300 tons. In considering these facts it must be borne in mind that the mines in question are not by any means without steam power. Twelve engines, varying in power from say 10 to 50 horse, were at work. To prove that this item of cost is unusually large in the Marquette region, I will give a few facts regarding the employment of live stock at mines, which have come under my notice elsewhere. While the cases cited do not present all circumstances like the Marquette mines, they are sufficiently near to afford interesting comparisons.

The Cornwall Ore Bank Co., Penn., shipped from their one immense deposit, in 1870, over 174,000 tons, employing no horses in the work. The ore was all handled by one locomotive, the cars being loaded by wheelbarrows. No pumping is required in this mine, and the facilities for reaching the ore with cars are unusually

good. The ore is quite soft, so that the blasting does not endanger the tracks.

The Iron mountain mine, Missouri, shipped in 1870 more ore than any one mine in the Marquette region. It employed during the winter 68, and during the summer a somewhat less number of horses, mules, and oxen. One animal moved about twelve tons per day, or 3,600 tons per year; but more than three-fourths of this stock was employed in getting "surface ore," a feature which does not exist in Marquette mining. The bluff (quarried) ore moved per horse employed was more than five times the above amount. No steam-engine or locomotive was in use at the mine.

At the Caledonia and Keene mines, St. Lawrence County, New York, in 1869, three horses handled 27,500 tons of ore and waste, the average haul being over 700 feet, all up grade, in places steep. This gives over 9,000 tons per head; steam was not employed for handling material at either mine.

The Sterling mine, Orange County, New York, shipped in 1869 40,000 tons of ore, which was handled under circumstances quite similar to those encountered in the Marquette region, by two horses and one small stationary engine, which gives 20,000 tons per animal employed. The system of tramways and sidings at this mine is very complete.

Passing from American to Swedish mines, which are far deeper, and in which there is a larger percentage of rock mixed with ore, we find that in the Persberg mines, in 1870 (see table), the total cost for handling ore and water drawing was 142 per cent. of the whole cost, or 33 cents per ton of ore; and this amount included the handling of all the rock and other waste material which in our table is embraced under *Dead-work*. If we take out of dead-work 10 cents for handling this waste and add it to the amount found above, we have 51 cents as total cost of handling Lake Superior ores, equal to twenty per cent. of the whole cost, or about fifty per cent. greater than in the Swedish mines, but there water was exclusively used.

It is not difficult to understand how horses* have come to play so important a part at our mines.

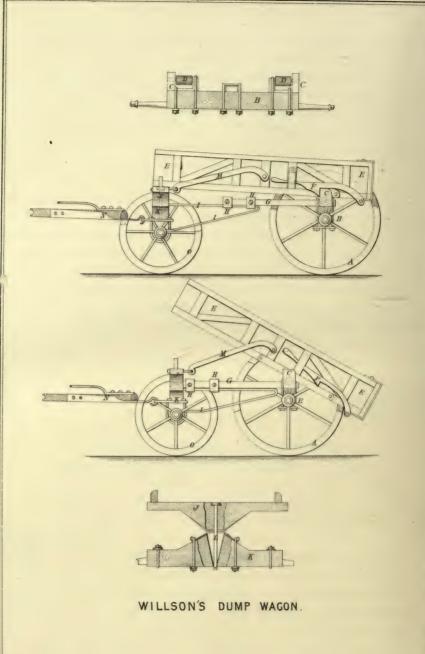
^{*} It should be noted that oxen have been in use for some time at the Lake Superior mine, but, so far as I am informed, at no other.

The first operation in opening a new mine is, usually, to strip off the earth and rock covering, which can be best accomplished with the horse and cart. On the ore face thus exposed, mining is begun, the ore being hauled to the cars (often not brought very near to the pit), and such rock as is mixed with the ore is sorted out and hauled in another direction. It is very convenient and economical to back a cart directly to the miners' hands, and this was done until it came to be regarded as the way to get out ore. There was certainly no better way at the start in many cases: but when horses come to be used on hauls of over 500 feet and up grades, in places as steep as I in IO, the operation costing 25 to 30 cents per ton, it may be worth while to ask if such ore had not better be left in the ground until machinery propelled by steam can be brought to bear on it. Another cause which conspired to prolong this expensive mode, was the great demand for ore during the war and the consequent high prices. Mine superintendents were given no time to plan nor make improvements looking to future economy. Mine owners did not then want surveys, nor machinery, nor tunnels, nor anything that had reference to the future; they only wanted ore, nor did they care much what it cost, nor what the quality was (so consumers say): it was ore, ore, ore! Wherever three men could be set at work, a cart was backed up to them and shipments began from a new pit.

On short hauls, smooth roads, and light grades, horses can be used to advantage, and will continue to be so used, especially where there is more or less uncertainty as to the quantity of ore in the pit worked, which is often the case. But where there is a large mass of ore, rock, or earth to be moved under any other circumstances, it will usually pay to bring steam-power to bear upon it. Portable, or easily-to-be-moved railroads, and small locomotives for long hauls are in much favor at this time, and would have the advantage of utilizing existing wagon-roads. But the first step in many cases is undoubtedly to lay horse railways on the present roads. As is shown above in the remarks on the use of horses in certain New York iron mines, one animal can move from ten to twenty thousand tons on such roads in one year. If the horses at our Marquette mines can be made to perform one third this amount of work, the present cost of hauling will be reduced fifty per cent.

Portable hoisting-engines are extensively used in New Jersey and





Pennsylvania; they can be set up quickly just where wanted, and handle material rapidly and with great economy. A thorough system of under-ground communications which would bring all or most of the material to the main hoisting-shaft is always to be aimed at, as in this way the dead lift may be made by steam. At present, owing to the continued pressure for ore, it is not uncommon to see ore and rock carted up-hill, over abominable roads, from pits which in a few months, perhaps, will or could be reached by drifts along which the ore could be cheaply trammed to a steam hoisting-shaft.

As may be supposed, this extensive use of draught animals has led to great perfection in the carts, wagons, and sleds. A dump-sled for winter use, contrived by Captain Merry, of the Jackson mine, is a perfect vehicle of its kind. I am unable to give drawings of but one, known as Daniel Willson's Patent Dump Wagon, of which over 50 are in use in the region. See Plate XVII.

While harnessed to the cart or wagon is the favorite mode of using the horse, it is by no means the only way. Some pits in the course of mining became too deep for cart roads; these were in many instances worked by *swing derricks*, horses being the power employed; the long booms of these derricks made it possible to drop the bucket in different parts of a wide pit. This method is, however, very expensive, as the following figures will show. The total lift from bottom of pit to bottom of cart was in one case 79 feet; the cost being as follows:—

2	men filli	ng		 \$4	00
I	man to 1	and		 2	00
2	derrick h	norses and	driver	 5	25
				\$11	25

This sum paid for hoisting 45 tons in 10 hours, is equal to 25 cents per ton. In one case, where the hoist was 55 feet, the cost was 16 cents per ton.

In another case, with the ordinary two-bucket horse-whim, the cost of hoisting 65 feet, and landing, was 6 cents per ton; this did not include filling the buckets. In another case the ore was hoisted 40 feet, and landed for 5 cents per ton, not including the filling. Estimating the filling at 10 cents, these facts show that it costs in

the cases cited an average of I cent to lift one ton of ore 7 feet, including the landing or dumping, which employs one man.

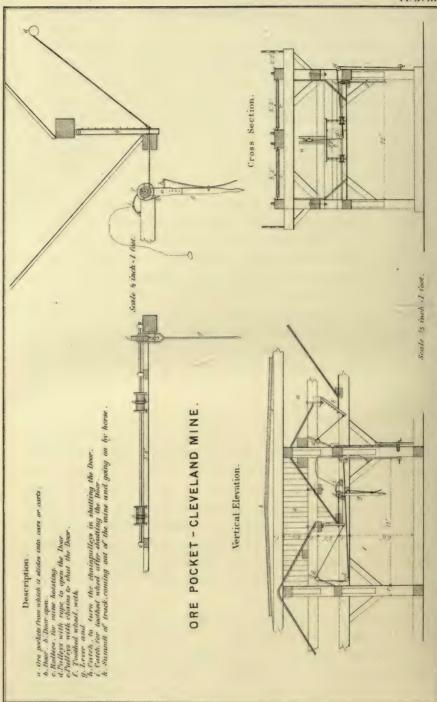
Without attempting to fully solve the important problem of the best mode of handling the material at Marquette mines, for that is beyond the scope of this report, I would suggest the following general policy as being safe for the mines to pursue:—

Let all large pits now worked, where a considerable amount of horse labor is required, be suspended until some form of steam machinery can be brought to bear on them. There are, of course, exceptions to this rule: for instance, where the other costs are unusually light, more money may be expended in handling the ore, as is often the case with the soft hematites; but the principle is, I think, correct. It would not be difficult to find many instances of this kind; for example, a given pit is worked, the ore being moved by horses, at a profit say of 50 cents per ton, which if left for one year could be reached by some tunnel or other improvement which would permit the same ore to be taken out at a profit of \$1.00 per ton; it would certainly pay to wait in such instances. In these cases it will usually be found that the superintendent has been persuaded into promising that his mine can be made to produce a certain amount of ore which may have been already sold, his attention being thereby fixed on a large product, rather than cheap mining. This subject will be considered more fully below. I will here only ask, if it is not better policy for a mine to net say \$50,000 on 50,000 tons of ore, than to make the same sum on 100,000 tons. If the mines were inexhaustible it might not make much difference, but as it is, it may make all the difference there is between a profitable business and an unprofitable one in the end. It must be borne in mind, that while the ore business has been on the whole profitable, there are large mines that have been producing ore for years that have never returned a dollar to their stockholders.

Among the mining appliances which have been brought to great perfection in the Marquette region, are the various forms of pockets and shoots for transferring the ore, first, from the mine cars, buckets, and carts to the railroad cars, and second, from these to the vessel.

The magnificent ore docks at Marquette, Escanaba, and L'Anse belong to the latter class, and are undoubtedly the best of the kind





in the United States if not in the world. They are described and illustrated in Chapter I., and in Appendix F. of Vol. II.

Of the first class there are numerous varieties, from the simple log crib built up alongside and above the track, into which the ore is dumped from elevated railways, and from the sloping bottom of which it is "shot" through holes closed by rods into cars at a cost of not over $3\frac{1}{2}$ cents per ton, to the more expensive and perfect contrivance employed at the Cleveland mine, which is shown in Plate XVIII.

The mine car in this case passes over the centre of the pocket, which dumps its ore in turn into a car or cart below, by an ingeniously arranged door which is shown on an enlarged scale.

- 4. Loading Ore from Stock Pile.—During the winter no shipments are made from the mines, hence the product has to be piled up. It is the policy of some mines, and I think it is the best, to do most of their dead work in the winter, hence to stock but little ore; others maintain nearly the same rate of production in proportion to the force employed, winter and summer. Stocked ore has to be loaded in cars by hand, which is always contract work and costs from 9 to 12 cents per ton, the mean being, say 11 cents, including all costs connected with it. This amount, distributed over the whole product for the year, was found to average for the cases inquired into, $3\frac{5}{10}$ cents per ton.
- 5. Machinery for Pumping and Hoisting.—Notwithstanding the great cost of the work of horses, a large amount of machinery, as has already been remarked, is now in use, as the following statements will prove:—

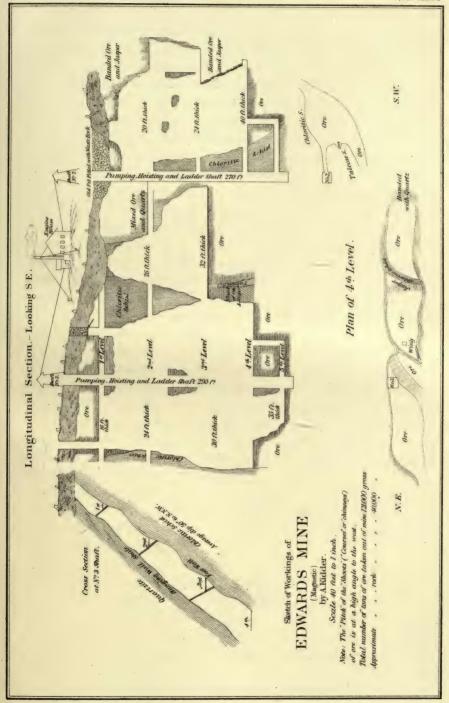
The introduction of machinery has so far seemed to make but little relative diminution in the number of horses employed, because of the greater amount of waste material which has to be moved in the later years. The amount given in the table, opposite this item, II $\frac{2}{10}$ cents, is designed to be an approximation to the cost of running the machinery of such mines as have plants distributed over the entire product of those mines. I estimate that less than one-half of the product of such mines was handled by machinery in 1870. The actual cost of moving the ore so handled, including the pumping, varied from 14 to 21 cents, the mean, as shown by my data, being about 18 cents. This cost is made up of wages of engineers and firemen, say fifteen per cent.; fillers, landers, and surface tram-

ming, sixty per cent.; fuel, repairs of machinery and supplies, say twenty-five per cent. This covers the cost from miners' hands to cars or stock pile.

While this sum is materially less than the cost of the same work by horses, it is much greater than in the Copper region of Lake Superior, where this work is brought to great perfection. Some of the appliances employed in the Copper region cannot be used at iron mines on account of the greater irregularity of the deposits. But time will introduce many economies which will reduce this item below the figures given. It must be borne in mind, in comparing the cost of steam machinery with horses, that in the case of the engines all the pumping is included, while the horses handle only the ore and rock. Making this correction, it is safe to say that it costs at least four times as much to handle the same material by horses as by machinery.

The following description of recently erected plants will give a good idea of the machinery now in use at the iron mines, it being essentially such as is employed at the copper mines.

The Macomber mine machinery consists of one steam-engine with cylinder 18 × 24 inches, with bed cast solid in one piece. Valve is of the kind known as the H valve, and is worked by link motion; steam pipe 4 inches in diameter; exhaust pipe 6 inches in diameter; engine supplied with the Judson governor. Pump for feeding boiler is worked from cross-head; also an auxiliary for fire protection, etc. Main shaft is 5 inches in diameter, of hammered iron, and 16 feet long. One boiler 48 inch shell, 26 feet long, with two 18-inch flues. Smoke-stack is 40 feet high and 24 inches in diameter. The winding drums are 4 feet in diameter, and of sufficient capacity to contain 525 feet of 11/4 inch wire rope. They are worked by a friction movement, thrown in and out of gear by means of eccentrics with lever attachments. The brakes are known as band-brakes, which clamp the entire surface of the drum, 5 inches in width, and are of sufficient power to hold a loaded skip at any point in case of accident. They are worked by levers with hand or foot, as may be desired. The drums make about 131/2 revolutions per minute, the engine making 80, which gives the skip a speed of a trifle less than 3 feet per second. The skips are of heavy boiler iron, each having four 12-inch wheels. The capacity of each is 35 cubic feet, equal to about 21/2 tons of ore. The pump is 10 inches in diameter by 6 feet stroke, capable of discharging 660 gallons of





water per minute. It is worked from a slotted crank arm, on end of main drum shaft, which admits of lengthening or shortening the stroke at pleasure. The pump is double acting, with single valve on a new plan. It is furnished with rods, travellers, connections, balance bobs, etc. This machinery was furnished complete in all its parts, and set up at the mine in working order for pumping and hoisting by the Iron Bay Foundry, Marquette, Mich., 1872.

The Barnum mine plant consists of one horizontal high pressure steam-engine of 20 inches diameter of cylinder and 30 inches stroke; steam furnished by two tubular boilers, each 48 inches in diameter and 14 feet long, and each containing 50 tubes, three inches in diameter. Maximum power of this engine is 120 horse, but is working at present at one-third its capacity. There are two winding drums, each 5 feet in diameter; speed of engine about 60 revolutions per minute, and of drums about 12. Drums are attached to main shaft by cone-gears, which are operated by steam cylinders and levers; screw-levers control the brakes and drums during the descent of the skip.

Engine is connected to the drum-shaft by spur-gearing in the proportion of one to five; speed of skip in shaft, about 3 feet per second; load of ore, 5,000 pounds; weight of skip, which is self-dumping, is 2,400 pounds, making the total load 7,400 pounds. Actual power employed, about 47 horse; engine also draws water with a 6-inch Cornish pump. Total weight of this machinery about 42 tons, and total cost about \$10,000. Built at the Michigan Iron Foundry, Detroit, in 1869.

The foregoing described plants, together with those given in the subjoined tabular statement (pages 280 and 281), embrace over three-fourths of all the machinery employed in hoisting and pumping in the entire region.

V. MANAGEMENT AND GENERAL EXPENSES.

This covers only such expenses as are incurred in the mining region, and not salaries of officers above the superintendent, nor the cost of selling the ore.

1, 2.—Salaries, Office Expenses, and Taxes.—This element of cost constitutes less than 5 per cent. of the whole cost of the ore,

DESCRIPTION OF STATIONARY ENGINES, WITH THEIR

NAME OF MINE.	Number or name of Pit or Shaft.	Size of Cylinder, Length and Dia- meter.	No. of Cylinders.	Number, size, and kind of Boiler.	Average working-pressure.	Nominal horse-power.	Kind of work, as pumping, hoisting, etc.	Height to which ore is lifted in feet.	Average number of tons hoisted in 24 hours.
	Pit No. 4.		2	Steam supplied for this double and single engine from two of Root's patent boilers, 50 horse- power each, con- nected together.	70 lbs.	140	Hoisting.	125 feet.	120
Jackson	Pit No. 6.	13"x 30", one 40 horse Root's pat'nt trunk engine	1	do. do. do.	70 lbs.	40	Hoisting and pumping.	From 80 ft. to 125 ft.	200
	Pit No. 7.	8"x 12"	2	One boiler, 42" dia- meter x 12 feet long, tubular, 40-3 in. flues.	70 lbs.	20	Hoisting and pumping.	50 feet.	50
	Pit No. 5.	5'' x 8''	2	Tubular boiler, 40 in. diam. x 13 ft. long, 40-3 in. tubes.	70 lbs.	8	Hoisting.	50 feet.	40 tons of ore, rock and water.
	Machine shop.	8 x 16 inches.	1	Tubular boiler, 60 in. diam. x 25 ft. long, 121-2 in. tubes.			Running ma- chinery in shop.		
Champion {	4 shafts now worked by main en- gine. 2 new shafts now being sunk.	One horizon- tal engine, 14 inches bore, 20 in. stroke of piston.		Two return flue boilers, 42 inches diam., 28 ft. long, 2 flues in each, 16 inches diameter. One locomotive boiler, 28 inches diam. of shell, 26 x 30 in. fire-box, 36-2 in. flues, at 1st level of No. 3 shaft.	to the sq. in.	hoist-	One 6-inch plunge pump. One No. 7 Earle pump, at 3d level of No. 3 shaft; elevat- ing water to surface; sup- plied with steam from boi- ler at 1st level.	180 feet.	400 from 4 shafts.
Edwards.	Nos. 2 & 3.	24 x 36 inches.	I	Two. 5 ft. diam., 27 ft. long, with return flues each.	70 lbs.	150	Pumping and hoisting.	300 feet.	200
		-							
Lake Ange-	2 Pits.	16 x 24 inches.	I	One. 42 in. shell, 20 ft. long, with 2-14 in. flues.		60	Pumping and hoisting.	75 feet.	150
Washing- ton.	At No. 7 op- ening, known as No. 1 & 2, skip roads.	inches.	1	One boiler, 2 flues, 24 ft. 6 in. length, 44 in. diameter.	90 lbs.	50	Hoisting,	No. 1 skip, 130 feet. No. 2 skip, 55 feet.	} 44 } 35
	Main shaft.	20 x 30 inches.	1	Two boilers, 3 flues, 43 x 6 feet.	60 lbs.		Hoisting and pumping.	160 feet.	350
Toles Su	Hematite.	12 X 20 inches.	I	One boiler, 2 flues, 3½ x 24 feet.	65 lbs.		Hoisting and pumping.	130 feet.	100
Lake Su- perior	Portable en- gine & boiler "Sect. 16."	10 x 18 inches.	1	One boiler, flue.	30 lbs.		Hoisting and pumping.	60 feet.	
	Sect. 21.	10 X 12 inches.	1	One boiler, upright flue.	35 lbs.		Hoisting.	60 feet.	

WORK, AT SIX MARQUETTE MINES, JANUARY, 1873.

Skip an	of Barrel inches.	·sdun,	evolutions of Engines per minute.	day that worked.	Shaft.	quantity used in	D
Kind of Skip and its load.	Diameter of Barrel pump in inches.	Kind of Pumps	Revolutions of gines per min	Hours per day that Pump is worked.	Vertical. Inclined.	Kind and quantity of fitel used in 24 hours.	Remarks.
Two 5 ft. drums, with 4 wheel, self - dumping. Skip-car 2½ tons.			75		2	One cord of wood	There are also two (2) 12-horse power locomotives, which are used for distributing cars in the tunnels during the shipping season.
Two 3 ft. drums, one hoisting skip-car 2½ tons, 4 hoisting patent dump buckets 1 ton each.	8	Cornish jack-head.	100	10	ıı	per day; don't run at night.	
Inside dump-car 3 tons.	8	Cornish jack-head.	100	12	I	One cord of wood per day; don't run at night.	in various parts of the
Bucket 1 ton.			100	10	I	¾ cord per day.	Earle steam pump; I No. 8 Knowles steam pump; I No. G Came- ron steam pump; I Worthington duplex
			80			One cord per day.	pump; also one 8 (eight) inch double acting bucket pump.
Wrought-iron skips, 42 inches long, 30 in width and depth, Hold 3,000 lbs. of ore.		Plunge-pump, 6 in. diameter of cylinder and 6 in. column, elevating the water to the surface 180 feet.	20	22	incl'd		Makers—Hodge 8 Christie, Detroit, Mich This one engine doe all the work of this mine.
Cornish skip 1½ tons.	one 7	Two 6-in. draw-lifts from 5th to 4th levels, at Nos. 2 and 3 shafts. One 8-in. draw-lift, at No. 2 shaft, from 5th level to 2d. One 7-inch plungerpole, from 4th level at No. 3 shaft, taking also No. 2 water to surface.	30	20	incl'd	Wood, six cords.	Makers—Hodge & Christie, Detroit. See plan of mine-Plate xix.
Cornish skip 11/2 tons.		One 10-inch double- acting pump.	30	10	incl'd	Wood, 2½ cords Coal, ¾ ton.	Maker—D. H. Merritt Marquette.
Iron self-dumper about 1 ton.		Earle. Nos. 4, 6, and 7.	see cata- logue.	3	vert'l	wood; about 1,50	ft Furnishing steam for Burleigh Drill Com; pressor, 3 Earle pump. (2 No. 4 & 1 No. 7), be sides to hoisting engits.
Iron skip 3 tons.	10	Plunger.	about 30	10	incl'd	Six cords wood.	Makers—Wash'n Iro Works, N'burgh, N. Y
Iron car 3 tons.	8	Bucket plunger.	about 60	14	incl'd	Three cords wood	
Iron skip 2 tons	, 6	Bucket plunger.	about 40	10	incl'd	Three cords wood	1.
Iron skip 2 tons			about 80		incl'd	Three cords wood	I.

amounting to about 12 cents per ton. I am happy to note here a much better showing than in the Persberg mines, Sweden, where this item, in 1870, cost $16\frac{1}{2}$ per cent. of the whole, or 36 cents per ton of ore; nearly three times its cost with us. I presume the excess of this item in Sweden may be largely due to heavier taxes, and smaller production.

CHAPTER X.

CHEMICAL COMPOSITION OF ORES.—ANALYSES.

THIS chapter contains the results of over one hundred and fifty analyses, more or less complete, of iron ores from the Upper Peninsula of Michigan, mostly from the Marquette region, together with five analyses of pig-iron produced from these ores; and several analyses of ores from other parts of the U. S., which are largely used with Lake Superior ores as mixtures. In order to bring out the variations in quality of the ores, and to obtain reliable practical averages, seldom less than two and in one instance eight samples were analyzed from the same mine.

By far the largest portion of the samples, the analyses of which appear in this Report, were selected by myself with a view to obtaining a fair and safe average of the ore sampled, one that would be borne out and confirmed by practically working the same ore in the furnace. I am well aware, from extended observation and practical experience, that a large majority of the published analyses of iron ores, not only have no practical value, but are positively detrimental to the best interests of the iron trade, representing as they so often do the ores to be richer in iron than they actually are, simply because the samples analyzed were not honestly or skilfully collected. Even the most skilful and conscientious men, if they err at all in collecting a sample from a new iron location, are almost sure to err on the side of finding too much, rather than too little iron. The chemist is often wrongly blamed for these false results. My experience with many analysts leads me to believe that they are, as a rule, thoroughly honest and painstaking men, who return correct results for the samples sent them; the trouble is with the samplers. This point receives further consideration under Explorations, Chapter VII.

In earnestly endeavoring to avoid this rock on which so many mining engineers and geologists have wrecked their reputations, I

may in some instances have gone to the opposite extreme and collected samples which were below the average richness—at least I am quite persuaded that I shall be charged with this—hence venture this explanation in advance of the charge. If such mistakes are found, I can only say myself and not the analysts are to blame, and I stand ready to make such corrections as lie in my power.

My method of sampling is as follows:—1st. To obtain an average of a producing mine; I found that the immense stock piles accumulated at Cleveland, Ohio, at the end of the shipping season, afforded excellent opportunities for sampling. The stock piles at the mines or a large number of loaded cars were often resorted to, and in many instances it was thought best to go into the mine and take the samples from the solid ledge or the loose ore as it was being taken out. In either case an ordinary shot bag, holding 4 or 5 pounds of ore, was filled with small fragments, varying from the size of a pea to that of a walnut, of all kinds of ore, from all parts of the pile, together with the rock, if any, which was found mixed with the ore. Some of these fragments were picked up and some were broken from larger pieces; the dust and mud over the ore made it often impossible to distinguish whether the pieces taken were ore or rock. These samples were all pulverized and thoroughly mixed, and from this the specimens were taken for the chemist, the same being forwarded by mail in small numbered tin tubes; and in each instance a pound or more of the pulverized ore was retained for future reference. The reserved portions are now in my safe in Marquette, from which samples will be furnished to any who may desire. 2d. To obtain an average sample from a new locality or from exploration pits is more difficult and unsatisfactory. This subject is fully treated under Explorations, Chapter VII.

With all this care my results varied, in extreme cases, from 10 per cent. below to 5 per cent. above the true average, but the common variation was not more than three per cent. Two or three of the extreme results, known to be wrong, are omitted from the tables. The name of the sampler is in every case given when known, and the circumstances of its collection are briefly stated in the notes. The samples collected by E. R. Taylor, of Cleveland, were, at my request, taken in accordance with the rules above given.

The surname of the chemists and date at which analysis was

made, as near as could be ascertained, are given under the result in every instance except one. The number of analyses made, with names in full and address of these gentlemen, are as follows:—

No.	Made.
Professor Oscar D. Allen, New Haven, Conn	17
Professor Geo. J. Brush, New Haven, Conn	I
J. Blodgett Britton, Philadelphia, Pa	56
A. A. Blair, St. Louis, Mo	2
Dr. C. F. Chandler, School of Mines, N. Y	8
Dr. C. F. Chandler and F. A. Cairns, School	
of Mines, N. Y	12
Chandler and Schweitzer	I
F. H. Emmerton, Chicago, Ill	I
F. B. Jenney, Marquette, Mich	8
Prof. Geo. W. Maynard, New York	5
Maynard and Wendel	3
Ed. R. Taylor, Cleveland, Ohio	14
Dr. A. Wendel, Troy, N. Y	20
Dr. Otto Wuth, Pittsburgh, Pa	30
Samuel Peters	1
T. G. Wormley	4

The metallic iron was usually determined by but one chemist, as the chances of difference on this element are small. Phosphorus determinations are more difficult, and considerable differences in the amount of this element found in the same sample by different chemists, will be observed. For this reason duplicates were often sent to two and sometimes to three; the results being given as returned by them. If any one supposes the differences to be due to errors in samples, which is improbable, I will gladly furnish duplicates for re-examination. The specific gravities of powder were mostly determined by Mr. Jenney, and not by the chemists over whose names they are sometimes placed.

The subjoined table contains an approximate general summary of the results, exhibiting the average composition of the four classes of ore now produced by the following mines:—

I. Red Specular Ores. Barnum, Cleveland, Jackson, Lake Superior, New York, Republic, and Kloman.

II. Black Magnetic and Slate Ores. Champion, Edwards, Michigan, Spurr, and Washington.

III. Soft Hematites. Foster, Lake Superior, Lake Angeline, Taylor, Macomber, New England, Shenango, S. C. Smith, and Winthrop.

IV. Flag Ore. Cascade.

Table No. XIII. of Atlas contains a somewhat similar summary so far as metallic iron and phosphorus are concerned. More facts are incorporated in this table, which has slightly changed the averages.

	I.	II.	III.	IV.
Protoxide of Iron Sesqui- or Peroxide of Iron. Oxide of Manganese. Alumina Lime Magnesia. Sulphur Phosphoric Acid Silicic Acid, Silica, or Insoluble Silicious Matter Water, Combined "Uncombined "Total Volatile Matter	90.52 Trace. 1.39 0.70 0.42 0.05 0.258 5.892	19.639 67.761 0.13 2.13 0.68 0.69 0.132 0.199 7.828	75.75 0.80 1.536 0.36 0.294 0.110 0.185 14.035 3.94 1.18	70.98 Trace. 2.01 0.45 0.20 0.03 0.13 25.12
	100,000	100,000	100.000	100.00
Metallic Iron Phosphorus. Sulphur Metallic Manganese Specific Gravity.	62.915 0.111 0.05 Trace. 4.74	62.930 0.085 0.132 0.091 4.59	52.649 0.078 0.110 0.56 3.88	49.332 0.053 0.03 Trace. 4.09

A glance at this table shows us that, except the soft hematite III., which contains about 5 per cent. of water, all the ores are essentially and chiefly composed of oxide of iron and silica or insoluble silicious matter. The other elements, viz., oxide of manganese, alumina, lime, magnesia, sulphur, phosphoric acid, and water amount in the aggregate to only about 5 per cent. in the I., II., and IV. classes. So constant is this ratio that a valuable determination of iron in a hard ore, and one sufficiently accurate for practical purposes, can be made by ascertaining the percentage of insoluble silicious matter, adding 5 to it and subtracting the sum from 100. The result is the iron oxide, which, multiplied by .70 for red, and .72 for black oxides, gives the metallic iron.

Regarding the percentage of metallic iron, consumers of Lake

Superior ores will at once note that their furnace books very often show a higher yield than 62.9 per cent., which is given in the table as the average percentage for first-class ores. This may not have been the case in exceptional years, like 1872, when the consumption so crowded the production that mines had not the time nor skilled labor to make such selection as they usually make. But that furnaces running on first-class ores usually make a better yield than that given, is shown by "Table of Metallurgical Qualities of certain Lake Superior Ores by Consumers," Plate No. XIII. of Atlas, where various consumers credited these ores, in 1870, with an average of over sixty-four per cent. of iron, as shown by their furnace-books. This discrepancy is easily accounted for; the chemist's result is in pure metallic iron, the furnace man's is in pig iron, which contains several per cent. of carbon and silicon, and other substances,—see subjoined analyses. Therefore the chemist should always find less iron than is shown by the furnace accounts if he has an average sample of the ore. Just what this difference is depends on the grade of iron made, on the waste in the slag, and other things: good authorities have placed it at 21 per cent.

Passing to a more detailed examination of the facts recorded in the table, we find, in descending order,—oxide of manganese has a maximum of nearly one per cent. in the hematite, and is nothing in the specular and flag ores. If the hematite was subdivided into manganiferous and non-manganiferous varieties, as suggested under Lithology, Chapter III., then one variety would contain only a minute quantity of manganese, while the other would reach an average of, say 3 per cent. of the oxide. The presence of manganese adds to the value of an ore, especially for making steel. Alumina reaches a maximum of over 2 per cent. in the magnetite ores, and is least in the specular ores. The earthy character of the hematites would lead one to expect more of this element in that class. Lime and magnesia aggregate a trifle over one per cent, in the high grade ore, and less than this amount in those of low grade. Sulphur is relatively most abundant in the magnetites; but, so far as I know, the minute quantity found has never been objected to by consumers of the ore. The quantity of phosphoric acid and phosphorus is of such moment in connection with the wants of the Bessemer steel manufacture, now rapidly developing in the West, that this subject will receive especial attention hereafter.

The distribution and relations of the *silicious matter* have been mentioned;—it has its maximum in the flag ores where it reaches one-fourth of the whole weight, and is least in the rich speculars, which contain only about 6 per cent. on the average.

The total water in the hard ores is only about I per cent. In the soft hematites it rises to an average of over 5 per cent., and, as will be seen in the subjoined analyses, increases in a few instances to about twice this amount, the greater part of which is combined with the limonite, which largely makes up the soft ore. An appreciable amount of volatile matter, supposed to be mostly carbonaceous, occurs only in the hematite ores. The specific gravities given will be observed to have a very significant relation to the amount of iron, which subject is considered fully in Chapter III.

Phosphorus in Lake Superior Ores.

Pig-iron intended for the use of *steel* makers must be remarkably free from phosphorus, *one-tenth of one per cent.*, according to some authorities, being the maximum amount allowable for many purposes. As it has been found impossible, up to this time, to eliminate this element from the metal either in the blast furnace or in any of the various processes for making steel, it is indispensable, in steel manufacture, that we start with an ore comparatively free from it; and for the best bar iron, only a very small amount of phosphorus is admissible,—its effect being to produce cold shortness.

It is a safe practical maxim of iron metallurgy that all the phosphorus contained in the coal, limestone, and ore charged into a blast furnace will be found in the resulting pig-iron, and that the conversion of such pig-iron into steel will increase the phosphorus just in the ratio in which the metal is wasted in the process. It is therefore very evident, if say one-tenth of one per cent. only is admissible in steel, not only our ores but fuel and flux must be very free from phosphorus at the start. In considering the facts regarding this element here given, it must be constantly borne in mind that a rich ore may contain more phosphorus than a lean ore, and yet produce a pig-iron containing less phosphorus than the other, because less of the rich ore is required to make a ton of iron.

To illustrate: an ore yielding $66\frac{2}{3}$ per cent. in the furnace, and containing .06 of phosphorus, will produce a pig containing .09 of phosphorus; while an ore containing but 50 per cent. of iron and .05 of phosphorus will produce a pig containing .10 of phosphorus; therefore the amount of iron in the ore must be always considered in comparing the amounts of phosphorus. Applying this rule to the facts given in the foregoing table, we shall find that the apparent greater freedom of the hematite and flag ores from phosphorus is nearly balanced by their comparative poverty in iron.

The distribution of phosphorus among the Lake Superior ores, so far as my facts go, follows no obvious law; it seems to have little, if any, relation to the kind of ore. Some of the hematite ores are among the lowest and others among the richest in this element, and so of the specular and magnetic ores.

A rule, to which there are, however, several exceptions, seems to be that the ores poor in iron and rich in silica, contain least phosphorus; but the analyses of the Republic mountain ore show more iron and less silica than in any other, and that it is also very low in phosphorus. The table of analyses, in Plate No. XIII. of Atlas, presents most of the facts in a compact form; but as this subject is of peculiar interest at this time in connection with the Bessemer steel manufacture, I venture to incorporate a second tabular statement here, in which the mines are arranged in order of the quantity of phosphorus, beginning with the lowest. No mine is included from which less than two samples have been analyzed. The deposits and mines marked with a * are new, and not sufficiently developed to enable me to say that an average sample of the ore was obtained.

Mine.	Kind of Ore.	Phosphorus.	Iron.
Lake Angeline	Jaspery Specular	0.031	53.83
Winthrop	Soft Hematite	0.037	54.63
Republic *	Specular and Magnetic	0.040	66.51
Michigamme*	Magnetic	0.041	64.388
Silas C. Smith	Hematite	0.047	49.70
Cascade	Flag	0.053	49.332
Menominee Iron reg'n*	Specular & Hematite.	0.054	48.209
Edwards	2d Class Magnetic	0.055	49.190
Macomber	Hematite	0.058	54.92
Cascade	Flag and Specular	0.061	51.253

Mine.	Kind of Ore.	Phosphorus.	Iron.	
Jackson	Specular	0.066 0.067 0.067 0.070 0.072 0.074 0.079 0.080 0.089 0.094 0.104 0.104 0.107 0.124 0.126	1ron. 63.715 54.72 61.60 56.315 63.55 44.29 50.70 48.24 63.55 52.27 63.81 62.11 52.88 57.155 61.092 54.19 52.40 61.69	
Washington	Magnetic	0.14I 0.224	61.30 5 61.74	

It has been stated that an inspection of the first table did not warrant us in asserting that either of the four classes of ore represented could be easily recognized as being comparatively free from phosphorus; so an examination of the above presentation of the facts forces us to the conclusion that the distribution is not geographical; for we here see widely-separated mines containing the same amount of phosphorus, whilst contiguous mines vary widely. In fact, in different parts of the same mine there is found a wide difference in the quantity of this noxious element; e.g.: The New York mine results show more than twice as much phosphorus in the ore from pit No. 1 as from pit No. 2; and the Lake Superior ore appears to contain less phosphorus than the Barnum, although they belong to one deposit. A part of this difference is undoubtedly due to errors in sampling and errors in the analysis; but the number of samples analyzed, the care taken in collecting them, and the reputation of the chemists, leave but little doubt that the relative and absolute average amounts of phosphorus in the ores from the developed mines are nearly expressed in the foregoing table.

At the suggestion of Mr. A. L. Holley, I selected, with much care, an average sample of the rock which occurs in the hard ores, more or less of which goes into the furnace, and had it analyzed; the result was less than the average amount of phosphorus. This fact, in connection with the low amount found in the second class and flag ores, leads me to believe that no care in selecting and sorting ore will diminish the quantity of phosphorus.

By way of verifying the amount of phosphorus in Lake Superior ores, here given, there are presented in the following table five analyses of pig-iron made from them with charcoal, and a flux containing no appreciable amount of phosphorus. They may, therefore, be said to indicate very accurately the amount of phosphorus in the ores, which, as will be seen, averages about the maximum amount given above as admissible in steel.

-	I	2	3	4	5	Average.
Magnesia. Silicic Acid, or Silica. Silicon Graphitic Carbon Combined " Metallic Iron. Phosphorus Sulphur Metallic Manganese	.138	1.16 3.72 0.30 0.104 0.045	0.47 1.83 3.35 0.00 93.49 0.082 trace.	3.21	2.91 3.61 .05 .092	2.28 3.39 .38 93.34 0.108 0.032 .174

No. I was chipped from many pigs of No. I gray foundry iron made at the Pioneer furnace Negaunee, of Jackson ore. Analysis by Dr. C. F. Chandler. No. 2 is a pig-iron made from assorted Lake Superior ores at the Appleton Furnace, Wisconsin. Analysis by Mr. Morrell. No. 3 is also a specimen of Appleton iron. Analysis by Dr. Wuth. No. 4 is No. I gray foundry iron made by the Jackson Iron Co. at Fayette, Michigan, of Jackson ore with charcoal, and is extensively used in the manufacture of Bessemer steel. Analysis by Mr. Morrell. No. 5 is a specimen of pig made by the Michigan Iron Co. in Marquette County, of a mixture of specular, magnetic, and hematite ores. Analysis by Mr. Morrell. The analysis of Pioneer pig was at the expense of the Survey; the others were furnished by Mr. Holley. It was proposed to

carry this work much further, but the limited means would not permit.

For contributions in money, and valuable suggestions and encouragement in obtaining the results set forth in this chapter, I am under especial obligation to John Fritz, of Bethlehem, Pa., and S. P. Ely, of Marquette; A. Pardee, Daniel J. Morrell, A. B. Meeker, and W. H. Barnum also contributed liberally towards paying for the chemical work, which cost nearly \$2,000.

The physical and mineralogical character of the following ores is given under Lithology, in Chapter III. For commercial statistics, and, incompletely, the metallurgical qualities, see Plates XII. and XIII. of Atlas.

(The mines are arranged alphabetically. The upper line gives the number of the sample.)

BARNUM MINE-Specular Ore.

	58.	14.*	14.*	262.*	262.*	277.*	277.*	Ave- rages.
Sesqui- or Peroxide of Iron Oxide of Manganese. Alumina Lime Magnesia Sulphur Phosphoric Acid	93.40 0.33 0.30 0.15 min'te trace.	1.92 0.64 0.53			0.288		0.640	
Silicic Acid, or Silica Water, Total Metallic Iron Phosphorus.	99.50 65.38	9.43 0.46 99.98 60.79		64.30	0.123	56.31		61.69 .134
Sulphur, Specific Gravity.	Sept.	ndler & 4, 1872. + 0 850.				Dec.	Feb.	.134
	Chemist, Allen. 2, 1871. Sampler, Brooks.	Chemists, Chand Cairns, Mar. 4, Sampler, Brooks.	Chemist, Wuth. Sampler,	Chemist, Taylor. Sampler, Taylor.	Chemist, Wendel. Sampler,	Chemist, Britton. 31, 1872. Sampler, Brooks.	Chemist, Wendel. 6, 1873. Sampler,	

Notes.—58. From three stock piles at mine. 14. Large stock pile at Cleveland. 262. Stock pile at Cleveland. 277. All parts of mine.

CLEVELAND MINE-Specular Ore.

	36.	36.	36.	260.	260,	271.	271.	272.	272.	273.	Ave- rages.
Sesqui- or Peroxide of Iron Oxide of Manganese. Alumina Lime. Magnesia Sulphur Phosphoric Acid Silicia Acid, or Silica. Water, Total.	88.50 trace. 1.84 0.89 0.75 0.01 0.46 6.40 1.23	0.229		0.178	0.14	0.343		0.218			
Metallic Iron Phosphorus Sulphur Metallic Manganese Specific Gravity	61.95 0.20 0.01 trace. 4.64	0,100	0.187	0.076	62.10	0.147		0.093	0.154	63.730 0.111 4.93	61.092 0.119
	Chemists, Chandler & Cairns. Mar. 9, 1872.	Chemist, Wuth, Sampler,	Chemist, Britton. Sampler,	Chemist, Wendel. Sampler, Taylor.	Chemist, Taylor. Sampler,	Chemist, Wendel. Feb. 6, 1873. Sampler, Brooks.	Chemist, Britton. Dec. 31, 1872. Sampler,	Chemist, Wendel. Feb. 6, 1873. Sampler, Brooks.	Chemist, Britton. Dec. 31, 1872. Sampler,	Chemist, Britton. Dec. 31, 1872.	

Notes.—36. Large stock pile in Cleveland. 260. Stock pile in Cleveland. 271. Laurie Genth's Pit, No. 3. 272. Swede's pit. 273. School House opening. The last three were from mine.

^{*} The occurrence of the same number more than once in this line, signifies that duplicates of the same sample were sent to different chemists.

CHAMPION MINE-Magnetic and Slate Ore.

	38.	38.	38.	227.	227.	228.	228.	Ave- rages.
Protoxide of Iron. Sesqui- or Peroxide of Iron. Oxide of Manganese. Alumina. Lime. Magnesia Sulphur. Phosphoric Acid. Silicic Acid, or Silica. Water, Total.	17.87 74.93 0.05 1.15 0.52 0.92 0.12 0.28 3.70 0.52	0.116		0.021	0.337	0.161	0.316	
Metallic Iron Phosphorus. Sulphur Metallic Manganese. Specific Gravity.	66.04 0.12 0.12 0.03 4.75	0.051	4.43	0.009	57.97	0.070 4.87	66.65	63.55
	Chemists, Chandler & Cairns. Mar. 9, 1872. Sampler, Brooks.	Chemist, Wuth. Sampler,	Chemist, Britton. Sampler,	Chemist, Wuth. Sampler, Brooks.	Chemist, Wendel, Feb. 6, 1873.	Chemist, Wuth. Sampler, Brooks.	Chemist, Wendel, Feb. 6, 1873.	

Notes.—38. Large stock pile in Cleveland, all varieties. 227. "Slate ore," Shaft No. 4. 228. "Black ore," Shafts Nos. 1, 2, and 3. The two last from mine.

CASCADE MINES-Flag Ore.

	17.	22.	22.	257.	257•	258.	258.	15.	Ave-
Sesqui- or Peroxide of Iron. Protoxide of Manganese.	71.98	trace.						66.20	
Lime Magnesia Sulphur Phosphoric Acid	0.16 0.06 0.04 0.07	0.75 0.34 0.03 0.24	0.248			••••		.14	
Silicic Acid, or Silica	25.26 1.03 0.71	0.87						31.02	
Metallic Iron Phosphorus Sulphur Specific Gravity	0.03	58.59	0.108	46.120 .042 3.95	0.043			46.450	49.332
Special Galaxy	Chemist, Britton, Aug. 18, 1870. Sambler, Brooks.	Chandler & Mar. 4, 1872. Brooks, J'y,'72	Chemist, Wuth, 1872. Sampler,	Chemist, Britton, 1872.	Chemist, Allen, 1872. Sampler,	Chemist, Britton.	Chemist, Allen. Sampler,	Chemist, Britton, Aug., 1870. Sambler. Brooks.	

Notes.—17. Selected bird's-eye slate ore. Exploration pit. 22. The richest pieces from a small stock pile in Cleveland. 257. Emma mine. 258. Bagley mine; bird's-eye slate ore. The two last were obtained from the mine workings. 15. Old opening, north face ridge, S.W. corner. Sect. 29.

CASCADE MINES-Flag and Specular Ore.

	259.	259.	266.	266.	256.	256.	Ave- rages.
Phosphoric Acid			0.16	0.096			
Metallic Iron Phosphorus Specific Gravity	50.820 0.078 4.13	0.073	44.00 0.069	0.041	58.940 0.055 4.44	0.055	51.253
	Chemist, Britton. 1872. Sampler, Brooks.	Chemist, Allen. 1872. Sampler,	Chemist, Taylor. Sampler, Taylor.	Chemist, Wendel. Sampler,	Chemist, Britton. 1872. Samplers, J. Fritz & A. Pardee.	Chemist, Allen. 1872. Sampler,	

Notes.—259. Saw-Mill opening, west of stream. 256. West End Mine (specular ore). Stock pile at Mine. 266. Stock pile at Cleveland.

CANADIAN ORES-Magnetic.

	1	1	1	I	1	1
	222.	217.	216.	220, 8.	220, a.	Ave- rages.
Bisulphide of Iron. Sesqui- or Peroxide of Iron. Proto-sesquioxide of Iron. Alumina Lime. Magnesia Sulphur. Phosphoric Acid. Silicia Acid, or Silica. Water, Total (moisture) Carbonate of Lime. Carbonate of Magnesia Oxygen with the Sulphur and loss.	92.19 .68 .28 .83 .78 .14 3.55 .48	0.21		78.03 1.17 0.077 6.10 13.71 0.91	2.19 84.38 2.86 0.74 5.61 0.087 4.13	
Metallic Iron. Phosphorus	66.86	51.40 0.092	45.20 .037	99-997 54.00 .033	99.997 60.00 0.037	55·49 0.052
	Chemist, Britton. Nov. 17, 1870. Sampler, Brooks.	Chemist, E.R. Taylor. Jan. 2, 1873. Sampler, Taylor.	Chemist, Taylor. Jan. 2, 1873. Sampler, Taylor.	Chemist, Wuth. Sampler,	Chemist, Wuth. Sampler,	

Notes.—222. Analysis furnished by Redington and Adams, Cleveland. 217. Stock pile at Cleveland. 220. Analyses furnished by Dr. Wuth. a. Magnetic ore after roasting: b. Red hematite. 222 and 217 are Forsyth ore. 216 and 220 are Marmora ore.

EDWARDS MINE-Magnetic Ore.

	41.	41.	41.	199.	Averages
Protoxide of Iron Sesqui- or Peroxide of Iron Oxide of Manganese Alumina	21.60 55.80 0.10 4.34	••••		9.98 85.41	
Lime Magnesia Sulphur Sulphuri Sulphuric Acid Phosphoric Acid	0.77 0.84 0.16	0.288		.03	
Silicic Acid, or Silica. Insoluble Silicious Matter Water, Total	0.81			2.43	
M tallic Iron Phosphorus Sulphur Metallic Manganese Specific Gravity	55.75 0.05 0.16 0.06 4.24	0,125	0.137	67.45	61.60
	is, Chandler & March 9, 1872.	z, Wuth.	Chemist, Britton. Sambler,	Chemist, Taylor. Jan., 1873. Sampler, Unknown.	
	Chemists, Caims. Sampler,	Chemist, Sampler,	Chemis	Chemist, 1873. Sampler,	

Note.-41. Large stock pile in Cleveland.

EDWARDS MINE-Second-class Magnetic Ore.

	265.	265.	286.	Averages.
Phosphoric Acid	0.10	0.136		
Metallic Iron Phosphorus	48.80	0.058	49.580 0.061	49.190 .055
	Taylor. Taylor.	Wendel.	Britton. 1873. Brooks.	
	Chemist, T. Sampler, T.	Chemist, W Sampler,	Chemist, B. Jan. 9, 18. Sampler, B.	

Notes.-265. Stock pile at Cleveland. 286. From mine.

FOSTER MINE-Hematite Ore.

	49•	87.	270.	270.	26.	Averages.
Sesqui- or Peroxide of Iron. Oxide of Manganese. Alumina. Lime. Magnesia Sulphuric Acid. Phosphoric Acid Silicic Acid, or Silica. Insoluble Silicious Matter Water, Combined Water, Total. Metallic Iron. Phosphorus	74.69 .42 .50 .37 .63 .18 16.44 	20.68 6.12	0.33 51.40 0.144	0.226	79.49 0.25 1.19 0.27 0.33 0.17 0.19 9.28 8.74 99.91 55.64 0.083	52.27
Sulphur	Chemist, Brush. July 5, 1871.	Chemist, Britton. Nov. 4, 1871. Sampler, Brooks.	Chemist, Taylor.	Chemist, Wendel.	Chemist, Chandler. o May 14, 1866. o Sampler, Brooks.	1094

Notes.—49. Stock pile at Pioneer Furnace, Negaunee, Mich. 87. From mine, numerous fragments. 270. Stock pile at Cleveland. 26. From mine when first opened.

JACKSON MINE -Specular Ore.

	24.	24.	24.	51.	230.	230.	230.	Ave- rages.
Sesqui- or Peroxide of Iron Oxide of Manganese. Alumina Lime Magnesia. Sulphur. Phosphoric Acid. Silicia Acid, or Silica. Water, Total. Alumina, Lime, Magnesia, Water, etc.	93.75 trace. 0.73 0.61 0.23 0.03 0.32 3.27 1.09	0.127		0.60 0.18 0.10 1.45 1.67				
Metallic Iron. Phosphorus. Sulphur. Specific Gravity.	0.14		0.069	0.04		61.810		63.71
	Chemists, Chandler & Cairns. Mar. 9, 1872. Sampler, Brooks.	Chemist, Wuth. Sampler. Brooks.	Chemist, Britton. Dec. 20, 1872. Sampler, Brooks.	Chemist, Chandler. July 13, 1871. Sampler, Brooks.	Chemist, Wuth. Sampler, Brooks.	Chemist, Britton, Feb. 18, 1873. Sampler,	Chemist, Allen. Sampler,	

Notes.—24. Large stock pile in Cleveland. 51. Stock pile at Pioneer Furnace, Negaunee, Mich. 230. Slate ore. West end of mine.

JACKSON MINE-Hematite and Jaspery Ores.

	231.	231.	231.	229.	229.	229.	Ave-
Phosphoric Acid	0.316	0.523			0.338	0.054	
Metallic Iron Phosphorus Specific Gravity	0.138	59.30 0.224	54.530 0.154	56.590 0.061 4.59	58.20 0.144	0.023	57.155
	Chemist, Wuth. 1872. Sampler, Brooks.	Chemist, Wendel. Feb. 6, 1873.	Chemist, Britton. Feb. 18, 1873.	Chemist, Britton. Feb. 18, 1873.	Chemist, Wendel. Feb. 6, 1873. Sampler,	Chemist, Wuth. Sampler,	

Notes.—231. Hematite ore—west part of mine. 229. Old Pioneer opening—Jaspery ore. The Hematite and Specular ores occur together in this mine.

KLOMAN MINE-Specular Ore.

	235.	225.	Averages.
Metallic Iron	Chemist, Britton. + 0.89 Sampler, Brooks. 6.655	Chemist, Allen. o. Sampler,	63.55 .089

Note.—235. Fragments broken from outcrop, before work began.

LAKE SUPERIOR MINE-Specular Ore.

	37-	37•	37-	261.	261.	44.	274.	Ave-
Sesqui- or Peroxide of Iron. Oxide of Manganese. Alumina Lime Magnesia Sulphur. Phosphoric Acid. Silicic Acid, or Silica. Water Total	trace. 1.64 0.57 0.24 0.02	0.075		0.24	0.239			
Metallic Iron Phosphorus Sulphur Specific Gravity	0.06	0.033	0.046	63.50 0.103		64.37	59.89 .065 4.69	
	Chemists, Chandler & Cairns, Mar. 9, 1872. Sampler, Brooks.	Chemist, Wuth. Sampler,	Chemist, Britton. Sampler,	Chemist, Taylor. Sampler, Taylor.	Chemist, Wendel. Sampler,	Chemist, Britton. Sampler, John Fritz.	Chemist, Britton, Dec., 1872.	

Notes.—37. Large Stock pile in Cleveland. 261. Stock pile at Cleveland. 44. Stock pile at Bethlehem Furnace. 274. Lower bed. Pit No. 1. Pennsylvania mine.

LAKE SUPERIOR MINE-Hematite.

	10.	10.	10.	269.	269.	276.	276.	87.	Ave-
Sesqui- or Peroxide of Iron. Oxide of Manganese Alumina Lime. Magnesia. Sulphur Phosphoric Acid. Insoluble Silicious Matter Silicic Acid, or Silica Water, Combined. "Uncombined."	79.80 0.10 2.05 0.45 0.53 0.03 0.30 12.52 4.11 0.14	0.104		0.24			0.668	15.42	
Metallic fron Phosphorus. Sulphur Metallic Manganese. Specific Gravity	55.86 0.13 0.03 0.07 4.12	0.045	0.066	52.00	0.101	0.131	0.286		54.28 .130
	Chemists, Chandler & Cairns, March 4, 1872. Sampler, Brooks.	Chemist, Wuth. Sampler,	Chemist, Britton. Sampler,	Chemist, Taylor. Sampler, Taylor.	Chemist, Wendel. Sampler,	Chemist, Britton. Dec. 31, 1872. Sampler, Brooks.	Chemist, Wendel. Feb. 6, 1873.	Chemist, Britton. Nov. 4. 1871. Sampler, Brooks.	

Notes.—10. Large Stock pile at Cleveland. 269. Stock pile at Cleveland. 276. Hematite workings of mine.

LAKE ANGELINE-Specular Ore (Jaspery).

	21.	21.	267.	267.	34.	Averages
esqui- or Peroxide of Iron	72.00				85.43	
xide of Manganese	trace.					
lumina	0.92				1.89	
ime	0.33				0.24	
lagnesia	0.34				0.13	
ulphur	0.02				none.	
hosphoric Acid	0.08	0.101	0.04	0.083	none.	
ilicic Acid, or Silica	25.09				12.31	
Vater, Combined	1.09	i.				
" Uncombined	0.12					
	99.99				100.00	
Tetallic Iron	50.40		52.00		59.08	53.83
hosphorus	0.03	0.044	0.017	0.036	none.	.031
ulphur	0.02					
pecific Gravity	3.97					
	8 %			.:		
	r & 1872.	00	1872.	1872.		
	2 E	1872.	90	H 00		
	Chandler farch 4, 1	H			5	
	nch ch cs.		ir.	e e	0	
	ha	유	oly y	ü	45 ST	
	Chandle March 4, Brooks.	Wuth.	Taylor. Taylor.	Wendel,	Wuth. 1865. Unknown.	
	10° 0		Fed C.			
	Chemists, Cairns. Sampler,	Chemist, Sampler,	Chemist, Sampler,	Chemist, Sampler,	Chemist, Dec. 29, Sampler,	
	iri	14	in	14	oc.	
	Cale	ie.	re rm	re cm	Do	

Notes.—21. Stock pile in Cleveland. 267. Stock pile in Cleveland.

LAKE ANGELINE MINE-Hematite.

	268.	268.	280.	Averages.
Phosphoric Acid	0.09	0.160		
Metallic IronPhosphorus.	Chemist, Taylor. 51872. Sampler, Taylor. 60.00.00.00.00.00.00.00.00.00.00.00.00.0	Chemist, Wendel. Sampler,	Chemist, Britton. 9 of Dec. 31, 1872.	50.70 .079

Notes,-268, Stock pile at Cleveland. 280, Stock pile at mine.

MICHIGAMME MINE-Magnetic Ore.

	I.	197.	225.	225.	225.	Averages
Protoxide of Iron Sesqui- or Peroxide of Iron Protoxide of Manganese. Allumina. Lime Sulphur. Sulphuric Acid Phosphoric Acid. Silicic Acid, or Silica Water, Total Organic or Carbonaceous Matter. Titanic Acid. Copper and Carbonic Acid.	1.01 2.12 .12 .05 3.06 .57	29.109 61.631 traces. 2.120 1.070 0.002 0.008 0.057 3.280 1.497 0.340 0.032 none.	0.067		0.392	
Metallic Iron. Phosphorus. Sulphur Specific Gravity	0.027	99.146 65.767 0.024 0.005	0.029	0.019	63.01	64.388
	Chemist, Britton. Sept. 21, 1870. Sampler, Brooks.	Chemist, Ralph Crooker, Boston. Sampler, Ralph Crooker.	Chemist, Wuth. 1872.	Chemist, Britton. 1872. Sampler, Brooks.	Chemist, Wendel. Feb. 6, 1873.	

Notes, -1. Drill mud from 3 holes. 197. Numerous fragments from Exploration pits. 225. Taken at mine, fragments after blasting. All were taken before mine was opened.

MACOMBER MINE-Hematite.

	35•	35•	87.	Averages.
Sesqui- or Peroxide of Iron. Oxide of Manganese Alkalies (by difference). Sulphur Phosphoric Acid. Silicie Acid, or Silica Insoluble Silicious Matter. Water, Combined Water, Total.		0.130	14.51 2.23	
Metallic Iron Phosphorus Sulphur Metallic Manganese	100.00 53.76 0.06 0.14 1.51	0.057	56.08	54.92 0.058
	Chemist, Chandler. Oct. 6, 1871. Sampler, Brooks.	Chemist, Wuth. Sampler,	Chemist, Britton. Nov. 4, 1871. Sampler, Brooks.	

Notes.—35. From two trains of 16 cars each, one month apart. 87. From Mine. Numerous fragments. This mine belongs to the Negaunee hematite group, and contains considerable manganese.

MAGNETIC MINE-Magnetic Flag Ore.

	69.	54+	232.	232.	Averages
Proto-sesquioxide of Iron Oxide of Manganese Alumina. Lime Magnesia. Sulphur Phosphoric Acid Insoluble Silicious Matter Soluble Silicious Matter Undetermined and Loss. Metallic Iron. Phosphorus Sulphur Specific Gravity.	Chemist, Jenney. 1872. Sampler, Brooks. 1872. O.524. O.524. O.525. O.526. O.5279. O.5279. O.5279.	Chemist, Britton. Nov. 22, 1870. Samplest, Brooks. Samplest, Brooks. Samplest, Brooks.	Chemist, Britton. 1872. + 0.55 Sampler, Brooks. & C. 80.80	Chemist, Allen. Jan. 1, 00. 1893.	54.72 .067

Notes.—69. From small stock pile at mine. 54. From layers of rich ore banded with rock. From outcrop. 232. Small stock pile at mine.

MENOMINEE IRON REGION-Specular Ores and Hematites.

	95.	98.	102.	246.	74.	254.	68.	68.
Sesqui- or Peroxide of Iron. Oxide of Manganese Alumina. Lime. Magnesia. Sulphur Phosphoric Acid. Silicic Acid, or Silica. Insoluble Silicious Matter. Water, Total.	50.22	0.53 0.17 0.044 19.52					81.35 1.32 trace. 0.41 0.237 0.14 12.043	0.260
Metallic Iron Phosphorus Metallic Manganese Specific Gravity		54.81			56.44	44.720	56.944	0.113
	Chemist, Jenney. 1872. Sampler, Brooks.	Chemist, Wuth. Sept. 19, 1871. Sampler, Brooks.	Chemist, Chandler. Sampler, R. Pumpelly.	Chemist, Britton. Sampler, Brooks.	Chemist, Jenney. 1872. Sampler, Brooks.	Chemist, Britton. Sampler, Brooks.	Chemist, Jenney. 1872. Sampler, Brooks.	Chemist, Wuth. Sampler,

Notes.—95. Average of prevailing variety of lean ore, Sect. 31, T. 42, R. 29. 98. Average of five of the richest pieces found, S. 31, T. 42, R. 29. 102. Average of 10 analyses for P. S. and L. S. Ship Canal Co., Sect. 31, T. 42, R. 29. 246. Same as 95. 74. Boulders at west ½ post, Sect. 10. T. 39, R. 29. 68. From outcrop in swamp, Sect. 13, T. 42, R. 23. 254. Slate ore south ½ post, Sect. 30, T. 40, R. 30.

MISSOURI-IRON MOUNTAIN MINE-Specular Ore.

	127.	127.	127.	128.	128.	128.	Averages
Sesqui- or Peroxide of Iron	93.57 0.76 0.08 0.46 0.23 0.035 4.75 0.12		0.008	95.42 0.86 0.06 0.32 0.21 0.036 3.02 0.07		.012 0.067	
Metallic Iron Phosphorus. Sulphur Metallic Manganese Specific Gravity	66.049 0.016 0.12 4.944	0.043	0.049	99.996 67.416 0.016 0.07 5.002	0.025	.029 .012	66.732 .029 .010
	Chemist, Wuth. April, 1872. Sampler, Brooks.	Chemist, Allen. May, 1872. Sampler, Brooks.	Chemist, A. A. Blair. 1872. Sampler, Brooks.	Chemist, Wuth. April, 1872. Sampler, Brooks.	Chemist, Allen. Sampler,	Chemist, A. A. Blair. Sampler,	

Notes.—127. "Quarry Ore." Chippings from all parts of the pit and Stock piles. 128. "Surface Ore" (Boulders). Chippings and pebbles from all the diggings and Stock piles.

NEW YORK MINE-Specular Ore.

	20.	20.	20.	237.	237.	238.	238.	Averages.
Sesqui- or Peroxide of Iron. Oxide of Manganese. Alumina. Lime. Magnesia. Sulphur. Phosphoric Acid. Silicic Acid, or Silica. Water, Total. Metallic Iron. Phosphorus. Sulphur.	90.00 trace. 1.87 1.20 0.60 0.03 0.57 4.72 0.98 99.97 63.00 0.22				0.151		0.326	61.74 .225
Specific Gravity	Chemists, Chandler & Cairns. March 4, 1872.	Chemist, Wuth.	Chemist, Britton. Sampler,	Chemist, Britton. Sep-	Chemist, Allen. Sampler,	Chemist, Britton. March, 1873.	Chemist, Allen. Sampler,	

Notes.—20. Large Stock pile at Cleveland—all varieties. 237. Great South Opening—Pit No. 1. 238. Beardsley's Pit—No. 2. The two last from mine.

NEW ENGLAND MINE-Soft Hematite.

	87.	239.	Averages
Sulphur Insoluble Silicious Matter. Water Combined. Volatile Matter (a little organic and water).	25.66	None. 23.30 2.69	
Metallic Iron Phosphorus Salphur Metallic Manganese Specific Gravity		46.84 0.08 none. 0.18 3.79	48.24
	Chemist, Britton. Nov. 4, 1871. Brooks.	Britton. De- 27, 1872. Brooks.	

Notes.—87. From mine, numerous fragments. 239. From cars and stock pile at mine. First-class specular ore was formerly mined here, but is not at present.

NEGAUNEE HEMATITES-Manganiferous Soft Hematite.

	243.	243.	243.	11.	II.	108.	116.	Ave-
Sesqui- or Peroxide of Iron Oxide of Manganese. Alumina Lime Magnesia. Sulphur Phosphoric Acid. Silicic Acid, or Silica Water Combined " Uncombined				6.71 1.46 0.45 0.66 0.04	0.171	65.48 1.54 29.25		
Metallic Iron Phosphorus Sulphur Metallic Manganese Specific Gravity	0.067		0.099	45.78	0.074		50.58	.07
	Chemist, Jenney. 1872.	Chemist, Britton. 1872. Sampler,	Chemist, Allen. 1872.	Chemists, Chandler & Cairns. Mar. 4, 1872.	Chemist, Wuth. 1872. Sampler,	Chemist, Chandler. June 29, 1872. Sampler, Brooks.	Chemist, Jenney. July 13, 1872. Sampler, Brooks.	

Notes.—243. From exploration pits. 11. Small stock pile at Cleveland. 108. Average of three analyses of ore from exploration pits. 116. Dark brown chalky ore. All from Sects. 6, 7, and 8, T. 47, R. 26.

NEW YORK STATE ORES (ST. LAWRENCE & WAYNE CO.)-Hematites.

	203.	206.	205.	204.	215.	209.
Protoxide of Iron. Sesqui- or Peroxide of Iron. Protoxide of Manganese. Alumina Lime. Magnesia. Sulphur. Phosphoric Acid Silicic Acid, or Silica., Insoluble Silicious Matter. Water, Total. Carbonic Acid	75.30 0.15 1.69 7.04 0.38 0.03 trace. 10.12	77.24 0.45 1.60 0.23 0.05 trace. 12.93 2.107	12.49 56.54 trace. 0.69 8.23 2.13 none. 0.36 4.28	12.72 57.93 0.07 4.54 2.32 0.85 0.07 0.16 10.97	63.31 6.03	1.49
Metallic Iron	100.12	94.607	99•73	100.00	44.31	41.80
Phosphorus			0.16	0.07	0.43	0.64
	Chemists, Maynard & Wendel. Mar., 1871. Sampler, Geo. W. Maynard.	Chemist, Jenney. 1872. Sampler, Brooks. Jan. 1871.	Chemists, Maynard & Wendel. Sampler, Geo. W. Maynard.	Chemists, Maynard & Wendel. Sampler, Geo. W. Maynard.	Chemist, Chandler. Sampler, Unknown.	Chemist, Taylor. Jan. 2, 1873. Sampler, Geo. R. Tut-tle.

Notes,—203. Sampled for John A. Griswold & Co., at mine. 204. Do. do. 205. Do. do. 206. From small stock pile at Cleveland. 215. Analysis furnished by H. B. Tuttle. 203 and 206 are from Keene Mine, 204 and 205 are from the Caledonia Mine, both owned by Rossie Iron Works. 209 and 215 are Wayne Co. ore.

NEW YORK (LAKE CHAMPLAIN REGION)-Magnetic.

					1
	288.	289.	290.	291.	292.
Protoxide of Iron Sesqui- or Peroxide of Iron. Protoxide of Manganese. Alumina Lime. Magnesia	26.69 59.84 •55 1.87	25.35 56.19 0.12 3.56 0.82	23.29 50.13 0.38 4.22 1.28 0.85	8.87 69.99 0.38 3.67 1.90 traces.	19.05 42.97 3.47 1.19 0.09
Sulphur. Phosphoric Acid Silicic Acid, or Silica. Water, Total Carbonate of Lime.	.20	trace. 12.34 0.47	trace. 20.02	0.24 0.07 14.60	trace. 32.94
Metallic Iron	100.56 62.61	98.85	100.17 53.21	99·72 55·91	99.71
	Chemist, Geo. W. Maynard. Sampler, Geo. W. Maynard.	Chemist, Geo. W. Maynard. Sampler, Geo. W. Maynard.	Chemist, Geo. W. Maynard. Sampler, Geo. W. Maynard.	Chemist, Geo. W. Maynard. Sampler, Geo. W. Maynard.	Chemist, Geo. W. Maynard. Sampler, Geo. W. Maynard.

Notes.—288. Wetherby, Sherman & Co., and Port Huron Iron Ore Co., No. 21. 289. New Bed; Wetherby, Sherman & Co. 290. Hammond, Crown Point. 291. Indian; Ferrona ore; Hassey, Wells & Co. 292. Fisher; Port Henry Iron Ore Co.

OHIO IRON ORES-Black Band and Kidney.

11.5	293.	294.	295.	296.	297.	298.	Ave- rages.
Protoxide of Iron Sesqui- or Peroxide of Iron Oxide of Manganese. Alumina Lime. Magnesia Sulphur Phosphoric Acid. Silicic Acid, or Silica. Water, Combined. Water, Total. Volatile Matter Carbonic Acid. Lime Phosphate. Lime Carbonate.	30.50	23.02 8.79 1.70 0.70 0.88 0.11 0.492 26.22 21.10	75.00 1.65 0.60 2.80 1.48 trace. 0.773 17.02	7.60 1.35 2.60 1 carbo. 1 6.50 0.18 0.863 8.96	75.00 1.85 0.60 5.94 3.64 0.12 1.26 8.46 2.28	12.34 1.70 0.50 (carbo. { 5.33 } trace. 11.94 .78	
Metallic Iron Phosphorus Specific Gravity	99.60	99.712 24.06 0.216 2.321	52.50	99.573 36.31 0.379 3.434	99.15 52.50 0.554 4.076	99.15 35.88 0.797 2.539	
	Chemist, T. G. Wormley.	Chemist, T. G. Wormley.	Chemist, T. G. Wormley.	Chemist, T. G. Wormley. Sampler,	Chemist. Sampler.	Chemist. Sampler.	

Notes.—293. Black Band, Mineral Ridge, Mahoning Co., O. 294. Black Band, Tuscarawas Coal and Iron Co., Tuscarawas Co., O. Raw. 295. Black Band, Tuscarawas Coal and Iron Co., Tuscarawas Co., O. Calcined. 296. "Shell" or "Kidney Ore," Tuscarawas Coal and Iron Co., Tuscarawas Co., O. Raw. 297. "Shell" or "Kidney Ore," Tuscarawas Coal and Iron Co., Tuscarawas Co., O. Calcined. 298. Nodular Ore, Washingtonville Co., Columbiana Co., O. For further analyses of Ohio Iron Ores, consult Geological Survey of Ohio, 1870, pp. 47, 48, 49, 219, 223.

REPUBLIC MINE-Specular and Magnetic.

	233.	233.	234.	234.	Ave-
Metallic Iron Phosphorus Specific Gravity	Chemist, Britton. 1872. G. O. Sampler, Brooks.		Chemist, Britton. 1872. 9 0 0 0 0 Sampler, Brooks.	Chemist, Allen. 1872.	

Notes.—233. Specular ore, First stock pile at opening of mine. 234. Magnetic ore. First stock pile at opening of mine.

SAGINAW MINE-Specular and Hematite.

	281.	282.	Averages.
Metallic Iron Phosphorus	Chemist, Britton. Jan. o.	Chemist, Britton. Jan. 9 & S9 1873.	52.40 .132

Notes,—281. Small stock pile (first mined) at mine. 282. Ditto. Both samples are soft hematite. By oversight no sample of the specular ore, which is first-class, was collected.

SHENANGO MINE-hematite.

	242.	242.	78.	Averages.
Sesqui- or Peroxide of Iron. Oxide of Manganese. Alumina Lime Magnesia. Phosphoric Acid. Silicic Acid, or Silica Water, Combined. Metallic Iron. Phosphorus, Specific Gravity.	55.140	Chemist, Allen. 1872. 9	Chemist, Wuth. 1872. Samplers, Davock, Davock, 10, 10, 10, 10, 10, 10, 10, 10, 10, 10	56.315 .070

Notes,—242. From small stock pile at mine. 78. "Taken from under snow, with no possible selection."—Letter from Davock, Glidden & Co.

SILAS C. SMITH MINE-Hematite.

	70.	70.	87.	Averages.
Sesqui- or Peroxide of Iron	71.70			
Oxide of Manganese	0.10			
Alkalies	2.03			
Sulphur	0.27			
Phosphoric Acid	0.09	0.127		
Silicic Acid, or Silica.	23.38			
Insoluble Silicious Matter			23.79	
Water, Total	2.43		2.43	
	100.00			
Metallic Iron	50.19		49.21	49.70
Phosphorus	0.04	0.055		.047
Sulphur	0.27			
	Chemist, Chandler. Sept. 7, 1871. Sampler, Brooks.	Chemist, Wuth. Sampler,	Chemist, Britton. Nov. 4, 1871. Sampler, Brooks.	

Notes.—70. From small stock pile at mine when first opened. 87. From mine when first opened, numerous fragments.

SPURR MOUNTAIN MINE-Magnetic Ore.

	1					
	2.	226.	226.	226.	97•	Averages.
Proto-sesquioxide of Iron. Oxide of Manganese. Alumina Lime. Magnesia Sulphur. Phosphoric Acid. Silicic Acid, or Silica	89.21 traces. 2.67 .67 0.19 0.35 trace. 6.28	0.259			92.36 0.15 1.66 0.73 0.75 0.221 4.31	
Metallic Iron Phosphorus Sulphur Specific Gravity	99·37 64.60 0.35	0.113 4.62	0.112	59.96	66.87 0.096	63.81
	Chemist, Chandler. Nov. 1868. Sampler, Brooks.	Chemist, Wuth. 1872. Sampler, Brooks.	Chemist, Britton. 1872. Sampler,	Chemist, Wendel. Feb. 6, 1873.	Chemist, Wuth. Sept. 14, 1872. Samplers, Morgan & Herne.	

Notes.—2. Numerous fragments broken from outcrops of ore. 226. Fragments broken from outcrop 97. Numerous fragments broken from outcrop. All before mine was opened.

TAYLOR MINE-L'ANSE RANGE-Hematite.

	88.	81.	89.	89.	87.	Averages.
Sesqui- or Peroxide of Iron. Oxide of Manganese. Alumina Lime. Magnesia. Sulphur Phosphoric Acid. Silicic Acid, or Silica Insoluble Silicious Matter. Water, Combined "Total.	82.664 0.894 0.312 0.226 0.090 0.236 6.180 9.438	10.75	62.25 1.87 3.028 0.31 0.26 trace. 0.31 21.30		5.29 8.70	
Metallic Iron. Phosphorus. Sulphur Metallic Manganese.	57.86 0.102 0.090 0.62	52.00	43.576 0.13	44.78	57.51	52.88
	Chemists, Chandler & Schweitzer. Sampler, Brooks,	Chemist, Britton. 1871. Sampler, Brooks.	Chemist, Jenney. 1872. Sampler, Brooks.	Chemist, Britton, Feb. 18, 1873.	Chemist, Britton. Sampler, Brooks.	

Notes.—88. From shaft 20 feet in ore. 81. From three trenches across ore deposit. 89. From all pits, shafts and trenches showing ore. 87. From mine, numerous fragments. All before mine was opened.

WASHINGTON MINE-Magnetic Ore.

	39•	39.	39•	264.	264.	284.	285.	Ave-
Proto-sesquioxide of Iron. Oxide of Manganese. Alumina Lime. Magnesia. Sulphur Phosphoric Acid. Silicic Acid, or Silica. Water, Total.	0.23 0.85 0.92 0.77 0.03 0.25	0.406		0.21	0.170			
Metallic Iron. Phosphorus Sulphur Metallic Manganese. Specific Gravity.	99.90 65.94 0.11 0.03 0.14 4.66	0.177	0.149	67.20			54.800 0.146	
	Chemists, Chandler & Cairns. March, 1872. Sampler, Brooks.	Chemist, Wuth. 1872. Sampler,	Chemist, Britton, 1872. Sampler,	Chemist, Taylor. 1872. Sampler, Taylor.	Chemist, Wendel. 1872. Sampler,	Chemist, Britton. Jan. 9, 1873.	Chemist, Britton. Jan. 9, 1873. Sampler, Brooks.	

Notes.—39. Large stock pile in Cleveland. 264. Stock pile at Cleveland. 284. Shafts Nos. 1 and 4 at mine, 285. Shafts Nos. 2 and 6 at mine,

WINTHROP MINE-Soft Hematite.

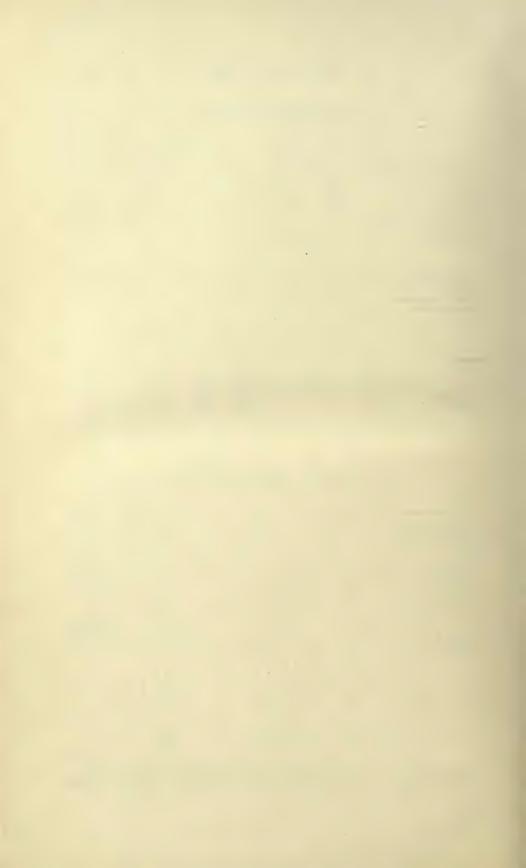
	240.	240.	287.	Averages.
Sesqui- or Peroxide of Iron			84.66 1.41 0.13 0.40 0.007 0.02 0.084 12.70	-
Metallic Iron Phosphorus Sulphur Metallic Manganese Specific Gravity	50.00 0.03 none. 0.53 4.03	0.045	59.26 0.037	54.63 .037
	Chemist, Britton. Dec. 26, 1872. Sampler, Brooks.	Chemist, Allen. Sampler,	Chemist, Fred. H. Emmerton. Feb. 13, 1873.	

Notes.-240. From all parts of mine. 287. Stock pile at a Chicago furnace.

WISCONSIN IRON ORES-Iron Ridge (Hematite).

	298.	Averages.
Pure Metallic Iron	56.44	
Oxygen with the Iron	24.91	
rotoxide of Manganese	2.30	
luminaime	2.28	
agnesia	.99	
ulphuric Acid	.15	
hosphoric Acid	1.10	
nsoluble Silicious Matter	3.57	
Vater and Carb. Acid.	6.30	
	99.26	
Metallic Iron	56.44	
Phosphorus.	.48	
ulphur	.06	
	1	
	n.	
	Britton. Hager-	
	Bri.	
	I	
	F	
	5.	
	hemist. nampler	
	Chemisi Sampler man.	
	m m	

Note.—298. The ore is a fossil ore, in grains about the size and shape of flax-seed; Dr. Topham calls it Oolitic ore. There are several old analyses showing no phosphorus, but they are not reliable.



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ADDENDA.

A

Extracted from the Marquette Mining Journal.

The following table exhibits in gross and net tons the amount of iron-ore shipped from the Lake Superior mines during the season of 1873, together with its total value at \$7 per ton, gross:

ackson New York Lleveland Lake Superior Champion	113,892 70,882 132,082 166,666 72,782	740 600 820	\$797,246 31 496,175 87
New York	70,882 132,082 166,666	600 820	496,175 87
New York	70,882 132,082 166,666	600 820	496,175 87
Cleveland	132,082 166,666	820	
Lake Superior	166,666		
Champion			924,576 56
	12,102	600 1,210	1,166,663 87 509,477 78
Vashington	38,014	1,530	266,102 78
Washington	105,452	2,160	738,170 75
Cloman	21,065	590	147,456 84
ascade	20,507	760	143,551 37
Barnum	48,076	1,260	336,535 93
oster	27,372	440	191,605 38
Rowland	1,404	1,390	9,832 34
ake Angeline	43,933	1,360	307,535 25
Pittsburgh & Lake Superior	21,498	1,400	150,490 37
Edwards	31,730	350	222,111 09
Spurr	31,933	1,560	223,535 87
Aichigamme	28,966	1,230	202,765 84
Michigan (Clarksburgh)	3,212		22,484 00
Keystone	10,426	700	72,984 18
AcComber	38,969	1,630	272,788 09
Iimrod	2,074	390	14,519 21
Marquette	2,148	210	15,036 61 234,826 29
Winthrop	33,546	1,370	
Shenango	8,658	910	60,608 85 8,320 29
Albion	1,188	1,370	11,591 59
Carr		1,100	8,935 44
Bagaley	1,239	1,500	8,677 69
Green Bay		1,890	6,648 91
Emma		1,740	49,964 44
Rolling Mill	11,319	740	79,235 31
Saginaw	37,139	60	259,973 18
Smith		1,860	65.301 81
Grand Central		1,150	46,406 59
Gribben	4,517	1,350	31,623 22
New England.,		410	1,268 28
Allen			3,570 00
Goodrich	3,258	830	22,808 59
Home	1,090	2,000	7,636 25
Magnetic	78	1,560	550 87
fron Mountain		1,510	788 72
Hungerford	145	440	1,016 37
Total	1,163,057	160	\$8,141,398 98

The following table shows the aggregate production of the several furnaces in the district for 1873, together with the value of the iron (\$45) at furnace:

•		
FURNACES.	Gross tons.	Value.
Pioneer Collins. Michigan Greenwood Bancroft Morgan Champion Deer Lake Fayette Bay Munising Grace Beecher Beecher (muck bar) Lake Superior Company's peat furnace Escanaba (shipped to November 19th)	7,098 2,000 4,467 4,416 4,100 6,324 3,949 3,447 10,696 8,760 2,237 7,800 710 428 500 2,175 2,400	\$319,410 90,000 201,015 198,720 184,500 284,580 177,705 155,115 481,320 394,200 100,665 351,000 31,950 25,680 22,500 97,875 108,000
Total	71,507	\$3,224,235

 \mathbf{B}

It has not been possible to obtain trustworthy statistics of the products of the Missouri mines and furnaces, but the following figures are given as an estimate:

Iron-Ore. Iron Mountain district. .350,000 Central Missouri .75,000 Pig-Iron (Coke and Coal). Vulcan Iron-Works. 16,300 South St. Louis Iron-Works. 9,300 Missouri Furnace Co. 12,000 Carondelet furnace 5,000

42,600

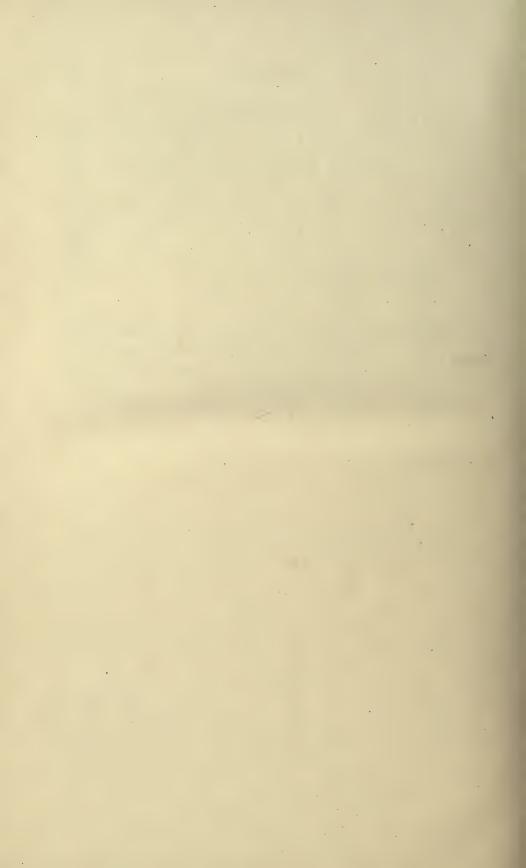
Pig-Iron (Charcoal).

Meramee Iron	n-Works	4,300	
Moselle		5,100	
		7,600	
		* *	
O		700	
Hamilton	66	700	
Iron Mountain furnaces			
Pilot Knob	46	5,000	
		33,400	

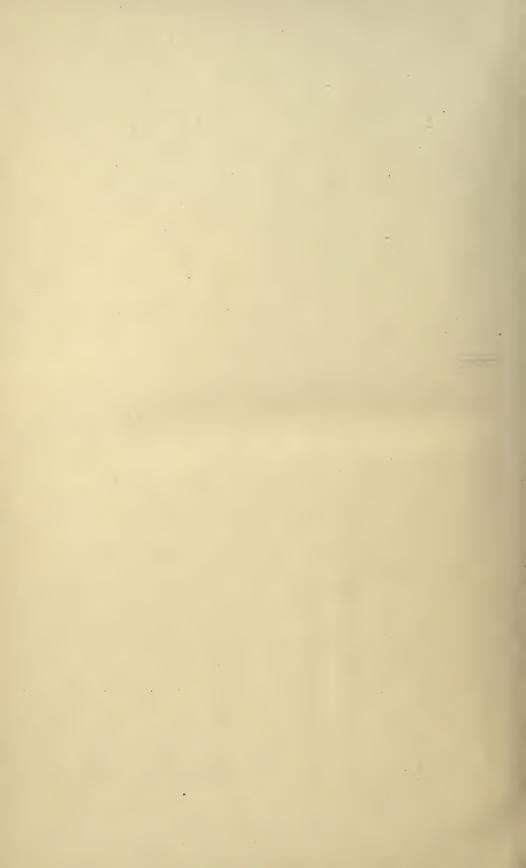
Total pig-iron, 75,000 gross tons.

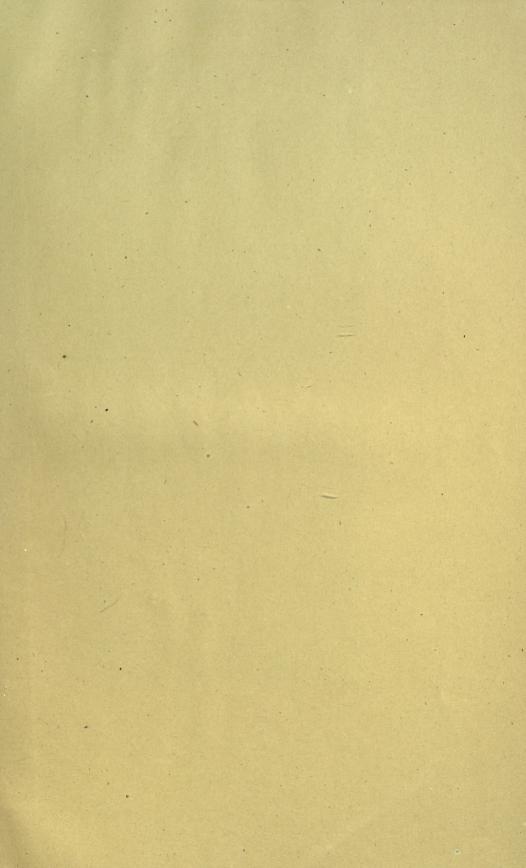
C

Explorations made during the summer of 1873 in the Menomonee district of Lake Superior have proved the existence of workable beds of ore of fine quality in T. 39, R. 28; T. 39, R. 29; T. 40, R. 30; T. 42, R. 28; T. 42, R. 29.









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